Technical note: Fourier approach for estimating the thermal attributes of streams

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10   **Abstract.** Temperature models that directly predict ecologically important thermal attributes across spatio-temporal scales are still poorly developed. This study developed an analytical method to estimate seasonal and diel periodicities as well as irregularities in stream temperature at data-poor sites based on Fourier analysis. We first quantified the thermal attributes of a glacier-fed stream in the Swiss Alps using 2-years of hourly-recorded temperature. Stream temperature was accurately decomposed to an average 3.8 °C, diel periodicity spanning 4.9 °C, seasonal periodicity spanning 7.5 °C, and an irregularity
15   having an average of 0.0 °C but spanning 9.7 °C. These thermal attributes then were used to estimate thermal attributes at spot-measured sites along the river, resulting in a different relative contribution (weighting) of attributes among sites. The results confirm that the developed method can infer stochastic behaviors in stream thermal attributes at spot-measured sites. Additional ways to further improve the methodological approach are discussed.

1 Introduction

20   Temperature is a fundamental determinant of physical and biogeochemical patterns and processes in ecosystems. Organisms respond to temperature via different adaptations and distinct life history strategies across various spatial and temporal dimensions (Cossins and Bowler, 1987). More broadly, global biodiversity has tracked external temperature transitions over millennia (Mayhew et al., 2012). In the Anthropocene, unprecedented rapid transitions in thermal regimes owing to climate change are a global concern (IPCC, 2014). Consequently, adequate characterization of thermal regimes is an important
25   prerequisite for understanding the role of temperature in ecosystems and for ecological conservation.

Periodicity and irregular extremes in temperature, as representations of variability, have attracted a growing interest in recent climate-relevant ecological studies (e.g., Thompson et al., 2013). The combination of seasonal and diel periodicity cycles in temperature promotes diverse behaviors and spatial distributions of exothermic organisms according to respective life history strategies and life cycle stages (e.g., Vannote et al., 1980). In contrast, irregular extremes such as heatwaves can
30   induce physiological exposure and vulnerability (Paaijmans et al., 2013) as well as cause abrupt shifts in biogeochemical

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processes (Frank et al., 2015) and ecological assemblages (Thompson et al., 2013). Hence, the variability in thermal periodicities and irregular extremes need to be distinguished from average temperatures.

In freshwaters, ecological responses to thermal regime shifts may be less understood than in marine and terrestrial ecosystems (Thompson et al., 2013), even though freshwater biodiversity has experienced major decreases in recent decades (WWF, 2014). Progress in understanding response patterns have been delayed partially because the quantification of thermal attributes is difficult for running waters. Physical disturbance in streams can be considerable due to periodic flow pulses (Poff et al., 1997). Particularly where flashy flows mobilize gravel riverbeds, e.g. braided rivers, the installation of data-loggers for long-term time series monitoring can be logistically difficult. For natural floodplains acting as biodiversity hotspots (Tockner and Stanford, 2002), frequent time-series temperature data (daily or sub-daily) recorded throughout a year or longer are difficult to acquire because the floodplain mosaic changes frequently and dramatically (Van der Nat et al., 2003). Often in these cases, researchers rely on spot-measures of temperature at study sites and thus lack time-series temperature, thereby limiting understanding of the ecological consequences of thermal attributes in freshwaters. Clearly, an estimate of the thermal attributes at spot-measured sites would benefit this understanding.

Given this background, there is still much room towards improving statistical models to predict stream temperature (see review in Benyahya et al., 2007). For instance, regression models employ correlative relationships with air temperature (e.g., Pilgrim et al., 1998) and streamflow (Webb et al., 2003). Autoregressive models take into account the autocorrelation structure within water temperature time-series data and also the correlation with external variables (Kothandaraman, 1971; Cluis, 1972; Long, 1972). To deal with seasonally changing parameters in regression and autoregressive models, periodic autoregressive models were introduced (Benyahya et al., 2007). However, temperature patterns at multiple temporal scales, particularly the combination of seasonal and diel periodicity patterns, are still rarely considered (Steel and Lange 2007). Moreover, temperature models that directly predict ecologically important thermal attributes (seasonal and diel periodicity patterns and irregular extremes) do not exist presently, as many models focus solely on temperature numeric values at a given time.

