

Interactive comment on “Is groundwater sufficient to support sustainable irrigation agriculture in a reclaimed wetland region?” by Z. Pang et al.

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Thank you, Dr. Currell, for your comments concerning our manuscript entitled “Is groundwater sufficient to support sustainable irrigation agriculture in a reclaimed wetland region?” (MS No. hess-2016-155). Your comments are very valuable and helpful for improving our manuscript. We have numbered your comments for clarity. Our responses are described one by one in the following:

Comment 1: Limitations of the methods used to arrive at estimates of groundwater age and categorisation of water into ‘modern’ and ‘pre-modern’ need to be acknowledged and discussed in much more detail.

Reply 1: You are right. Due to the fact that tritium in precipitation in most parts of the world has returned to natural background, it is becoming more and more difficult to use

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it as an age dating tool for groundwater. In this study, we have used the 3H content versus 14C activity plot to determine the initial 14C activity. It gives a good indication of the initial 14C activity. The 3H contents of groundwater in the confined aquifers are very low (<2.2TU) and the 14C ages are thousands of years, which indicates that the groundwater in the confined aquifers contains no modern component, referred to as pre-modern water. However, groundwater in the unconfined aquifers shows a wide range of 3H values. Groundwater samples near the river contain higher levels of 3H (6.5-71.3TU), indicating the existence of modern water component.

Comment 2: The links between groundwater and surface water also need to be more clearly demonstrated with reference to the data and figures (such as maps and spatially referenced comparisons between surface water and groundwater levels).

Reply 2: There are strong evidences to support the existence of linkages between groundwater and surface water. Isotopic composition of groundwater in the unconfined aquifers is more enriched than that of groundwater in the confined aquifers and groundwater samples in the unconfined aquifers are plotted on the local evaporation line (Fig. 7 and Fig. 8). High NO₃⁻ and Ca²⁺ concentrations can be found in the shallow unconfined groundwater (Districts I and III) (Fig. 3 and Fig. 4), indicating the interactions with irrigation water, while those in the confined area are generally low. Groundwater in the unconfined aquifers shows a wide range of 3H values. Groundwater samples near the river contain high levels of 3H (6.5-71.3TU) indicating their linkage with river water. In a word, the groundwater in the unconfined aquifer has strong links with the surface water, while such links in the confined aquifers are very weak, if any, because there is a thick clay layer on top of it that separates them.

Comment 3: The link between groundwater age and recharge mechanism, and groundwater sustainability is not explained clearly enough. Are the authors proposing that low recharge rates and a lack of tritium indicate ‘pre-modern’ water in the confined aquifer, and thus that there is a limit to the sustainable extraction rate from this system? If so, this should be carefully explained and the potential for ‘capture’ of water from other

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areas (and release of water from aquitards) explored.

Reply 3: The low 3H values in the confined area indicate that the groundwater is pre-modern and high 14C ages show that the recharge rate is very low. The fast decline of water table is also an indication that the aquifers get very limited recharge and potential other sources, e.g. release of aquitards, is also limited. It is therefore justifiable to conclude that groundwater alone is not sufficient to support sustainable irrigation agriculture. Integrated use of groundwater and surface water is a better solution to sustain irrigation agriculture in such regions.

Comment 4: There may be water quality implications for high rates of groundwater extraction also, as documented in Currell et al. *Journal of Hydrology* 385 pp 216-225.

Reply 4: We agree with you. As shown in Fig. 8, the water table in the confined aquifers shows a trend of gradually declining while that in the unconfined aquifers remains relatively stable. Some samples with high NO₃⁻ and Ca²⁺ concentrations (especially NO₃⁻) in the confined area may indicate leakage from the unconfined groundwater similar to that reported in Currell et al., 2010. If the pumping in the confined area continues at or increases from the present levels, the groundwater table declining will continue and the water quality will deteriorate, which is indeed unsustainable.

Comment 5: With regard to the unconfined aquifer, it appears that there is active recharge, on the basis of high nitrate and tritium concentrations observed in shallow groundwater. Is this attributed to recharge through irrigation return leakage, rainfall recharge, surface water leakage, or some combination of the three? Is groundwater quality a limiting factor for the utilisation of the unconfined aquifer groundwater (e.g. because of the high nitrate concentrations)? These issues should all be clearly explained with reference to the data and more detailed discussion of the trends observed in different parts of the study area. I think some further figures such as maps showing the distribution of tritium and perhaps nitrate in the aquifers will be illustrative of the areas where recharge is actively occurring.

