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Technical Note: Multiple wavelet coherence for untangling scale-specific and localized multivariate relationships in geosciences

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Abstract

The scale-specific and localized bivariate relationships in geosciences can be revealed using simple wavelet coherence. The objective of this study is to develop a multiple wavelet coherence method for examining scale-specific and localized multivariate relationships. Stationary and non-stationary artificial datasets, generated with the response variable as the summation of five predictor variables (cosine waves) with different scales, were used to test the new method. Comparisons were also conducted using existing multivariate methods including multiple spectral coherence and multivariate empirical mode decomposition (MEMD). Results show that multiple spectral coherence is unable to identify localized multivariate relationships and underestimates the scale-specific multivariate relationships for non-stationary processes. The MEMD method was able to separate all variables into components at

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the same set of scales, revealing scale-specific relationships when combined with multiple correlation coefficients, but has the same weakness as multiple spectral coherence. However, multiple wavelet coherences are able to identify scale-specific and localized multivariate relationships, as they are close to 1 at multiple scales and locations corresponding to those of predictor variables. Therefore, multiple wavelet coherence outperforms other common multivariate methods. Multiple wavelet coherence was applied to a real dataset and revealed the optimal combination of factors for explaining temporal variation of free water evaporation at Changwu site in China at multiple scale-location domains. Matlab codes for multiple wavelet coherence are developed and provided in the supplement.

1. Introduction

Geoscience data such as topography, climate, and ocean waves usually present cyclic patterns, with high-frequency (small-scale) processes being superimposed on low-frequency (large-scale) processes (Si, 2008). More often than not, geoscience data is non-stationary, consisting of a variety of frequency regimes that may be localized in space or time (Torrence and Compo, 1998; Si and Zeleke, 2005; Graf et al., 2014). The wavelet method is a common tool for detecting multi-scale and localized features of non-stationary processes in geosciences. Simple wavelet coherency has been widely used for untangling scale-specific and localized relationships for non-stationary processes in areas including geophysics (Lakshmi et al., 2004; Müller et al., 2008), hydrology (Labat et al., 2005; Das and Mohanty, 2008;

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Tang and Piechota, 2009; Carey et al., 2013; Graf et al., 2014), soil science (Si and Zeleke, 2005; Biswas and Si, 2011), meteorology (Torrence and Compo, 1998), and ecology (Polansky et al., 2010). This method, however, is limited to two variables. Processes in geosciences are usually complex and may be affected by more than two environmental factors. A method is needed for analyzing multivariate (>2 variables) and localized relationships at multiple scales.

Several methods have been used for characterizing multivariate relationships. For example, multiple spectral coherence (MSC) has been used to explore the scale-specific relationships between soil saturated hydraulic conductivity (K_s) and multiple soil physical properties (Koopmans, 1974; Si, 2008), but requires a stationary data series which is rare in geosciences. Multivariate empirical mode decomposition (MEMD), a data-driven method, decomposes each variable into different components (intrinsic mode functions (IMFs)) with each IMF corresponding to a "common scale" inherent in multiple variables (Rehman and Mandic, 2010). The MEMD method is meritorious due to its ability to deal with both non-stationary and nonlinear systems. The combination of squared multiple correlation coefficient and MEMD (MCC_{memd}) has been used to explore the multivariate control of soil water content or saturated hydraulic conductivity at multiple scales (Hu and Si, 2013; She et al., 2013, 2015; Hu et al., 2014). However, the sum of variances from different components typically does not equal the total variance of the original series, which may result in misleading MCC_{memd} results. In addition, in geosciences, multivariate relationships are most likely to change with time or space due to non-stationarity of the processes involved.

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However, localized multivariate relationships are not available using any of the existing multivariate methods. Therefore, it is required to extend the wavelet coherence from two variables to multiple variables.

An attempt to extend wavelet coherence from two to three variables has been made by Mihanović et al. (2009). Their method was also applied later in the marine sciences (Ng and Chan, 2012a, b). Limitations arise when using three variable wavelet coherence: first, only two predictor variables are considered; second, the two predictor variables must be orthogonal. Otherwise, extremely high or low (spurious) coherence (>>1 or <0) may be produced. This spuriousness is inconsistent with the definition of coherence and may limit the application of this method in geosciences, where environmental variables are usually cross-correlated. Therefore, a robust method for calculating MWC, which produces coherence in the closed interval of [0, 1], is needed.

The objective of this paper is to develop an MWC that applies to cases where there are multiple environmental variables of which may be cross-correlated. This method is first tested with artificial datasets to demonstrate its advantages over existing multivariate methods. It is then applied to a temporal series of evaporation (E) from free water surface and meteorological factors at Changwu site in Shaanxi, China.

2. Theory

Simple wavelet coherence can be understood as the traditional correlation coefficient localized in the scale-location domain (Grinsted et al., 2004). Just as

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correlation coefficients can be extensions from two variables to multiple (>2) variables, wavelet coherence between two variables may also be extended to multiple variables. Similar to simple wavelet coherence, MWC is based on a series of auto- and cross-wavelet power spectra at different scales and spatial (or temporal) locations for the response variable and all predictor variables.

Following Koopman (1974), a matrix representation of the smoothed auto- and cross-wavelet power spectra for multiple predictor variables X ($X=\{X1, X2, \ldots, Xq\}$) can be written as

$$\overleftrightarrow{W}^{X,X}(s,\tau)=\begin{bmatrix} \overleftrightarrow{W}^{X1,X1}(s,\tau) & \overleftrightarrow{W}^{X1,X2}(s,\tau) & \cdots & \overleftrightarrow{W}^{X1,Xq}(s,\tau) \\ \overleftrightarrow{W}^{X2,X1}(s,\tau) & \overleftrightarrow{W}^{X2,X2}(s,\tau) & \cdots & \overleftrightarrow{W}^{X2,Xq}(s,\tau) \\ \vdots & \vdots & & \vdots \\ \overleftrightarrow{W}^{Xq,X1}(s,\tau) & \overleftrightarrow{W}^{Xq,X2}(s,\tau) & \cdots & \overleftrightarrow{W}^{Xq,Xq}(s,\tau) \end{bmatrix}, \quad (1)$$

where $\overleftrightarrow{W}^{Xi,Xj}(s,\tau)$ is the smoothed auto-wavelet power spectra (when *i*=*j*) or cross-wavelet power spectra (when $i \neq j$) at scale s and spatial (or temporal) location τ respectively. For the detailed calculation of smoothed auto- and cross-wavelet power spectra, see Supplement, Sect. S1.

