Interactive comment on "Recent changes and drivers of the atmospheric evaporative demand in the Canary Islands" by S. M. Vicente-Serrano et al. Anonymous Referee #1

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The manuscript presents a trend analysis of the FAO-56 reference evaporation using meteorological data from 8 sites at the Canary Islands. Interestingly the results show a remarkable heterogeneity in both the drivers of ET0 and its trends, which I did not expect due to the maritime climate. The most consistent effect is the decrease in relative humidity at most sites and thus an increase of the aerodynamic component of ET0. Generally the paper is well written, the data analysis is comprehensive and well designed. The topic of observed ET0 changes and choice of the sites are relevant and well suited for publication in HESS. Although I have a some remarks I am positive that the authors can implement these and recommend minor revisions.

We would like to thank the Reviewer#1 for his/her positive assessment of our manuscript. We are also very grateful with the detailed revision of the manuscript and the constructive comments raised with the purpose of improving our research article. They have been discussed below and implemented in this revised version.

Comments and remarks

The trends in ET0 seem to be significant because of the low values in the beginning of the chosen period. The results of the two sites with longer coverage show no significance. Thus the trend seems to be rather an effect of decadal scale variability. Please indicate this within the discussion of the results.

We thank for this point and fully agree with this suggestion. This has been included in the discussion section of the revised manuscript:

"In any case, we must also stress that trends in ETo at the regional scale are mostly significant because of the low values in the beginning of the study period starting in the 1960s. Thus, the results of the two sites with longer temporal coverage (i.e., Izaña and Santa Cruz de Tenerife) do not show significant trends. This makes necessary to consider these trends with caution since they could be driven by variability processes at the decadal scale."

Abstract: L16 ET0 is not explained, please state here in the abstract that you estimate AED by the FAO-56 reference evaporation equation

The Reviewer#1 is right and we have included this information in the abstract:

"Overall, the annual ETo, which was estimated by means of the FAO-56 Penman-Monteith equation, increased significantly by…"

L17-18 The sentence "The radiative component . . . did" can be removed because this is again stated in the next sentence. Also explain the meaning of the two components.

The sentence has been removed as suggested and we have explained the meaning of aerodynamic and radiative components as follows:

"In this study we analysed the contribution of (i) the aerodynamic (related to the water vapour that a parcel of air can store) and (ii) radiative (related to the available energy to evaporate a quantity of water) componets to the decadal variability and trends of ETo."

Introduction: The main research hypotheses should be clearly formulated

Thanks for this suggestion. The main hypothesis of our study has been included in the introduction section:

"The main hypothesis of the study is that in opposition to other continental temperate regions of the North Hemisphere, the warm and humid climate of the subtropical Canary Islands provides the water supply to the atmosphere needed to maintain the AED constant under the current global warming scenarios; consequently, only wind speed and solar radiation could affect the observed decadal variability and trends of the AED."

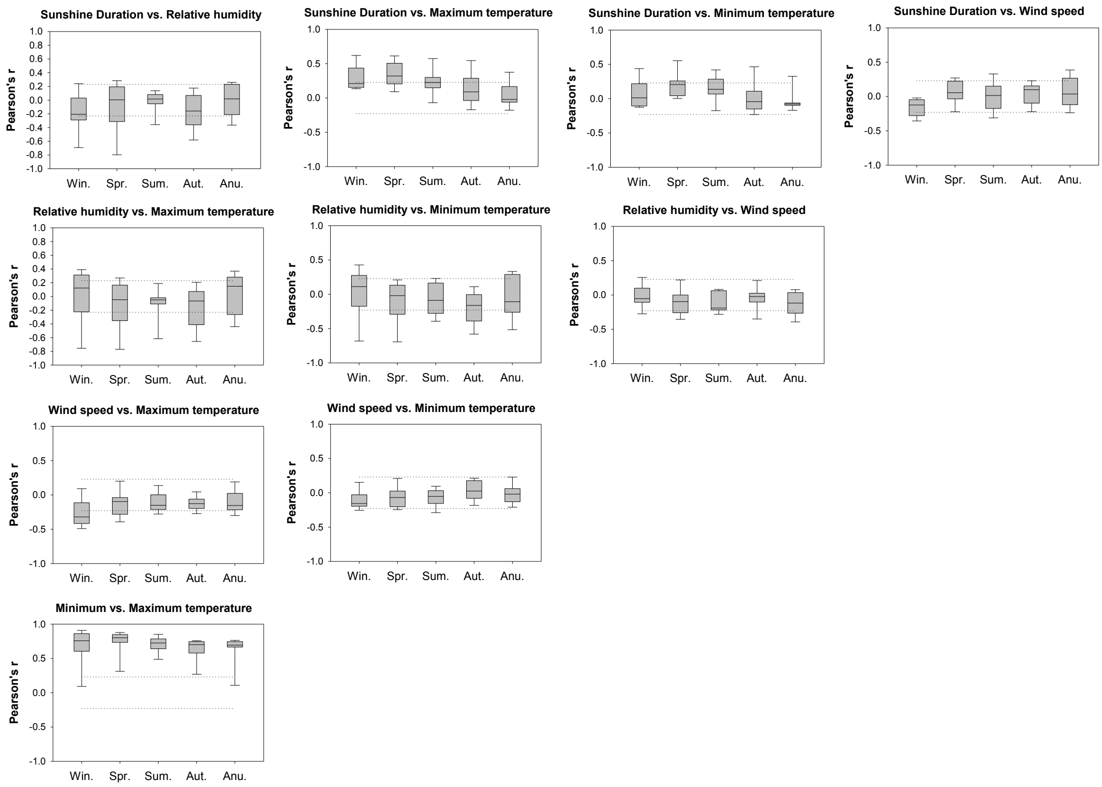
Section 2.1 L107-116: The homogenisation alters the original data and can affect the detection of trends. To achieve reproducibility of the results I recommend to provide an overview about data gaps, breakpoints and corrections which should be added as supplement.

The Reviewer#1 is right when stating that homogenisation alters the original data, but this "alteration" is really necessary in any study that focuses on climate variability and trends using observed datasets. On the contrary, there is a substantial risk on the robustness of the obtained climate trends based on inhomogeneous or not quality-controlled and tested series. For this reason, independently of the alteration of the series after the homogenisation tests, the application of homogenisation methods is strictly mandatory in any climate study aimed at retrieving long-term trends.

Moreover, we completely agree with the Reviewer#1 on the need of including further information on the data gaps and homogeneities found in the different variables. We have included a new table in the manuscript indicating these issues (please see below).

Section 2.3 L160-171 I do like the simple yet illustrate way to determine the effects of single variables on the detected trend. By design this is done as a local sensitivity analysis where one variable is changed holding the others fixed. However, it is not a global sensitivity analysis and co-variation of the forcing variables is neglected. Especially for the meteorological variables used here, I suspect that the variables and eventually their trends in time do co-vary - e.g. temperature and relative humidity. Did you consider such effects and are they important to understand the long-term variability?

We agree that co-variation of the forcing variables is not considered here because co-variation between meteorological variables in the Canary Islands is really low. This is illustrated in a set of box-plots below in which the correlation between the seasonal and annual series of the meteorological variables in the eight meteorological stations is shown. With the exception of the high positive correlations found between maximum and minimum air temperatures, the correlation among the other variables is really low and mostly non-significant. Only in winter and spring there are dominant significant correlations between the sunshine duration and the maximum air temperature. Given the strong independence in the variability of the different climate variables the co-variation study suggested by Reviewer#1 would not provide any new result in comparison with that applied here. In any case, we much appreciate this raised comment.



L195 . . . the aerodynamic component (Eq. 3).

Replaced.

Discussion: L303- 305 differences in ET0 trends across sites . . . "must be considered either due to random effects and uncertainty at various levels or due to microgeographic effects . . . " I think the differences of the trends and the different strength of the aerodynamic / radiative components at the sites deserve more attention in the discussion. The results are presented already in a detailed manner and these aspects should be discussed. That means is there a link of the different strength of certain forcing variables and the magnitude of ET0 trend at a given site.

We have included some discussion about this issue in the revised manuscript:

"There is not a general pattern that may connect the observed trends in a certain forcing variable with the observed trend of ETo in each of the eight analysed stations although those that showed a higher increase in ETo (i.e., Lanzarote, Los Rodeos and

Gran Canaria) displayed a higher increase in the aerodynamic component; a process which is in agreement with the significant reductions observed in relative humidity."

L514: in preparation

Replaced in the revised manuscript.

Tables: Please add a Site Metadata Table with site name, WMO-ID, LAT, LON, Height, gaps, relocation, corrections, available variables

We have included the suggested table in the revised version of the manuscript:

Code	Longitude	Latitude	Name	relocation	Relative humidity		Sunshine duration		Wind speed		maximum temperature		minimum temperature	
					data gaps	Inhom.	data gaps	Inhom.	data gaps	Inhom.	data gaps	Inhom.	data gaps	Inhom.
C029O	-13.60	28.95	Lanzarote/Airport	1972	2.20%	1998	0.78%	1978-2002	0.47%	1971	1.23%	2004	1.23%	1988
C139E	-17.75	28.61	La Palma/Airport	1970	0.94%		2.51%		0.47%	1976	0.37%		0.37%	1997
C249I	-13.85	28.43	Fuerteventura/Airport	1969	0.15%	2000	1.25%	1995-2005	0.15%		0.23%	1983	0.23%	1977
C430E	-16.48	28.30	Izaña	--	1.72%	1999	7.40%	2005	6.91%		5.20%	1985	5.20%	
C447A	-16.31	28.46	Los Rodeos	--	0.31%		1.10%	1966	0.15%	1970	0.30%	2005	0.30%	2005
C449C	-16.25	28.45	Santa Cruz de Tenerife	--	0%		0.94%		0%	1987	0%		0%	1994
C649I	-15.38	27.91	Gran canaria/Airport	--	0.15%	1981-1994	2.67%	1978	0.31%	1972	0.20%	1984	0.20%	1994
C659P	-15.41	28.15	San Cristóbal	1994	11%		1.88%	1980	10.50%	1994	5.30%	1966	5.30%	

Figure 1: Please add LAT - LON coordinates as a grid and a scale for the distance

Figure 1 has been replaced following the Reviewer's suggestion.

Figure 4 and 5: The grey lines are not very informative. Please adapt the figures, using different colors or line types for the sites. It might be also useful to demean the time series for the display. In the annual panel the bold line is missing.

We have removed grey lines and only included the two meteorological stations with longest records (Izaña in green, and Santa Cruz de Tenerife in brown)

Fig. 5 the labels are too small to be readable

Labels have been replaced to be readable.

Finally, we wish to thank the Reviewer#1 for reviewing our paper and for your useful comments/suggestions.

Interactive comment on "Recent changes and drivers of the atmospheric evaporative demand in the Canary Islands" by S. M. Vicente-Serrano et al. Anonymous Referee #2

The manuscript deals with an analysis of the atmospheric evaporative demand (AED) over the Canarian Island for the period 1961-2013. Basis are meteorological data (monthly, p4l96) from 8 stations which are used as inputs for the FAO-56 PenmanMonteith equation to derive monthly AED. While the paper is generally well written, I feel there are a number of conceptuals issues that need to be resolved and addressed before a possible acceptance.

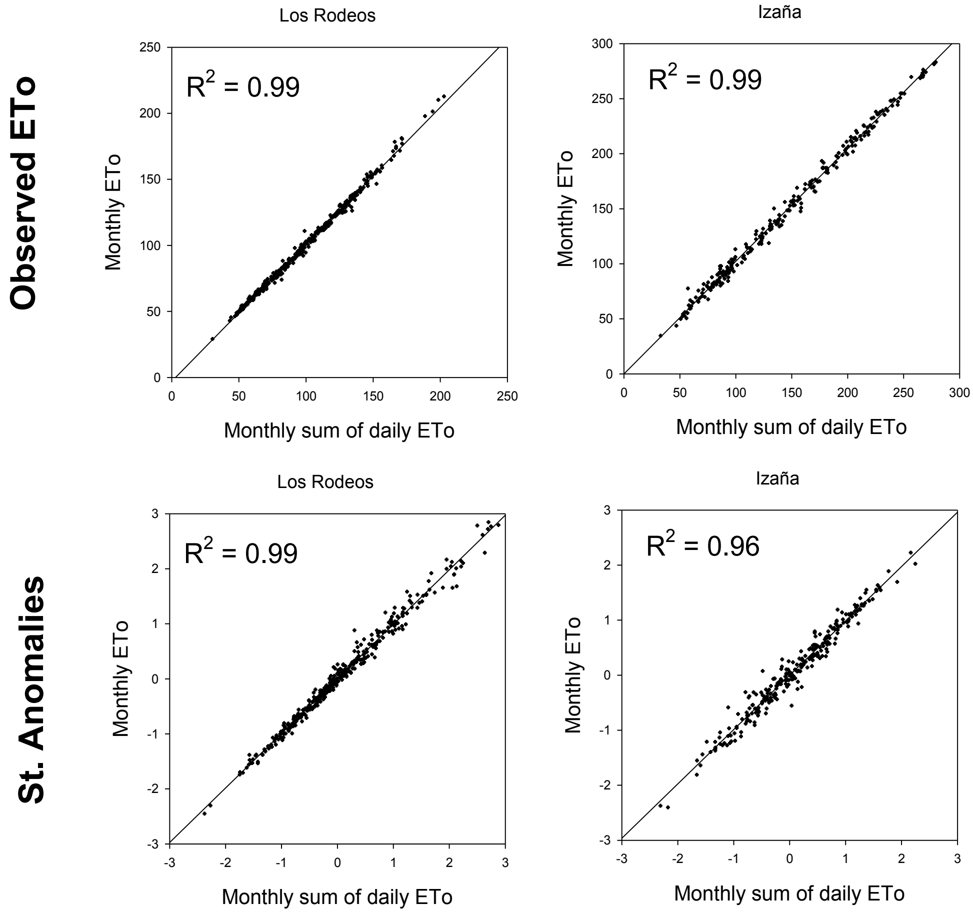
Thanks for your positive assessment of our manuscript and strong effort providing constructive comments. Addressing your comments has helped improved our revisions. Please find below our answers to each comment and if you have any residual concerns please feel free to raise those points.

- As the FAO-56 is a non-linear equation that has been developed for daily inputs, how do authors justify the application of monthly average input values?

The FAO-56 equation can be obtained from daily and monthly records, as Allen et al., 1998 stated: "the FAO Penman-Monteith equation requires air temperature, humidity, radiation and wind speed data for daily, weekly, ten-day or monthly calculations".

