



1 **Coupled hydro-meteorological modelling on HPC platform for high**
2 **resolution extreme weather impact study**

3 Dehua Zhu^{[1][2]}, Shirley Echendu^[1], Yunqing Xuan^[1], Mike Webster^[1], Ian Cluckie^[1]

4 [1] College of Engineering, Swansea University Bay Campus, Swansea, SA1 8EN, UK

5 [2] Current Affiliation: School of Hydrometeorology, Nanjing University of

6 Information Science and Technology. Nanjing, 210044, China.

7 **Abstract:** High performance computing (HPC) has long been used in the disciplines of
8 atmospheric and oceanic sciences, and remains the main tool of choice to extract
9 numerical solutions to complex geophysical problems on the global scale, often
10 accompanied with very large numbers of degrees of freedoms. However, with the
11 growing recognition that the spatially distributed feedback from the land surface is
12 important to weather and the climate system, representation of the land surface is
13 established with increasingly complex (and physically complete) models, which often
14 leads to the coupling of heterogeneous models such as numerical weather prediction
15 (NWP) models and hydrological models. As a result, the spatial grids and the temporal
16 resolutions have become finer and thereby computers with far greater computational
17 and storage capacity are in great demand than those used in the past. Additionally,
18 impact-focused studies that require coupling of accurate simulations of
19 weather/climate systems as well as impact-measuring hydrological models that
20 demand larger computer resources in its own right.

21 In this paper, we present a preliminary analysis of an HPC-based hydrological
22 modelling approach, which is aimed at utilising and maximising HPC power resource,
23 to support the study on extreme weather impact due to climate change. Here, two
24 case studies are presented through implementation on the HPC Wales platform of the
25 UK mesoscale meteorological Unified Model (UM) UKV, alongside a Linux-based
26 hydrological model, HYdrological Predictions for the Environment (HYPE). The results
27 of this study suggest that high resolution rainfall estimation produced by the UKV has
28 similar performance to that of NIMROD radar rainfall products as input in a
29 hydrological model, but with the added-value of much extended forecast lead-time.

30 **Key words:** HPC, extreme weather impact, Unified Model, UKV, HYPE, coupled hydro-
31 meteorological modelling

32 **1. Introduction**

33 Extreme precipitation with great intensity and the subsequent flash flooding events
34 arising from rivers and mountainous watersheds often lead to considerable economic
35 damage and casualty, because water levels can react extremely quickly within rather



36 limited warning lead-time (flash flooding). Therefore, the evaluation of potential
37 flooding risks in extreme weather conditions, and the corresponding protection
38 measures required, demand accurate short-term flood forecasting, and more often
39 very short lead-time forecasting – termed ‘nowcasting’ (Cloke and Pappenberger,
40 2009).

41 Understandably, hydrological models together with hydraulic models play a key role
42 in predicting runoff, river flow as well as possible inundations. However, the lead time,
43 which is crucial for hazard mitigation and evacuation, is often highly limited in such a
44 classic model chain configuration, since basically then, the lead-time is the travelling
45 time of flood water. It is therefore, other means of providing rainfall estimates with
46 extra lead-time (e.g., weather radar observations), have become increasingly essential
47 in flood forecasting under extreme weather conditions (Zhu et al. 2014). However, it
48 has also been recognised, e.g., by Golding (1998) and Smith and Austin (2000), that
49 the performance of radar-based rainfall nowcasting deteriorates rapidly when the
50 lead-time goes beyond 0.5 hr. Then, the combination of radar nowcasting and
51 hydrological forecasting is reduced to a normal model, or even worse. In fact, early
52 attempts, whilst using the NIMROD radar rainfall product, already introduced a rainfall
53 forecast from numerical weather prediction models to compensate for this
54 shortcoming.

55 The fast development of HPC, as well as that of NWP (Numerical Weather prediction)
56 models, has since given rise to the use of NWP, either directly or indirectly in
57 hydrological simulations, in an effort to push hydrological forecasting beyond the limit
58 of the rainfall-observation time-horizon. This link between two different modelling
59 disciplines is often referred to as model coupling. The resulting coupled
60 meteorological-hydrological models appeared from the beginning of the 21st century,
61 being initially focused on flash flood forecasting, and later extended to handle
62 climatic-hydrological coupling. This has facilitated many climate-change impact
63 studies on water resources that rely heavily on the use of climate projections or
64 simulations. Nevertheless, the linkage between the meteorological and hydrological
65 models is scientifically challenging due to differences in model structures and issues
66 of incompatible units (use of different scales in time and space). This is encapsulated,
67 in particular, in the task of how best to transform and regionalise global climate
68 scenarios, with spatial resolutions of 1,000-10,000 km², to hydrological mesoscale
69 catchments of (10 – 1,000 km²).

