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Climatic controls on diffuse groundwater recharge across Australia

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Abstract

Reviews of field studies of groundwater recharge have attempted to investigate how climate characteristics control recharge, but due to a lack of data have not been able to draw any strong conclusions beyond that rainfall is the major determinant. This study has used numerical modeling for a range of Köppen-Geiger climate types (tropical, arid and temperate) to investigate the effect of climate variables on recharge for different soil and vegetation types. For the majority of climate types the total annual rainfall had a weaker correlation with recharge than the rainfall parameters reflecting rainfall intensity. In regions with winter-dominated rainfall, annual recharge under the same annual rainfall, soils and vegetation conditions is greater than in regions with summer-dominated rainfall. The relative importance of climate parameters other than rainfall is higher for recharge under annual vegetation, but overall is highest in the tropical climate type. Solar radiation and vapour pressure deficit show a greater relative importance than mean annual daily mean temperature. Climate parameters have lowest relative importance in the arid climate type (with cold winters) and the temperate climate type. For 75 % of all considered cases of soil, vegetation and climate types recharge elasticity varies between 2 and 4, indicating a 20 % to 40 % change in recharge for a 10 % change in annual rainfall Understanding how climate controls recharge under the observed historical climate allows more informed choices of analogue sites if they are to be used for climate change impact assessments.

1 Introduction

Diffuse groundwater recharge is strongly influenced by local vegetation and climate characteristics, which are largely dependant on the climate types. The major climate types, as classified based on the Köppen-Geiger method (Peel et al., 2007), are Tropical (A), Arid (B), Temperate (C), Cold (D) and Polar (E) with further sub-division based on annual and seasonal rainfall and temperature distribution. Based on such a broad classification it is likely that the relationships between groundwater recharge and climate characteristics under these climate types are significantly different. This is due to variation in precipitation, its type (snow or rain), seasonality and intensity but also a result from the effect of vegetation on groundwater recharge. The latter can influence seasonal plants water demands and water use within various climatic conditions and therefore a proportion of rainfall which becomes recharge.

A recent review of field based recharge estimates in Australia investigated the influence of Köppen-Geiger climate types on the relationship between rainfall and recharge (Crosbie et al., 2010a). That study found that there was a significant difference between recharge under winter dominated rainfall when compared to equiseasonal rainfall but the results were confounded by differing soil types and so no strong conclusions could be drawn. It was also found that temperature distinctions within the Köppen-Geiger climate types had no influence over recharge.

Considering the difficulties in designing a field based experiment to investigate the controls that climate characteristics have on recharge, modeling is the preferred method for investigation. Climate change impact studies have provided some information on climatic controls on recharge but generally consider only a small area that has a limited range of climate types. It has been shown that in general rainfall is the most important climate parameter influencing recharge (Allen et al., 2004; Serrat-Capdevila et al., 2007) However many exceptions to this pattern have been reported (Crosbie et al., 2010b; Döll, 2009; Rosenberg et al., 1999) which indicates that other factors can influence this relationship. Among them are the frequency and seasonality of rainfall. Vivoni et al. (2009) demonstrated for a catchment in New Mexico that either an increase in the intensity of summer rainfall, or an increase in the frequency of winter rainfall, can lead to an increase in recharge. In semi-arid areas, higher intensity rainfall can lead to higher episodic recharge even under projections of decreased total rainfall (Crosbie et al., 2012a; Ng et al., 2010)

Most studies that have reported an influence of temperature upon recharge have been for cold climates (energy limited regions), and are associated with variations in snowfall, snowmelt and frozen ground under different temperature conditions (Eckhardt and Ulbrich, 2003; Jyrkama and Sykes, 2007; Okkonen et al., 2010; Vivoni et al., 2009). For warmer climates (water limited regions), Rosenberg et al. (1999) found that recharge could decrease with an increase in rainfall due to higher temperatures and higher evapotranspiration rates in the Ogallala Aquifer. In the Upper Nile Basin it was found that 3 °C increase in temperature along with an increase in rainfall led to reduction in recharge. This was attributed to the effect of higher temperature on evapotranspiration (Kingston and Taylor, 2010).

Furthermore, changes in recharge induced by changes in solar radiation, vapor pressure deficit and CO_2 concentration as they influence water use efficiency by plants and reduce the plants water demands (McCallum et al., 2010). The natural vegetation is largely adapted to local climatic conditions, and some titles of climate types reflect this (e.g. tundra or steppe or savannah). Changes in climatic conditions, which lead to changes in the vegetation and/or their water use efficiency, can have a follow up impact on recharge. For instance, increased recharge was simulated in parts of the Murray-Darling Basin despite a decrease in rainfall, which was attributed to a reduction in transpiration, i.e. the transpiration reduced due to the effect of temperature on vegetation when the optimum temperature for vegetation growth was exceeded (Crosbie et al., 2010b)

Understanding of the climate types effect on recharge as a renewable groundwater resource is particularly important in larger countries (such as Australia, USA, China or Russia) or regions where groundwater management are set to be undertaken by a group of the countries (such as European Union) where the managed areas extends across a number of climate types. As climate change is a growing concern in water management globally, the knowledge of the processes influencing water resources in individual climate zones and also their potential changes are important for development of an adequate climate adaptation strategy.

In addition the effect of climate change is likely to vary in different climate types (IPCC, 2007), so the follow on impact on groundwater resources is likely to be specific in the region of their occurrence. In addition the projected changes in temperature and rainfall may lead to climate type shift, causing further changes in land cover or land use. This may further impact on groundwater resources.