Several previous studies focused on drought using ETo at larger spatial scales have also computed the Penman Monteith ETo using monthly values for some variables. Using monthly averages instead of higher temporal resolution of data (e.g., daily) has not a relevant influence on the ETo estimations. An example is showed below for the ETo calculations in two stations of the Canary Islands (Los Rodeos and Izaña) for the 1978-2010 period based on the raw climatic data. The figure shows the relationship between the monthly ETo sum from daily measurements and the calculations from the average of monthly climate variables, which justifies the equality of applying both procedures. This is clearly observed for the ETo monthly values (including seasonality) but also considering monthly standardized anomalies in which seasonality is removed.

In addition, we would like to remark that obtaining high quality and homogeneous time series of the necessary variables for calculating ETo on a daily basis is highly problematic since there are not robust methodologies to homogenize climate variables at the daily scale, whereas testing and correcting homogeneity using monthly records is reliable. Given high agreement of both daily and monthly ETo estimations but stronger robustness and homogeneity of monthly series, it seems recommendable to use monthly climate records in climate change studies.



- As some of the input variables (Rn) have to be estimated from other parameters, some of the discussion about these relationships (p.14) need to be provided earlier in the text.

The estimation of the Rn is the exception in relation to the other variables needed to apply the FAO-56 Penman-Monteith equation. Rn is indirectly estimated by means of sunshine duration records. We agree with the Reviewer#2 and the implications of this method are discussed in depth in the discussion section. Furthermore, in the revised manuscript we have moved some of this discussion to the section of methods.

- While in general there are many graphical illustration for plenty of aspects, I actually miss graphs with the temporal dynamics and developments of input variables into the FAO-56 equation. Where can I see the trend for Rn, T, wind speed, rH? This would be important as they are the controlling variable in the equation.

This is included in Table 7 (Table 8 in the revised manuscript). This table shows the magnitude of change for air temperature, relative humidity, sunshine duration and wind speed in ºC, %, hours and m s^{-1} decade-1, respectively, over the 1961-2013 period. This is analysed for the different available meteorological stations but also for the regional series. In addition, Table 8 also includes the statistical significance of the observed changes at the confidence 95% confidence level.

- Why are authors relating calculated ET0 with variables that have been used to calculate ET0 before (or used to derive inputs from where ET0 is calculated) - see for example Fig. 6. Why

don't authors simply calculate the sensitivities (partial derivatives) of FAO-56 with respect to the driving variables. I simply did that and only from using a temperature increase of 0.6 ◦C (keeping specific water content constant) and some realistic Rn, T , ra, rs – values (I used the original PM formula) I could derive the changes in ET0 stated by the authors. I feel a sensitivity study in this way including trend analysis of the inputs would be more compact and informative for the readers.

In the manuscript we already combined these two suggested approaches. On the one hand, we determined the relationship between ETo calculations and the interannual variability of the different meteorological variables; on the other hand, we also followed the approach using the PM equation, including trend analysis. This was explained in the methods section:

"…we applied the PM equation while holding one variable as stationary (using the average from 1961 to 2013) each time. This approach provided five simulated series of ETo, one per input variable, which could be compared to the ETo series computed with all the data to determine the isolated influence of the five variables. Significant differences between each pair of ETo series (i.e., the original one and the alternative one in which one variable was kept constant) were assessed by comparing the slopes of the linear models, with time as the independent variable. A statistical test for the equality of regression coefficients was used following Paternoster et al. (1998). The significance of the difference was assessed at a 95% confidence level (p<0.05)."

Figures 7 and 8 show the results of this analysis.

- Authors state they applied the Mann-Kendall – did they check and correct for autocorrelation?

We considered the autocorrelation in the trend analysis applied to the series of (i) ETo, (ii) aerodynamic and radiative components of the ETo, and (iii) the series of the different climate variables (i.e., air temperature, relative humidity, wind speed and sunshine duration). This was applied using the FUME R package, which performs the modified Mann-Kendall trend test, returning the corrected p-values after accounting for temporal pseudorreplication (Hamed and Rao, 1998; Ye and Wang, 2004). This has been detailed in the revised manuscript.

Hamed, K.H. and A.R. Rao, (1998). A modified Mann Kendall trend test for autocorrelated data. Journal of Hydrology 204, 182-196.

Yue, S. and C. Wang (2004). The Mann-Kendall Test Modified by Effective Sample Size to Detect Trend in Serially Correlated Hydrological Series. Water Resources Management 18, 201-218.

Overall, I feel there is still a large potential to improve the overall structure/concept of the manuscript as outlined above. As a result I suggest major revisions to the manuscript before publication.

Finally, we thank Reviewer#2 for the revision task and the useful comments raised for improving the results presented in this manuscript. Hopefully we have answered to all these major and minor concerns satisfyingly; otherwise we are available for further clarifications.

Recent changes and drivers of the atmospheric evaporative demand in the Canary Islands

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Abstract

We analysed recent evolution and meteorological drivers of the atmospheric evaporative demand (AED) in the Canary Islands for the period 1961 -2013. We employed long and high quality time series of meteorological variables to analyze current AED changes in this region and found that AED has increased during the investigated period. Overall, the annual ETo, which was estimated by means of the FAO-56 Penman-Monteith equation, increased significantly by 18.2 mm decade^{-1} on average, with a stronger trend in summer (6.7 mm decade^{-1}). ~~The radiative component showed much lower temporal variability than the aerodynamic component did. Thus, more~~In this study we analysed the contribution of (i) the aerodynamic (related to the water vapour that a parcel of air can store) and (ii) radiative (related to the available energy to evaporate a quantity of water) componets to the decadal variability and trends of ETo. More than 90% of the observed ETo variability at the seasonal and annual scales can be associated with the variability of the aerodynamic component. The variable that recorded more significant changes in the Canary Islands was relative humidity, and among the different meteorological factors used to calculate ETo, relative humidity was the main driver of the observed ETo trends. The observed trend could have negative consequences in a number of water-depending sectors if it continues in the future.

Key-words: Reference Evapotranspiration, Aerodynamic component, Radiative component, Temporal changes, Potential Evapotranspiration, Global warming, Canary Islands.

1. Introduction

The atmospheric evaporative demand (AED) is one of the key variables of the hydrological cycle (Wang and Dickinson, 2012), with multiple implications for agriculture, hydrology and the environment (Allen et al., 2015). Several studies have indicated that current global warming is increasing the intensity of the hydrological cycle, mainly as a consequence of an intensification of the AED (Huntington, 2006). Sherwood and Fu (2014) suggested that mechanisms driving the AED over land regions could be the main driver of increasing climate aridity in world semi-arid regions under a global warming scenario.

Warming may play an important role in increasing the AED via the aerodynamic component (McVicar et al., 2012a). Following the Clausius-Clapeyron relationship, the quantity of water vapour that a given mass of air can store increases exponentially with the air temperature. Nevertheless, there are other climate variables whose temporal evolution could compensate the increased AED induced by increasing air temperature, such as wind speed and vapour pressure deficit (McVicar et al., 2012a). In addition, the radiative component of the AED, which is related to the available solar energy that transforms a unit of liquid water into vapour, may compensate or accentuate the increase in AED associated with warming. Wild et al. (2015) noted that solar radiation increased over large regions since the 1980s as a consequence of changes in cloud cover and/or atmospheric aerosol concentrations.

These large number of variables interact in a non-linear manner to determine the AED (McMahon et al., 2013), so assessing recent changes in the AED and defining their determinant factors is not an easy task. For this reason, while several studies analysed the AED at the global scale using different datasets and methods, there is no general consensus on the recent AED evolution (Sheffield et al., 2012; Matsoukas et al., 2011; Wang et al., 2012; Dai, 2013). In this context, the few existing direct AED observations, based on evaporation pans, show a decrease since the 1950s at the global scale (Peterson et al. 1995; Roderick and Farquhar 2002 and 2004), a finding that adds more uncertainty regarding the behaviour of the AED under current global warming. These issues stress the need for new studies that employ high quality datasets to assess the time evolution of the AED at the regional scale.

There are a number of studies published in the last decade that analysed the AED evolution across different regions of the World. Some of them are based on AED estimated using empirical formulations, mostly based on air temperature data (e.g., Thornthwaite, 1958; Hargreaves and Samani, 1995). However, to adequately quantify the AED evolution it is necessary to use long-time series of the meteorological variables that control its radiative and aerodynamic components (e.g. air temperature, vapour pressure deficit and wind speed). Although these variables are generally poorly measured and highly inhomogeneous over both space and time, numerous regional studies analysed the evolution of the AED by means of the robust Penman-Monteith (PM) equation using long times series of these variables. The available regional studies show quite contradictory results, where some studies showed AED negative trends, including those in China (Xu et al., 2006; Ma et al., 2012; Zhang et al., 2007; Liu et al., 2015) and northwest India (Jhajharia et al., 2014). In contrast, other regional studies found positive trends in AED, including those in central India (Darshana et al., 2012), Iran (Kousari and Ahani, 2012; Tabari et al., 2012), Florida (Abtew et al., 2011), continental Spain (Espadafor et al., 2011; Vicente-Serrano et al., 2014a; Azorin-Molina et al., 2015), France (Chaouche et al., 2010) and Moldova (Piticar et al., 2015).

The contrasted trends among world regions would be a consequence of the evolution of the different meteorological variables that control the AED. Specifically, some studies suggest that temporal variability and changes in the AED are related to changes in the relative humidity, mainly in semi-arid regions (Wang et al., 2012; Espadafor et al., 2011; Vicente-Serrano et al., 2014b), whereas others stress the importance of solar radiation (Roderick and Farquhar, 2002; Roderick et al., 2007; Ambas and Baltas, 2012; Fan and Thomas, 2013) or wind speed (McVicar et al., 2012b). Among these studies, few analyzed the AED variability and trends and their possible drivers in the eastern North Atlantic region (Chaouche et al., 2010; Vicente-Serrano et al., 2014a; Azorin-Molina et al., 2015). Nevertheless, there are no studies about this issue in the sub-tropical areas of the north Atlantic region. In this area, there are very few meteorological stations measuring long-term series of the variables necessary to make robust calculations of the AED. This uneven distribution of meteorological observatories constrains the high interest to know the evolution of atmospheric processes in this region, where climate variability is strongly controlled by changes in the Hadley circulation (Hansen et al., 2005) that affects the position and intensity of the subtropical anticyclone belt. Knowing the evolution of AED and its main drivers in this region is highly relevant given the general climate aridity of the region and the low availability of water resources (Custodio and Cabrera, 2002). In this work we analyze the recent evolution and meteorological drivers of the AED in the Canary Islands. ~~The~~The main hypothesis of the study is that in opposition to other continental temperate regions of the North Hemisphere, the warm and humid climate of the subtropical Canary Islands provides the water supply to the atmosphere needed to maintain the AED constant under the current global warming scenarios; consequently, only wind speed and solar radiation could affect the observed decadal variability and trends of the AED. Thus, the availability of long and high quality time series of meteorological variables in the Canary Islands provides an opportunity to analyze current AED changes in the sub-tropical northeastern Atlantic region and the role played by different meteorological variables.

2. Methods

2.1. Dataset

We used the complete meteorological records of the Spanish National Meteorological Agency (AEMET) in the Canary Islands for the following variables at the monthly scale: maximum and minimum air temperature (308 stations), wind speed (99), sunshine duration (42) and mean relative humidity (139). A majority of the stations cover short periods or are affected by large data gaps. As the number of meteorological stations before 1961 was very little for several variables we restricted our analysis to the period between 1961 and 2013. Specifically, only 8 meteorological stations had data gaps of less than 20% of the months in all the necessary variables. As illustrated in Figure 1, these stations are distributed between the Islands of Tenerife (3 stations), Gran Canaria (2), La Palma (1), Lanzarote (1) and Fuerteventura (1). Given that some series included records for a longer period (e.g., Izaña from 1933 and Santa Cruz de Tenerife from 1943), neighbouring stations with shorter temporal coverage were used to reconstruct the existing data gaps in the selected observatories, using a regression-based approach. Details of the site names, coordinates, relocations, data gaps and inhomogeneities of the selected meteorological stations can be found in Table 1.
Then, the time series were subject to quality control and homogenization procedures. The quality control procedure was based on comparison of the rank of each data record with the average rank of the data recorded at adjacent stations (Vicente-Serrano et al., 2010). A relative homogeneity method was applied to identify possible inhomogeneities. For this purpose, we used HOMER (HOMogenization software in R), which compares each candidate series with a number of available series (Mestre et al., 2013). The method provides an estimation of break points in the time series relative to other stations, indicating high probabilities of the presence of inhomogeneities. This method was applied to the different variables and time series following Mestre et al. (2013). Finally, a single regional series for the different variables was obtained using a simple arithmetic average of data values at the available eight stations.

2.2. Calculation of ETo

The Penman-Monteith equation (PM) equation is the standard technique for calculation of ETo from climatic data (Allen et al.,1998), and it is the method officially adopted (with small variations) by the International Commission for Irrigation (ICID), the Food and Agriculture Organization (FAO) of the United Nations, and the American Society of Civil Engineers (ASCE). The PM method can be used globally, and has been widely verified based on lysimeter data from diverse climatic regions (Allen et al., 1994; Itenfisu et al., 2000; López-Urrea et al., 2006). Allen et al. (1998) simplified the PM equation, developing the FAO-56 PM equation, and defined the reference surface as a hypothetical crop with assumed height of 0.12 m, surface resistance of 70 s m^{-1} and albedo of 0.23 that had evaporation similar to that of an extended surface of green grass of uniform height that was actively growing and adequately watered. The ETo FAO-56 PM is expressed as:

$$ET_o = \frac{0.408 \cdot \Delta \cdot (R_n - G) + \gamma \cdot \frac{900}{T + 273} \cdot u_2 \cdot (e_s - e_a)}{\Delta + \gamma \cdot (1 + 0.34 u_2)}$$

$$ET_o = \frac{0.408 \cdot \Delta \cdot (R_n - G) + \gamma \cdot \frac{900}{T + 273} \cdot u_2 \cdot (e_s - e_a)}{\Delta + \gamma \cdot (1 + 0.34 u_2)} \qquad (1)$$

where *ETo* is the reference evapotranspiration (mm day^{-1}), R_n is the net radiation at the crop surface (MJ m^{-2} day^{-1}), G is the soil heat flux density (MJ m^{-2} day^{-1}), T is the mean air temperature at 2 m height (ºC), u_2 is the wind speed at 2 m height (m s^{-1}), e_s is the saturation vapour pressure (kPa), e_a is the actual vapour pressure (kPa), e_s-e_a is the saturation vapour pressure deficit (kPa), Δ is the slope of the vapour pressure curve (kPa ºC^{-1}), and γ is the psychrometric constant (kPa ºC^{-1}). Thus, the monthly ETo can be calculated from data of the monthly averages of five meteorological parameters: maximum and minimum air temperature, relative humidity (which allows calculating the vapour pressure deficit), wind speed at a height of 2 m, and daily sunshine duration (which allows estimating the net radiation). García et al. (2014) compared the capability of sunshine duration series to reconstruct long term radiation in the observatory of Izaña (Tenerife), showing very good temporal agreement between sunshine duration and radiation, independently of the season of the year. Further details on the required equations to obtain the necessary parameters from meteorological data can be consulted in Allen et al. (1998).