70 Simulation with meteorological–hydrological coupling in high spatial and temporal
71 resolution is a comparatively new field of hydrological research, yet some pioneering
72 work has recently appeared. In order to analyse the prediction of selected events
73 characterized by peak flows, Westrick et al. (2002) proposed a hydrometeorological
74 forecasting system for mountainous watersheds by coupling the Penn State–NCAR



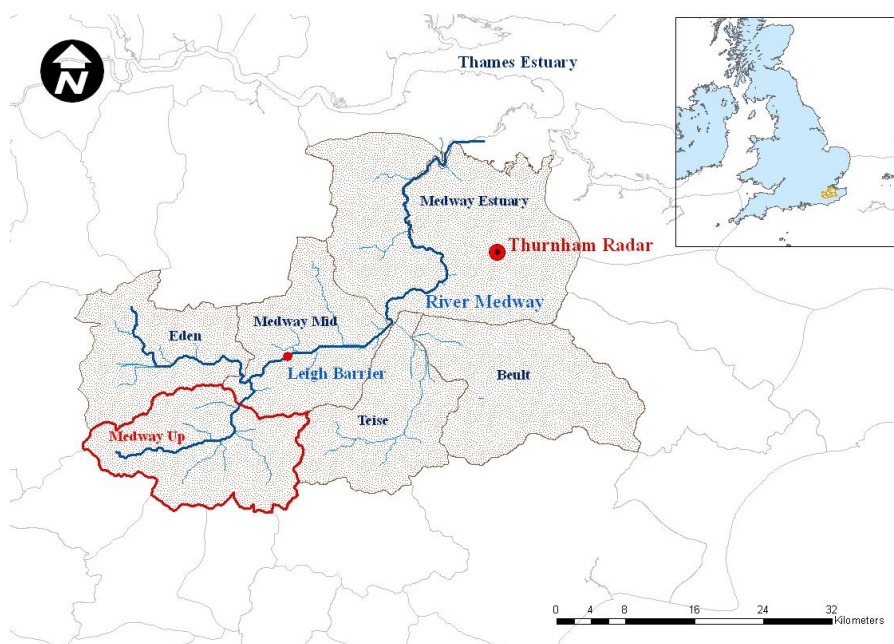
75 Mesoscale Meteorological Model (also known as MM5 for brevity; Dudhia, 1993; Grell
76 et al., 1995) in 4*4 km² resolution and the distributed hydrological model DHVSM
77 (Wigmosta et al., 1994). Jasper et al. (2002) compared the hydrological performance
78 of radar and gauge measurements, with five different high-resolution numerical
79 weather prediction (NWP) models and grid-cell sizes between 2 and 14 km. This work
80 covered the prediction of peak-flows on the alpine Ticino-Toce watershed, using the
81 distributed hydrological model WaSiM (Schulla and Jasper, 2000). The results suggest
82 that, the accuracy and consistence of NWP rainfall in hydrological applications heavily
83 depend on their process modelling at all scales of model nesting. Particularly so, as
84 inaccuracies introduced by downscaling of precipitation from NWP models can lead
85 to large differences in the predicted hydrological results, especially during extreme
86 convective storm periods. Kunstmann and Stadler (2005) coupled (in 1-way manner)
87 the mesoscale meteorological model MM5 with the distributed hydrological model
88 WaSiM. The meteorological re-analysis data were dynamically downscaled with MM5
89 grid cell sizes from 100 km to 2 km using four nests. Findings show that the MM5-
90 based interpolation of precipitation yielded 21% less total yearly precipitation in the
91 catchment area, compared to the station-based interpolation. Yarnal et al. (2000)
92 linked a high-resolution meteorological model (MM5 at 4 km resolution) and a suite
93 of coupled hydrological models (HMS) in the Susquehanna River Basin Experiment
94 (SRBEX). This work points out that the coupled model has to confront several issues,
95 such as, physics and parameterizations for a mesoscale atmospheric model to match
96 the time-scales of climate coupled to the hydrological process models, meteorological
97 and climatological process models with different scales; and the immense
98 computational needs accordingly. Xuan et al (2009) also indicted that the inaccuracies
99 and uncertainties in NWP could propagate to the downstream hydrological models,
100 and they proposed to use an ensemble-based approach, together with effective bias
101 correction, to mitigate this problem.

102 The majority of the studies cited above have been relying on the use of the so-called
103 downscaling of large-scale NWP results using regional meteorological model such as
104 MM5. These studies are often conducted in an off-line manner where hydrological
105 modellers have hardly any control of NWP except the meso-scale one used for
106 downscaling. However, the work presented in this paper, not only focuses on the
107 performance of coupled high-resolution meteorological–hydrological simulations for
108 extreme storm events on a HPC platform, it is also aimed at exploring the potential of
109 building and running fully-coupled NWP-hydrological forecasts on a single computer
110 platform; and therefore is able to obtain first-hand knowledge on fully integrated
111 hydro-meteorological modelling. As such, we did not apply the meteorological model
112 in forecasting mode, but used hindcasting mode instead, to test the model
113 performance and benchmarking over several selected historical events.