Australia has a highly varied and variable climate, with large regions experiencing arid and semi-arid climates. This has resulted in increasing pressures on groundwater resources as the population has grown and development has taken place (DEWHA, 2010). The reduction in rainfall in many Australian regions over recent years has seen both a reduction in recharge and an increase in groundwater use as surface water resources have become scarce (CSIRO, 2009c). Limited water resource availability already constrains regional development in many parts of Australia, for both industrial and agricultural activities. Understanding how climate characteristics control recharge under the observed historical climate is the key to making projections of how recharge will change under a future climate.

Through numerical modeling, this paper aims to investigate the control that climate characteristics have over groundwater recharge at a point scale and at a continental scale. Specifically it will:

- examine the influence of rainfall, including total rainfall, rainfall intensity and rainfall seasonality,
- examine the influence of other climate variables such as vapour pressure deficit, temperature and solar radiation,
- make recommendations for climate change impact studies.

In the following discussion the term “climate type” will be used when the climate characteristics are discussed, while the term “climate zone” will be used when the geographical extent of a climate type occurrence is referred to.

2 Climate types in Australia

There are 15 Köppen-Geiger climate types in Australia, shown in shown in Fig. 1 along with the key climate characteristics which were used for climate-type identification. However the largest areas are associated with 8 climate types summarized in Table 1: Tropic Savannah (Aw), three Arid types (Bwh, Bsh and Bsk) and four Temperate types (Cfa, Cfb, Csa and Csb).

Climate types are closely linked to rainfall distribution, rainfall intensity and seasonality across the continent (BoM, 2011a). The northern regions (under climate type Aw and partly the northern part of arid Bsh zone; see Table 1) are greatly influenced by monsoons and tropical cyclones, both of which bring heavy rains during summer months. Rainfall in the south-west (Csa/Csb climate types) is dominated by heavy rain events occurring during winter. In the south and south-east, frontal weather systems and east coast lows during winter, in combination with localised troughs, bring occasional heavy rains with prolonged periods of lower intensity rainfall, mainly during winter. The sub-tropical ridge brings dry and stable conditions to large parts of middle Australia. To some extent, the approximate position of the ridge separates the summer-dominated rainfall in the north from the winter-dominated rainfall in the south. The southern regions also experience an overall cooler climate.

General characterization of the main climate types in Table 1 shows a large range of mean annual rainfall within each zone (as a spatial distribution of mean annual rainfall within each climate zone), in which the maximum can be seven times the minimum such as under temperate Cf climates. The lowest range in mean annual rainfall is in Arid Bsk climate. The range of mean temperatures across individual climate zones could be greater than 10 °C as under Temperate Cfb climate but also arid Bwh and Bsh climates, which have a large longitudinal and latitudinal extent. The lowest range in mean temperatures is in Arid Bsk climate which are less than 4 °C.

3 Methods

The methodology of the investigation on the effect of climate characteristics and their combination within individual climate types on diffuse recharge in Australia was based on statistical analysis of the climate data and corresponding modelled recharge data for the regions with the main climate types.

3.1 Recharge modeling

The climatic controls on recharge analysis was conducted using the results of potential recharge estimated using the WAVES model (Zhang and Dawes 1998), the modeling is described in more detail in Crosbie et al. (2012b). This model is a soil-vegetation-atmosphere-transfer model that accounts for the components of an unsaturated zone water balance at a point scale. The WAVES model requires three main datasets: climate parameters, soil and vegetation.

In this study a soil profile of 4 m depth was modelled with a free-draining lower boundary condition. It was assumed that the drainage through the bottom of the model (deep drainage) was potential groundwater recharge. The assumption was made that diffuse recharge in dryland areas is independent of the depth of groundwater below ground surface; this assumption results in errors where the watertable is close to the surface or where the tree roots are deeper than 4 m.

Soil data, including hydraulic characteristics were derived from the ASRIS v1 database (Johnston et al., 2003). The soil profile was modelled as a two-layer system, with 0.5 m topsoil and 3.5 m subsoil with topsoil typically being more permeable than subsoil.

The recharge modelling was undertaken for three vegetation classes: annuals, perennials and trees. The vegetation parameters required by WAVES were taken from the User Manual (Dawes et al., 2004). The annuals (including crops) were modelled as annual pasture, the perennials were modelled as perennial pasture and the trees (including forestry) were modelled as an overstorey of eucalypts with an understorey of perennial grasses. Each climate zone used different parameters for each of the three vegetation types modelled to account for different species present in each climate zone (Crosbie et al., 2012b)

It was observed that the modelled recharge across the continent showed a similar trend in relationship between soils permeability and recharge. As illustrated in Fig. 2 the range of the estimated recharge is the greatest for soils with hydraulic conductivity less than 1 m d^{-1}, expressed as a weighted mean of the topsoil and subsoil hydraulic conductivity (K). For all considered combination of soils and vegetation types, the estimated recharge was not significantly different for soil with K greater than 1.5 m d^{-1}. For this reason and despite a large variety of soils modelled, the results were presented only for three soil types with low, medium and high hydraulic conductivity, approximately defined as 0.01, 0.1 and 1.0 m d^{-1}, respectively. This allowed a better comparison of recharge for different climate types.

The climate data that WAVES requires is rainfall, maximum temperature, minimum temperature, vapour pressure deficit and solar radiation supplied to the model at a daily step. The historical climate data was extracted from SILO (Jeffrey et al., 2001) for the 80 yr period from 1930 to 2009. This period was chosen as it was reported to be the most reliable for climate impact assessments in many regions of Australia (CSIRO, 2008, 2009a,b). The recharge modelling was undertaken using observed climate data at 100 points across Australia, selected to reflect the rainfall gradient within individual climate zones Their position was biased toward areas where groundwater is used from unconfined aquifers

It is important to note that the modelled recharge was only partially validated with field data (Crosbie et al., 2010a, 2011). Therefore the estimates are only considered as an approximation of the actual recharge and so the results are mainly reported in relative terms.