We also calculated the evolution of the radiative (Eq.2) and the aerodynamic (Eq.3) components (Eq.3) of the ETo, as follows:

$$ETo(r) = \frac{[0.408\Delta(Rn - G)]}{[\Delta + \gamma(1 + 0.34u_s)]} ETo(r) = \frac{[0.408\Delta(Rn - G)]}{[\Delta + \gamma(1 + 0.34u_s)]}$$

(2)

$$ETo(a) = \frac{\left[\gamma\left(\frac{900}{T+272}\right)u_2(e_s - e_a)\right]}{[\Delta + \gamma(1 + 0.34u_s)]} ETo(a) = \frac{\left[\gamma\left(\frac{900}{T+272}\right)u_2(e_s - e_a)\right]}{[\Delta + \gamma(1 + 0.34u_s)]}$$

(3)

2.3. Analysis

Using the time series of ETo, we determined the seasonal (winter: December–February; spring: March–May; summer: June–August; autumn: September–November) and annual ETo averages. To analyze changes in ETo we used the nonparametric Mann-Kendall statistics that measures the degree to which a trend is consistently increasing or decreasing. The Mann-Kendall statistic is advantageous compared to parametric tests as it is robust to outliers and it does not assume any underlying probability distribution of the data (Zhang et al., 2001). For these reasons, it has been widely used for trend detection in a wide range of hydrological and climatological studies (e.g., Zhang et al., 2001; El Kenawy and McCabe, 2015). To assess the magnitude of changeAutocorrelation was considered in the trend analysis applied to the series of ETo, the series of the aerodynamic and radiative components of the ETo and the series of the different climate variables (temperature, relative humidity, wind speed and sunshine duration). This was applied using the FUME R package, which performs the modified Mann-Kendall trend test, returning the corrected p-values after accounting for temporal pseudorreplication (Hamed and Rao, 1998; Ye and Wang, 2004). To assess the magnitude of change in ETo, we used a linear regression analysis between the series of time (independent variable) and the ETo series (dependent variable). The slope of the regression indicated the amount of change (ETo change per year), with higher slope values indicating greater change. We also calculated the trend observed in the different meteorological variables (air temperature, relative humidity, sunshine duration and wind speed) at both the seasonal and annual scales.

To get insight into the influence of changes in the different meteorological variables on ETo, we related the evolution of ETo with relative humidity, maximum and minimum air temperature, wind speed and sunshine duration by means of correlation analyses. To assess the importance of trends in the different meteorological variables on the observed trends in ETo between 1961 and 2013, we applied the PM equation while holding one variable as stationary (using the average from 1961 to 2013) each time. This approach provided five simulated series of ETo, one per input variable, which could be compared to the ETo series computed with all the data to determine the isolated influence of the five variables. Significant differences between each pair of ETo series (the original one and the alternative one in which one variable was kept constant) were assessed by comparing the slopes of the linear models, with time as the independent variable. A statistical test for the equality of regression coefficients was used (Paternoster et al., 1998). The significance of the difference was assessed at a confidence interval of 95% ($p<0.05$).

3. Results

3.1. Average ETo values

Figure 2 shows a box-plot with the seasonal and annual values of ETo in the different meteorological stations across the Canary Islands, which are also summarized in Table ~~1~~2. There were strong seasonal differences in ETo, as all different meteorological stations show their maximum values in summer and minimum in winter, albeit with strong differences among them. In winter, the highest average values were recorded in the most arid islands (i.e., Fuerteventura and

Lanzarote) and in the station of Los Rodeos (North Tenerife). In summer, the stations of Izaña and Los Rodeos showed the highest average values (663.8 and 612.9 mm, respectively). The lowest summer ETo averages were recorded at the stations of Gran Canaria island (San Cristóbal and Gran Canaria/Airport). At the annual scale, there were very few differences in the average values between the stations of Los Rodeos, Izaña, Fuerteventura and Lanzarote, with very high ETo values ranging between 1693 and 1784 mm (Table ~~1~~2). The observatory with the lowest ETo values is located in Gran Canaria Airport, although the observatory of San Cristóbal (also in the Gran Canaria island) records the minimum values in summer. The magnitude of the differences can be quite important (up to 34%) between the highest ETo values recorded in Los Rodeos, Izaña, Fuerteventura and Lanzarote and the lowest ETo values (Gran Canaria and San Cristóbal). In general, variability, as revealed by the coefficient of variation, was higher in the meteorological stations that recorded the highest ETo values at the annual scale, but there was no clear spatial pattern at the seasonal scale as different stations showed few differences in terms of the coefficients of variation (Table ~~1~~2).

In the majority of weather stations the seasonal and annual ETo magnitude was mostly driven by the aerodynamic component. The average aerodynamic fraction was higher than the radiative fraction in the weather stations that record the highest ETo values (Los Rodeos and Izaña) in all seasons around the year (Figure 3). In other weather stations (Sta. Cruz de Tenerife and San Cristóbal), the ETo associated with the radiative component was much higher than that observed for the aerodynamic component (Table ~~2~~3). The temporal variability in the aerodynamic component was much higher than that observed in the radiative one, regardless of the season of the year or the meteorological station.

3.2. Long-term evolution of ETo

The regional ETo series for the whole Canary Islands (Figure 4) shows a significant increase at the annual scale (18.2 mm decade^{-1}), which is stronger in summer (6.7 mm decade^{-1}) (Table 34). Nevertheless, there was a strong variability between the different meteorological stations, since most meteorological stations experimented significant increases of ETo between 1961 and 2013. The largest annual increase was recorded in Los Rodeos (34.8 mm decade^{-1}), La Palma (29.8 mm decade^{-1}) and Lanzarote (29.7 mm decade^{-1}). Considering a longer period (1933-2013 for Izaña, and 1943-2013 for Santa Cruz de Tenerife), the changes are not statistically significant, although it was not possible to check the homogeneity of the climate records prior to 1961 and thus the results for the longer period must be carefully considered. For the period 1961-2013, there is no general spatial pattern in the observed changes, thus some differences can be observed. For example, in the Gran Canaria island, San Cristóbal station shows a statistically non-significant negative change in ETo on the order of -8.4 mm decade^{-1}, while there is a general significant increase of 28.4 mm decade^{-1} in the Gran Canaria Airport.

Trends in the aerodynamic and radiative components showed clear differences among stations and for the average Canary Islands (Figure 5). Main changes were recorded in the aerodynamic component. The regional series showed an increase of 16.2 mm decade^{-1} in the aerodynamic component, but it only showed an increase of 2 mm decade^{-1} in the radiative component (Table 45). This can be translated to an average increase in the ETo of 89% over the whole period due to changes in the aerodynamic component, and of 11% due to changes in the radiative component. However, there are spatial differences between the meteorological stations, since the aerodynamic component showed a decrease of 21 mm decade^{-1} in San Cristóbal, compared to an increase of 44.6 mm decade^{-1} in Los Rodeos. On the contrary, the radiative component showed lower differences among stations, with values ranging from -9.9 mm decade^{-1} in Los Rodeos to 12.7 mm decade^{-1} in San Cristóbal. Nevertheless, and regardless of the observed trends, the results indicate that the inter-annual variability of ETo between 1961 and 2013 was mainly driven by the aerodynamic component, independently of the season or the meteorological station considered (Table ~~5~~6). The temporal correlation between ETo and the aerodynamic component was statistically significant for the different meteorological stations in the seasonal and the annual series, with correlation coefficients higher than 0.95 in most cases. The correlation for the regional series was also strong and statistically significant. In contrast, the correlation coefficients calculated between ETo and the radiative component were much lower, and generally non-significant ($p<0.05$). Los Rodeos is the unique weather station where the correlation between ETo and the radiative component was statistically significant at both the seasonal and annual scales, but showing a negative correlation. Overall, the results show that the correlation between the annual radiative component and the total annual regional series of ETo is statistically non-significant.

3.3. Drivers of ETo variability and trends

Table ~~6~~7 shows the correlation between the different meteorological variables and ETo at the seasonal and annual scales in the eight meteorological stations. Maximum and minimum air temperatures were positively correlated with ETo and this relationship was statistically significant in some stations, and the correlation coefficients tended to be higher for maximum air temperature. In Los Rodeos and La Palma, the ETo variability could not be explained by the variability in air temperature, with correlation coefficients weaker than 0.3. Overall, the results indicate that the seasonal and annual series of ETo were significantly correlated with variations of sunshine duration and wind speed, suggesting that these two variables are the key drivers of ETo variability in the Canary Islands. The variable that showed the strongest correlation with the evolution of ETo in the seasonal and annual series of the different meteorological observatories was relative humidity, with negative coefficients. Only in the annual series of Santa Cruz de Tenerife the correlation was non-significant. Moreover, there were no significant differences in the magnitude of correlations among seasons.

The regional series summarise the pattern observed in the individual meteorological stations (Figure 6). In winter, relative humidity had the strongest correlation with ETo (r=-0.85), with a mostly linear relationship. Minimum air temperature and sunshine duration showed significant positive correlations with ETo (r=0.40 and 0.36, respectively). Maximum air temperature and wind speed showed weaker correlation with the winter ETo. In spring, the magnitude of the correlations was similar among the different variables, and the highest correlation corresponded again to relative humidity (r=-0.72). A similar pattern was found in summer, where relative humidity showed the strongest correlation (r=-0.74) followed by maximum and minimum air temperature. In autumn, relative humidity also showed the strongest correlation and wind speed showed more importance than both maximum and minimum air temperature. As expected, relative humidity showed the strongest correlation with ETo (r = -0.83) at the annual scale, followed by wind speed (r = 0.62). On the contrary, the correlation with maximum air temperature was statistically non-significant.

The general increase observed in ETo in the Canary Islands was largely determined by changes in the different meteorological variables (Table ~~7~~8). The maximum air temperature does not show noticeable changes, with the exception of Gran Canaria/Airport, Lanzarote and San Cristóbal stations where significant trends were found. The regional average did not show significant changes. On the contrary, the minimum air temperature showed an average increase of 0.12 ºC decade^{-1} in summer and 0.09 ºC decade^{-1} at the annual scale between 1961 and 2013. The significant increase recorded in summer was found in six meteorological stations, with a maximum of 0.25º C decade^{-1} in Izaña. Changes in relative humidity were also significant. There was a significant decrease in winter, summer and annually, which represent a decline of 0.47% decade^{-1}, although there were differences among stations. Sunshine duration and wind speed did not show noticeable changes, and the unique remarkable pattern was the significant increase of the summer sunshine duration at the regional scale (0.12 hours decade^{-1}) and the significant increase of wind speed in the station of Los Rodeos in the four seasons and also annually.

With respect to the sensitivity of changes in ETo to its five driving meteorological drivers (Figure 7), substantial differences were found between variables. The differences between observed ETo and simulated ETo with average maximum and minimum air temperature were small irrespective of the season, indicating a low sensitivity to these two variables. In contrast, ETo was more sensitive to setting sunshine duration and wind speed at their mean values. Thus, in the station of Los Rodeos, the predicted magnitude of change in winter, autumn and annually was different from the observed magnitude of change. The highest sensitivity was, however, to relative humidity. In general, the different meteorological stations showed an important increase in observed ETo with respect to predicted ETo keeping relative humidity as constant. This was observed at the seasonal and annual scales. Thus, in three meteorological stations the observed magnitude of change on annual basis is between two and three times higher than that predicted considering relative humidity as stationary. This pattern was also found in the regional series (Figure 8). Considering air temperature, sunshine duration and wind speed as constant, there were no statistical differences between the observed and predicted magnitudes of change, both seasonally and annually. On the contrary, leaving relative humidity as constant, the magnitude of the trend was quite different to the observations, and temporal trends would not be statistically significant. Thus, the magnitude of change of ETo, considering relative humidity as constant, is significantly different from the ~~the~~ observed magnitude of change in winter and annually.

4. Discussion

This work analyses the recent evolution (1961-2013) of reference evapotranspiration (ETo) in the Canary Islands and its relationship with the evolution of its atmospheric drivers. We analysed the time evolution of ETo in eight meteorological stations in which the necessary meteorological variables for calculation of the ETo were available. The results showed a general increase in ETo, although different magnitudes of change were found between the different meteorological stations.

These differences did not follow any specific geographic pattern, so they must be considered either due to random effects and uncertainty at various levels or due to micro-geographic effects that were not considered in this study. There is not a general pattern that may connect the observed trends in a certain forcing variable with the observed trend of ETo in each of the eight analysed stations although those that showed a higher increase in ETo (i.e., Lanzarote, Los Rodeos and Gran Canaria) displayed a higher increase in the aerodynamic component; a process which is in agreement with the significant reductions observed in relative humidity.

Nevertheless, with the exception of the observatory of San Cristóbal in the north of Gran Canaria Island, other meteorological observatories showed positive changes in ETo, with annual trends statistically significant in six stations. In any case, we must also stress that trends in ETo at the regional scale are mostly significant because of the low values in the beginning of the study period starting in the 1960s. Thus, the results of the two sites with longer temporal coverage (i.e., Izaña and Santa Cruz de Tenerife) do not show significant trends. This makes necessary to consider these trends with caution since they could be driven by variability processes at the decadal scale.