115 2. Materials and Methods

116 In this study, the principal goal of the experiment has been to simulate the river basin
117 response to extreme storm events, by linking a semi-distributed hydrological HYPE
118 model to the UK Met Office Unified Model (UM) at a much finer spatial scale (1.5km).
119 The combined high-resolution one-way driven model experiments generate runoff
120 hydrographs for three extreme flood events, which occurred in the Upper Medway
121 catchment (220 km²) located to south of London in the UK (see Figure 1). The
122 properties of this catchment regarding the topography, vegetation and soil types, as
123 well as the availability of a hydrological data set, have been detailed in Zhu and Cluckie
124 (2012) and Zhu et al. (2014).



125

126 Figure 1. Location of UpperMedway Catchment in the UK

127 The catchment elevation varies between 30 m and 220 m above mean sea-level and
128 the majority of slope ranges from 2 degrees to 8 degrees, which makes up around 70%
129 of the whole catchment. This suggests that the main scenery of the Upper Medway
130 Catchment is small hills surrounding the flat, little relief low-lying area without much
131 variation of elevation. The land use in the catchment can be simplified and described
132 as permanent grass (over 95%). The catchment is characterized by a mixture of
133 permeable (chalk) and impermeable (clay) geologies and the dominant aquifers
134 consist of the Ashdown Formation and the Tunbridge Wells Formation of the Hastings
135 Group. The saturation-excess mechanism is the major runoff generation process in the
136 catchment.



137 In such model experiments, two different sets of meteorological input data were used:
138 (1) surface observation data from station measurements and from weather radar
139 estimation, and (2) forecast rainfall data UKV from high-resolution numerical weather
140 prediction (NWP) models with grid-cell sizes 1.5 km. The experiments were designed
141 as: (1) selecting representative storms and hydrographs for simulation; (2) simulating
142 these storms using the high-resolution UM-UKV model, and (3) modelling the river-
143 basin response to the simulated storm events using the HYPE hydrological model.
144 Moreover, it is worth noting that this experiment of a fully-coupled NWP-hydrological
145 forecast is preliminarily designed to be a one-way coupling system, which will form
146 the basis for extension into a two-way coupling system, which will be developed
147 further in the future.

148 One notes that Met Office has used the operational high-resolution UK 1.5 km model
149 (UKV) under the New Dynamics algorithm specification. This introduces nested
150 operations, through parallel suites PS30 and the time periods of interest. As such, this
151 consists of a Global 25km simulation, followed by a North Atlantic European 12 km
152 simulation, and finally, a UKV 1.5 km simulation. Each such simulation stage provides
153 the necessary lateral (spatial-temporal) boundary conditions for the regionally-refined
154 subsequent stage.

155 **2.1 UKV model configuration and implementation**

156 The unified model (UM) is an atmospheric predictive numerical modelling software,
157 offered by the UK Met Office and written in FORTRAN. Here, its output is coupled with
158 a hydrological model for the purpose of accurate flood and extreme storm prediction.
159 The UM was built on Archer hardware, with specification as a Cray XC30 MPP
160 supercomputer, with up to 4920 compute nodes, each having a two 12-core Intel Ivy
161 Bridge series processor, providing a total of 118,080 processing cores. Each node has
162 64GB memory, with a subset of large memory nodes possessing 128 GB.

163 The high-resolution simulation was achieved using the UKV suite which is a regional
164 configuration of the UM, derived from operations through Parallel Suites (PS30),
165 which consist of three nested domain simulations: Global 25 km simulation, North
166 Atlantic European (NAE) 12km simulation, and UKV 1.5 km simulation. The Global
167 N512L70 problem suite is discretized into approximately 25 km mid-latitudes, upon a
168 1024x769 grid. There are 70 model levels vertically and a time-step of 10 mins is used.
169 The regional NAE problem suite has a resolution of 12km, across a 600x360 grid. The
170 NAE suit also has 70 vertical levels but the time-step choice is 5 mins. Finally, the
171 regional UKV is set at 1.5 km resolution over a 622x810 grid with a time-step of 50sec.

172 UKV model implementation requires a few events for model run. This includes an
173 initialisation date and a number of subsequent time-duration periods i.e. 3 days, 6
174 Days, 8 Days and 12 days. Each event is specified to run at different lead times. A



175 selection of start dumps for the different lead times are requested from ECMWF or
176 the Met. Office. The Met. Office holds start dumps to a back-date of up to five years
177 only; prior to that, start dumps would need to be obtained from other sources.

178 The UKV model outputs were also on a rotated lon-lat grid, whose resolution is not
179 constant, with small deviation from 1.5 km depending on the locations. The data was
180 further projected onto the National Grid Reference Grid to become comparable with
181 other sources of data, such as the weather radar rainfall observation from the
182 NIMROD system. A nearest-neighbour interpolation was used to produce the evenly
183 distributed grid data after projecting.