3.2 Climate parameters and diffuse groundwater recharge

A set of analyses were undertaken to investigate the effect of historical climate characteristics on recharge and to clarify the relative importance of annual rainfall, temperature, solar radiation and vapor pressure deficit (VPD) in recharge estimation; the sensitivity of recharge to change in rainfall (recharge elasticity) and the set of rainfall parameters which better define the effect of rainfall intensity on recharge.

3.2.1 Time-series data preparation

The daily series of modeled recharge were aggregated to an annual time series, with the start date being dependent on rainfall seasonality. In areas that experienced summer-dominated rainfall, the aggregation period began in September; otherwise the aggregation period began in March. The Köppen-Geiger climate type was calculated for each of the 100 points using the definitions of Peel et al. (2007).

Rainfall and recharge were aggregated by summation, while other climate variables were averaged. Additional methods for aggregating rainfall were also used to investigate the effect of rainfall intensity, and these involved summation of rainfall events that were:

- above a threshold value (5 mm, 10 mm, 20 mm, 40 mm, 60 mm),
- within a range (0–20 mm, 20–40 mm, 40–60 mm),
- larger than the 95th or 99th percentile rainfall event,
- above a threshold using a moving-average approach.

The moving-average approach was applied to account for the effect of a prolonged rainfall period on recharge. This involved calculating the moving average of the daily rainfall series applying a 7, 14, and 21 day window, and then aggregating the averaged rainfall values that were above a threshold of 2.5 mm d^{-1} or 5 mm d^{-1} (Table 2).

Two measures were used to quantify the independent effect of climate variables on recharge; these were Pearson's Product Moment Correlation Coefficient, and the slope of the linear regression between the climate variable and recharge. Relative importance measures (*Ri*) were then used to assess the contribution of climate variables to explain variance in recharge (within the context of a multiple regression model) (Gromping, 2006). The relative importance measure used in this case was the approach proposed by Lindeman et al. (1980), as recommended by Gromping (2006).

3.2.2 Recharge elasticity

Similar to the concept of the runoff elasticity to precipitation (P), the relationship of elasticity of recharge (R) to P can be estimated as (Schaake 1990):

$$\varepsilon(P,R) = \frac{dR}{dP}\frac{P}{R}. \tag{1}$$

The modeled recharge data and observed rainfall data was used for recharge elasticity analysis. Similar analysis has been done for surface runoff elasticity estimation (Chiew, 2006), but is not commonly considered for groundwater recharge characterization.

4 Results

The modeling results indicate that the recharge values within the same climate type can vary by orders of magnitude depending on specific combinations of soil type and land cover. The difference in recharge can be more than 25-fold due to changes in land cover, i.e. vegetation, and more than 400-fold under various soils. At the same time, annual recharge percentage in rainfall can vary from less than one percent under trees and low permeability soils, to more than 50 % under annuals and highly permeable soils.

However besides soil and vegetation influence, the climate types in Australia have an effect on the relationship between rainfall and recharge. To allow a comparative analysis of the climate type effect on recharge, nine combinations of soils and land cover are presented below, including three soil permeability types (low $K = 0.01\,\mathrm{m\,d^{-1}}$, medium $K = 0.1\,\mathrm{m\,d^{-1}}$ and high $K = 1\,\mathrm{m\,d^{-1}}$) with land covers of annual crops, perennial vegetation and trees.

4.1 Relative importance of climate characteristics in recharge estimation

For most analysed data, rainfall (P) had higher relative importance (Ri^{P}) than other climate variables, including mean annual temperature (T), vapour pressure deficit (VPD) and solar radiation (SR) ($\sum Ri^{\mathrm{SR,VPD},T}$). Figure 3 shows the combined relative importance ($\sum Ri$) of the four climate variables for selected soils and land cover averaged for all climate types. $\sum Ri$ indicates the degree of inter-annual recharge variability explained by variability in climate characteristics and $\sum Ri$ varies from more than 95 % (or 0.95 in Fig. 3) to less than 30 % (or 0.3 in Fig. 3). In all climate types the relative importance of annual climate characteristics reduces from annual to perennial vegetation and even less for trees, but also under soils with lower permeability. The lower $\sum Ri$ indicates that other variables not incorporated into annual means have a greater impact on recharge in such conditions.

The only climate type where Ri of rainfall is lower than $\sum Ri^{\mathrm{SR,VPD},T}$ is the Arid (Bsk) climate under low permeable soils and perennial vegetation or trees. Under this climate type $\sum Ri$ in recharge estimation are the overall lowest across all other climate types (less than 70 % or 0.7 as shown in Fig. 3).

The highest values of $\sum Ri^{\mathrm{SR,VPD},T}$ are related to i) the climates with summer dominated rainfall (Aw and Cfa) and ii) the cooler climate types (Cfb and Csb) where the range of the solar radiation and temperature are greater than elsewhere and their mean values are lower among all considered climate types.

Compared to VPD and solar radiation, mean annual temperature has the lowest relative importance under all climate types, which is likely to be due to a relative consistency of mean annual temperature within individual climate type over the simulation period (BoM, 2011b) However the high values of T relative importance is related to a cooler climate (Cfb), climate with winter dominated rainfall (Csa) but also in Arid climate (Bsh). The latter climate zone covers the large area with a greatest variability in annual temperature as shown in Fig. 1b. On average climate variables other than rainfall explain 15 % of the variability of recharge with a maximum of 30 %.