The few existing studies in Northwest Africa (Ouysse et al., 2010; Teken and Kropp, 2012) are not comparable with our findings, since the variables required to apply the Penman-Monteith equation were not available. Instead, these studies relied on simplified methods that just employ air temperature records. Despite the difference in methods, these studies also found a general increase in the ETo. The closest region in which it is possible to make a direct comparison using the same method is the Iberian Peninsula, where a general increase of 24.5 mm decade^{-1} was found between 1961 and 2011 (Vicente-Serrano et al., 2014a). This study also found that the variability and trends in the aerodynamic component determined most of the observed variability and the magnitude of change of ETo in a majority of the meteorological stations in the Iberian Peninsula. The radiative component showed much lower temporal variability than the aerodynamic component did. Thus, more than 90% of the observed ETo variability at the seasonal and annual scales can be associated with the variability of the aerodynamic component. This is in agreement with the results obtained in previous studies. For example, Wang et al. (2012) showed that recent ETo variability at the global scale was mainly driven by the aerodynamic component. Equally, other studies in Southern Europe indicated a higher importance of the aerodynamic component (Sanchez-Lorenzo et al., 2014; Azorin-Molina et al., 2015). It could be argued, however, that quantification of the radiative component in our study was based on a simplified assumption since it was calculated from sunshine duration that is mostly determined by the cloud coverage (Hoyt, 1978). Nevertheless, it is also worth noting that global radiation measurements, sunshine duration records contain a signal of the direct effects of aerosols (Sanroma et al., 2010; Sanchez-Romero et al., 2014; Wild, 2015) in the Canary Islands. Nevertheless, the Canary Islands is a region mostly free of anthropogenic aerosols given the large frequency and intensity of trade winds (Mazorra et al., 2007), and it is not expected that the frequency of Saharan dust events, that could affect incoming solar radiation, has noticeably changed over the last decades (Flentje et al., 2015; Laken et al., 2015). Consequently, in the Canary Islands we can consider high accuracy determining the radiative component using sunshine duration series. ~~García et al. (2014) compared the capability of sunshine duration series to reconstruct long term radiation in the observatory of Izaña (Tenerife), showing very good temporal agreement between sunshine duration and radiation, independently of the season of the year.~~ In continental Spain, Azorin-Molina et al. (2015) also found strong positive correlations between interannual variations of solar radiation and sunshine duration in different meteorological stations. Overall, in the Canary Islands there is a positive and significant correlation between inter-annual variations of ETo and sunshine duration, although this correlation did not explain the observed trends of ETo in the region.

We showed that the temporal variability of ETo is strongly controlled by the temporal variability of relative humidity. Specifically, seasonal and annual series of ETo in the different stations showed very strong negative and significant correlations with those of the relative humidity. Thus, the magnitude of correlations were much higher than those obtained for other meteorological variables, and this finding was common to the whole set of meteorological stations. This strong control of relative humidity on the temporal variability of ETo has been already identified in some studies in the Iberian Peninsula (Vicente-Serrano et al., 2014b; Azorin-Molina et al., 2015; Espadafor et al., 2013).

Among the variables that control the aerodynamic component, wind speed and maximum air temperature did not show significant trends at the regional scale and only few stations recorded significant trends in these variables, either at the seasonal or the annual scales. Significant trends were obtained for minimum air temperature, mainly in summer. Recently, Croper and Hanna (2014) analysed long term climate trends in the Macaronesia region, and for the Canary Islands they showed an increase in air temperature during summer for the period 1981-2010. Martín et al. (2012) analysed air temperature changes in the Tenerife Island from 1944 to 2010 and they also showed that night-time air temperature increased rapidly compared to daytime temperature. Nevertheless, they found strong spatial contrasts between the high mountains, that showed a higher increase, and the coastal areas in which the air temperature regulation of the ocean could be reducing the general air temperature increase.

In any case, the variable that recorded more significant changes in the Canary Islands was relative humidity, and among the different meteorological variables used to calculate ETo, relative humidity was the main driver of the observed ETo trends. Significant negative humidity trends were recorded in winter, summer and autumn, but also annually. Thus, simulation of ETo series considering the different meteorological variables as constant produced few differences in relation to the observed evolution of ETo, with the exception of the relative humidity. Leaving relative humidity as constant for the period 1961-2013 showed no significant ETo changes at seasonal and annual scales and also statistically significant differences with changes obtained from observations. In continental Spain, Vicente-Serrano et al. (2014b) showed a general decrease of relative humidity from the decade of

1960, mainly associated with a general decrease of the moisture transport to the Iberian Peninsula as well as a certain precipitation decrease. Similarly, Espadafor et al. (2011) and Vicente-Serrano et al. (2014b) showed that the strong increase in ETo in the last decades is associated with the relative humidity decrease due to air temperature rise~~.~~, which caused more severe drought events (Coll et al., 2016; Lorenzo-Lacruz and Morán-Tejeda, 2016; Peña-Gallardo et al., 2016). In the Canary Islands, no precipitation changes have been identified during the analyzed period (Sánchez-Benitez et al., ~~2015~~2016). Therefore a lower moisture supply from the humidity sources to the islands should explain the observed pattern toward a relative humidity decrease. Sherwood and Fu (2014) suggested that differences in the air temperature increase between oceanic and continental areas could increase land aridity, as a consequence of the sub-saturation conditions of the oceanic air masses that come to the land areas, given higher warming rates in maritime regions in comparison to continental areas. The results of this study confirm this pattern in the Canary Islands, since this region should not be constrained by constant moisture supply from the surrounding warm Atlantic Ocean. Overall, Willett et al. (2014) recently found a general decrease in relative humidity at the global scale, including several islands and coastal regions in which the moisture supply was expected to be unlimited. This finding suggests that contrasted mean air temperature and trends between land and ocean areas could also play an important role in explaining this phenomenon, even at local scales.

5. Conclusions

We found that the reference evapotranspiration ETo increased by 18.2 mm decade^{-1} -on average- between 1961 and 2013 over the Canary Islands, with the highest increase recorded during summer. Although there were noticeable spatial differences, this increase was mainly driven by changes in the aerodynamic component, caused by a statistically significant reduction of the relative humidity.

This study provides an outstanding example of how climate change and interactions between different meteorological variables drive an increase of the ETo event in a subtropical North Atlantic Islands. Given the general aridity conditions in most of the Canary Islands and the scarcity of water resources, the observed trend could have negative consequences in a number of water-depending sectors if it continues in the future.

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References

Abtew,W., Obeysekera, J., Iricanin, N., (2011): Pan evaporation and potential evapotranspiration trends in South Florida. Hydrol. Process. 25, 958–969.

Allen, R. G. L. S. Pereira, D. Raes, and M. Smith (1998), Crop evapotranspiration: Guidelines for computing crop water requirements, Food and Agricultural Organization (FAO) Irrig. Drain. pap. 56, Rome.

Allen, C.D., Breshears, D., McDowell, N.G., (2015): On underestimation of global vulnerability to tree mortality and forest die-off from hotter drought in the Anthropocene. Ecosphere, 6: 129.

Ambas, V. T., and E. Baltas (2012), Sensitivity analysis of different evapotranspiration methods using a new sensitivity coefficient, Global Nest J., 14, 335–343.

Azorin-Molina, C. Vicente-Serrano, S.M. ; Arturo Sanchez-Lorenzo; Tim R. McVicar ; Enrique Morán-Tejeda ; Jesus Revuelto ; Ahmed El Kenawy, Natalia Martín-Hernández ; Miquel Tomas-Burguera. Atmospheric evaporative demand observations, estimates and driving factors in Spain (1961-2011). Journal of Hydrology 523: 262-277.

Chaouche, K., Neppel, L., Dieulin, C., Pujol, N., Ladouche, B., Martin, E., Salas, D., Caballero, Y., 2010. Analyses of precipitation, temperature and evapotranspiration in a French Mediterranean region in the context of climate change. Compt. Rendus Geosci. 342, 234–243.

Coll, J.R., Aguilar, E., Prohom, M., Sigró, J., (2016): Long-term drought variability and trends in Barcelona (1787-2014). Cuadernos de Investigación Geográfica, 42, DOI: 10.18172/cig.2927
Cropper, T.E., Hanna, E., (2014): An analysis of the climate of Macaronesia, 1865-2012. International Journal of Climatology, 34: 604-622.
Custodio, E., Cabrera, M.C. (2002): ¿Cómo convivir con la escasez de agua? El caso de las Islas Canarias. Boletin Geologico y Minero 113: 243-258.
Dai, A., (2013): Increasing drought under global warming in observations and models. Nature Climate Change 3, 52–58.
Darshana, A., Pandey, R., Pandey, P., (2012): Analysing trends in reference evapotranspiration and weather variables in the Tons River Basin in Central India. Stoch. Env. Res. Risk A. http://dx.doi.org/10.1007/s00477-012-0677-7.
El Kenawy A and McCabe MF (2015) A multi-decadal assessment of the performance of gauge- and model-based rainfall products over Saudi Arabia: climatology, anomalies and trends, Int. J. of Climatol., doi:10.1002/joc.4374.
Espadafor, M., Lorite, I.J., Gavilán, P., Berengena, J., (2011): An analysis of the tendency of reference evapotranspiration estimates and other climate variables during the last 45 years in Southern Spain. Agric. Water Manag. 98, 1045–1061.
Fan, Z.-X., and A. Thomas (2013), Spatiotemporal variability of reference evapotranspiration and its contributing climatic factors in Yunnan Province, SW China, 1961–2004, Clim. Change, 116, 309–325.
Flentje, H., Briel, B., Beck, C. et al. (2015): Identification and monitoring of Saharan dust: An inventory representative for south Germany since 1997. Atmospheric Environment 109: 87-96.
García, R.D., et al. (2014): Reconstruction of global solar radiation time series from 1933 to 2013 at the Izaña Atmospheric Observatory. Atmospheric Measurement Techniques 7: 3139-3150.
Hamed, K.H. and A.R. Rao, (1998). A modified Mann Kendall trend test for autocorrelated data. Journal of Hydrology 204, 182-196.
Hargreaves, G.L., Samani, Z.A., (1985): Reference crop evapotranspiration from temperature. Appl. Eng. Agric. 1, 96–99.
Hoyt, D.V., (1978): Interannual Cloud-Cover Variations in the Contiguous United States. J. Appl. Meteor., 17, 354–357.
Huntington, T.G., (2006): Evidence for intensification of the global water cycle: Review and synthesis. Journal of Hydrology 319: 83-95.
Itenfisu, D., Elliott, R.L., Allen, R.G., Walter, I.A., 2000. Comparison of Reference Evapotranspiration Calculations across a Range of Climates. Proceedings of the 4th National Irrigation Symposium. ASAE, Phoenix, AZ.
Jhajharia, D., Kumar, R., Dabral, P. P., Singh, V. P., Choudhary, R. R. and Dinpashoh, Y. (2015), Reference evapotranspiration under changing climate over the Thar Desert in India. Met. Apps, 22: 425–435. doi: 10.1002/met.1471.
Kousari, M.R., Ahani, H., (2012): An investigation on reference crop evapotranspiration trend from 1975 to 2005 in Iran. Int. J. Climatol. 32, 2387–2402.
Laken, B.A., Parviainen, H., García-Gil, A., Muñoz-Tuñón, C., Varela, A.M., Fernandez-Acosta, S., Pallé, P., (2015): Thirty years of atmospheric extinction from telescopes of the North Atlantic Canary Archipelago. Journal of Climate. doi: http://dx.doi.org/10.1175/JCLI-D-14-00600.1
Liu, T., Li, L., Lai, J., Liu, C., Zhuang, W. (2015): Reference evapotranspiration change and its sensitivity to climate variables in southwest China. Theoretical and Applied Climatology. In press.

López-Urrea, R., F. Martín de Santa Olalla, C. Fabeiro, and A. Moratalla (2006), Testing evapotranspiration equations using lysimeter observations in a semiarid climate, Agric. Water Manage., 85, 15–26.

Lorenzo-Lacruz, J., Morán-Tejeda, E., (2016): Spatio-temporal patterns of meteorological droughts in the Balearic Islands (Spain). Cuadernos de Investigación Geográfica, 42, DOI: 10.18172/cig.2948.

Ma, X., Zhang, M., Li, Y.,Wang, S., Ma, Q., Liu, W., (2012): Decreasing potential evapotranspiration in the Huanghe River Watershed in climate warming during 1960-2010. J. Geogr. Sci. 22, 977–988.

Martín, J.L., Bethencourt, J., Cuevas-Agulló, E., (2012): Assessment of global warming on the island of Tenerife, Canary Islands (Spain). Trends in minimum, maximum and mean temperatures since 1944. Climatic Change, 114: 343-355.

Matsoukas, C., N. Benas, N. Hatzianastassiou, K. G. Pavlakis, M. Kanakidou, and I. Vardavas (2011), Potential evaporation trends over land between 1983–2008: Driven by radiative fluxes or vapour-pressure deficit?, Atmos. Chem. Phys., 11, 7601–7616.

Mazorra, L., Diaz, F., Navarro, P., Deniz, F. (2007): Accumulated frequency estimation for daily clearness index. ISES Solar World Congress 2007, ISES 20074, pp. 2632-2635.

Mestre O, Domonkos P, Picard F, Auer I, Robin S, Lebarbier E, Böhm R, Aguilar E, Guijarro J, Vertacnik G, Klancar M, Dubuisson B, Stepanek P (2013) HOMER: HOMogenisation softwarE in R- methods and applications. Idöjárás 117: 47-67.

McMahon, T.A., Peel, M.C., Lowe, L., Srikanthan, R., McVicar, T.R., (2013): Estimating actual, potential, reference crop and pan evaporation using standard meteorological data: a pragmatic synthesis. Hydrol. Earth Syst. Sci. 17 (1), 1331–1363.

McVicar, T.R., Roderick, M.L., Donohue, R.J., Van Niel, T.G., (2012a): Less bluster ahead? ecohydrological implications of global trends of terrestrial near-surface wind speeds. Ecohydrology 5 (4), 381–388.

McVicar, T.R., Roderick, M.L., Donohue, R.J., et al., (2012b): Global review and synthesis of trends in observed terrestrial near-surface wind speeds: implications for evaporation. J. Hydrol. 416–417, 182–205.

Ouysse, S., Laftouhi, N.-E., Tajeddine, K., (2010): Evaluation of evapotranspiration variation in the Draa basin using statistical and empirical methods (South-Eastern Morocco). XXXVIIIIAH Congress Groundwater Quality Sustainability. Krakow, 12–17 September 2010.

Paternoster, R., Brame, R., Mazerolle, P., and Piquero, A. R. (1998). Using the Correct Statistical Test for the Equality of Regression Coefficients. Criminology, 36(4), 859–866.

Peña-Gallardo, M., Gámiz-Fortis, S.R., Castro-Díez, Y., Esteban-Parra, M.J., (2016): Análisis comparativo de índices de sequía en Andalucía para el periodo 1901-2012. Cuadernos de Investigación Geográfica, 42, DOI: 10.18172/cig.2946.

Peterson, T. C., V. S. Golubev, and P. Y. Groisman (1995), Evaporation losing its strength, Nature, 377, 687–688.

Piticar, A., Mihăilă, D., Lazurca, L.G., et al. (2015): Spatiotemporal distribution of reference evapotranspiration in the Republic of Moldova. Theoretical and Applied Climatology. In Press.

Roderick, M. L., and G. D. Farquhar (2002), The cause of decreased pan evaporation over the past 50 years, Science, 298, 1410–1411.

Roderick, M. L., and G. D. Farquhar (2004), Changes in Australian pan evaporation from 1970 to 2002, Int. J. Climatol., 24, 1077–1090.