The "well-worn" Fourier analysis (Fourier, 1878) is well-suited for analyzing combined multi-temporal patterns and predicting thermal attributes. Unfortunately, the application of Fourier analysis for assessing stream temperature patterns has slowed since its early emergence in 1970 (Kothandaraman, 1971; Cluis, 1972; Long, 1972; but see Maheu et al., 2015). This trend is surprising when considering its high potential for detecting periodicity patterns at multiple scales in time-series data, water temperature in particular. Therefore, we felt it timely to further investigate the methodological application of Fourier analysis for assessing thermal patterns in running waters with the intent to reintroduce its use in ecological field studies.

Our primary aim was to develop an analytical method for estimating average, seasonal and diel periodicities as well as irregular extremes of stream temperature at data-poor sites (i.e., spot measures) based on Fourier analysis. As a case study, we first quantified these thermal attributes using 2-years of hourly-recorded time-series temperature data in an alpine glacial-fed stream in Val Roseg, Switzerland (see Ward and Uehlinger (2003) for a synthesis of research conducted in this catchment). Using those results, we then estimated thermal patterns at sites on the same stream where monthly spot-measures of temperature were taken during the same 2-years.

2 Methods

2.1 Compositional variables in stream temperature

We assume that hourly stream temperature $T(t)$ at a given time t is generally composed of the average value $\bar{T}$, seasonal periodicity pattern $S(t)$, and diel periodicity pattern $D(t)$. The remaining component, which is unexplained by these three

5 components, results from multiple external factors such as weather conditions (sunshine, rain, wind, etc.): we call this component an irregularity $\varepsilon(t)$. A distinctive high/low value in the irregularity indicates thermal extremes that strongly disturb the periodicity patterns in temperature. Thereby, hourly stream temperature is expressed as:

$$T(t) = \bar{T} + S(t) + D(t) + \varepsilon(t) \tag{1}$$

2.2 Fourier analysis for temperature decomposition

Fourier transformation converts a function of time $T(t)$ into a function of frequency $G(f)$ by transforming time-series data into

10 a combination of trigonometric curves. Both functions are identically reversible:

$$G(f) = \int_{-\infty}^{\infty} T(t)e^{-i2\pi ft}dt \tag{2}$$

$$T(t) = \int_{-\infty}^{\infty} G(f)e^{i2\pi ft}df \tag{3}$$

As measured stream temperature is a discrete variable, we applied the fast Fourier transformation algorithm. The algorithm searches for a single solution to identically explain a time-series variable $T(t)$ by summing the trigonometric curves ($e^{i2\pi ft}$) of different frequency f amplified by a corresponding spectral intensity $G(f)$:

$$G(f) = \sum_{t=0}^{N-1} T(t)e^{-i2\pi ft} \tag{4}$$

$$T(t) = \frac{1}{N}\sum_{t=0}^{N-1} G(f)e^{i2\pi ft} \tag{5}$$

where N is the data length equivalent to the number of summed curves. As seen in Eq. (5), the time-series variable, hourly

15 stream temperature $T(t)$, is formulated with summed trigonometric curves (frequency $f = 0, 1, 2, …, N–1$) whose vibration magnitudes are amplified by a corresponding spectral intensity $G(f)$. A strong spectral intensity indicates that the corresponding curve dominates the composed time-series variable. As a consequence, stable patterns of seasonal/diel periodicities in time-series temperature data are detected by trigonometric curves with remarkably strong spectral intensities.

The mismatch between Eq. (1) and Eq. (5) needs to be solved to express time-series stream temperature as in Eq. (1). Although

20 our intention is to explain time-series temperature by the four components above, Fourier analysis decomposes it into N components based on a series of trigonometric curves. This mismatch in the number of components requires summing some of the trigonometric curves to best represent seasonal and diel periodicity patterns. However, it is unknown which curves at what frequencies and spectral intensities are required to sufficiently express these periodicity patterns. For selecting curves to

be summed, we considered the condition that the seasonal and diel periodicities consist of their period lengths (inverse of frequency, $1/f$) within the ranges of 1–365 days and 1–24 hours, respectively. Moreover, we introduced a threshold value in spectral intensity for selecting only the dominant components as well as to avoid mixing noises with the periodicity patterns. In so doing, the shape of functions $S(t)$ and $D(t)$ are dependent on the spectral threshold value. We set the threshold value at

5   0.1 °C, a minimal unit of temperature measurement in the case study system (see sect. 2.4). The analysis was performed using the fast Fourier transform (*fft*) function of the "stats" library in R 3.1.2 (R Core Team, 2014).