When individual points within each climate type are considered (Fig. 4), the rainfall importance in recharge estimation reduces under lower annual rainfall conditions within each individual climate type but also across the entire data set. For the latter Ri^P was found to be lowest for areas with annual rainfall less than 700 mm. In agreement with the discussion above the relative importance of rainfall under all soil/vegetation is the lowest under the arid climate Bsk, but also in the areas neighbouring with the Bsk climate zone. This includes the western areas of the temperate climate zones (Cf) and the southern areas of the Bsh arid climate zone. When annual rainfall is lower than 400 mm, it appears that the relative importance of rainfall increases, e.g. desert climate type (Bwh).

Within individual climate types, Ri^P changes are related to the distribution of annual mean rainfall within the zone and are influenced by the other climate types located in the neighbouring regions. For instance, temperate climate Cfa covers the eastern regions of the country stretching from the north-east to south-east. It is characterised by the greatest variation in rainfall and its relative importance in recharge estimation. Figure 4 shows that the extreme Ri^P values within this climate type is similar to Ri^P within tropical climate (Aw) for the most northern modelled points, as well as Ri^P within arid climate (Bsk) for the most southern modelled points.

Another example is related to Ri^P variability within the arid climate Bsh, which is also greatly influenced by the position of this climate zone in relation to other climate zones. When neighbouring with tropical climate (Aw), Ri^P has the higher values, when neighbouring with other arid climates (Bsh) – Ri^P has the lowest values.

The variability in Ri^P, indicated by a spread of points in Fig. 4 increases from more to less permeable soil (note that heavier soils with $K = 0.01\,\mathrm{m\,d^{-1}}$ are not present in Cfb and Csa/b climate zones). The annual rainfall, which corresponds to the minimum Ri^P, is lower under tree land cover, which is about 400 mm against 500 mm under annuals. Under the same annual rainfall, Ri^P in recharge estimation is greater for temperate climate types with winter dominated rainfall (Cs).

The reduction in annual rainfall and its relative importance leads to an increase in recharge sensitivity to other climate parameters considered in this study (Fig. 4). Under similar annual rainfall $\sum Ri^{\mathrm{SR,VPD},T}$ is greater for the climate types with summer dominated rainfall or cooler climate.

As rainfall is the major factor in recharge, the following Sects. 4.1, 4.2 and 4.3 examine the relationship between rainfall and recharge only.

4.2 Relationship between annual rainfall and modelled recharge

Reflecting the high importance of rainfall in recharge estimation, the correlation is strong between rainfall and recharge for the majority of cases; $R_P^2 > 0.7$ for 82 % cases of soil/vegetation/climate type combinations. The strongest correlation between modelled recharge and annual observed rainfall is in the areas with high annual rainfall under tropical savannah (Aw) and temporal climate types without dry season (Cf) (Fig. 6). Correlation between rainfall and recharge becomes weaker under climate types with overall lower annual rainfall as its importance in recharge estimation also drops for rainfall below 700 mm (Fig. 4). As a result arid climate types, particularly Bsk, are characterized by the overall weakest relationship between rainfall and recharge.

Correlation between rainfall and recharge is generally weakened under perennial vegetation and trees and under soil with lower permeability, where recharge is relatively low. Under similar annual rainfall R_P^2 is greater in the climate types with winter dominated rainfall (Cs) for all combinations of soil and vegetation.

The stronger correlation between rainfall and recharge was found when the higher percent of annual rainfall becomes recharge (Fig. 6). The general pattern for all combinations of soil and vegetation was a reduction in R/P with a reduction in annual mean rainfall. As expected, the percentage of the annual rainfall which becomes recharge reduces under soils with lower permeability and land cover from annual to perennial to trees. However this general pattern is also influenced by the climate type. Under annual crop R/P is greater in the tropics (Aw) and Arid climate type (Bsh), neighboring with Aw climate zone. However under perennial vegetation and trees the high percentage of the annual rainfall which becomes recharge is associated with the climate types where winter rainfall dominates.

Sensitivity of recharge to changes in rainfall, defined as recharge elasticity ε_R, increases under conditions which cause overall less recharge. This includes low rainfall, low soil permeability and under perennial and tree land cover. However for 75 % of all combinations of soil, vegetation and climate types ε_R varies between 2 and 4, indicating a 20 % to 40 % change in recharge for a 10 % change in annual rainfall (Fig. 8).

The exception to this pattern is related to the cases where recharge was estimated under soil with particularly low hydraulic conductivity ($K = 0.01\ \mathrm{m\,d^{-1}}$) and perennial vegetation or trees as a land cover. In such conditions, the overall recharge is low, and only significant changes in rainfall can lead to changes in recharge.

4.3 Rainfall seasonality

The effect of the rainfall seasonality on recharge rates is illustrated in Fig. 9. It shows the relationship between annual rainfall and annual modeled recharge for all locations which fall within the temperate climate with winter dominated rainfall (Cs) and within the tropical climate with summer dominated rainfall (Aw) for perennial vegetation and soil with K of $1\ \mathrm{m\,d^{-1}}$ On average the estimated recharge is greater under Cs climate type for similar values of annual rainfalls. Similarly for Cf climate types the recharge is greater for Cfb where under overall equiseasonal conditions the proportion of winter rainfall is greater. This relationship may also be also influenced by overall cooler conditions under Cfb climate (Table 1).

The effect of rainfall seasonality of the estimated recharge is also evident in Fig. 10 showing the mean annual recharge values under two soil types three land cover types and all climate types as a function of the annual rainfall and the proportion of annual rainfall which falls during summer. For perennial vegetation and trees as a land cover the annual recharge values reduce, when the proportion of summer rainfall increases. However under annual vegetation the trend is reversed: the annual recharge values increase along with an increase in the proportion of summer rainfall.

Figure 9 also shows that the relationship between recharge and rainfall is not linear, the percentage of annual rainfall which becomes recharge increases during the years with the high annual rainfall. This is due to higher rainfall intensity or duration of wet periods during the wetter years, which is explored further in the following section.