184 **2.2 The configuration and calibration of hydrological model - HYPE**

185

186 Whilst many hydrological models could have been selected, e.g., see Zhu et al (2014),
187 an open source model – HYPE (HYdrological Predictions for the Environment) has been
188 selected in this study to avoid reliance on commercial modelling packages. HYPE is
189 developed at Swedish Meteorological and Hydrological Institute (SMHI) with a focus
190 on integrating water and water quality throughout the model compartments,
191 predictions in ungauged catchments with large model set-ups, e.g. across Europe. It is
192 a dynamical model forced with time series of precipitation and air temperature,
193 typically on a daily time-step. Forcing in the form of nutrient loads is not dynamical.
194 Examples of HYPE applications include atmospheric deposition, fertilizers and waste
195 water.

196

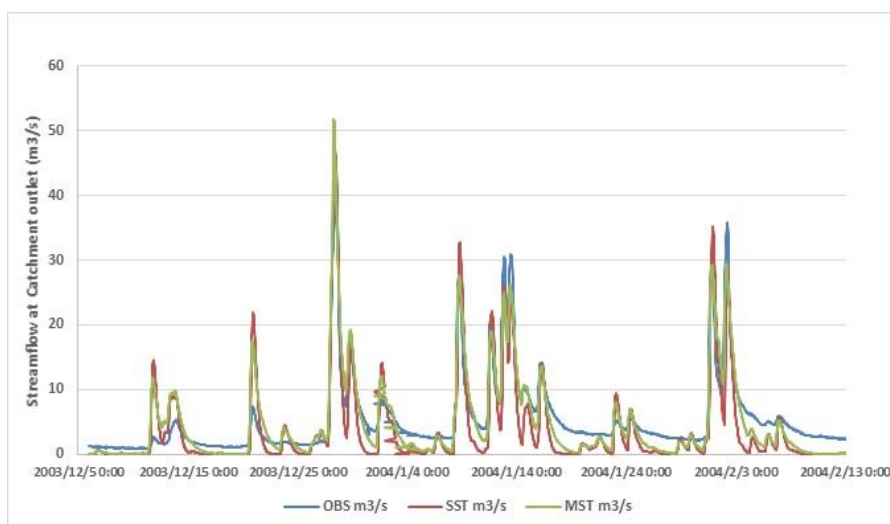
197 The HYPE model is able to predict water and nutrient concentrations in the landscape
198 at the catchment scale. Its spatial division is related to sub-catchments and
199 corresponding characteristics, including land use, vegetation, soil type and elevation.
200 Within a particular catchment, the model will simulate water content in different
201 compartments, including soil moisture, shallow groundwater, rivers and lakes.

202 The default time-step in HYPE is daily, but it can be reduced to hourly, which is
203 normally specified in the input dataset, such as precipitation. Since there is no 2-D
204 surface runoff algorithm built in the HYPE model, it is in principle a lumped model.
205 However, spatial variations can be accounted for by portioning the catchment into
206 smaller sub-catchments. In this respect, the simulated precipitation was processed as
207 the catchment average before being fed to the HYPE model.

208 The winter flood event started from 06/12/2003 to 28/02/2004 was used for model
209 calibration, carried out using 1-hour time step rain-gauge measurements and
210 parameterised with the streamflow observation at the catchment outlet. In order to
211 achieve the best fit between observed and modelled flow, the model parameters were
212 calibrated in simulation mode using a mixture of manual and automatic parameter



213 adjustment, according to their functionalities in the model. First, all the parameters
214 went through an initial manual sensitivity analysis, to decide those worthy of further
215 automatic parameterisation. This then involved progressive Monte Carlo simulation,
216 with parameter space reduced in stages.
217



218
219
220

Figure 2. The comparison of model calibration with different soil settings

221 The soil properties setting is critical in HYPE model. Figure 2 shows the model
222 calibration performance with single soil type (SST) and multiple soil type (MST)
223 settings. This data clearly indicates that the recessions period with SST setting was
224 much faster than the observation, possibly due to the less resilience from a single soil
225 type setting and the shallow depth of soil layer in the model. Consequently, multiple
226 soil types and the increment depth of the soil layer were introduced to the model
227 while the recession of flood was improved. Additionally, the most critical performance
228 criteria for the model, as in the Nash-Sutcliffe Efficiency (NSE), increases from 0.68(SST)
229 to 0.82(MST).

230

231 **2.3 The settings of a coupled UKV-HYPE case study**

232

233 The UKV model is set to make 36-hr forecast with a high resolution inner domain (1.5
234 km grid boxes) over the area of forecast interest, separated from a coarser grid (4 km)
235 near the boundaries by a variable resolution transition zone. This variable resolution
236 approach allows the boundaries to be moved further away from the region of interest,
237 reducing unwanted boundary effects on the forecasts.