The matrix of smoothed cross wavelet power spectra between response variable Y and predictor variables Xi can be defined as

$$\overleftrightarrow{W}^{Y,X}(s,\tau)=\begin{bmatrix} \overleftrightarrow{W}^{Y,X1}(s,\tau) & \overleftrightarrow{W}^{Y,X2}(s,\tau) & \cdots & \overleftrightarrow{W}^{Y,Xq}(s,\tau) \end{bmatrix}, \quad (2)$$

where $\overleftrightarrow{W}^{Y,Xi}(s,\tau)$ is the smoothed cross-wavelet power spectra between Y and Xi at scale s and spatial (or temporal) location τ.

The smoothed wavelet power spectrum of response variable Y is $\overleftrightarrow{W}^{Y,Y}(s,\tau)$.

Following Koopmans (1974), the MWC at scale s and location τ, $\rho_m^2(s,\tau)$, can

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be written as

$$\rho_m^{2}(s,\tau)=\frac{\overleftrightarrow{W}^{Y,X}(s,\tau)\,\overleftrightarrow{W}^{X,X}(s,\tau)^{-1}\,\overline{\overleftrightarrow{W}^{Y,X}(s,\tau)}}{\overleftrightarrow{W}^{Y,Y}(s,\tau)}. \tag{3}$$

When only one predictor variable (e.g., *X1*) is included in *X*, Eq. (3) is the equation for simple wavelet coherence, $\rho_s^{2}(s,\tau)$, between two variables (Torrence and Webster, 1999; Grinsted et al., 2004):

$$\rho_s^{2}(s,\tau)=\frac{\overleftrightarrow{W}^{Y,X1}(s,\tau)\,\overline{\overleftrightarrow{W}^{Y,X1}(s,\tau)}}{\overleftrightarrow{W}^{X1,X1}(s,\tau)\,\overleftrightarrow{W}^{Y,Y}(s,\tau)}. \tag{4}$$

Therefore, simple wavelet coherence is consistent with multiple wavelet coherence if only one predictor variable is included. In addition, the wavelet phase between a response variable (*Y*) and a predictor variable (*X1*) is

$$\phi(s,\tau)=\tan^{-1}\left(\mathrm{Im}\left(W^{Y,X1}(s,\tau)\right)/\mathrm{Re}\left(W^{Y,X1}(s,\tau)\right)\right), \tag{5}$$

where Im and Re denote the imaginary and real part of $W^{Y,X1}(s,\tau)$ respectively. Note that the phase information between a response variable *Y* and multiple predictor variables *X* cannot be obtained.

Multiple wavelet coherence at the 95% confidence level is calculated using the Monte Carlo method (Grinsted et al., 2004). Surrogate spatial series (i.e., red noise) of all variables are generated with a Monte Carlo simulation based on their first-order autocorrelation coefficient (AR1). The MWC at each scale and location is calculated using the simulated spatial series. This is repeated an adequate number of times (e.g., 1000) (Grinsted et al., 2004). At each scale, MWCs at all locations outside the cones of influence from all simulations are ranked in ascending order. The value at the 95th

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percentile represents the 95% confidence level for the MWC at that scale. The Matlab codes and user manual document for calculating MWC and significance level are provided in the Supplement (Sect. S2–S4).

3. Data and analysis

3.1 Artificial data for method test

The method is tested using a stationary and non-stationary artificial dataset generated following Yan and Gao (2007). The response variable (y for the stationary case and z for the non-stationary case) encompasses five cosine waves (y1 to y5 for the stationary case and z1 to z5 for the non-stationary case) with different dimensionless scales (Fig. 1). For the stationary case, $y1=\cos(2\pi x/4)$, $y2=\cos(2\pi x/8)$, $y3=\cos(2\pi x/16)$, $y4=\cos(2\pi x/32)$, and $y5=\cos(2\pi x/64)$, where x=0, 1, 2, …, 255. There is one regular cycle every 4, 8, 16, 32, and 64 locations, representing dimensionless scales of 4, 8, 16, 32, and 64 for y1, y2, y3, y4, and y5 respectively (Fig. 1a). The regular cycles make each predictor and response series stationary. For the non-stationary case, $z1=\cos(500\pi(x/1000)^{0.5})$, $z2=\cos(250\pi(x/1000)^{0.5})$, $z3=\cos(125\pi(x/1000)^{0.5})$, $z4=\cos(62.5\pi(x/1000)^{0.5})$, and $z5=\cos(31.25\pi(x/1000)^{0.5})$, where x=0, 1, 2, …, 255. The equation with the square root of the location term results in the gradual change in frequency (scale), with the greatest dimensionless scales of 4, 8, 16, 32, and 64 at the right hand side for z1, z2, z3, z4, and z5 respectively (Fig. 1b). The average scales for these predictor variables are 3, 5, 9, 17, and 32 respectively. The location-varying scales make each predictor and response variable non-stationary.

For both the stationary and non-stationary series, the variance of the response variable is 2.5. The predictor variables, each with a variance of 0.5, are orthogonal to each other, and contribute equally to the total variance of the response variable. The cosine-like artificial datasets mimic many time series such as seismic signals, turbulence, air temperature, precipitation, hydrologic fluxes, and the El Niño-Southern Oscillation. They also mimic spatial series such as ocean waves, seafloor bathymetry, land surface topography, and soil water content along a hummocky landscape in geosciences. Therefore, they are representative of a geoscience data series and are suitable for testing the new method.

Multiple wavelet coherence between the response variable y (or z) and two (y2 and y4, or z2 and z4) or three (y2, y3, and y4, or z2, z3, and z4) predictor variables were calculated. The advantage of the artificial data is that the known scale- and localized features for all variables, and the known relationships between the response and each predictor variable are exact. By definition, the coherence is 1 at scales corresponding to that of included predictor variables and 0 at other scales.

To demonstrate the advantages of MWC in dealing with abrupt changes (a type of transient and localized feature), the second half of the original series of y2 (or z2) or y4 (or z4) is replaced by 0, and MWC between the response variable and new set of predictor variables is calculated. We anticipate that the coherence changes from 1 to 0 at the location where the new predictor variable becomes 0.

Predictor variables may not be as regular as that shown in Fig. 1 and may also be cross-correlated to one another. For these reasons, white noises with a mean of 0 and a standard deviation of 0.3, 1, and 4 are generated and added to predictor variables of y2 (or z2) and y4 (or z4). The resulting noised series are termed weakly, moderately, and highly noised series respectively, and have a correlation coefficient of 0.9, 0.5, and 0.1 respectively, with their original predictor variable. Multiple wavelet coherences between the response variable and different predictor variables (original and noised series) are calculated to demonstrate the performance of MWC when noised or correlated predictor variables are involved. Only the non-stationary case will be demonstrated because the performances of MWC for stationary and non-stationary cases are similar.