4.4 Rainfall intensity

Despite overall high correlation between annual rainfall and recharge, it is often weaker than the correlation between recharge and the sum of high intensity rainfall on an annual basis. Figure 11 shows relationship between R_P^2 and R_{Pi}^2, where R_P^2 is the coefficient of correlation between recharge and total annual rainfall (P); and R_{Pi}^2 is the coefficient of correlation between recharge and the annual aggregation of rainfall with higher intensity (Pi), when different approach to intensity assessment was used, including the threshold value (5 mm, 10 mm, 20 mm, 40 mm, 60 mm) (Fig. 11a), rainfall bands (0–20 mm, 20–40 mm, 40–60 mm) (Fig. 11b), percentile of the rainfall events (Fig. 11c) and a moving-average approach (Fig. 11d).

The highest overall correlation was found to be between annual recharge and the annual aggregation of rainfall using a moving average daily rainfall approach. The latter accounts for the high intensity rainfall events and the prolonged periods of smaller rainfall events simultaneously It was found that aggregated 95 percentile daily rainfalls on annual basis show better correlation with recharge than 99 percentile daily rainfalls. However $R^2_{p_{99}}$ was greater than R^2_P under arid climate types, but also under highly permeable soils and tree land cover in Csa and Cfb. When the rainfall intensity thresholds are considered the daily rainfall greater than 20 mm aggregated on an annual basis provides a better correlation with annual recharge than total annual rainfall.

It appears that there is less improvement in correlation between recharge and the annual aggregation of rainfall with higher intensity compared to R^2_P when the latter is particularly high ($R^2_P > 0.95$) or low ($R^2_P < 0.3$). At the high R^2_P values the annual rainfall is overall characterised by higher intensity, while at the lower R^2_P rainfall is likely to be of lower intensity or other than rainfall climate variables play a more important role in recharge estimation.

When individual climate types are considered, the difference between R^2_P and R^2_{Pi} is not significant when recharge is estimated under soils with lowest hydraulic conductivity ($K = 0.001\ \mathrm{m\,d^{-1}}$) for all climate types and vegetation covers. For other soil types the least difference between R^2_P and R^2_{Pi} (on averaged $< 5\,\%$) is under annual vegetation (with exception of Csa climate zone), but also under tropical Aw and arid Bsk climate types (Fig. 12a,b). The latter represent the extremes in rainfall intensity across all considered climate types with the highest intensity being typical for the tropic Aw climate type, and the lowest – for the arid Bsk climate type (Fig. 13a). This also reflects a general trend in reduction of rainfall intensity from the north to the south of the country (Fig. 13b).

5 Discussion

5.1 Climate controls on recharge

Climatic controls on recharge have not been explicitly addressed in the literature on recharge estimation. This is due to the limitation of the available recharge data as well as its dependency on the techniques used for recharge estimation (Crosbie et al., 2010a; Petheram et al., 2002).

It appears that there are certain trends in the relationship between recharge and climate characteristics and some of these trends are equally relevant across all climate types, but others are more specific for the individual conditions.

In agreement with other published data (Petheram et al., 2002), total annual rainfall was found to be the main factor influencing diffuse recharge across all considered climate types. In general, a reduction in rainfall tends to weaken the correlation between rainfall and recharge as well as reduces the rainfall's importance in recharge estimation. Under low rainfall the importance of other climate parameters on recharge rises. Among them VPD and solar radiation appear to be the dominant factors, while annual mean temperature has the lowest importance in recharge estimation within individual climate types.

An increase in rainfall intensity leads to an increase in (i) recharge, (ii) a proportion of rainfall that becomes recharge, (iii) the relative importance of rainfall, and a reduction in the relative importance of other climate parameters in recharge estimation. Reduction in annual rainfall commonly coincides with a reduction in rainfall intensity, which has a profound impact on recharge. This is true for both individual modelling locations and climate types at the continental scale.

However there are some exceptions to this trend which are related to rainfall seasonality and cooler climate types. For the climate types with summer dominated rainfall (e.g., Aw) the main recharge season is coincidental with higher vegetation water demand and as such leads to a greater influence of VPD, solar radiation and temperature on recharge. As a result the proportion of rainfall that becomes recharge is generally lower under summer dominated rainfall for areas under perennials and tree type vegetation. In the climate types with winter dominated rainfall (such as Cs) the rainfall period coincides with low water demand and vegetation growth, and as a result R/P under perennial vegetation and trees is higher within this climate type compared to Aw climate type. However this trend is not observed under annual vegetation, where R/P in

Aw is still greater, which is likely to be related to high rainfall intensity under this climate type.

Under the cooler climates (such as Cfb) vegetation growth and water use is sensitive to the changes in annual minimal temperatures and as such this is the only climate type where relative importance of temperature was higher than relative importance of vapour pressure deficit and solar radiation.

Changes in recharge are largely proportional to changes in rainfall but not equal. It appears that changes in annual rainfall lead to 2- to 4-fold greater changes in recharge. The recharge elasticity also increases when land cover include perennial vegetation and trees. However depending on the climate type the sensitivity of the recharge to changes in rainfall may vary. Relative changes in recharge (ε_{R}) are greater under the conditions which generally cause a reduction in recharge, and hence the changes in rainfall may have a greater impact on recharge in climates with lower rainfall (e.g., Bsh, Bsk or Bwh) and in the regions with higher proportion of rainfall during the summer season (e.g., Bsh, Cf).

The increase in $\sum Ri^{\mathrm{SR,VPD},T}$ in the climate zones with rainfall less than 700 mm indicates the importance of vegetation in controlling recharge, as these climate variables influence the water use by vegetation.

5.2 Implications for groundwater resources management

Understanding of the relationship between recharge and climatic conditions, in addition to soil type and land cover type, have a regional application to groundwater management.