2.3 Extrapolation to spot-measured temperature data

Assuming that thermal attributes auto-correlate along the river continuum, we can extrapolate the thermal attributes based on time-series data to spot-measured data at sites along the same stream. For simplicity, we do not include external information

10   (e.g., discharge and air temperature). After decomposing the time-series temperature T_0 as Eq. (1) at a reference site, the decomposed factors are used to estimate temperature T_a at spot-measured site A along the same stream network. We assume that the four temperature components in Eq. (1) are linearly correlated between sites when other factors (e.g., major tributary or groundwater inputs) affecting stream temperature along the network are low enough to maintain the spatial-autocorrelation temperature patterns between sites. The temperature at site A is formulated as:

$$T_a(t) = \bar{T}_a + S_a(t) + D_a(t) + \varepsilon_a(t) \tag{6}$$
$$= \beta_1 \bar{T}_0 + \beta_2 S_0(t) + \beta_3 D_0(t) + \beta_4 \varepsilon_0(t) + \beta_5 \tag{7}$$

15   where coefficients β_{1-4} are the weighting parameters for each component (> 0), and β_5 is a parameter to adjust systematic bias. The following procedures are performed to estimate each parameter. First, linear regression between T_a and T_0 corresponding to temperatures at the measurement time of T_a is conducted to estimate β_1 (slope) and β_5 (intercept). Second, β_2 and β_3 are numerically estimated by minimizing the mean square error value based on the linear regression between an estimated $\beta_4 \varepsilon_0 ($ $= T_a - \{\beta_1 \bar{T}_0 + \beta_2 S_0(t) + \beta_3 D_0(t) + \beta_5\})$ and ε_a at the corresponding measurement times. Third, β_4 is numerically estimated

20   similarly as for the second step based on T_a and the estimated T_a.

To highlight the benefits of the extrapolation method, we compared the component extrapolation approach to a linear regression model that simply extrapolates time-series temperature based on the linear regression between T_a and T_0. Importantly, if the coefficients β_{1-4} have the same value as the slope of the linear regression and β_5 is the intercept, the Fourier approach is equivalent to the linear regression model.

2.4 The case study: an alpine glacier-fed stream (Val Roseg, Switzerland)

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The method was applied in the Roseg catchment, an alpine valley located in the Bernina massif of the Swiss Alps (Fig. 1). The catchment area is 66.5 km^2 and ca. 30% glaciated (Swiss National Hydrological and Geological Survey (OFEV); year of record 2010). Elevations range from 1766 to 4049 m a.s.l. The Roseg River is fed by meltwaters of the Tschierva and Roseg glaciers. The Roseg glacial runoff first drains into the pro-glacial lake Roseg before merging with the flow from the Tschierva glacier

30   (Fig. 1). Mean annual discharge at the end of the catchment was 2.8 m^3 s^{-1} (discharge record averaged for 1955–2013) and

daily discharge ranged from 3.3 to 19.6 m^3 s^{-1} in July and August and from 0.2 to 2.0 m^3 s^{-1} between November and March 2013 (OFEV). The study system comprises (i) a long proglacial reach below the Tschierva glacier, exhibiting extremely low temperatures due to glacial runoff (kryal), (ii) a single-thread channel downstream of the confluence of the proglacial reach and the Roseg lake outlet, (iii) a complex braided floodplain, and (iv) a constrained reach extending to the end of the catchment where the Pontresina hydrological station is located (see Tockner et al., 1997; Uehlinger et al., 2003 for a detailed description). Hourly time-series temperature was recorded at a reference site, the Pontresina hydrological station (site R in Fig. 1: 46°29'23.6"N, 09°53'53.3"E; 1766 m a.s.l.) in 2012 and 2013 (Fig. 2a: provided by OFEV). Spot-measured water temperature was taken monthly at two sites, one located within the proglacial reach (site A: 46°24'38.3"N, 09°51'31.2"E; 2106 m a.s.l.) and one below the lake outlet confluence reach (site B: 46°25'05.9"N, 09°51'27.1"E; 2054 m a.s.l.) (Fig. 1). For both sites, stream temperature was measured using a conductivity meter (WTW LF323, Weilheim, Germany) at different daily times on each visit from April to October 2012 and 2013 (in total 14 times).