238



239 Part of the motivation of using such resolution is to improve forecasts of convective
240 rainfall. The variable resolution model with 1.5 km grid length over the UK, but
241 increasing to 4km at the edges of the domain, enables the boundaries of the model to
242 be pushed further away from the area of interest at lower cost, and also to reduce the
243 resolution mismatch with the driving (12 km) model. The UKV rainfall estimation
244 produced by the Unified Model is used as the input for the HYPE model, which
245 provides the cornerstone to the coupled UKV-HYPE model.

246
247 As the HYPE model is currently in a lumped formation, the gridded UKV rainfall data
248 was averaged over the catchment area and interpolated into 1-hour time-steps. Two
249 selected historical flood event is simulated as a case study, with the coupled UKV-HYPE
250 model. The first event, started from 2007/01/05 to 2007/01/13, was a relatively small
251 flood event, considering the precipitation depth over the catchment is less than 25mm
252 in nine days and producing the highest peak flow of around 10 m³/s. The second
253 event was a flash flood with sudden high peak flow and short period, started from
254 2007/07/15 to 2007/07/23, especially during the period from 2007/07/20/ 08:00 to
255 11:00 when over 30mm precipitation fell on the catchment in three hours, detected
256 from the raingauge network.

257

258 **3. Results and Discussions**

259

260 Rainfall observations from weather radars were also introduced in this study to check
261 the UKV output, since rain-gauges are point-based and the radar rainfall can provide
262 well represented rainfall distribution. Moreover, the comparison with UKV input
263 through a hydrological model can be drawn, in terms of lead-time difference.

264

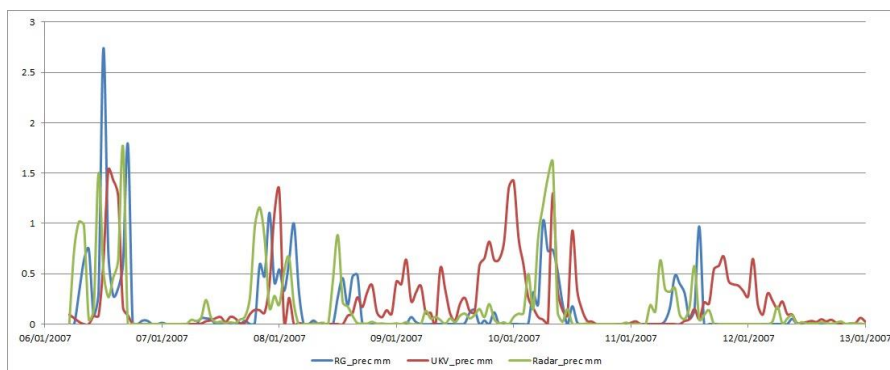
265 The radar rainfall estimates used in this study is extracted from the UK NIMROD
266 composite dataset. This has been provided and quality controlled by the UK Met Office
267 using the lowest available scan. It has been adjusted against available rain-gauge
268 measurement and undergone extensive processing to correct for various sources of
269 radar error. Such radar error would include noise, clutter, anomalous propagation,
270 attenuation, occultation, “bright band” and orographic enhancement. Therefore,
271 these high-resolution radar composite rainfall estimates incorporate the latest UK Met
272 Office processing algorithms to account for the different sources of errors in the
273 estimation of precipitation using weather radars (Harrison et al., 2000). This implies
274 that this data-set is the best possible estimate of rainfall available at the ground-level
275 in the UK (most error-free).

276

277 The processed UKV rainfall data was compared with other rainfall data from rain-
278 gauge and NIMROD radar. All types of rainfall data are interpolated and processed,

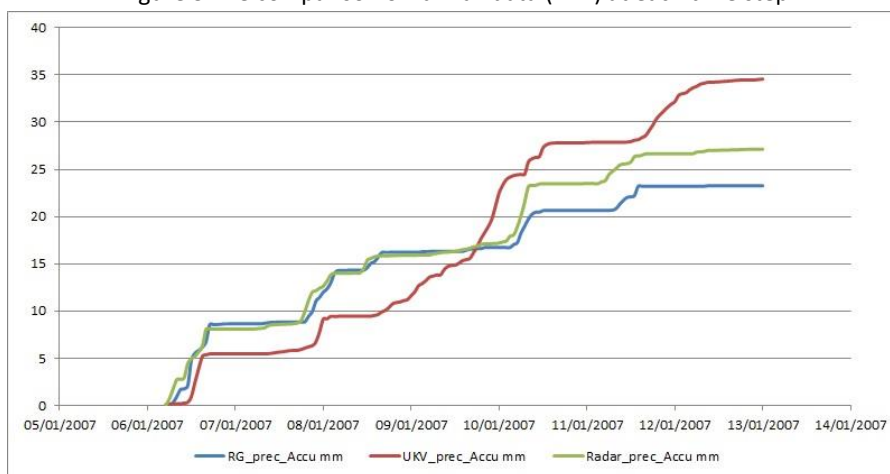


279 sharing the same spatial and temporal resolution, which is 1km and 1 hour, for
280 comparison purpose.
281