Groundwater management in Australia is largely characterised by a focus on the establishment of groundwater extraction limits for Groundwater Management Units (GMUs) based on sustainable yield estimates for defined groundwater systems. Such estimates are largely based on an assessment of the proportion of renewable groundwater resources which can be abstracted for consumptive use. Where groundwater models are not available, and that is in the vast majority of cases, sustainable yield assessment is commonly based on expert estimation of renewable groundwater resources, defined as a proportion of rainfall. A constant proportion is commonly set for an aquifer regardless of inter-annual variation in rainfall (DEWHA, 2009). Observations and modelling, however, have shown that this method has a number of shortcomings. In particular, this analysis has revealed a non-linearity in the recharge to rainfall ratio (R/P) for any given location due to variability in rainfall intensity, or the number of consecutive rain days. Furthermore inter-annual rainfall variability is magnified 2- to 4-times in recharge variability. The results of the current analysis indicate that assuming a constant $R:P$ will lead to an overestimation of renewable groundwater resources for lower annual rainfall periods and their underestimation for higher annual rainfall periods. This in turn indicates that for an adequate water resources assessment there is a need to account for historical variability of climatic conditions and their effect on renewable groundwater resources.

5.3 Implications for climate change studies

Under changing climate conditions, which are mainly caused by global warming, it is likely to expect that climate zones will shift with a follow-on effect on renewable groundwater resources via recharge.

High relative importance of rainfall indicates that the changes in rainfall may have a greater impact particularly if climate change leads to changes in rainfall intensity and seasonality. A shift in rainfall seasonality may cause a reduction in annual recharge in the climate types dominated by winter rainfall as has been observed in the south-west of the continent (Charles et al., 2010) or an increase in recharge in the climate types dominated by summer rainfall. Projected changes in rainfall seasonality under a future climate have been shown to produce recharge projections in accordance with that described above (Vivoni et al., 2009). The elasticity of the rainfall-recharge relationship and the non-linear nature of the annual R/P relationship are mirrored in the results of climate change impact studies (Barron et al., 2011; Crosbie et al., 2010b, 2012a). These observations demonstrate that understanding the climatic controls on recharge under the historical climate is the key to making projections of recharge under a future climate.

Some researchers have adopted the concept of potential changes in climate type for analysis of the climate change impact on groundwater recharge (Leterme et al., 2012). Within this approach the actual meteorological data from instrumental analogue stations are used for the future climate projection in the areas where the climate type shift is likely to introduce the climate type currently occurring within the region of the analogue stations. The outcome of undertaken analysis may be useful for such applications, indicating that relationship between recharge and climate parameters may be quite similar in the neighbouring climate zones.

6 Conclusions

The reported results of the carried out analysis allow defining certain trends in a climatic control on diffuse groundwater recharge across Australia.

- Annual rainfall is a major factor influencing recharge. However, for the majority of the considered climate types the total annual rainfall had a weaker correlation with recharge than the rainfall parameters reflecting rainfall intensity.
- Annual recharge is more sensitive to daily rainfall intensity in regions with winter-dominated rainfall, where it is also less sensitive to absolute changes in annual rainfall.
- In regions with winter-dominated rainfall, annual recharge under the same annual rainfall and soils conditions is less than in regions with summer-dominated rainfall for perennial vegetation and trees as land cover. However this trend is not observed under annual vegetation.
- Relative importance of annual rainfall in recharge estimation reduces under lower rainfall conditions, and along with that there is an increase the relative importance

of other climate parameters in recharge estimation (temperature, solar radiation and vapour pressure deficit). The effect of climate parameters other than rainfall on recharge is greater under climate types with summer dominated rainfall and under cooler climate types.

- An increase in rainfall intensity leads to an increase in recharge, a higher proportion of rainfall that becomes recharge, an increase in the relevant importance of rainfall, and a reduction in the relevant importance of other climate parameters in recharge estimation.
- There is a non-linear relationship between recharge and rainfall, which is likely due to the effect of rainfall intensity or duration of consecutive days with rainfall. Therefore, a proportion of recharge in annual rainfall (R/P) is not likely to be a constant – even under the same land cover and soil type.
- Annual changes in recharge are largely proportional to annual changes in rainfall but are not equal. It has been demonstrated that changes in annual rainfall lead to 2- to 4-fold greater changes in recharge.

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Table 1. Characteristics of selected climate zones.

Climate types		Rainfall		Rainfall seasonality: summer rainfall as proportion of annual		Mean temperature	
		Annual (mm)	Range (mm)	Annual	Range	Annual (˚C)	Range (˚C)
Tropical savannah	Aw	1125	758–2038	0.92	0.67–0.96	26.7	22.3–29.5
Arid desert hot	BWh	254	138–417	0.67	0.26–0.88	22.5	18.0–28.2
Arid steppe hot	BSh	483	225–870	0.75	0.15–0.96	23.4	18.0–29.7
Arid steppe cold	BSk	342	235–498	0.44	0.26–0.69	16.9	14.2–18.0
Temperate without dry season with hot summer	Cfa	762	439–3493	0.63	0.37–0.79	18.6	14.1–23.4
Temperate without dry season with warm summer	Cfb	953	433–3219	0.49	0.33–0.72	12.8	6.7–18.5
Temperate with dry hot summer	Csa	557	341–1517	0.22	0.15–0.77	17.5	14.8–21.5
Temperate with dry warm summer	Csb	665	347–1200	0.30	0.15–0.40	15.0	9.3–17.3

Table 2. The minimum daily rainfall event fully accounted in moving-average analysis over the set of considered periods and daily thresholds.