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Reply 5: Yes, there is a need to add more discussion. Based on the enriched 18O and 2H isotopic compositions, high NO₃⁻ and Ca²⁺ concentrations (especially NO₃⁻) (Fig. 3 and Fig. 4) and high 3H values near the river in the unconfined aquifers, we believe that the combination of irrigation return leakage, rainfall and surface water leakage have all contributed to the active recharge in the unconfined aquifers. On the basis of high NO₃⁻ concentrations, groundwater quality is indeed a limiting factor in the long term for the utilization of the unconfined groundwater. Furthermore, groundwater quality will also be a limiting factor for the sustainable utilization of it in the confined area as shallow groundwater leakage will occur in the future if the pumping in the confined area continues at or increases from the present levels.

Comment 6: The editor has noted that the authors need to provide background on the purpose of each analysis and more detail about the stable isotope evaporation model and tritium decay model. These areas have been addressed to some degree. However, I question how robust the use of the Ottawa tritium in precipitation record is for the study area, given there have been only 4 samples collected at the local IAEA station for comparison.

Reply 6: Atmospheric circulation in the stratosphere is controlled by latitude, resulting in the zonal distribution of tritium ("latitude effect" in the manuscript) in precipitation (Eriksson, E.: An account of the major pulses of tritium and their effects in the atmosphere, *Tellus*, 17, 118-130, doi: 10.1111/j.2153-3490.1965.tb00201.x, 1965; Clark and Fritz, 1997). On the other hand, tritium levels recorded near the study area, at the Qiqihar GNIP station (latitude 47°23'0"), show a similar trend to that in Ottawa, Canada (latitude 45°23'0"), during an overlapping period, and the two stations are at similar latitude. So we infer that the precipitation tritium record of Ottawa can be used to approximate the input function of tritium in precipitation in the Sanjiang Plain.

Comment 7: The 'latitude effect' is not clearly explained; do you mean that because the two stations are at similar latitude we can infer the tritium records are expected to be approximately the same? Some explanation and one or more references for this

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assumption is needed here.

Reply 7: See "Reply 6".

Comment 8: It should also be made clear that the age estimation using tritium is only a semi-quantitative tool (as the 3H-He method is not adopted).

Reply 8: We agree with you. The age estimation using tritium is only a semi-quantitative tool. We wanted to determine if there is modern water component in the groundwater of confined or unconfined aquifers. It seems sufficiently accurate to serve our purposes in this study.

Comment 9: The abstract needs some more reference to the data and more context. e.g. Is groundwater quality the limiting factor for groundwater utilisation from the unconfined aquifer? If so, then what particular aspects of water quality are important? What is the link between groundwater age and recharge mechanism, and sustainability of groundwater usage? For example, groundwater extraction from the confined aquifers will induce flow and leakage from other areas, is the quality of the induced flow a potential limiting factor (as in other areas in China)?

Reply 9: Yes, we can add more context to the abstract. Although active recharge occurs in the unconfined aquifers and the water table is relatively stable, the water quality with relatively high NO₃⁻ concentrations is a limiting factor for sustainable groundwater utilization. With weak interactions with surface water and lateral flow as the main recharge source, the groundwater age is older in the confined aquifers. The continuously declining water table demonstrates that groundwater is not sufficient for sustainable irrigation agriculture in terms of water quantity, though it is not severe at present. Furthermore, groundwater quality will also be a limiting factor for the sustainable utilization of groundwater in the confined aquifers as shallow groundwater leakage will occur in the future if the pumping in the confined aquifers continues at the present or increased levels.

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Comment 10: Methods Further information is needed on the sample collection methods for groundwater and surface water. Are the groundwater samples from production wells, or monitoring wells? What is the range of sample depths and screened intervals? For surface water, were the samples 'grab samples'? If so, at what time of year were they taken? This may impact whether the samples represent recent runoff, snow melt and/or water impacted by evaporative enrichment.