3 Results

3.1 Thermal attributes at the reference site

By converting the time-series temperature data at the reference site R (Fig. 2a) to its frequency domain, some frequency ranges had strong spectral pulses (Fig. 3). Nine trigonometric curves exceeded the threshold of 0.1°C (Table 1). Regarding period length (Table 1), the curves were allocated to seasonal and diel components: three curves for seasonal periodicity patterns (Fig. 2b); six curves for diel periodicity patterns (Fig 2c). Thereby, the time-series temperature at the reference site R (Fig. 2a) was decomposed into an average of 3.8 °C, a seasonal cycle spanning 7.5 °C (Fig. 2b), a diel cycle spanning 4.9 °C (Fig. 2c), and an irregularity spanning 9.7 °C with an average of 0.0 °C and standard deviation of 0.92 °C (Fig. 2d). Hourly stream temperatures excluding the irregularity (i.e., $\bar{T} + S(t) + D(t)$; Fig. 2e) explained 92% of the data variability (r^2), indicating a successful decomposition of the time-series data and a high reliability of the approach to characterize stream temperature components. The irregularity had a normal distribution (no inferred bias), indicating that the seasonal and diel periodicity patterns were extracted accurately from the original time-series data.

3.2 Estimating time-series temperature for spot-measured stream sites

Temperature at site A was approximately twofold lower than at site B during the study period. Temperature at site A had an average of 2.3 °C and spanned 0.4–4.8 °C; temperature at site B had an average of 5.4 °C and spanned 0.8–9.2°C. Temperatures at site A and B were linearly correlated to temperature at the reference site R (r^2 = 0.60 and 0.92, respectively; Fig. 4). This result indicates that spatial autocorrelation of the thermal attributes exists between them, even though a thermal influence from the pro-glacial lake differentiates the thermal patterns between sites A and B (Fig. 1). For Eq. (7), the proposed method estimated that the weighting parameters at site A (β_{1-4} = {0.32, 0.32, 0.50, 0.70}) showed higher heterogeneity in the relative contributions of the components than those at site B (β_{1-4} = {0.97, 0.97, 0.99, 0.93}). For

site A, the relative contribution of diel variability and irregularity to the estimated temperature ($\beta_3 = 0.50$, $\beta_4 = 0.70$) was higher than the average and seasonal variability ($\beta_1 = \beta_2 = 0.32$). At site B, the parameter composition was less variable (0.93–0.99), and therefore the performance of the decomposition approach was similar to the performance of the linear model. Using the developed method, the temperature at site A was estimated better than the estimated temperature using linear regression ($r^2 =$

5   0.66 and 0.60, respectively). For site B, both approaches performed equivalently ($r^2 = 0.92$). These differences between sites A and B indicate that the thermal attribute at site A was more different from the reference site R than that at site B.

Based on the obtained parameters, time-series temperatures at sites A and B were inferred (Fig. 5). At site A, due to relatively high contributions of diel cycles and irregularities ($\beta_3 = 0.50$, $\beta_4 = 0.70$), higher hourly variability was estimated by the proposed approach than the linear regression (Fig. 5a). The maximum values estimated by the proposed approach and linear regression

10   were clearly distinct at 5.9 °C and 4.0 °C, respectively, while an average of 1.1 °C was equivalent for both estimates. The linear model clearly underestimated (< 4.0°C) the high temperatures recorded at 4.4 °C on 30 May 2012 and 4.8 °C on 5 June 2013. In contrast, the proposed approach inferred a possibility of temperature reaching 5.9 °C, having a 97.5% percentile value of 4.0 °C. This difference between approaches indicates the ability of the proposed approach to estimate extreme thermal pulses and their occurrence probability.