Moving average period and threshold (mm d^{-1})		Daily rain as a single event over the defined period (mm)
7 days	2.5	17.5
	5.0	35
14 days	2.5	35
	5.0	70
21 days	2.5	52.5
	5.0	105

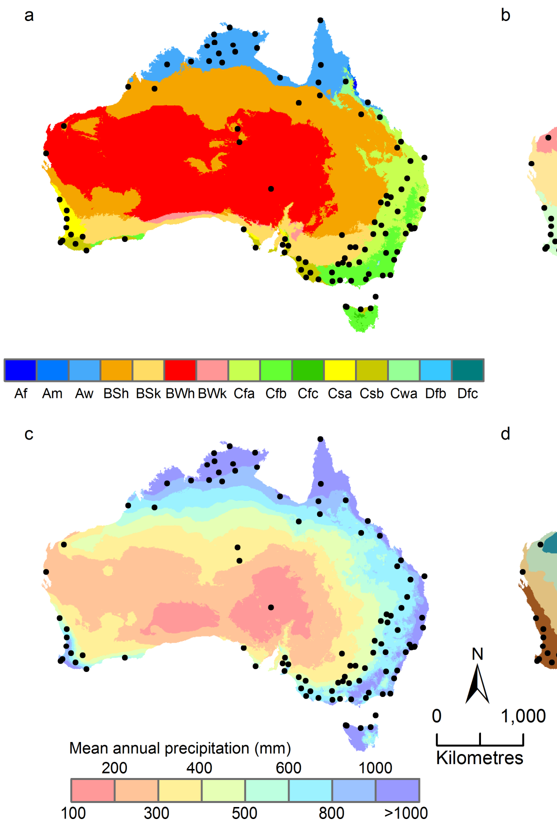

Fig. 1. Current climate parameters: **(a)** Köppen-Geiger climate zones; **(b)** mean annual temperature; **(c)** mean annual precipitation; **(d)** proportion of summer precipitation; the points indicate the location where climate data was obtained for recharge modeling. Adapted from Barron et al. (2011).

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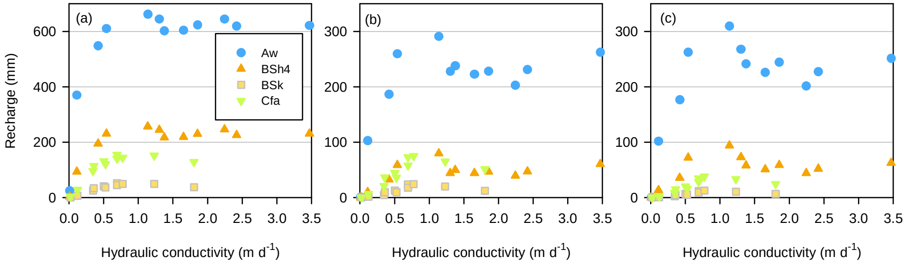


Fig. 2. Relationship between modelled recharge and soil hydraulic conductivity (as geometric mean between hydraulic conductivity of topsoil and subsoil).

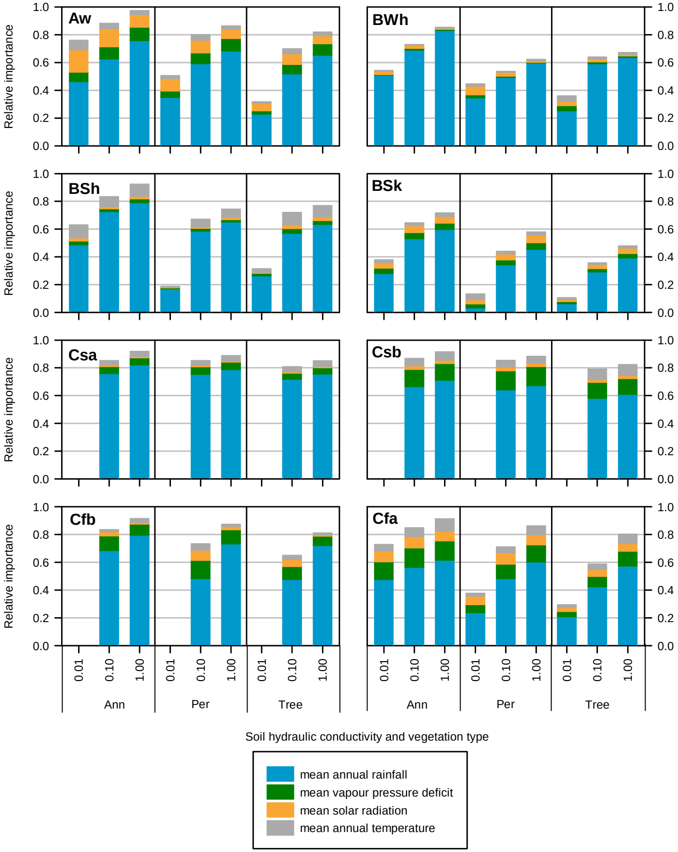


Fig. 3. Relative importance of climate characteristics within considered climate types under various soil and vegetation (note that in the area of Csa, Csb and Cfb clay reach soils do not present).