The MWC is compared to the MSC (Koopmans, 1974; Si, 2008) and MCC_{memd} (Hu and Si, 2013). The MSC is calculated based on the calculated auto- and cross- power spectra using an equation similar to Eq. (3). The detailed introduction of this method can be found in Si (2008). For the calculation of MCC_{memd}, a set of response and predictor variables form a multivariate data series for MEMD. The MEMD is a data driven method and has the ability to align "common scales" present within multivariate data. Please refer to Rehman and Mandic (2010) and Hu and Si (2013) for the MEMD analysis and the website (http://www.commsp.ee.ic.ac.uk/~mandic/research/emd.htm) for the related Matlab codes. The original series of response and predictor variables can be decomposed into different components (IMFs) with different scales by the MEMD. For IMFs at the same scale, multiple stepwise regressions are conducted between response and

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predictor variables, and the multiple correlation coefficients for each scale-specific IMF are calculated.

3.2 Real data for application

Daily evaporation (*E)* from free water surfaces of E601 evaporation pan (pan diameter of 61.8 cm) and other meteorological factors (i.e., relative humidity, mean temperature, sun hours, and wind speed) were collected from January 1, 1979 to December 31, 2013 at Changwu site in Shaanxi, China. The Changwu site is a transition area between semi-arid and subhumid climate where water limits agricultural productivity. Monthly averages of all variables were used in this study because we are mainly interested in seasonal and inter-annual variability.

4. Results and discussion

4.1 MWC with orthogonally predictor variables

For the stationary data, there are two narrow horizontal bands (red color) representing an MWC value of around 1 at the respective scales of 8 and 32 for all locations (Fig. 2a). These two bands also correspond to the scales of 8 and 32 respectively, for the two predictor variables. When an additional predictor variable with the scale of 16 is introduced, a wide band from 6 to 40 appears, signifying that the MWC equals approximately 1 at all locations at the scales of 8, 16, and 32. As anticipated, when all five predictor variables with scales ranging from 4 to 64 are included, coherence values of close to 1 are found in the whole scale-location domain (data not shown).

The application of MWC to the non-stationary datasets shows that the scales with significant MWC values gradually increase with the increase in distance. This increase in the scales is due to the non-stationarity of the variables (Fig. 2b). For example, when predictor variables of z2 and z4 are included, scales of the two bands corresponding to MWC around 1 increase from 4 to 8 and from 8 to 32, respectively. Furthermore, as expected, for only one predictor variable (stationary and non-stationary), MWC reduces to simple wavelet coherence; there is only one band of coherence around 1, which corresponds to the scale of that predictor variable (data not shown). Note that the significant MWC values for both stationary and non-stationary cases are not exactly 1 at all scales or locations due to the smoothing effect along both scales and locations. However, the mean MWC values of the significant bands are very high (i.e., 0.94 – 1.00) and the MWC values at the centre of the significant band are 1, which corresponds to the exact scale of a predictor variable.

When the point values in the second half of the data series of a predictor variable is replaced by 0, the MWC in that half is almost 0 at scales corresponding to that predictor variable (Fig. 3). For the stationary case, when the point values in the second half of the data series of predictor variable y2 (or y4) is replaced by 0, the MWC is around 1 at the scale of 8 (or 32) in the first half of the transect and 0 in the second half (Fig. 3a). Similar results were also found for the non-stationary case (Fig. 3b). This is expected because the constant series of 0 is not correlated to the response variables at any scale. Much like simple wavelet coherence, the MWC method is able to detect abrupt changes in the data series and has the advantages of dealing with localized multivariate relationships.

4.2 MWC with noised and correlated predictor variables

When z2 and a noised series derived from z2 are included as predictor variables, there is only one band of coherence close to 1 at scales corresponding to z2, irrespective of the correlation between z2 and a noised series of z2 (Fig. 4a). When z2 and a noised series of z4 are included as predictor variables, the coherence depends on the degree of the noise (Fig. 4b). For weakly noised series, there are two bands of coherence of around 1 corresponding to the scales of z2 and z4 respectively. The PASC is 23%, which equals that of when z2 and z4 are included. With the increase of noise, the coherence and corresponding PASC at the scales corresponding to z4 decrease. When z2 and a strongly noised series of z4 are considered, the band of coherence around 1 at scales corresponding to z4 disappears.

The inclusion of a third noised z4 variable substantially increases the area with high coherence (in red) as compared to the case when only z2 and z4 are included (Fig. 4c). This indicates that MWC will increase with the increase in the number of predictor variables, with the highest coherence less or equal to 1, irrespective of the number of predictor variables. However, the area of significant coherence may not necessarily increase (Ng and Chan, 2012a). In fact, the PASC values for three predictor variables (19-20%) are lower than for only two predictor variables (23%). This indicates that, in this case, two predictor variables are better than three in terms of explaining the variations of the response variable. This is because the variance of the response variable explained by the noised variable is already accounted for by other variables. Therefore, only an additional variable that can independently explain a fair amount of variance could contribute significantly to explaining variations of a response variable (Fig. 4b). This can also explain why there is only one band of coherence around 1 at scales corresponding to z2, when z2 and a noised series of z2 are included (Fig. 4a). This information is helpful in choosing predictor variables for developing scale-specific predictions, especially when predictor variables are correlated.

4.3 Comparison with other multivariate methods

4.3.1 MSC

The MSC as a function of scale is shown in Fig. 5a. For the stationary case, when y2 and y4 are included as predictor variables, there are two plateaus centered at the scales of 8 and 28 representing a coherence of 1. As expected, when an additional predictor variable y3 is added, the corresponding scale of 16 also shows coherence of 1. The MSC produces similar scale-specific relationships as MWC does for a stationary dataset except that the centered scale (i.e., 28) with coherence of 1 deviates from the expected value (i.e., 32) for predictor variable y4. For the non-stationary case, however, the MSC is much lower than 1 for the predictor variables of z2 and z4; MSC of 1 is present only at the scale of 8 when an additional predictor variable z3 is added. Obviously, the MSC underestimates the multivariate relationships and is not suitable to non-stationary processes (Si, 2008) due to its inability to deal with localized features. The MSC at a specific scale provides the average of multivariate relationships across all locations. Because the scale of a predictor variable changes with location for the non-stationary case, the MSC deviates greatly from 1.