Sánchez-Benítez, A., García-Herrera, R., Vicente-Serrano, S.M., ~~García-Herrera, R. (2015). Precipitation~~ (2016). Revisiting precipitation variability, trends and ~~influence of atmospheric circulation~~drivers in the Canary Islands~~ (Spain). In preparison.~~. Submitted to the International Journal of Climatology

Sanchez-Lorenzo, A., Vicente-Serrano, S.M., Wild, M., Calbó, J., Azorin-Molina, C., Peñuelas, J., (2014) Evaporation trends in Spain: a comparison of Class A pan and Piché atmometer measurements. Climate Research. 61: 269–280.

Sanchez-Romero, A., A. Sanchez-Lorenzo, J. Calbó, J. A. González, and C. Azorin Molina (2014), The signal of aerosolinduced changes in sunshine durationrecords: A review of the evidence, J. Geophys. Res. Atmos., 119, doi:10.1002/2013JD021393.

Sanroma, E., Palle, E., and Sanchez-Lorenzo, A., (2010): Long-term changes in insolation and temperatures at different altitudes. Environmental Research Letters, 5, 2.

Sheffield, J., Wood, E.J., Roderick, M.L., (2012): Little change in global drought over the past 60 years. Nature 491, 435–438.

Sherwood, S., Fu, Q. (2014): A drier future? Science 343: 737-739.

Tabari, H., Nikbakht, J., Talaee, P.H., 2012. Identification of trend in reference evapotranspiration series with serial dependence in Iran.Water Resour. Manag. 26, 2219–2232.

Tekken, V., Kropp, J.P. (2012): Climate-driven or human-induced: Indicating severe water scarcity in the Moulouya river basin (Morocco). Water, 4: 959-982.

Thornthwaite, C.W., 1948. An approach toward a rational classification of climate. Geogr. Rev. 38, 55–94.

Vicente-Serrano, S.M., Beguería, S., Juan I. López-Moreno, Miguel A. García-Vera y P. Stepanek. (2010): A complete daily precipitation database for North-East Spain: reconstruction, quality control and homogeneity. International Journal of Climatology. 30, 1146-1163.

Vicente-Serrano, S.M., Cesar Azorin-Molina, Arturo Sanchez-Lorenzo, Jesús Revuelto, Juan I. López-Moreno, José C. González-Hidalgo, Francisco Espejo. (2014a) Reference evapotranspiration variability and trends in Spain, 1961–2011. Global and Planetary Change, 121: 26-40.

Vicente-Serrano, S.M., Cesar Azorin-Molina, Arturo Sanchez-Lorenzo, Enrique Morán-Tejeda, Jorge Lorenzo-Lacruz, Jesús Revuelto, Juan I. López-Moreno, Francisco Espejo (2014b): Temporal evolution of surface humidity in Spain: recent trends and possible physical mechanisms. Climate Dynamics. 42:2655–2674

Wang, K., Dickinson, R.E., (2012): A review of global terrestrial evapotranspiration: observation,modeling, climatology, and climatic variability. Rev. Geophys. 50. http://dx.doi.org/10.1029/2011RG000373.

Wang, K., Dickinson, R.E., Liang, S., 2012. Global atmospheric evaporative demand over land from 1973 to 2008. J. Clim. 25 (23), 8353–8361.

Wild, M. (2015): Decadal changes in radiative fluxes at land and ocean surfaces and their relevance for global warming. WIREs Clim Change. doi: 10.1002/wcc.372.

Willett, K.M., et al. (2014): HadISDH land surface multi-variable humidity and temperature record for climate monitoring. Climate of the Past 10: 1983-2006.

Xu, Hh.-Y.u., Gong, L., Jiang, T., Chen, D., Singh, V.P., (2006): Analysis of spatial distribution and temporal trend of reference evapotranspiration and pan evaporation in Changjiang (Yangtze River) catchment. J. Hydrol. 327, 81–93.

Yue, S. and C. Wang (2004). The Mann-Kendall Test Modified by Effective Sample Size to Detect Trend in Serially Correlated Hydrological Series. Water Resources Management 18, 201-218.

Zhang X, Harvey KD, Hogg WD, Yuzyk TR (2001) Trends in Canadian streamflow, Water Resources Research, 37(4), 987-998.

Zhang, Y., Liu, C., Tang, Y., Yang, Y., (2007): Trends in pan evaporation and reference and actual evapotranspiration across the Tibetan Plateau. J. Geophys. Res. D: Atmos. 112 (Article numberD12110).

~~Table 1~~

Table 1: Site names, coordinates, relocations, data gaps and inhomogeneities of the selected meteorological stations in the Canary Islands

Code	Longitude	Latitude	Name	relocation	Relative humidity		Sunshine duration		Wind speed		maximum temperature		minimum temperature	
					data gaps	Inhom.	data gaps	Inhom.	data gaps	Inhom.	data gaps	Inhom.	data gaps	Inhom.
C029O	-13.60	28.95	Lanzarote/Airport	1972	2.20%	1998	0.78%	1978-2002	0.47%	1971	1.23%	2004	1.23%	1988
C139E	-17.75	28.61	La Palma/Airport	1970	0.94%		2.51%		0.47%	1976	0.37%		0.37%	1997
C249I	-13.85	28.43	Fuerteventura/Airport	1969	0.15%	2000	1.25%	1995-2005	0.15%		0.23%	1983	0.23%	1977
C430E	-16.48	28.30	Izaña	--	1.72%	1999	7.40%	2005	6.91%		5.20%	1985	5.20%	
C447A	-16.31	28.46	Los Rodeos	--	0.31%		1.10%	1966	0.15%	1970	0.30%	2005	0.30%	2005
C449C	-16.25	28.45	Santa Cruz de Tenerife	--	0%		0.94%		0%	1987	0%		0%	1994
C649I	-15.38	27.91	Gran canaria/Airport	--	0.15%	1981-1994	2.67%	1978	0.31%	1972	0.20%	1984	0.20%	1994
C659P	-15.41	28.15	San Cristóbal	1994	11%		1.88%	1980	10.50%	1994	5.30%	1966	5.30%	

Table 2: Seasonal and annual averages (mm) and coefficients of variation of ETo in the eight meteorological stations, averaged over the period 1961-2013.

	Sta. Cruz de Tenerife	Gran Canaria/Airp.	Los Rodeos	Izaña	Fuerteventura	La Palma	Lanzarote	San Cristóbal	Regional Mean
	Average								
Winter	222.0	181.5	297.5	250.2	298.1	251.6	294.5	217.7	251.6
Spring	390.1	302.2	468.8	414.1	460.8	361.5	468.7	342.3	401.1
Summer	512.7	415.5	612.9	663.8	560.2	438.7	586.1	383.0	521.6
Autumn	311.8	273.9	401.8	364.5	384.6	316.4	393.8	278.8	340.7
Annual	1435.5	1175.0	1784.4	1692.6	1702.0	1372.7	1741.0	1219.4	1515.3
	Coefficient of variation								
Winter	0.05	0.11	0.12	0.18	0.10	0.11	0.09	0.11	0.06
Spring	0.04	0.10	0.07	0.12	0.08	0.10	0.06	0.08	0.05
Summer	0.03	0.12	0.07	0.07	0.07	0.08	0.07	0.07	0.04
Autumn	0.03	0.10	0.10	0.10	0.07	0.11	0.07	0.08	0.05
Annual	0.02	0.07	0.06	0.07	0.07	0.08	0.06	0.05	0.04

Table ~~2~~3: Seasonal and annual averages (mm) and coefficients of variation of aerodynamic and radiative components of ETo in the eight meteorological stations.In bold the values greater than 50% of the total ETo of the station

	Aerodynamic								
	Sta. Cruz de Tenerife	Gran Canaria/Airp.	Los Rodeos	Izaña	Fuerteventura	La Palma	Lanzarote	San Cristóbal	Mean
	Average								
Winter	101.6	**98.8**	**198.8**	**198.8**	**195.9**	**153.2**	**190.4**	108.1	**155.7**
Spring	130.5	137.5	**287.2**	**271.0**	**251.1**	174.3	**262.0**	134.7	**206.0**
Summer	146.2	195.6	**394.7**	**424.7**	**288.5**	201.7	**328.1**	143.1	**265.3**
Autumn	109.3	133.4	**249.1**	**263.6**	**211.7**	157.6	**225.9**	102.0	**181.6**
Annual	487.5	568.0	**1134.4**	**1158.6**	**945.8**	**690.7**	**1004.4**	485.5	**809.4**
	Coefficient of variation								
Winter	0.12	0.19	0.22	0.23	0.18	0.19	0.16	0.27	0.11
Spring	0.11	0.18	0.15	0.17	0.16	0.20	0.12	0.26	0.09
Summer	0.13	0.24	0.12	0.14	0.15	0.18	0.12	0.20	0.08
Autumn	0.13	0.21	0.20	0.14	0.14	0.20	0.15	0.25	0.10
Annual	0.09	0.16	0.13	0.12	0.14	0.16	0.11	0.17	0.07
	Radiative								
	Sta. Cruz de Tenerife	Gran Canaria/Airp.	Los Rodeos	Izaña	Fuerteventura	La Palma	Lanzarote	San Cristóbal	Average
	Average								
Winter	**120.4**	82.7	98.6	51.4	102.2	98.4	104.1	**109.6**	95.9
Spring	**259.7**	**164.7**	181.5	143.1	209.7	**187.2**	206.7	**207.6**	195.0
Summer	**366.5**	**220.0**	218.3	239.1	271.7	**237.0**	258.0	**240.0**	256.3
Autumn	**202.4**	**140.5**	152.8	100.9	172.9	**158.8**	167.9	**176.8**	159.1
Annual	**948.1**	**607.0**	650.0	534.0	756.3	682.0	736.7	**734.0**	706.0
	Coefficient of variation								
Winter	0.05	0.08	0.10	0.12	0.08	0.08	0.09	0.08	0.06
Spring	0.06	0.07	0.08	0.09	0.06	0.07	0.06	0.08	0.05
Summer	0.04	0.06	0.07	0.08	0.05	0.09	0.06	0.10	0.04
Autumn	0.05	0.05	0.08	0.07	0.05	0.06	0.06	0.06	0.04
Annual	0.03	0.04	0.07	0.06	0.04	0.05	0.04	0.06	0.03

Table ~~3~~4: Magnitude of change (mm. decade^{-1}) of ETo in each meteorological station and the average of the eight stations over the period 1961-2013. Statistically significant at the 95% confidence level are given in bold. Numbers between brackets refer to the magnitudes of change for the periods 1933-2013 for Izaña and 1943-2013 for Santa Cruz de Tenerife.

	Sta. Cruz de Tenerife	Gran Canaria/Airp.	Los Rodeos	Izaña	Fuerteventura	La Palma	Lanzarote	San Cristóbal	Mean
Winter	**2.7**(0.31)	1.7	**11.3**	4.8 (-0.42)	3.2	**9.1**	**7.1**	**-5.1**	**4.3**
Spring	0.1 (-0.55)	**7.7**	**7.1**	-0.1 (-1.27)	3.9	**7.2**	4.0	**-5.8**	3.0
Summer	1.1 (-1.36)	**16.0**	**7.6**	6.0 (-0.64)	0.0	**7.7**	**10.1**	**5.0**	**6.7**
Autumn	**2.0**(0.62)	3.6	**11.2**	3.7 (0.30)	-0.2	**9.9**	4.8	**-5.0**	**3.8**
Annual	**7.3**(-1.95)	**28.4**	**34.8**	14.9 (-0.67)	9.2	**29.8**	**29.7**	-8.4	**18.2**

Table 45: Magnitude of change (mm. decade^{-1}) of both aerodynamic and radiative components of ETo in each meteorological station and the average of the eight stations over the period 1961-2013. Statistically significant at the 95% confidence level are given in bold. Numbers between brackets refer to the magnitudes of change for the periods 1933-2013 for Izaña and 1943-2013 for Santa Cruz de Tenerife.

	Sta. Cruz de Tenerife	Gran Canaria/Airp.	Los Rodeos	Izaña	Fuerteventura	La Palma	Lanzarote	San Cristóbal	**Mean**
Aerodynamic									
Winter	**3.7** (0.09)	2.9	**14.8**	5.1 (-0.96)	4.6	**10.1**	**9.1**	**-5.8**	**5.5**
Spring	-1.3 (-1.84)	**7.8**	**8.9**	0.1 (-3.39)	2.4	3.3	2.7	**-11.8**	1.5
Summer	0.1 (-2.95)	**16.8**	**9.9**	6.7 (-3.38)	-1.1	2.5	**8.1**	-1.5	**5.2**
Autumn	2.1 (-0.51)	**5.2**	**14.5**	3.7 (-1.03)	-1.1	**7.9**	4.6	-3.8	**4.1**
Annual	4.7 (-6.25)	**31.2**	**44.6**	15.6 (-6.93)	6.5	19.8	**28.0**	**-21.2**	**16.2**
Radiative									
Winter	-1.0 (0.22)	**-1.2**	**-3.5**	-0.4 (0.51)	-1.4	-1.0	-2.0	0.8	**-1.2**
Spring	1.4 (1.28)	-0.1	-1.8	-0.3 (2.12)	1.4	**3.9**	1.3	**6.1**	1.5
Summer	1.0 (1.58)	-0.8	-2.3	-0.7 (2.74)	1.1	**5.1**	2.0	**6.5**	1.5
Autumn	0.0 (1.13)	**-1.6**	**-3.3**	0.1 (1.34)	0.9	**2.0**	0.2	-1.2	-0.4
Annual	2.7 (4.29)	-2.8	**-9.9**	-0.7 (6.25)	2.7	**10.0**	1.7	**12.7**	2.0

Table ~~5~~6. Seasonal and annual Pearson's coefficients between the evolution of ETo and the evolution of aerodynamic and radiative components in the eight meteorological stations and the average. Statistically significant at the 95% confidence level are given in bold