15 **4 Discussion**

Currently, no stream temperature models explicitly predict ecologically important thermal attributes (seasonal and diel periodicity patterns and irregular extremes) because of the difficulty in capturing the combined patterns at multiple temporal scales in stream environments (Benyahya et al., 2007). This study developed a Fourier transformation approach that predicts these thermal attributes at data-poor sites based on the pre-analysis of time-series temperature data at a data-rich reference site

20   along the stream. The approach could estimate the relative composition of thermal attributes from a limited number of spot-measured data (see Eq. 7), while linear regression weighted the composition equally. The results emphasized the significance to develop further ecological-based thermal prediction models, aiming at deeper understandings in ecological responses to thermal attributes (Paaijmans et al., 2013; Thompson et al., 2013; Frank et al., 2015).

The developed prediction method confirmed its potential to evaluate the relative contribution of thermal attributes at a data-

25   poor site. The method is somewhat comparable with multiple linear regressions ($Y = \beta_0 + \beta_1 X_1 + \ldots$) in terms of the assumption that the thermal attributes are independent of each other and auto-correlated between sites, and their relative contributions change according to location. The unique difference is that the developed method can directly use the periodicity patterns and irregularities for prediction as Eq. (7). This type of model was not addressed in a recent review on temperature models (Benyahya et al., 2007).

30   The advantage of Fourier decomposition analysis at multiple temporal scales is the ability to evaluate the possible range maxima for each thermal attribute of time-series data (Fig. 2). Fourier analysis clearly detected strong periodicity patterns in time-series temperature data (Fig. 3) and represented them with a small number of trigonometric curves (Table 1). By doing

so, we found that the irregularity had broader range (9.7 °C) than the spanning ranges of seasonal and diel periodicity patterns (7.5 °C and 4.9 °C, respectively) at the study stream. This high irregularity may be an important thermal attribute in glacier-fed streams. Further, our results clearly showed that seasonal trends existed in the amplitudes of diel periodicities and irregularities, both having lower variability in winter and greater variability in spring (Figs. 2c-d) (Hopkins, 1971). Importantly,

5 these characteristics are difficult to detect without coupling the periodicity patterns at multiple temporal scales. By decomposing the time-series temperature data at the reference site as a pre-analysis for prediction, this decomposition method (see sect. 2.1 and 2.2) can be useful for quantifying thermal attributes in ecological studies.

The proposed method is highly promising towards evaluating potential changes in thermal variability due to climate change or anthropogenic thermal effluents in rivers and streams (Caissie, 2006; Webb et al., 2008). For instance, environmental change

10 could modify each component of stream temperature with a different degree of severity such as increasing the average, shifting the peak timing in seasonality, reducing diel variability, or amplifying the irregularity. In addition, as thermal regimes play a key role in aquatic ecosystem structure and functioning (e.g., environmental niches determining community composition, biogeochemical cycles, Vannote et al., 1980), each component of stream temperature can be used to better understand ecosystem dynamics and responses (Thompson et al., 2013).

15 While demonstrating the broad applicability of this approach, we caution over some of its limitations. First, the analysis connecting Eqs. (1) and (7) is threshold dependent. In addition to 0.1 °C, we tried other values including 0.05 and 0.2 °C but these could not clearly separate periodicity patterns and irregularities (results not shown). Therefore, a threshold value must be carefully chosen and needs to be evaluated whether the irregularity attribute is unbiased. Second, although the concept is feasible, complete validation was limited due to the low number of sample sites. To obtain more robust estimation, the model

20 needs additional validation as well as calibration. Third, if a systematic shift in thermal pattern is observed during a target period, the method would need adjustment. Our target 2-year period was short enough to neglect long-term trends in temperature. For example, some studies using long-term records (>20 years) detected an inter-annual increasing trend in mean temperature due to anthropogenic thermal discharge and land-use change (Beschta and Taylor, 1988; Hostetler, 1991). In such cases, the proposed equation in Eq. (1) must be modified because it assumes constant average temperature over a target period

25 for analysis; i.e., $\bar{T}$ should be a function of time as $\bar{T}(t)$. Last, the assumption of linearly-related components between the reference site and other sites is not always met in all cases. For instance, glacier-fed and groundwater-fed streams can be composed of different seasonal and diel periodicity patterns (Brown and Hannah, 2008). In our target system, flow from the Roseg pro-glacial lake differentiated thermal patterns of sites A and B (Fig. 1). Even though our analysis confirmed that the linear assumption was apparent in this system ($r^2 = 0.60$; Fig. 4), incorporation of the thermal pattern at the lake outlet into the

30 analysis would certainly increase estimation accuracy.