The inability of the MSC to deal with localized features is demonstrated further by the decrease of MSC at scales when the second half of the included predictor variable series are replaced by 0 for both the stationary and non-stationary series (Fig. 5b). For example, when the second half of the y4 series is replaced by 0 for the stationary case, the MSC at scales around 32 decreases from 1 to 0.52. Although the MSC can detect the decrease of coherence at the scales corresponding to the 0 values throughout the second half of the series, the exact locations for the decrease cannot be identified. In fact, the coherence decreases only in the second half of the series, and does not change in the first half of the series. The location for the decrease can be easily identified by the MWC, but not by MSC.

4.3.2 MCC_{memd}

Five intrinsic mode functions (IMFs) with non-negligible variance are obtained for multivariate data series. While the obtained scales for the response variable y are in agreement with the true scales for the stationary case, the obtained scales (i.e., 3, 6, 11, 21, and 43) for the response variable z deviate slightly from the average scales for the non-stationary case. For the response variable, the contribution of IMFs to the total variance generally decreases (20% to 13% for stationary and 27% to 11% for non-stationary) from IMF1 to IMF5, which disagrees with the fact that each scale contributes equally (i.e., 20%) to the total variance. In addition, the sum of variances over all IMFs for each variable is less than 100% (ranging from 84% to 93%), indicating that MEMD cannot capture all the variances. For the detailed results of MEMD, see Supplement, Sect. S5.

The MCC_{memd} as a function of scale is shown in Fig. 6a. For the stationary case, when predictor variables of y2 and y4 are included, the MCC_{memd} values are 0.98 and 0.93 respectively, at scales corresponding to that of y2 and y4. When a predictor variable of y3 is included, the MCC_{memd} values are 1.00, 1.00, and 0.96 respectively, at scales corresponding to that of y2, y3, and y4. For the non-stationary case, the corresponding MCC_{memd} values are 0.80 and 0.85 for the two predictor variable case, and 0.95, 0.99, and 0.91 respectively, for the case of three predictor variables. Therefore, the MCC_{memd} can be used to determine the scale-specific multivariate relationships. Similar to MSC, however, the MCC_{memd} underestimates the multivariate relationships, especially for the non-stationary case with less predictor variables. On the contrary, the MCC_{memd} can also overestimate the multivariate relationships. For example, when only predictor variables corresponding to scales of 8, 16, and 32 are considered, the MCC_{memd} value for the stationary case is 0.47 at the scale of 64, which deviates much from the expected MCC_{memd} value of 0 (Fig. 6a). The possible underestimation and overestimation by the MCC_{memd} may come from the decomposition errors inherent in the MEMD algorithm (Rehman and Mandic, 2010).

Similar to MSC, the localized multivariate relationships cannot be obtained from MCC_{memd}. This can be better explained by the decrease of MCC_{memd} when half of the series of the predictor variables are replaced by 0 (Fig. 6b). For the stationary case, the MCC_{memd} values at the scales corresponding to y2 (or y4) decrease from 0.98 to 0.49 and from 0.93 to 0.62 when the second half of the y2 (or y4) series are replaced by 0.

As explained above, the MWC has advantages in untangling localized multivariate relationships as compared to the common multivariate methods. It is important to reveal the multivariate relationships, which vary with time or space that are associated with different processes. For example, discharge usually happens on knolls, while recharge usually happens in neighboring depressions (Gates et al., 2011). Therefore, the controlling factors of soil water storage may vary with the land element characteristics of a location. For example, local controls may be more important on knolls, while non-local controls may be more important in depressions (Grayson et al., 1997). In a temporal domain, vegetation transpiration contributes more to the evapotranspiration in the growing seasons, which may result in the changes of environmental factors explaining temporal variations of evapotranspiration in different seasons.

4.4 Application of the MWC

Each meteorological factor was significantly correlated to the *E*, but the dominant factors explaining variation in *E* differed with scale. For example, the relative humidity was the dominating factor at small (2–8 months) and large (>32 months) scales, while temperature was the dominating factor at the medium (8–32 months) scales. Overall, the relative humidity corresponded to the greatest mean MWC (0.62) and PASC value (40%) at multiple scale-location domains. For the detailed relationships between *E* and each factor, see Supplement, Sect. S6.

The MWC analysis shows that the combination of relative humidity and mean temperature produced the greatest mean MWC (0.82) and PASC (49%) among all two-factor cases, indicating that they are the best to explain variations in *E* at multiple scale-location domains (Fig. 7a). However, adding an additional factor such as sun hours, which was the best among all three-factor cases, increased the average coherence (0.91), but slightly decreased the PASC to 48% (Fig. 7b). This indicated that sun hours was not significantly different from red noise in explaining additional variation in *E*. Similar results were found when the wind speed was added. The reason for this was that most areas with significant coherence between *E* and sun hours or wind speed, were a subset of areas with significant coherence between *E* and relative humidity or mean temperature (see Supplement, Sect. S3). Therefore, relative humidity and mean temperature were adequate to explain the temporal variation of *E* at various scales at this site. This is consistent with Li et al. (2012), who indicate that relative humidity and mean temperature are the two main contributors to the temporal change of potential evapotranspiration on the Chinese Loess Plateau.

5. Conclusions

Multiple wavelet coherence is developed to determine scale-specific and localized multivariate relationships in geosciences. The new method is tested and compared with exiting multivariate methods using an artificial dataset. The new method can be used to determine the proportion of the variance of a response variable that is explained by predictor variables at a specific scale and location (spatially or temporally). As compared with simple wavelet coherence, more variation may be explained at multiple scale-location domains by the MWC. Including more variables is only beneficial if the variables are not strongly cross-correlated and can independently explain a fair amount of variability in a response variable. Therefore, the best combinations of variables that explain multivariate spatial or temporal variability at multiple scales can be determined. This is important for optimizing variables for developing scale-specific prediction. The MSC and MCC_{memd} can determine multivariate relationships at multiple scales, but localized multivariate relationships are not available and both MSC and MCC_{memd} are likely to underestimate the degree of multivariate relationships for non-stationary processes. In addition, the performance of MCC_{memd} relies on the performance of MEMD, which needs further development. Application of the MWC into the real dataset indicates that the combination of relative humidity and mean temperature are the optimal factors to explain temporal variation of *E* at the Changwu site in China.