282

283 Figure 3 The comparison of rainfall data (mm) at each time step



284

285 Figure 4. The comparison of accumulative catchment average rainfall (mm)

286

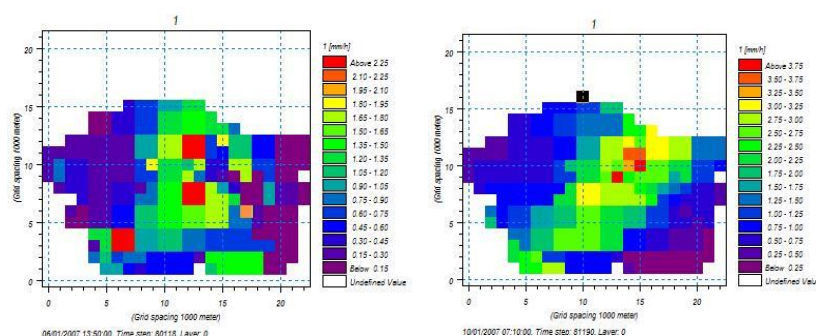
287 Figure 3 and Figure 4 show the rainfall comparison on each hourly step and the
288 accumulation of catchment average rainfall over the flood period, respectively. Trends
289 on the rainfall data at each time step of Figure 3 are reasonably good across all three
290 data-sets, with capture of the main data-peaks. The UKV rainfall data does however
291 pick up some exaggerated noisy peaks over the period between days 09/01/2007 to
292 10/01/2007 (see below to cumulative rainfall data).

293 Figure 4 shows that the NIMROD radar data produced more rainfall depth over the
294 catchment than rain-gauge measurements, but less than the UKV rainfall. In addition,
295 it shows similar rising cumulative rainfall for this event between all three data-sets,
296 and particularly between rain-gauge measurements and radar rainfall estimation up



297 to 10 Jan. In contrast, on cumulative rainfall of Figure 4, one notes that the UKV rainfall
 298 underestimates rain-gauge and radar data-sets before 10 Jan, but with a similar rising
 299 trend. Departure arises subsequently between all three data sets, with UKV rainfall
 300 providing the most extreme outcome.

301 Figure 5 shows the rainfall distribution from 1km radar NIMROD rainfall estimation
 302 over the study catchment on the midday of 6th Jan (06/01/2007 13:50) and the first
 303 half day of 10th Jan (10/01/2007 07:10)
 304

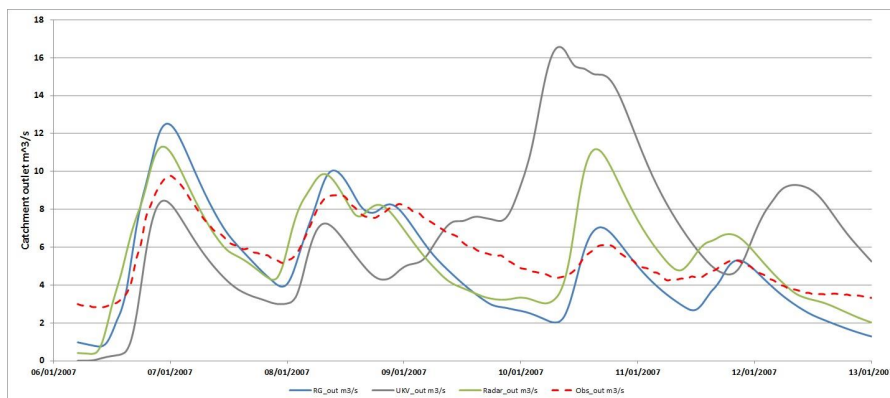


305
 306

307 Figure 5. The snapshots of rainfall distribution extracted from radar images over the
 308 catchment (left was stamped on 13:50 06/01/2007, right was stamped on 07:10
 309 10/01/2007)

310 The simulation of HYPE model with rain-gauge measurement, radar rainfall estimation
 311 and UKV rainfall is illustrated in Figures 6 and 7, for two events occurred in Jan 2007
 312 and July 2007, respectively.