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Table 1: Frequency *f*, period 1/*f*, and spectral intensity *G(f)* of trigonometric curves (threshold of *G(f)* >0.1) composing the thermal

attributes: average $\overline{T}$, seasonal periodicity pattern *S(t)*, and diel periodicity pattern *D(t)*.

| Attribute	*f*	1/*f*		*G(f)*
	[Hz]	[hrs]	[days]	[°C]
$\overline{T}$	0	–		3.78
S(t)	2	8760	365	2.01
	4	4380	183	0.11
	6	2920	122	0.19
D(t)	728	24.1	1	0.27
	730	24	1	0.59
	732	23.9	1	0.28
	734	23.9	1	0.10
	1458	12.0	0.5	0.12
	1460	12	0.5	0.21
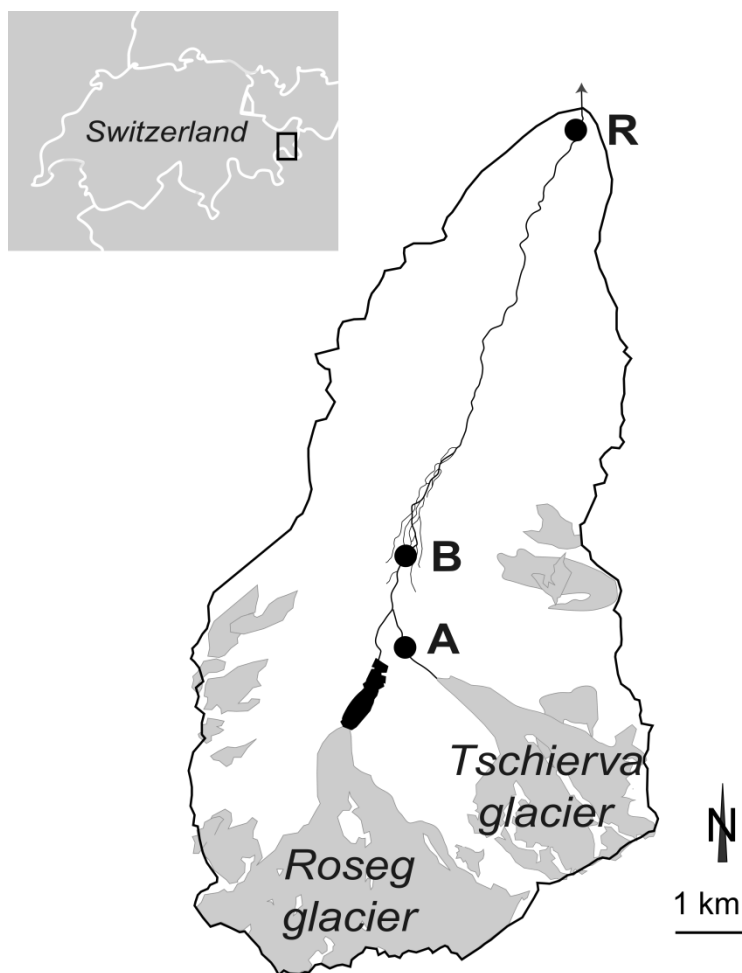
Figure 1: Location of the Val Roseg catchment in Switzerland. Study and reference (i.e., Pontresina hydrological station) sites are indicated by