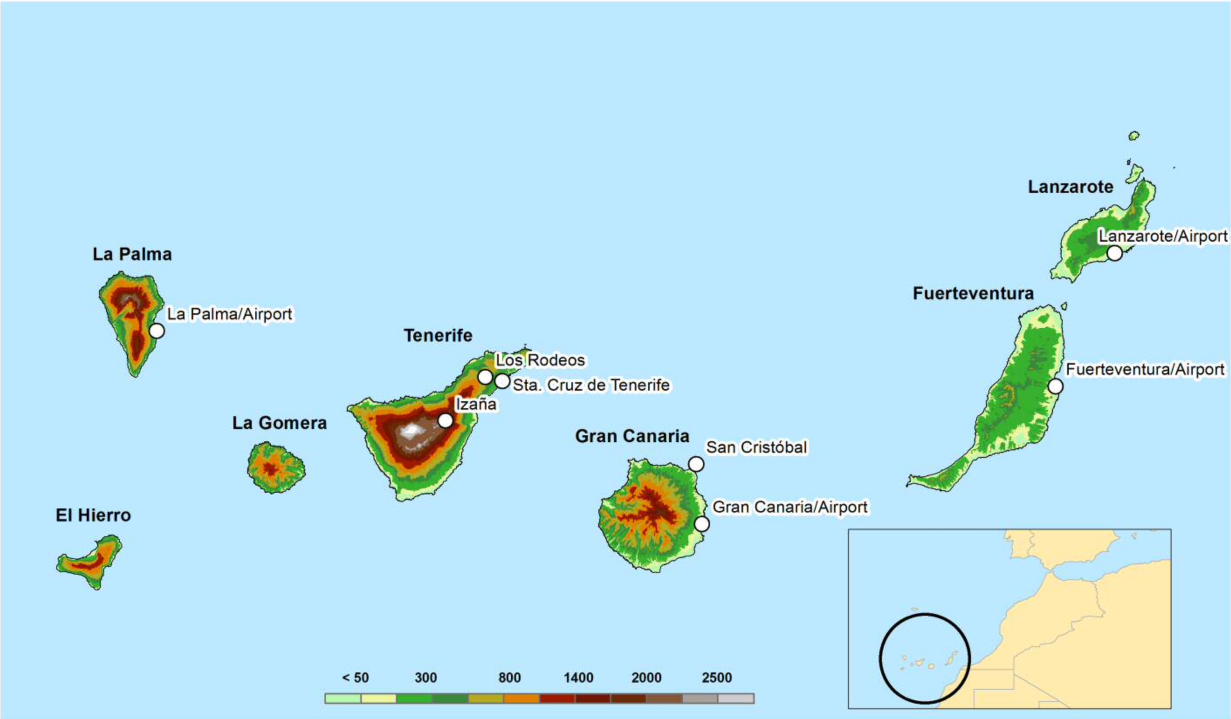
	Sta. Cruz de Tenerife	Gran Canaria/Airp.	Los Rodeos	Izaña	Fuerteventura	La Palma	Lanzarote	San Cristóbal	**Mean**
	Aerodynamic								
Winter	**0.88**	**0.95**	**0.99**	**0.99**	**0.98**	**0.97**	**0.97**	**0.96**	**0.93**
Spring	**0.65**	**0.93**	**0.95**	**0.96**	**0.95**	**0.93**	**0.93**	**0.88**	**0.87**
Summer	**0.74**	**0.96**	**0.96**	**0.97**	**0.94**	**0.84**	**0.94**	**0.63**	**0.85**
Autumn	**0.75**	**0.96**	**0.98**	**0.98**	**0.96**	**0.96**	**0.97**	**0.90**	**0.95**
Annual	**0.78**	**0.97**	**0.98**	**0.97**	**0.97**	**0.95**	**0.96**	**0.88**	**0.95**
	Radiative								
Winter	0.05	**0.37**	**-0.75**	**0.18**	**-0.62**	-0.22	**-0.44**	**-0.46**	-0.02
Spring	**0.38**	**0.52**	**-0.51**	**0.36**	-0.25	0.14	0.07	-0.17	**0.28**
Summer	0.05	**0.28**	**-0.37**	**-0.62**	-0.12	0.23	0.08	**0.41**	**0.29**
Autumn	0.14	0.09	**-0.67**	-0.01	-0.23	**0.43**	**-0.45**	-0.05	0.05
Annual	-0.05	-0.20	**-0.73**	**-0.36**	**-0.46**	0.04	**-0.28**	**-0.29**	-0.15

Table ~~6~~7. Seasonal and annual Pearson's coefficients between the time series of ETo and the different meteorological variables in the eight meteorological stations, calculated for the period 1961-2013. Statistically significant at the 95% confidence level are given in bold

	Sta. Cruz de Tenerife	**Gran Canaria/Airp.**	**Los Rodeos**	**Izaña**	**Fuerteventura**	**La Palma**	**Lanzarote**	**San Cristóbal**
				Maximum air temperature				
Winter	**0.32**	**0.51**	-0.12	**0.89**	-0.23	-0.01	-0.23	0.26
Spring	**0.46**	**0.69**	0.02	**0.90**	0.18	0.01	**0.62**	**0.42**
Summer	**0.48**	**0.80**	0.10	0.18	**0.33**	0.27	**0.51**	**0.44**
Autumn	0.18	**0.64**	0.04	**0.71**	**0.29**	0.12	0.09	**0.43**
Annual	0.17	**0.41**	-0.11	**0.64**	0.01	-0.03	0.16	**0.46**
				Minimum air temperature				
Winter	0.15	**0.50**	0.13	**0.83**	-0.24	0.17	-0.13	0.01
Spring	0.24	**0.53**	0.19	**0.83**	0.12	0.19	**0.49**	0.10
Summer	0.24	**0.55**	0.11	0.23	0.16	**0.33**	**0.55**	0.17
Autumn	0.21	**0.56**	**0.36**	**0.63**	0.20	**0.32**	**0.26**	0.21
Annual	0.04	**0.47**	0.13	**0.54**	-0.11	**0.30**	0.27	-0.07
				Relative humidity				
Winter	**-0.52**	**-0.91**	**-0.57**	**-0.83**	**-0.92**	**-0.92**	**-0.89**	**-0.72**
Spring	**-0.34**	**-0.89**	**-0.70**	**-0.90**	**-0.89**	**-0.90**	**-0.77**	**-0.82**
Summer	**-0.35**	**-0.93**	**-0.83**	**-0.46**	**-0.90**	**-0.89**	**-0.80**	**-0.61**
Autumn	**-0.30**	**-0.94**	**-0.55**	**-0.74**	**-0.90**	**-0.91**	**-0.78**	**-0.76**
Annual	-0.18	**-0.93**	**-0.62**	**-0.59**	**-0.93**	**-0.94**	**-0.85**	**-0.86**
				Sunshine duration				
Winter	**0.48**	**0.48**	0.16	**0.63**	0.01	**0.33**	0.18	0.06
Spring	**0.72**	**0.71**	0.08	**0.70**	0.27	**0.50**	0.25	0.21
Summer	**0.45**	**0.62**	0.20	0.18	**0.32**	**0.41**	**0.35**	**0.61**
Autumn	**0.47**	**0.38**	0.20	**0.53**	0.14	**0.69**	0.16	**0.34**
Annual	**0.40**	**0.30**	-0.01	**0.40**	0.15	**0.48**	0.08	-0.09
				Wind speed				
Winter	**0.61**	-0.01	**0.84**	**0.29**	**0.54**	**0.29**	**0.35**	**0.62**
Spring	**0.47**	0.18	**0.62**	**0.33**	**0.52**	0.22	0.24	**0.44**
Summer	**0.65**	**0.37**	**0.48**	**0.77**	**0.39**	-0.01	**0.33**	0.26
Autumn	**0.62**	0.22	**0.78**	**0.48**	**0.31**	0.27	**0.62**	**0.48**
Annual	**0.73**	**0.47**	**0.72**	**0.69**	**0.50**	0.25	**0.34**	**0.38**

Table ~~7~~8. Magnitude of change (ºC, %, hours and ms^{-1} decade^{-1}) of the different meteorological variables over the period 1961-2013. In bold statistically significant trends at the 95%.

	Sta. Cruz de Tenerife	**Gran Canaria/Airp.**	**Los Rodeos**	**Izaña**	**Fuerteventura**	**La Palma**	**Lanzarote**	**San Cristóbal**	**Mean**
					Maximum air temperature				
Winter	-0.06	-0.09	-0.05	-0.01	-0.08	-0.08	**-0.18**	**-0.18**	-0.09
Spring	-0.08	0.03	-0.02	-0.12	-0.02	-0.02	0.08	0.14	0.00
Summer	-0.06	**0.20**	0.00	-0.07	0.00	0.00	0.07	0.12	0.04
Autumn	-0.06	-0.08	-0.08	-0.04	-0.10	-0.06	-0.11	**-0.17**	-0.09
Annual	-0.05	0.03	-0.01	-0.05	-0.03	-0.02	-0.01	0.00	-0.02
					Minimum air temperature				
Winter	-0.02	-0.01	0.02	0.16	-0.02	0.02	-0.02	0.14	0.03
Spring	0.02	0.03	0.03	0.18	0.04	0.04	0.05	0.09	0.06
Summer	0.08	**0.12**	**0.10**	**0.25**	**0.11**	0.07	**0.10**	**0.13**	**0.12**
Autumn	0.07	0.01	**0.09**	**0.19**	0.05	0.09	0.09	0.08	**0.09**
Annual	0.05	0.05	0.08	**0.20**	0.06	0.07	0.08	**0.12**	**0.09**
					Relative humidity				
Winter	**-0.51**	-0.51	-0.22	-1.11	-0.81	**-1.53**	**-1.56**	-0.18	**-0.80**
Spring	0.18	**-1.06**	-0.22	0.20	-0.76	-0.96	**-0.88**	**0.90**	-0.33
Summer	0.39	**-1.58**	-0.16	-0.91	-0.06	-0.72	**-0.99**	0.45	**-0.45**
Autumn	0.02	-0.72	0.01	-0.26	-0.29	**-1.65**	**-0.99**	0.31	**-0.45**
Annual	0.02	**-0.89**	-0.03	-0.52	-0.49	**-1.05**	**-1.11**	0.32	**-0.47**
					Sunshine duration				
Winter	0.02	**-0.10**	-0.04	0.02	-0.12	0.08	-0.05	**-0.11**	-0.04
Spring	0.08	0.11	0.08	0.06	0.03	**0.22**	-0.06	0.05	0.07
Summer	0.06	**0.15**	0.05	-0.03	0.00	**0.25**	0.09	**0.35**	**0.12**
Autumn	0.03	-0.04	0.03	0.08	0.00	0.19	0.03	**-0.16**	0.02
Annual	0.06	0.03	0.03	0.04	-0.01	**0.18**	0.02	0.04	0.05
					Wind speed				
Winter	**0.04**	0.04	**0.33**	0.01	0.00	**0.07**	0.02	**-0.18**	0.04
Spring	-0.01	0.08	**0.19**	0.07	-0.08	-0.08	**-0.13**	**-0.24**	-0.03
Summer	0.02	**0.21**	**0.24**	-0.01	-0.05	-0.11	-0.06	0.01	0.03
Autumn	0.03	0.07	**0.33**	0.03	-0.07	-0.05	-0.04	-0.06	0.03
Annual	0.02	**0.10**	**0.27**	0.02	-0.04	-0.04	-0.04	**-0.12**	0.02



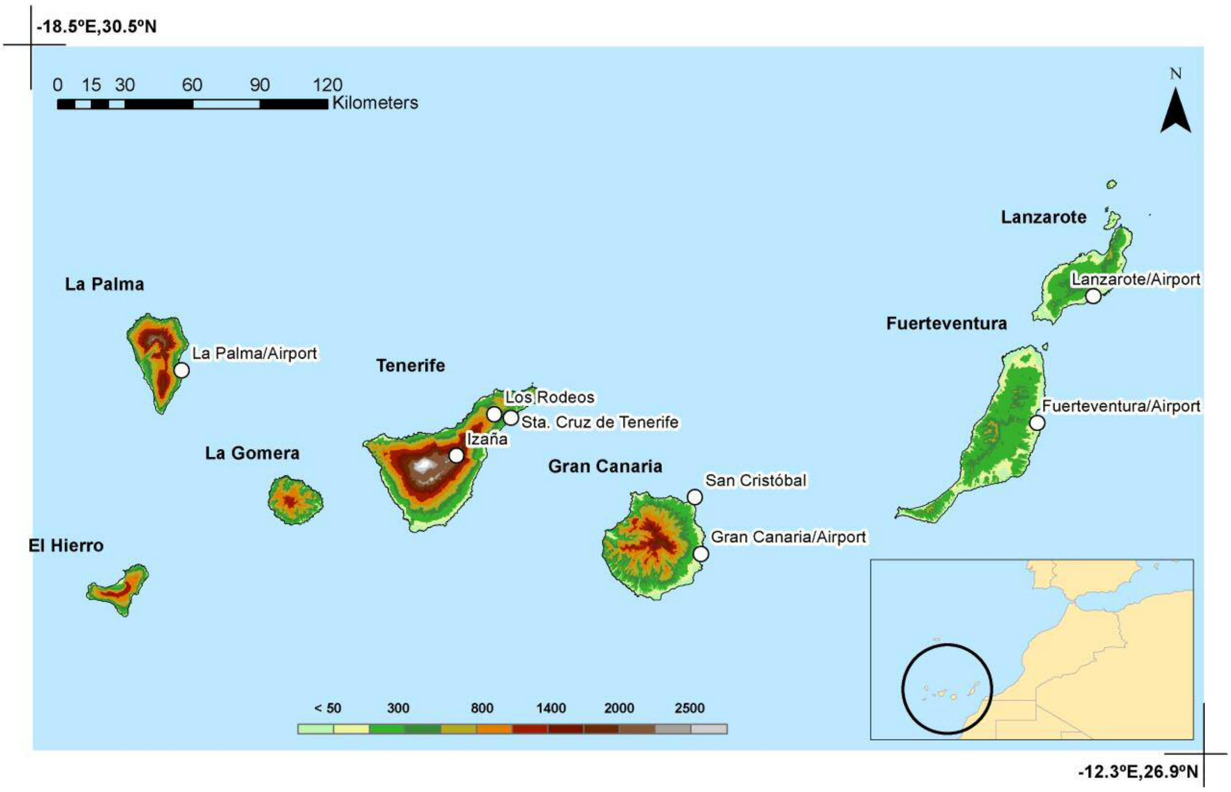


Figure 1: Location and relief of the Canary Islands and meteorological stations used in the study. Altitude is given in meters.