313

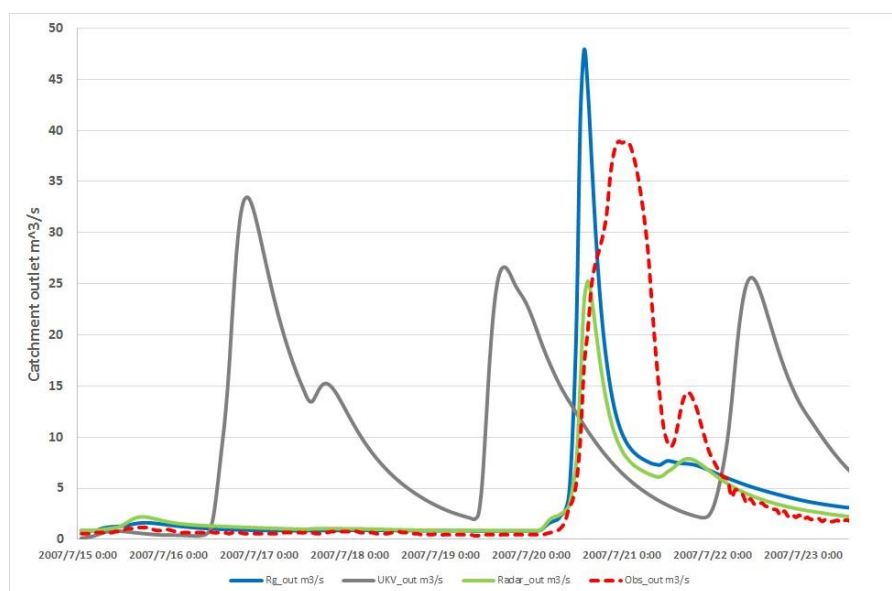


314

315 Figure 6. The comparison of flow simulation in HYPE (Event January 2007)



316 The performance of UKV rainfall in HYPE model simulation for this January 2007 flood
317 event of Figure 6 shows that the peaks and troughs are reasonably well represented
318 against the observed data up to 10 Jan, after which the fourth peak is overestimated,
319 and thus so is the final peak. The radar data suffers likewise, over the final two peaks,
320 which are better captured by the rain-gauge data. The rain-gauge data does however
321 underestimate the observed data output over this period.



322

323 Figure 7. The comparison of flow simulation in HYPE (Event July 2007)

323

324 The second event of July 2007, in Figure 7, would appear to pick up an exaggerated
325 peak in the UKV rainfall through HYPE model simulation after the first day
326 (16/07/2007), which is not reflected in the other data-sets. This early disturbance
327 influences the early undershoot of the observed-data first-peak (at 21/07/2007), and
328 the overshoot of the observed-data second-peak (before 22/07/2007). Notably, rain-
329 gauge data output overshoots the observed-data first-peak, whilst NIMROD radar
330 data output provides an undershoot; both undershoot the observed-data second-
331 peak. This is rather a testing event with only one single main flood event to sharply
332 capture. Clearly, one would need to investigate further in this event instance as to why
333 the early disturbance has arisen for UKV output in this case, and provide more data
334 evidence to prove or refute this particular finding. Further case study events would
335 help clarify this issue, as the Jan 2007 event did not show this up.

336

337

338



339 **4. Conclusions**

340

341 This paper describes a recent effort of integrating both the driver NWP models and
342 the impact analyser – hydrological model on a single HPC platform to support better
343 and refined studies on extreme weather impacts. What distinguishes this study from
344 others is it is first time that modellers are able to simulate the entire system, ranging
345 from the global circulation down to a target catchment for impact study. This study
346 also explores the feasibility of building weather/climate services together with the
347 impact oriented analysis on a single platform; and what can be done if otherwise, for
348 example, how computing resources can be re-arranged for that matter.

349

350 The study finds that when running the entire system on a reasonably powerful HPC
351 platform, the overall time frame does not yet allow for a real-time simulation even
352 without the most complex and demanding data simulation part. It is therefore
353 suggested that the components responsible for large scale simulation, such as global
354 and the European area should remain at national weather service centre where
355 dedicated HPC resources can well deal with the demand as they already have been
356 doing. However, it is still possible to have a high resolution version with less
357 geographical coverage running on a general purpose HPC platform together with the
358 impact analysing model such as a hydrological model and further inundation models.
359 This configuration also allows for finer control and/or tune the models to fit various
360 purposes.

361

362 The other main purpose of this study is to gain the sight of how a common hydrological
363 model can utilise the high resolution precipitation (among others) forecast and
364 simulation in an impact study of extreme weather events. It is encouraging to find that
365 event without fine-tuning, such as using various parameterisation schemes, the
366 coupled hydro-meteorological was still able to capture the major flood peaks with
367 much longer lead time compared with the conventional gauge- or radar-driving
368 forecast (2-3 days vs 2-3 hours). The high resolution UKV rainfall shows some
369 promising agreement with rain-gauge measurements and radar estimation in the first
370 2-3 days in this flood event, both in the catchment average rainfall amount and
371 hydrological simulation in HYPE.

372

373 The study also identified uncertainties associated with precipitation forecast,
374 particularly will increase as the forecast horizon goes beyond 3 days. For example, the
375 latter part of the flood event was not represented well by the HYPE model simulation
376 using the UKV rainfall, compared with those using other sources of rainfall, e.g., radar
377 and raingauges. This is, however, understandable and consistent with our previous
378 studies using other model, see, e.g., Seyoum et al (2013). Apparently, other more



379 complicated uncertainty-aware technique needs to be applied in this model coupling
380 configuration, which, in fact, is the key research topic for further studies.

381

382 Consequently, the following recommendations for future work are made:

383 1. The study needs to be repeated and extended, as more data-sets become available
384 from UKV.