Reply 10: We will add the information about water samples collection. The groundwater samples are from production wells. The sample depths are presented in Table 1 and groundwater samples are the mixing of water within the well depth intervals. The surface water samples S1, S2, S3 were taken at 2011 from rivers along the transect A-A' in Fig. 1. The other surface water samples were taken at 2009 from paddy fields, drainage channels and rivers within farms HH, QF and QS in the northeast part of the Sanjiang Plain (Fig. 1).

Comment 11: The LMWL should be calculated using a weighted regression method, as described in Hughes and Crawford, Journal of Hydrology 464-465 pp 344-351 (2012), rather than simple linear regression.

Reply 11: We have calculated using the PWLSR method as described in Hughes and Crawford, Journal of Hydrology, 464-465, 344-351 (2012), and the LMWL is $y=(7.39\pm 0.14)x-0.88$ ($R^2=0.99$), which is very similar to " $y=7.51x-0.92$ " in the manuscript. We will replace it.

Comment 12: Results The relationship between lithology and ion composition (e.g. Ca) should be discussed and examined in more detail. Are carbonate minerals in the soil and/or aquifer the likely source of Ca? Is fertiliser a potential source also? A plot of the Ca vs ¹³C isotopes would be helpful in this context. You may also consider including and discussing the full dataset on water major ion chemistry, and discuss TDS distribution in the aquifers.

Reply 12: We think a plot of ¹³C vs ¹⁴C is a good indication of carbonate minerals dis-

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solution. As shown in Fig. 10, there is no significant negative correlation relationship between ^{13}C and ^{14}C , indicating carbonate minerals dissolution is not the main source of Ca in the aquifers. However, Ca contents of the soil in Sanjiang Plain are high. The unconfined aquifer is influenced by vertical infiltration, and the interaction with the calcium-rich soil leads to the high calcium concentrations in the shallow groundwater. In the confined aquifer, lateral flow dominates the groundwater recharge. Lack of carbonate in the confined aquifers results in the relatively low Ca contents. Fertilizer is not a potential source of Ca as no calcium fertilizer has been used.

Comment 13: Plotting tritium and carbon-14 data vs sample depth would be useful, and also plotting tritium concentrations on a map. This would allow better assessment of where spatially the recent and 'pre-modern' water samples are distributed with respect to current agricultural irrigation areas, and it will help to better identify areas of 'active recharge' as distinct from those not receiving such recharge.

Reply 13: We presented ^3H and ^{14}C distributions in different areas with distinct hydrogeological conditions. Groundwater at QS farm (in unconfined area) shows a wide range of tritium levels (<1.0 - 71.3TU), corresponding to the different sampling locations. Samples with high levels of tritium (6.5 - 71.3TU) are from shallow groundwater collected near rivers, and those collected away from the river are from deeper groundwater showing low levels of tritium ($<1.0\text{TU}$). On the other hand, groundwater samples in the unconfined area generally have enriched ^2H and ^{18}O isotopic compositions and high NO_3^- and Ca^{2+} concentrations. These indicate that groundwater in the unconfined area have strong links with surface water and active recharge occurred. However, in the confined area (Districts II), low tritium levels and NO_3^- and Ca^{2+} concentrations, depleted ^2H and ^{18}O isotopic compositions and high ^{14}C age indicate groundwater has weak interactions with surface water without modern water recharge.

Comment 14: -When discussing 'vertical infiltration' as a recharge mechanism (e.g. p. 10) you should distinguish between recharge due to rainfall infiltration and/or irrigation return-flow, and recharge from surface water bodies such as rivers.

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Reply 14: As discussed in the manuscript, the samples near the rivers have high ^3H values indicating recharge from rivers. Those samples with high NO_3^- concentrations and enriched ^2H and ^{18}O isotopic compositions indicate vertical infiltration from rainfall and irrigation water.

Comment 15: -The use of the tritium/radiocarbon plot to estimate initial activities of ^{14}C has some merit, however it should be conducted more rigorously, explained in more detail, and used with some caution. Are you using a linear extrapolation between 'modern' and tritium free water in the various samples to arrive at the initial pMC of approximately 80? What about the influence of mixing between 'young' water and older water (which should produce a straight line relationship, as opposed to a decay-based curve)? Decay and mixing will produce different patterns in ^3H and ^{14}C and this needs to be carefully analysed. For further detail refer to Cartwright et al, Journal of Hydrology 380 pp. 203-221 (2010), particularly Figure 8. The use of this method does not discount the need to assess other potential sources of DIC and influences on initial ^{14}C activities. A plot of the ^{13}C vs ^{14}C data is needed, as is some further analysis of the ion chemistry (e.g. Ca vs ^{13}C) to shore up this area.