black dots and labeled A, B and R, respectively. Glaciers are shown as shaded areas with the pro-glacial lake shown below the Roseg glacier (modified from Uehlinger et al. 2003).
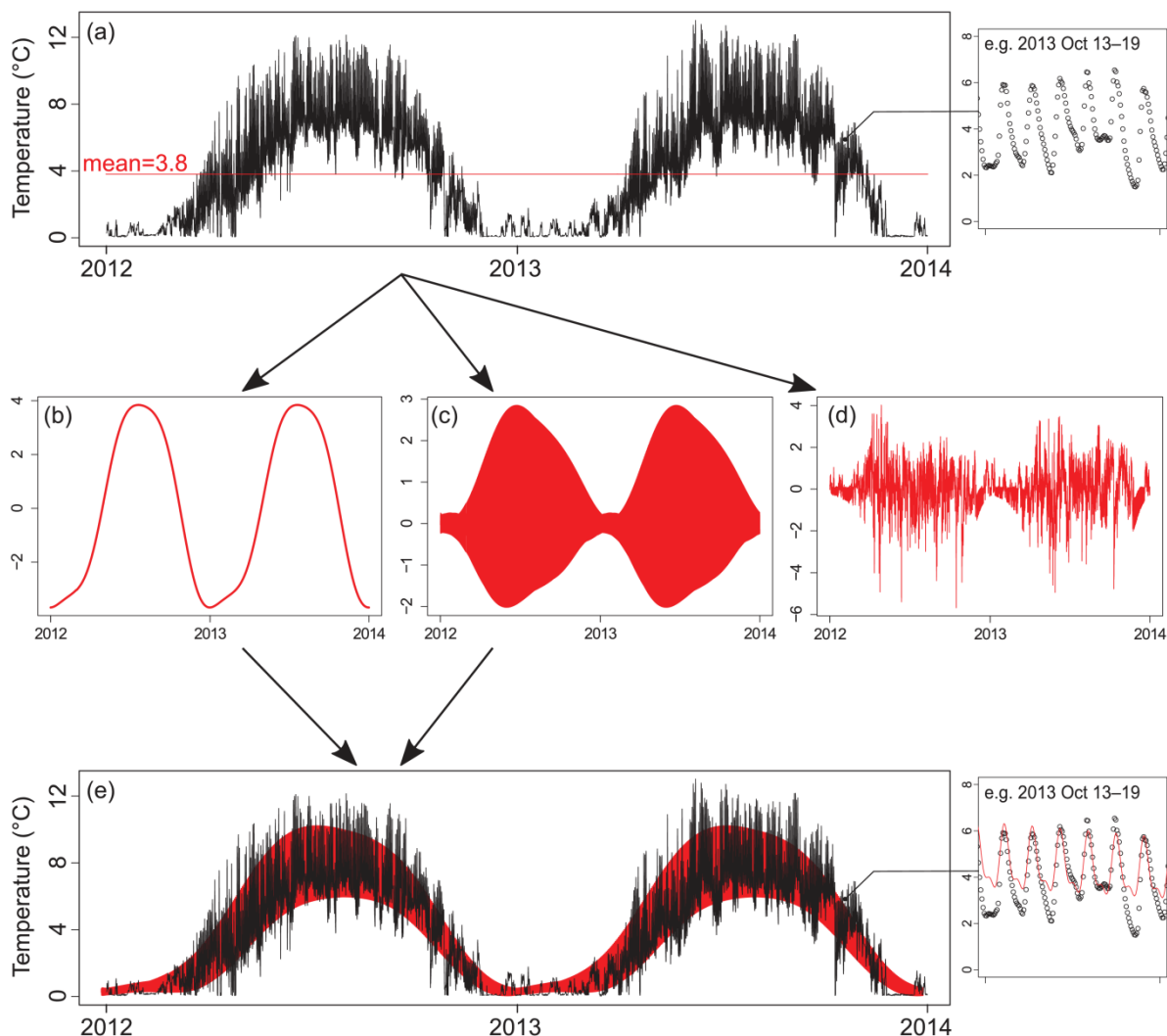
Figure 2: Hourly time-series temperature at the reference site decomposed using Fourier decomposition analysis. (a) The observed record T_a in 2012–2013 is decomposed into an average value (a), seasonal periodicity pattern S_a (b), diel periodicity pattern D_a (c), and the irregularity ε_a (d). The comparison between the observed and measured temperature excluding ε_a at the reference site (e).
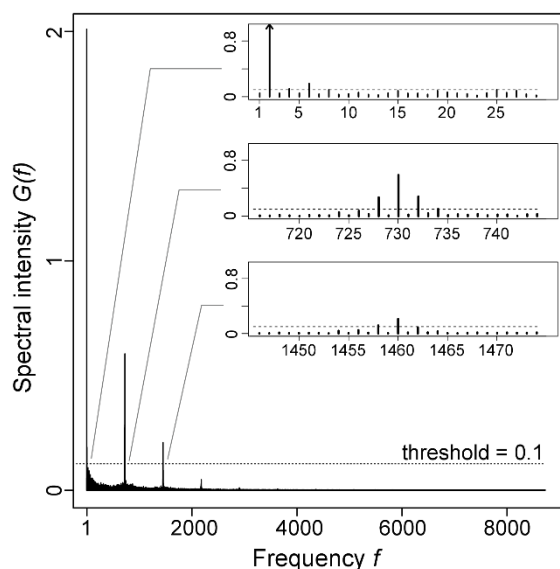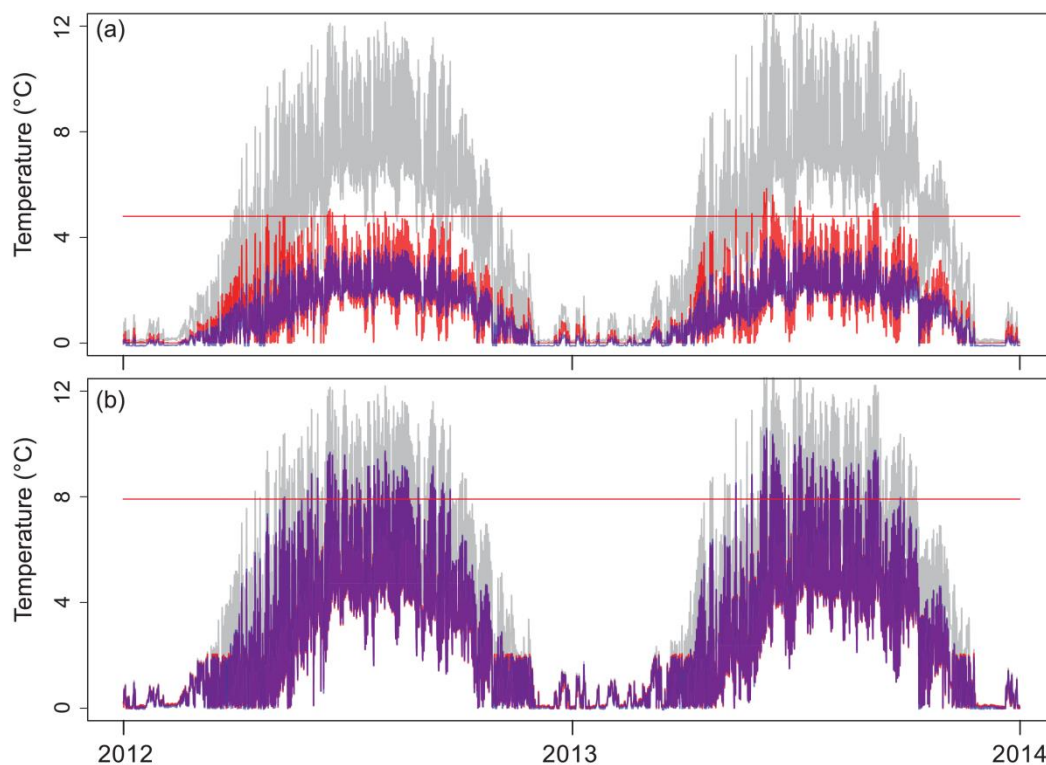