In summary, multiple wavelet coherence has advantages over existing multivariate methods, and provides an effective vehicle for untangling complex spatial or temporal variability for multiple controlling factors at multiple scales and locations. It may also be used as a data-driven tool for modeling and predicting various processes in the area of geosciences such as precipitation, drought, soil water dynamics, stream flow, and

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atmospheric circulation.

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References

Biswas, A. and Si, B. C.: Identifying scale specific controls of soil water storage in a hummocky landscape using wavelet coherency, Geoderma, 165, 50–59, doi: 10.1016/j.geoderma.2011.07.002, 2011.

Carey, S. K., Tetzlaff, D., Buttle, J., Laudon, H., McDonnell, J., McGuire, K., Seibert, J., Soulsby, C., and Shanley, J.: Use of color maps and wavelet coherence to discern seasonal and interannual climate influences on streamflow variability in northern catchments, Water Resour. Res., 49, 6194–6207, doi: 10.1002/wrcr.20469, 2013.

Das, N.N. and Mohanty, B. P.: Temporal dynamics of PSR-based soil moisture across spatial scales in an agricultural landscape during SMEX02: A wavelet approach, Remote Sens. Environ., 112, 522–534, doi:10.1016/j.rse.2007.05.007, 2008.

Gates, J. B., Scanlon, B. R., Mu, X. M., and Zhang, L.: Impacts of soil conservation on groundwater recharge in the semi-arid Loess Plateau, China, Hydrogeol. J., 19, 865–875, 2011.

Graf, A., Bogena, H. R., Drüe, C., Hardelauf, H., Pütz, T., Heinemann, G., and Vereecken, H.: Spatiotemporal relations between water budget components and soil water content in a forested tributary catchment, Water Resour. Res., 50, 4837–4857, doi:10.1002/2013WR014516, 2014.

Grayson, R. B., Western, A. W., Chiew, F. H. S., and Blöschl, G.: Preferred states in spatial soil moisture patterns: local and nonlocal controls, Water Resour. Res., 33, 2897–2908, doi: 10.1029/97WR02174, 1997.

Grinsted, A., Moore, J. C., and Jevrejeva, S.: Application of the cross wavelet transform and wavelet coherence to geophysical time series, Nonlinear Proc. Geoph., 11, 561–566, 2004.

Hu, W., Biswas, A., and Si, B. C.: Application of multivariate empirical mode decomposition for revealing scale- and season-specific time stability of soil water storage, Catena, 113, 377–385, doi:10.1016/j.catena.2013.08.024, 2014.

Hu, W. and Si, B. C.: Soil water prediction based on its scale-specific control using multivariate empirical mode decomposition, Geoderma, 193–194,180–188, doi: 10.1016/j.geoderma.2012.10.021, 2013.

Koopmans, L. H.: The spectral analysis of time series, Academic Press, New York, 1974.

Labat, D.: Recent advances in wavelet analyses: Part I. A review of concepts, J. Hydrol., 314, 275–288, doi: 10.1016/j.jhydrol.2005.04.003, 2005.

Lakshmi, V., Piechota, T., Narayan, U., and Tang, C. L.: Soil moisture as an indicator of weather extremes, Geophys. Res. Lett., 31, L11401, doi:10.1029/2004GL019930, 2004.

Li, Z., Zheng, F. L., and Liu, W. Z.: Spatiotemporal characteristics of reference evapotranspiration during 1961–2009 and its projected changes during 2011–2099 on the Loess Plateau of China, Agric. For. Meteorol., 154-155, 147–155, doi:10.1016/j.agrformet.2011.10.019, 2012.

Mihanović, H., Qrlić, M., and Pasrić, Z.: Diurnal thermocline oscillations driven by tidal flow around an island in the Middle Adriatic, J. Marine Syst., 78, S157–S168, doi: 10.1016/j.jmarsys.2009.01.021, 2009.

Müller, W. A., Frankignoul, C., and Chouaib, N.: Observed decadal tropical Pacific–North Atlantic teleconnections, Geophys. Res. Lett., 35, L24810, doi:10.1029/2008GL035901, 2008.

Ng, E. K. W. and Chan, J. C. L.: Geophysical applications of partial wavelet coherence and multiple wavelet coherence, J. Atmos. Ocean. Tech., 29, 1845–1853, doi: 10.1175/JTECH-D-12-00056.1, 2012a.

Ng, E. K. W. and Chan, J. C. L.: Interannual variations of tropical cyclone activity over the north Indian Ocean, Int. J. Climatol., 32, 819–830, doi: 10.1002/joc.2304, 2012b.

Polansky, L., Wittemyer, G., Cross, P. C., Tambling, C. J., and Getz, W. M.: From moonlight to movement and synchronized randomness: Fourier and wavelet analyses of animal location time series data, Ecology, 91, 1506–1518, doi: 10.1890/08-2159.1,

Hydrol. Earth Syst. Sci. Discuss., doi:10.5194/hess-2016-154, 2016
Manuscript under review for journal Hydrol. Earth Syst. Sci.
Published: 27 April 2016

2010.

Rehman, N. and Mandic, D. P.: Multivariate empirical mode decomposition, Proc. R. Soc. A., 466, 1291–1302, doi:10.1098/rspa.2009.0502, 2010.

She, D. L., Liu, D. D., Peng, S. Z., and Shao ,M. A.: Multiscale influences of soil properties on soil water content distribution in a watershed on the Chinese Loess Plateau, Soil Sci., 178, 530–539, doi: 10.1097/SS.0000000000000021, 2013.

She, D. L., Zheng, J. X., Shao, M. A., Timm, L. C., and Xia, Y. Q.: Multivariate empirical mode decomposition derived multi-scale spatial relationships between saturated hydraulic conductivity and basic soil properties, Clean-Soil Air Water, doi: 10.1002/clen.201400143, 2015.

Si, B. C.: Spatial scaling analyses of soil physical properties: A review of spectral and wavelet methods, Vadose Zone J., 7, 547–562, doi: 10.2136/vzj2007.0040, 2008.

Si, B. C. and Zeleke, T. B.: Wavelet coherency analysis to relate saturated hydraulic properties to soil physical properties, Water Resour. Res., 41, W11424, doi:10.1029/2005WR004118, 2005.

Tang, C. L. and Piechota, T. C.: Spatial and temporal soil moisture and drought variability in the Upper Colorado River Basin, J. Hydrol., 379, 122–135, doi: 10.1016/j.jhydrol.2009.09.052, 2009.

Torrence, C. and Compo, G. P.: A practical guide to wavelet analysis, Bull. Am. Meteorol. Soc., 79, 61–78, doi: 10.1175/1520-0477(1998)079<0061:apgtwa>2.0.co;2, 1998.