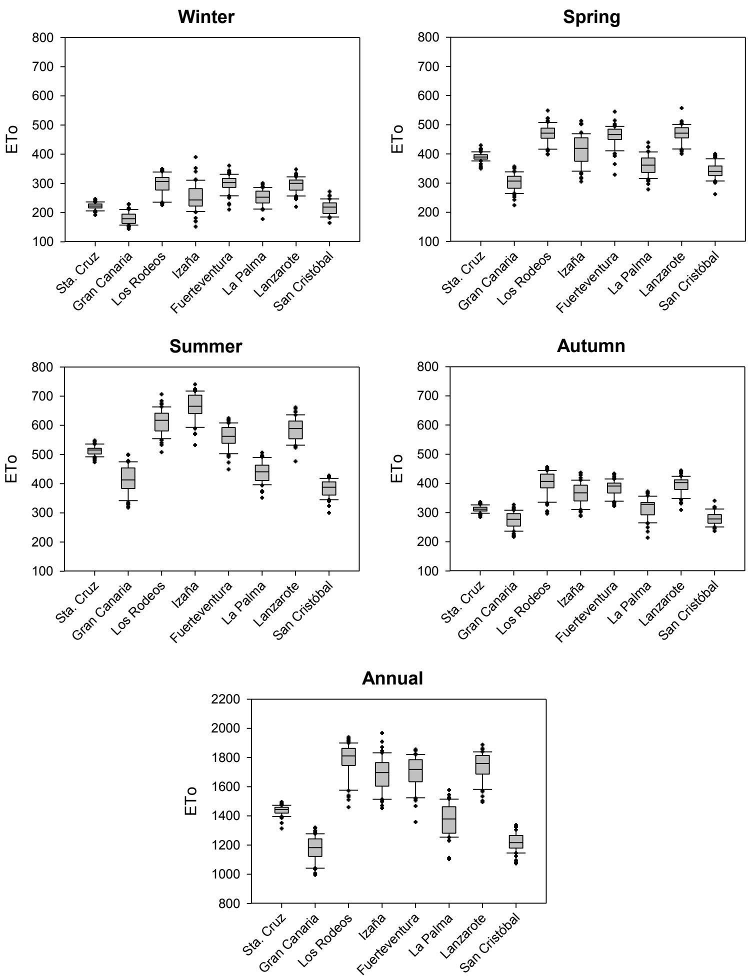


Figure 2: Box-plot with the annual and seasonal ETo values in the eight meteorological stations used in this study. The vertical lines of each plotted boxplot illustrate the 10th, 25th, 75th and 90th quantiles, respectively. The interquartile spread is represented by the range between the 25th and 75th quantiles. The dots show the highest and lowest values.

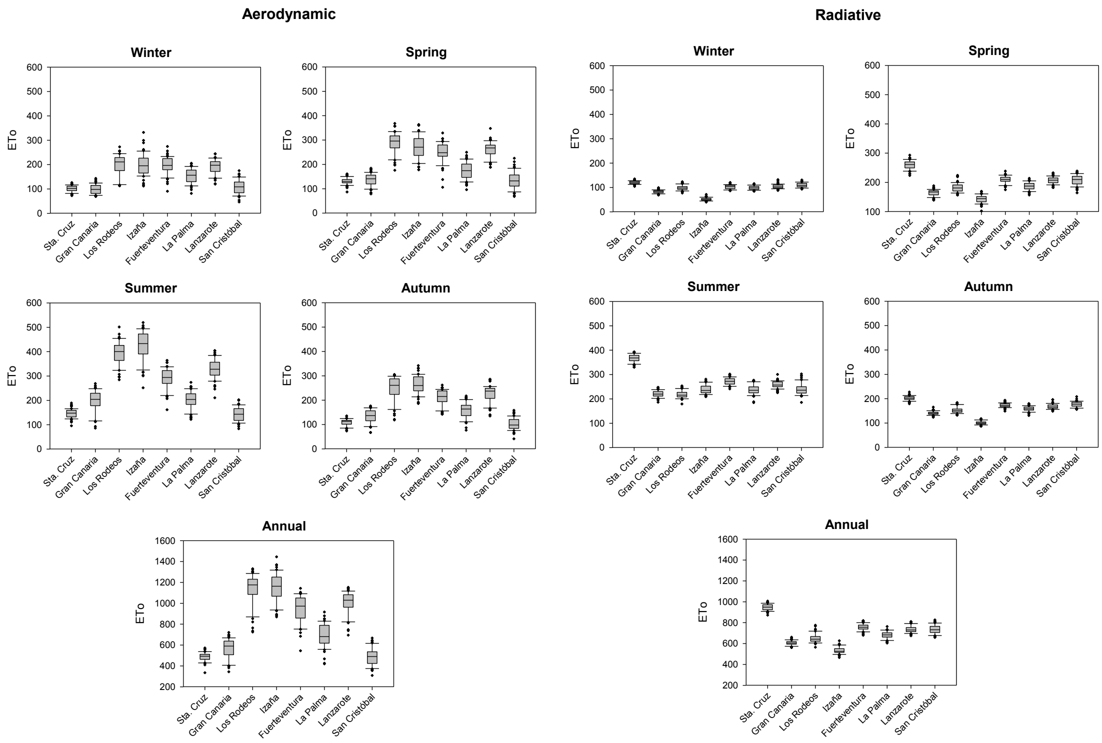
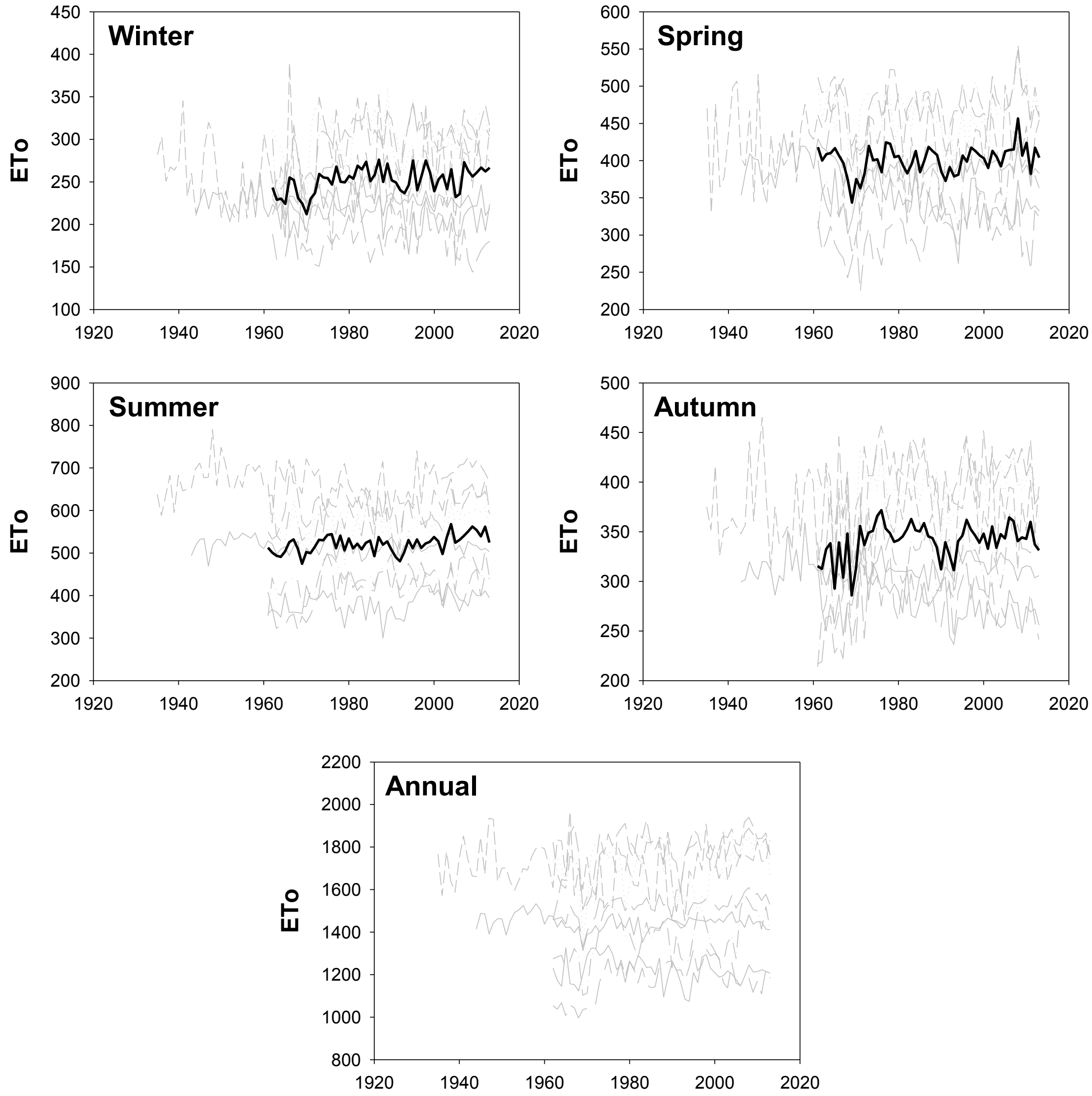


Figure 3: Box-plot with the annual and seasonal aerodynamic and radiative components of ETo in the eight meteorological stations used in this study


Winter
Spring
Summer
Autumn
Annual
ETo
1920
1940
1960
1980
2000
2020

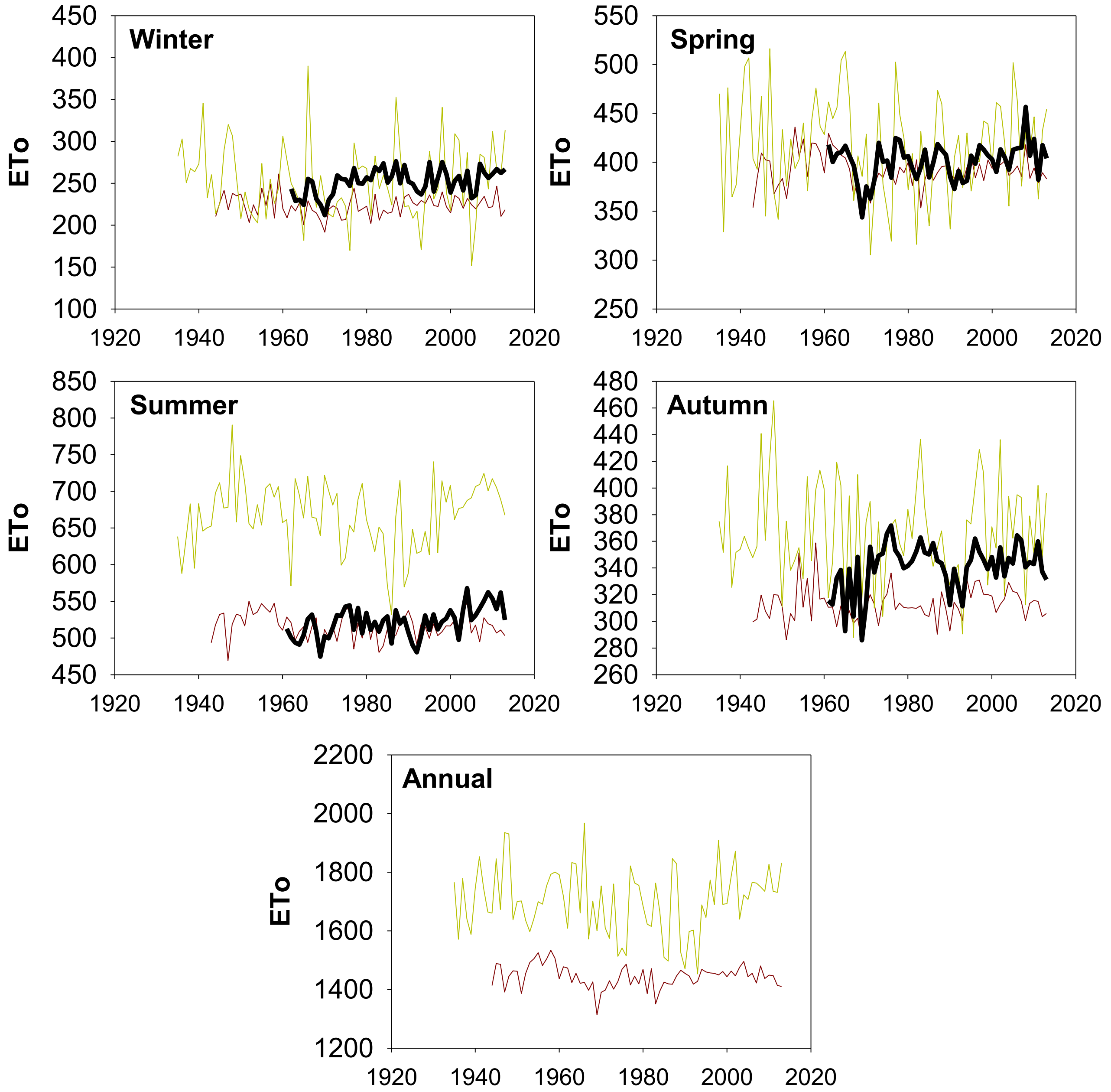
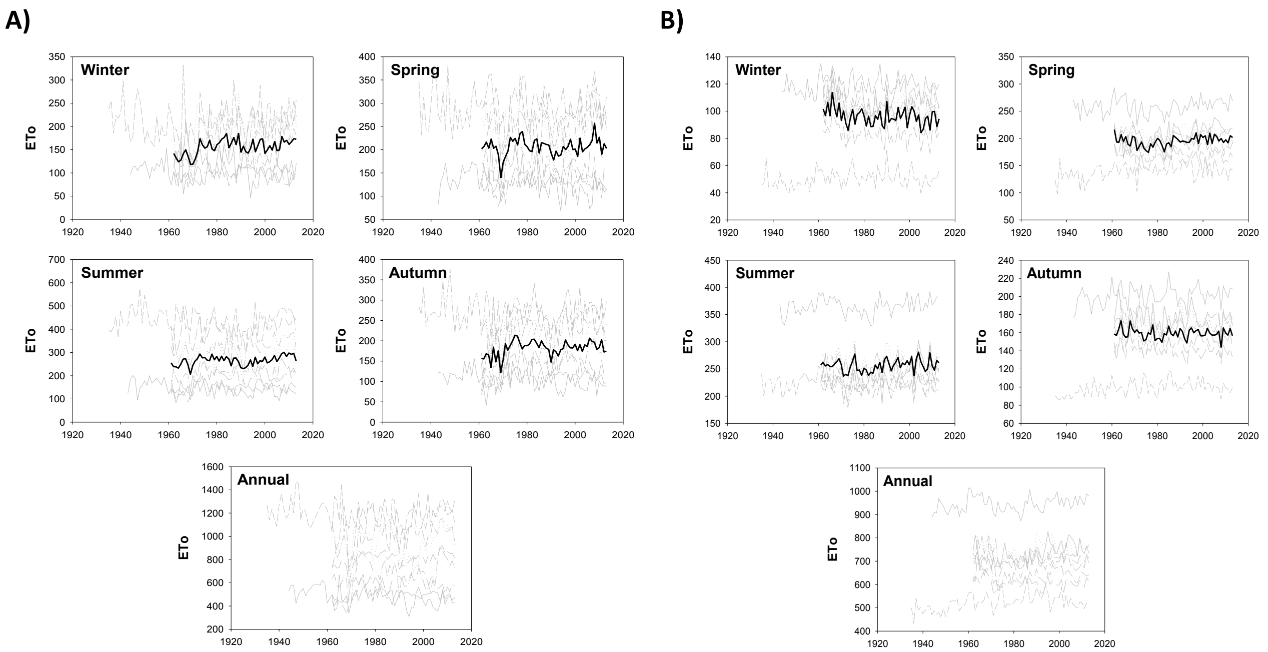


Figure 4: Evolution of seasonal and annual ETo in the eighttwo meteorological stations (grey lineswith longest records (Izaña, green and Santa Cruz de Tenerife, brown) and the average of the eight stations (black lines) from 1961 to 2013.