Figure 3: Frequency and spectral intensity of trigonometric curves converted from the 2-year temperature data at the reference site. Trigonometric curves whose spectral intensity was higher than the threshold value are found in Table 1.

5
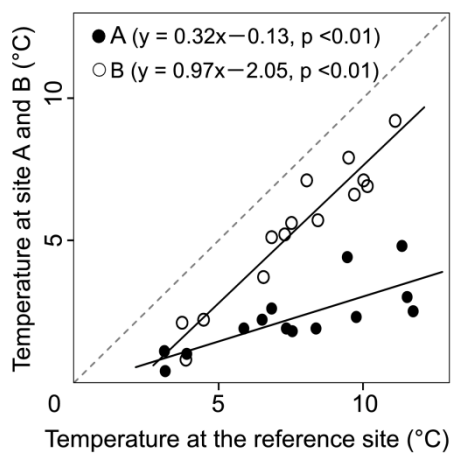
10   **Figure 4:** Linear regression of stream temperature at site A and B with the reference site temperature R.

Figure 5: Hourly stream temperatures estimated using Fourier decomposition (red) and linear regression (purple) at site A and B (plots a and b, respectively) and temperature at the reference site (grey). Horizontal red line represents the maximum temperature recorded from the monthly spot measures (4.8 °C on 5 June 2013 at site A, and 7.9 °C on 9 July 2013 at site B).