Torrence, C. and Webster, P. J.: Interdecadal changes in the ENSO-monsoon system, J.

Clim., 12, 2679–2690, doi: 10.1175/1520-0442(1999)012<2679:ICITEM>2.0.CO;2, 1999.

Yan, R. and Gao, R. X.: A tour of the Hilbert–Huang transform: an empirical tool for signal analysis, IEEE Instrum. Meas. Mag., 10, 40–45, doi: 10.1109/MIM.2007.4343566, 2007.

Hydrol. Earth Syst. Sci. Discuss., doi:10.5194/hess-2016-154, 2016
Manuscript under review for journal Hydrol. Earth Syst. Sci.
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Figure captions

Figure 1. (a) Stationary and (b) non-stationary series of response variables (y for stationary and z for non-stationary case) encompassing five cosine waves (y1 to y5 for stationary and z1 to z5 for non-stationary case) with different dimensionless scales.

Figure 2. Multiple wavelet coherence (a) between response variable y and predictor variables y2 and y4; (b) between response y and predictors y2, y3, and y4; (c) between response z and predictors z2 and z4; and (d) between response z and predictors z2, z3, and z4. The artificial data series (y) encompasses five cosine waves (y1, y2, y3, y4, and y5) with different scales for the stationary case, and the artificial data series (z) encompasses five cosine waves (z1, z2, z3, z4, and z5) with different scales for the non-stationary case. The predictor variables, connected by a hyphen, are shown in the top right corner of each subplot. Thin solid lines demarcate the cones of influence, and thick solid lines show the 95% confidence levels.

Figure 3. Multiple wavelet coherence (a) between y and y2h0 and y4; (b) between y and y2 and y4h0; (c) between z and z2h0 and z4; and (d) between z and z2 and z4h0. The artificial data series (y) encompasses five cosine waves (y1, y2, y3, y4, and y5) with different scales for the stationary case and the artificial data series (z) encompasses five cosine waves (z1, z2, z3, z4, and z5) with different scales for the non-stationary case. The variables y2h0 (or z2h0) and y4h0 (or z4h0) refer to the new series of y2 (or z2) and y4 (or z4), in which the second half is replaced by 0. The

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Figure 4. Multiple wavelet coherence of an artificial data series (z) encompassing five cosine waves (z1, z2, z3, z4, and z5) with different scales and (a) z2 and noised z2, (b) z2 and noised z4, and (c) z2, z4, and noised z4 for the non-stationary case. The predictor variables are connected by a hyphen and shown in the top right corner of each subplot. z2wn (z4wn), z2mn (z4mn), and z2sn (z4sn) indicate weakly, moderately, and strongly noised z2 (z4) series respectively. Weakly, moderately, and strongly noised series are correlated with original series, having correlation coefficients of 0.9, 0.5, and 0.1 respectively. Thin solid lines demarcate the cones of influence and thick solid lines show the 95% confidence levels.

Figure 5. Multiple spectral coherence (MSC) of an artificial data series (y or z) encompassing five cosine waves (y1 to y5; or z1 to z5) with different scales and (a) two (y2 and y4; or z2 and z4) or three (y2, y3, and y4; or z2, z3, and z4) data series, and (b) two (y2 and y4; or z2 and z4) data series when the second half of one data series is replaced by 0. The variables y2h0 (or z2h0) and y4h0 (or z4h0) refer to the new series of y2 (or z2) and y4 (or z4) in which the second half is replaced by 0.

Figure 6. Multiple correlation coefficient between multivariate empirical mode decomposition (MCC_{memd}) of an artificial series (y or z) and (a) two (y2 and y4; or z2 and z4) or three (y2, y3, and y4; or z2, z3, and z4) data series, and (b) two (y2 and y4; or z2 and z4) data series when the second half of one data series is replaced by 0. The

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Figure 7. Multiple wavelet coherence between evaporation (*E*) from water surfaces and meteorological factors ((a) relative humidity and mean temperature and (b) relative humidity, mean temperature, and sun hours) at Changwu site in Shaanxi, China. Thin solid lines demarcate the cones of influence, and thick solid lines show the 95% confidence level.

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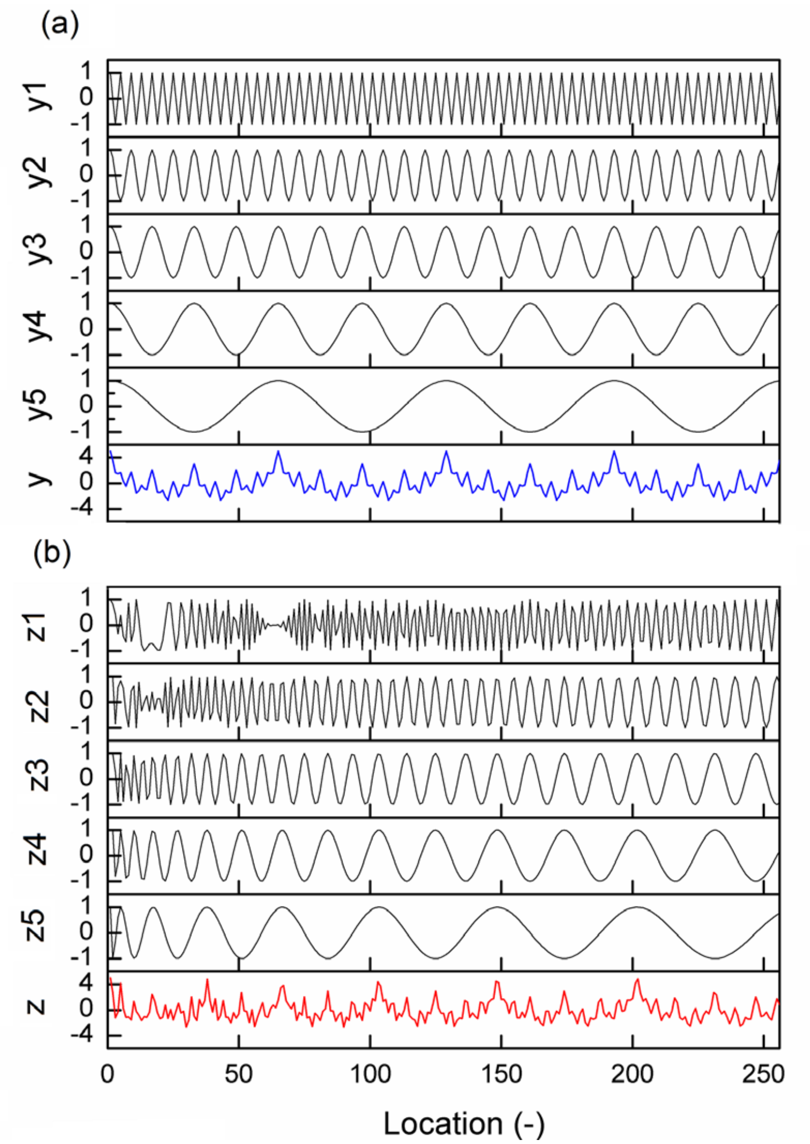