Reply 15: We agree with you. As shown in Fig. 10, there is a decay-based curve relationship rather than straight line relationship between ^{14}C and ^3H , which is similar to that shown in Cartwright et al, 2010. So the mixing between modern water and older water is precluded. To arrive at the initial ^{14}C activity, the decay-based curve relationship was used rather than straight line relationship between ^{14}C and ^3H . Based on the decay-based curve relationship between ^{14}C and ^3H , when ^3H value is below the ^3H detection limit, or in other words, the water is Tritium-free, the corresponding ^{14}C activity can be considered to be the initial ^{14}C value. Carbonate minerals dissolution should be considered when using ^{14}C age. We think a plot of ^{13}C vs ^{14}C is a good indication of carbonate minerals dissolution. As shown in Fig. 10, there is no significant negative correlation relationship between ^{13}C and ^{14}C , indicating carbonate minerals dissolution does not dominate in the aquifers.

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Comment 16: -As indicated above, the link between groundwater age, recharge and ground-water sustainability is not explained clearly enough. You need to put more work into defining (on your maps) where groundwater is influenced by direct vertical recharge, river recharge and lateral recharge, and discuss the water quality implications of these different mechanisms. Where in particular do you think the extraction rates for groundwater are likely to be much greater than recharge? What is the likely response of the aquifer(s) to extraction and is there any water level data to show what is happening currently? What are the likely water quality implications of extraction from different aquifers and depths (see previous comments)?

Reply 16: As is discussed in the manuscript, in the unconfined aquifers (Districts I and III), the samples near the rivers contain high $3H$ values, indicating recharge from rivers. Those samples with high NO_3^- concentrations and enriched $2H$ and $18O$ isotopic compositions indicate vertical infiltration from rainfall and irrigation water. Although active recharge occurred in the unconfined area and the groundwater table is relatively stable (Fig. 8), the water quality with high NO_3^- concentrations is a limiting factor for sustainable groundwater utilization. In the confined area (Districts II), with weak interactions with surface water and lateral flow as the main recharge source, the groundwater age is older. The continuous declining groundwater table (Fig. 8) demonstrated that groundwater is not sufficient for sustainable irrigation agriculture in terms of water quantity. Furthermore, groundwater quality will also be a limiting factor for the sustainable utilization of groundwater in the confined area as shallow groundwater leakage will occur in the future if the pumping in the confined area continues at the present or increased levels.

Comment 17: -Overall the discussion is too brief, and further discussion of limitations of your isotopic data, and alternative explanations need to be explored and discounted.

Reply 17: We will rewrite the discussion section with more details discussed above.

Comment 18: p2 Line 10 'Recharge and regime', do you mean 'recharge and ground-

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water flow patterns'?

Reply 18: We mean the groundwater recharge and the change of groundwater table.

Comment 19: Lines 11 & 12: Grammar is poor. Do you mean 'with ages over 50y is recharged by lateral flow.as evidenced by depleted heavy isotopes'? Which isotopes (I assume $18O$ and $2H$)?

Reply 19: We will rewrite the sentence as follows: Groundwater in the confined Quaternary aquifer with ages over 50 years and evidenced by depleted $18O$ and $2H$ isotopic compositions is recharged by lateral flow from nearby mountains.

Comment 20: Line 27: Citation (Assessment, 2005) is incorrect. A suggested citation format is given in the front matter of this report.

Reply 20: We will rewrite it as follow: Millennium Ecosystem Assessment: Ecosystems and human well-being: wetlands and water, World Resources Institute, Washington, DC, 2005. The citation in the text is "(Millennium Ecosystem Assessment, 2005)".

Comment 21: P4. Line 22: Suggest using ML rather than mega-L.

Reply 21: We will use ML instead of mega L.

Comment 22: P5. Line 2: 'hydrogeology' should be 'hydrogeological'.

Reply 22: We will use hydrogeological instead of hydrogeology.

Comment 23: P6. Line 19-20: Can remove the statement 'our current efforts...tracers'. It is better to clearly outline your study aims and scope in the introduction section.

Reply 23: We agree with you and we will remove the statement from the "Sampling and analyses" section.

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