385 2. The impact of the high resolution new radar data needs to be explored in the
386 context of distributed hydrological modelling.

387 3. The UKV rainfall needs to be fully assessed by various lead-times and ensemble
388 simulations, that encapsulate uncertainty generation and propagation through
389 complex 'cloud to catchment' or 'Whole Systems Modelling' concepts.

390

391 **Acknowledgements and declaration**

392 This study was partly-funded under the Knowledge Transfer Partnership scheme,
393 between Fujitsu Europe Laboratories Ltd and Swansea University, acting through
394 support from Innovate UK (KTP009201) and Welsh Government (GON CFI 468). The
395 authors Dr. Zhu and Dr. Echendu would like to gratefully acknowledge this support.

396

397 The study is also partly supported under the on-going research initiative, 'Welsh
398 Extreme Weather Study'. The HPC resources were provided by HPC Wales (HPCW179).
399 Access to the UM/UKV was supported by UK Met Office, under license agreement UM-
400 L0030, which is also gratefully acknowledged. The NIMROD data was provided by the
401 UK Met Office. The catchment soil data was provided by National Soil Resources
402 Institute. The HYPE model and continuous modelling technical support was provided
403 by Swedish Meteorological and Hydrological Institute. Finally, our thanks go to the
404 National Centre for Atmospheric Sciences for their technical support and access to
405 HPC resources.

406

407 **Reference**

408 Cloke, H.L., Pappenberger, F., 2009. Ensemble flood forecasting: A review, *Journal of*
409 *Hydrology*, 375, 613-626.

410

411 Dudhia, J., 1993. A non-hydrostatic version of the Penn State/NCAR mesoscale model:
412 validation tests and simulation of an Atlantic cyclone and cold front. *Mon. Weather*
413 *Rev.*, 121.

414

415 Golding B.W., 1998. NIMROD: a system for generating automated very short range
416 forecasts. *Met Appl*, 5(1), 1-16.

417



- 418 Grell, G., Dudhia, J., Stauffer, D. 1994. A description of the fifth generation Penn
419 State/NCAR Mesoscale Model (MM5), NCAR Technical Note, NCAR/TN-398CSTR, p.
420 117.
421
- 422 Jasper, Karsten, Joachim Gurtz, and Herbert Lang. 2002. Advanced flood forecasting
423 in Alpine watersheds by coupling meteorological observations and forecasts with a
424 distributed hydrological model. *Journal of hydrology* 267(1), 40-52.
425
- 426 Kunstmann, H., Stadler, C., 2005. High resolution distributed atmospheric-hydrological
427 modelling for Alpine catchments. *Journal of Hydrology* 314, 105–124.
428
- 429 Schulla J., Jasper K. 2000. Model Description WASIM-ETH (Water Balance Simulation
430 Model ETH), ETH-Zurich, Zurich.
431
- 432 Seyoum, M., S. van Andel, Y. Xuan and K. Amare (2013): Precipitation Forecasts for
433 Rainfall Runoff Predictions: A case study in poorly gauged Ribb and Gumara
434 catchments, upper Blue Nile, Ethiopia, *Physics and Chemistry of the Earth*, doi:
435 10.1016/j.pce.2013.05.005
436
- 437 Smith, K.T., Austin, G.L., 2000. Nowcasting precipitation — a proposal for a way
438 forward, *Journal of Hydrology*, 239, 34-45.
439
- 440 Wigmosta, M.S., Vail, L.W., Lettenmaier, D.P., 1994. A distributed hydrology-soil-
441 vegetation model for complex terrain. *Water Resour. Res.* 30, 1665–1679.
442
- 443 Westrick, K., Storck, P., Mass, C., 2002. Description and evaluation of a
444 hydrometeorological forecast system for mountainous watersheds. *Weather Forecast*
445 17, 250–262.
446
- 447 Yarnal, B., Lakhtakia, M. N., Yu, Z., White, R. A., Pollard, D., Miller, D. A., & Lapenta, W.
448 M., 2000. A linked meteorological and hydrological model system: the Susquehanna
449 River Basin Experiment (SRBEX). *Global and Planetary Change*, 25(1), 149-161.
450
- 451 Zhu, D.; Cluckie, I. D., 2012. A preliminary appraisal of Thurnham Dual Polarisation
452 Radar in the context of hydrological modelling structure. *Journal of Hydrology*
453 Research 43(4-5), 736-752.
- 454 Zhu, D., Xuan, Y. and Cluckie, I. D., 2014. Hydrological appraisal of operational weather
455 radar rainfall estimates in the context of different modelling structures. *Hydrol. Earth*
456 *Syst. Sci.*, 18, 257-272.



- 457 Xuan, Y., Cluckie, I. D. and Wang, Y. (2009) Uncertainty analysis of hydrological
458 ensemble forecasts in a distributed model utilising short-range rainfall prediction,
459 Hydrology and Earth System Sciences, 13, 293-303.