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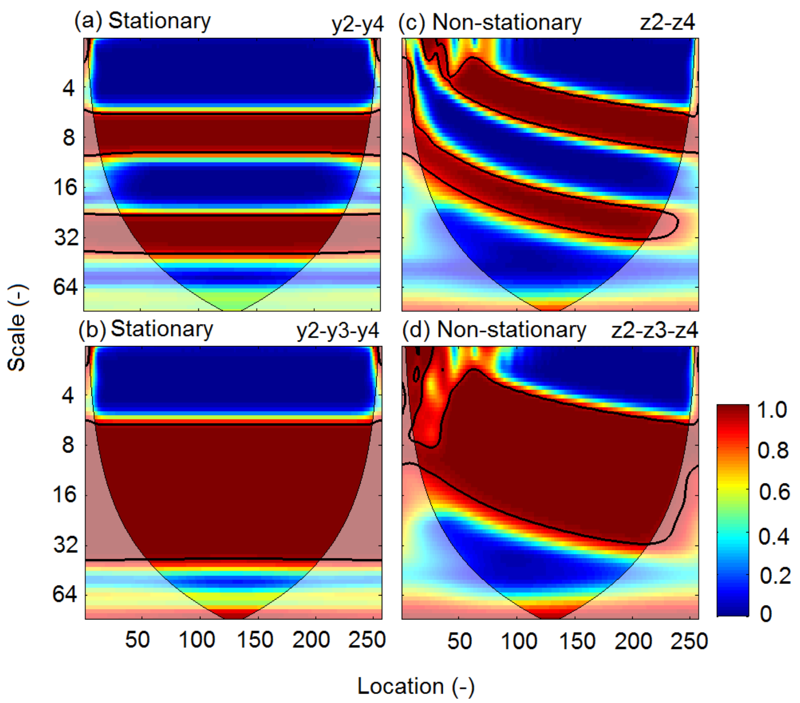

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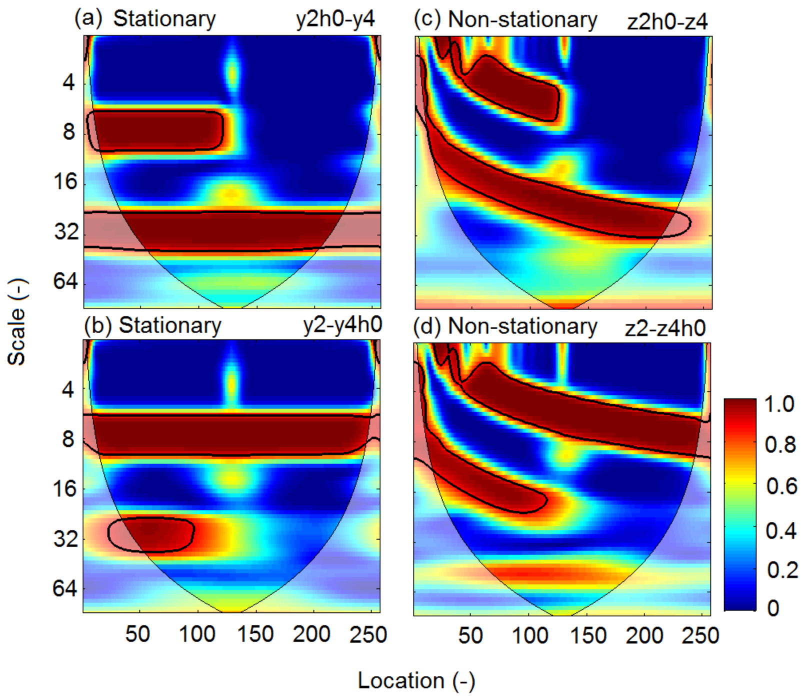

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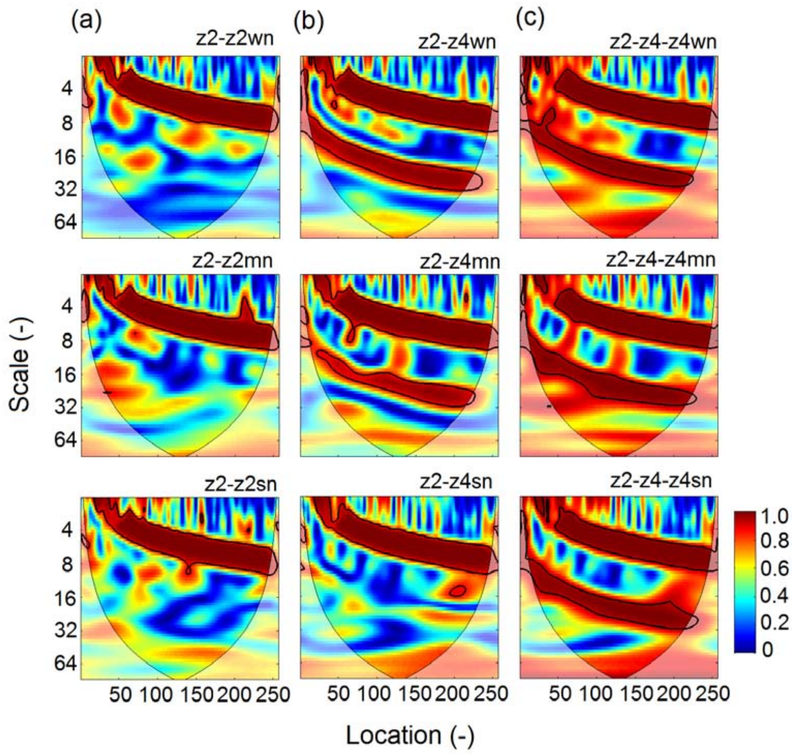