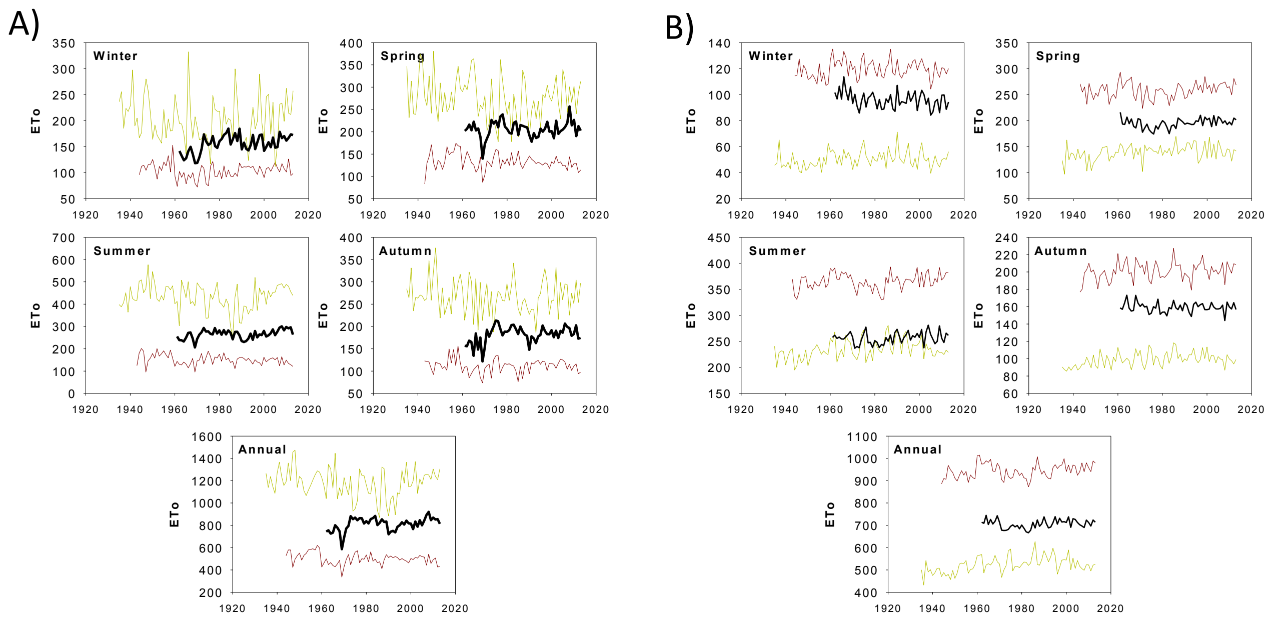


Figure 5: Evolution of seasonal and annual aerodynamic (A) and radiative (B) components of the ETo in the ~~eight~~two meteorological stations ~~(grey lines~~with longest records (Izaña, green and Santa Cruz de Tenerife, brown) and the average of the eight stations (black lines) from 1961 to 2013

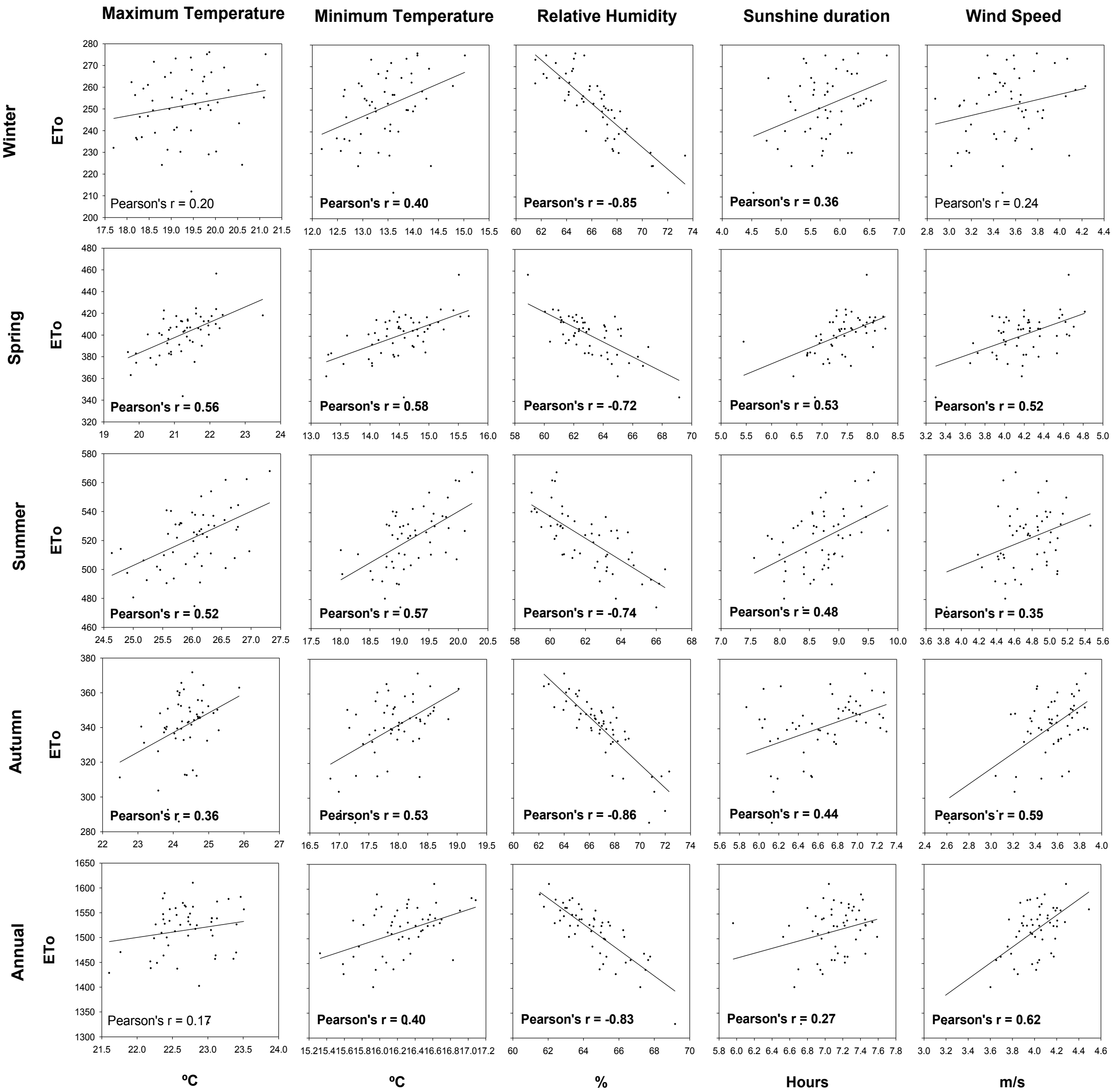


Figure 6. Relationship between the regional annual and seasonal ETo and the regional series of the different meteorological variables. Pearson's coefficients are included in each plot. In bold the coefficients statistically significant at the 0.95 confidence level

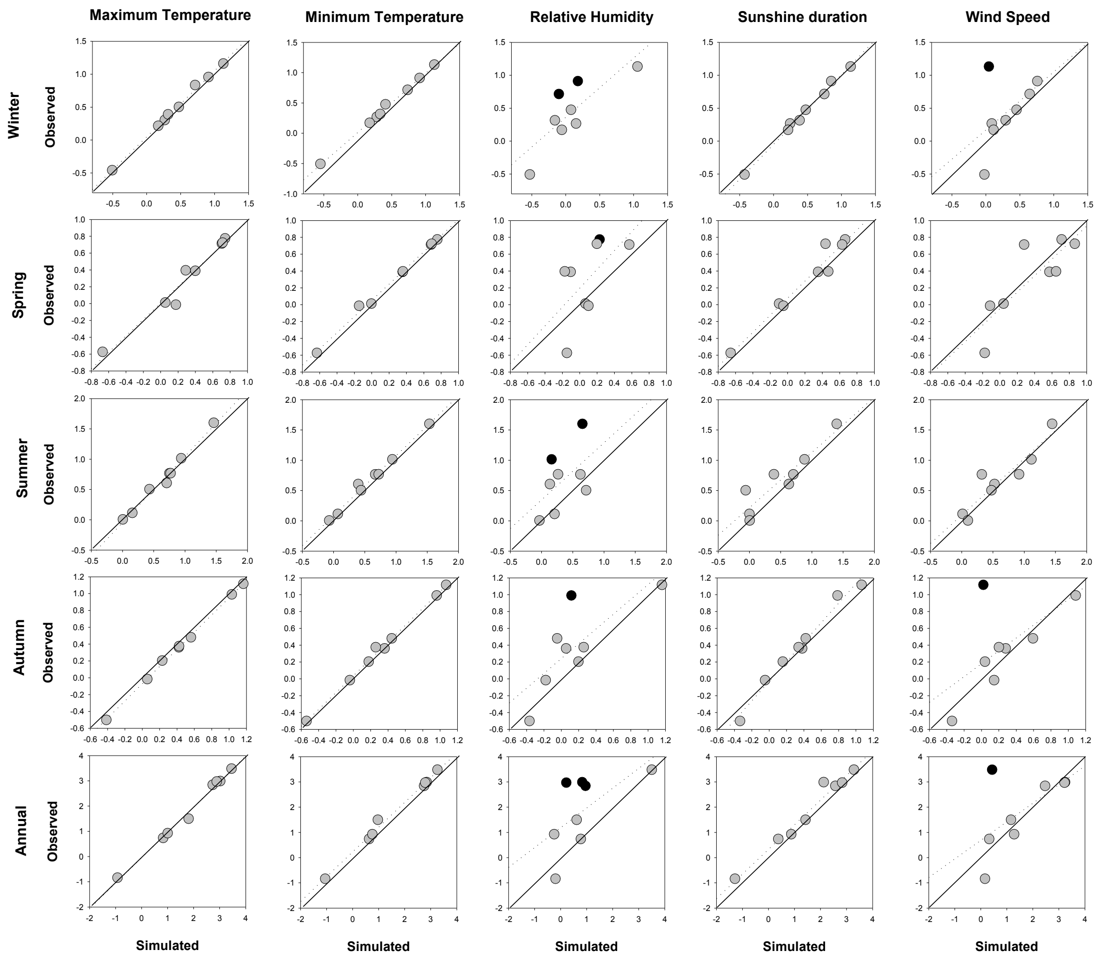


Figure 7: Relationship between the seasonal and annual observed magnitude of change of ETo (mm. year-1) in each meteorological station and the simulated magnitude of change maintaining each meteorological variable as constant. Black dots indicate significant differences in the trends.

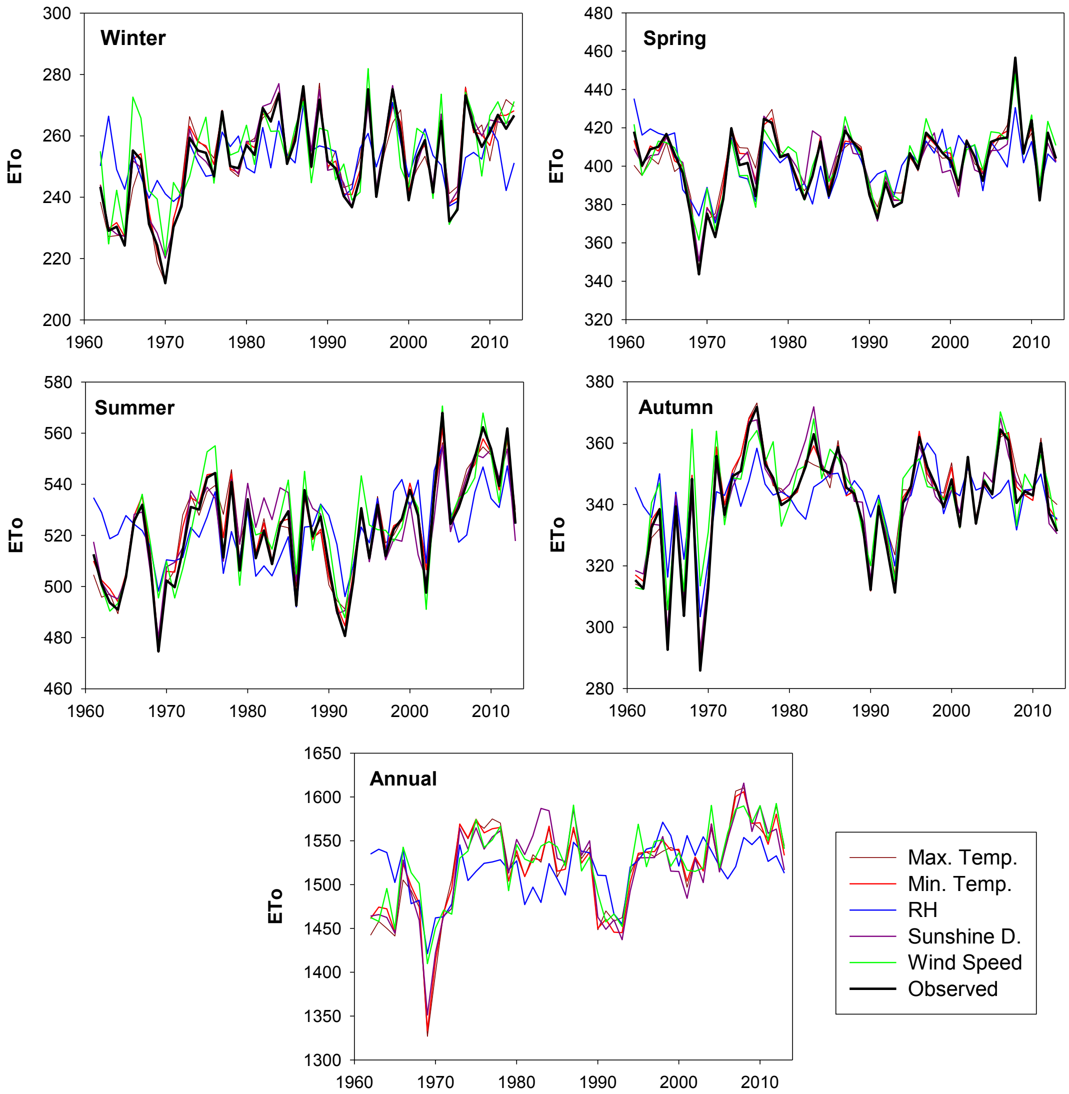


Figure 8: Seasonal and annual evolution of the observed regional ETo compared to the simulated ETo considering no temporal changes in each one of the meteorological variables from 1961 to 2013.