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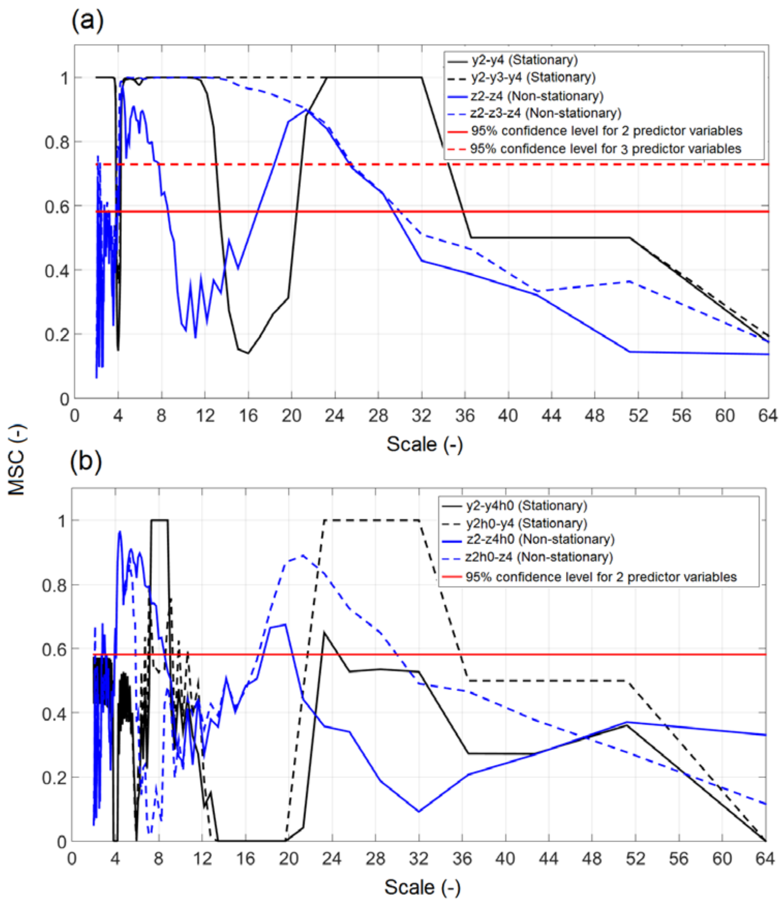

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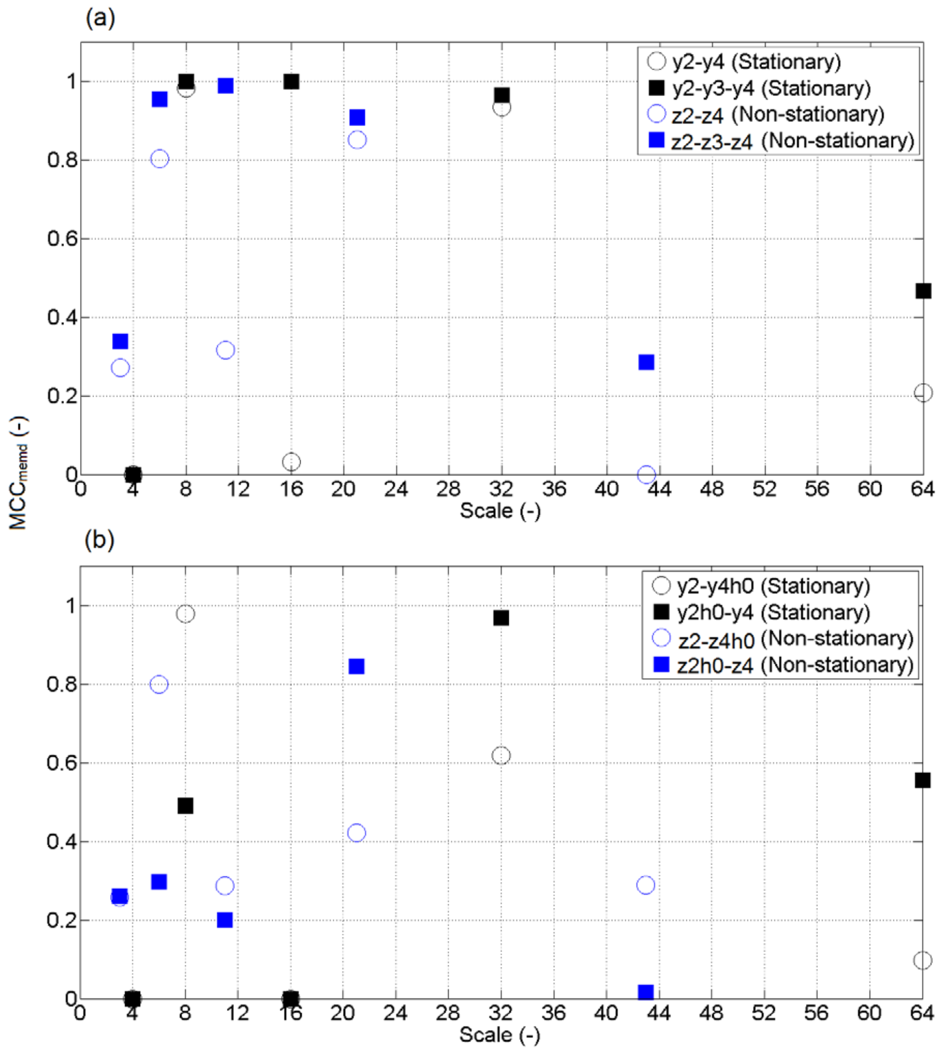

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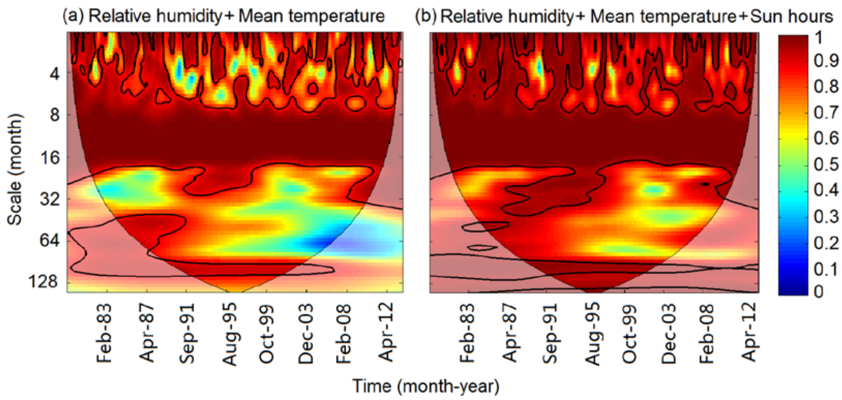

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