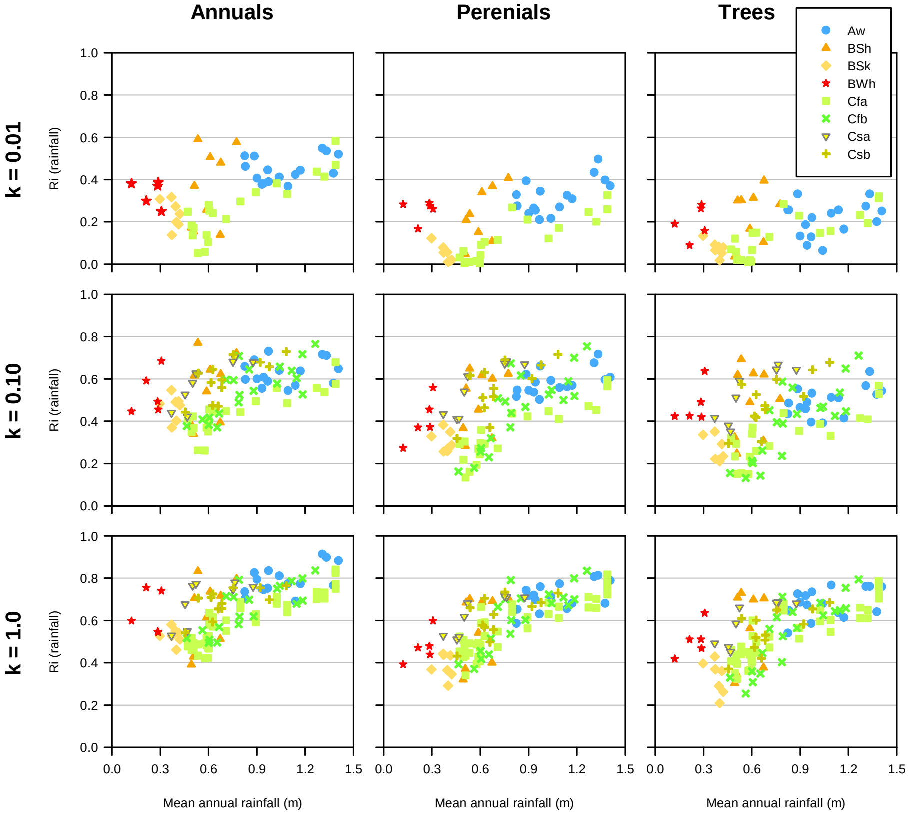


Fig. 4. Relationship between mean annual rainfall and its relative importance in recharge estimation within considered climate types for perennial vegetation and soil with $K \sim 1\,\mathrm{m\,d^{-1}}$.

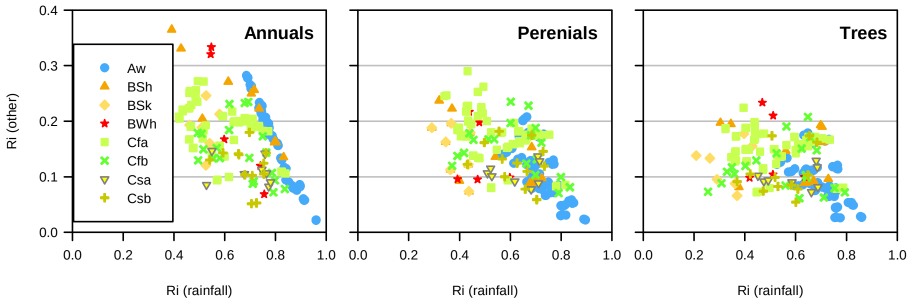


Fig. 5. Relationship between mean annual rainfall and relative importance of T, VPD and solar radiation (cumulatively) within considered climate types for soil with $K \sim 1\,\mathrm{m\,d^{-1}}$.

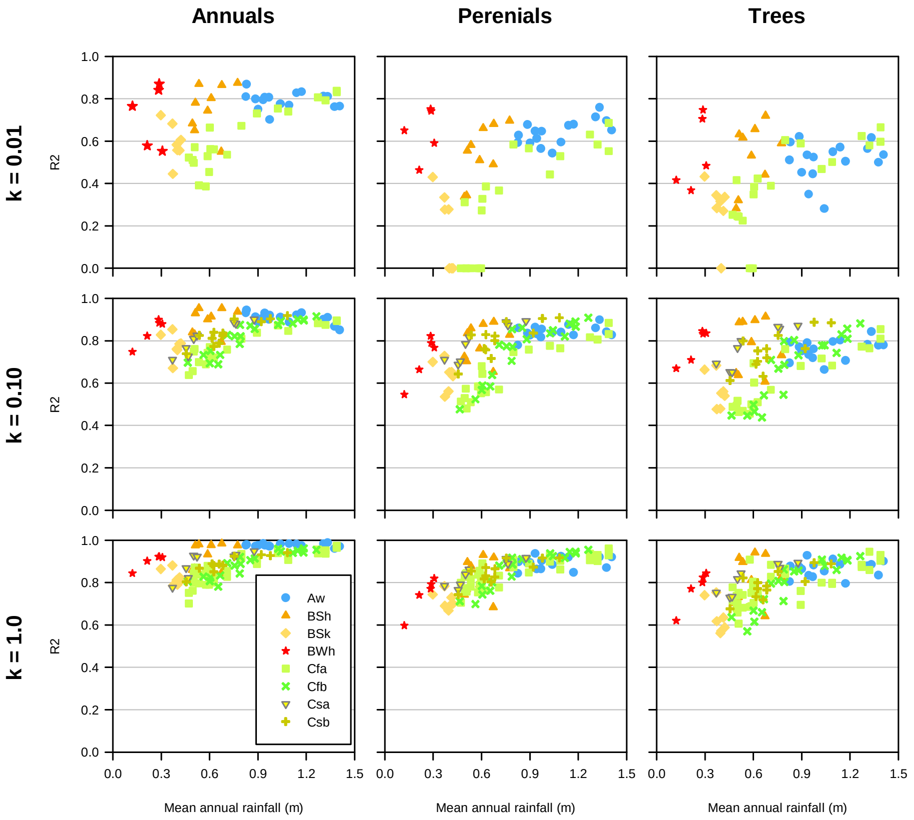


Fig. 6. Relationship between mean annual rainfall (*P*) and the coefficient of correlation between mean annual recharge (*R*) and *P* for all modelled locations.

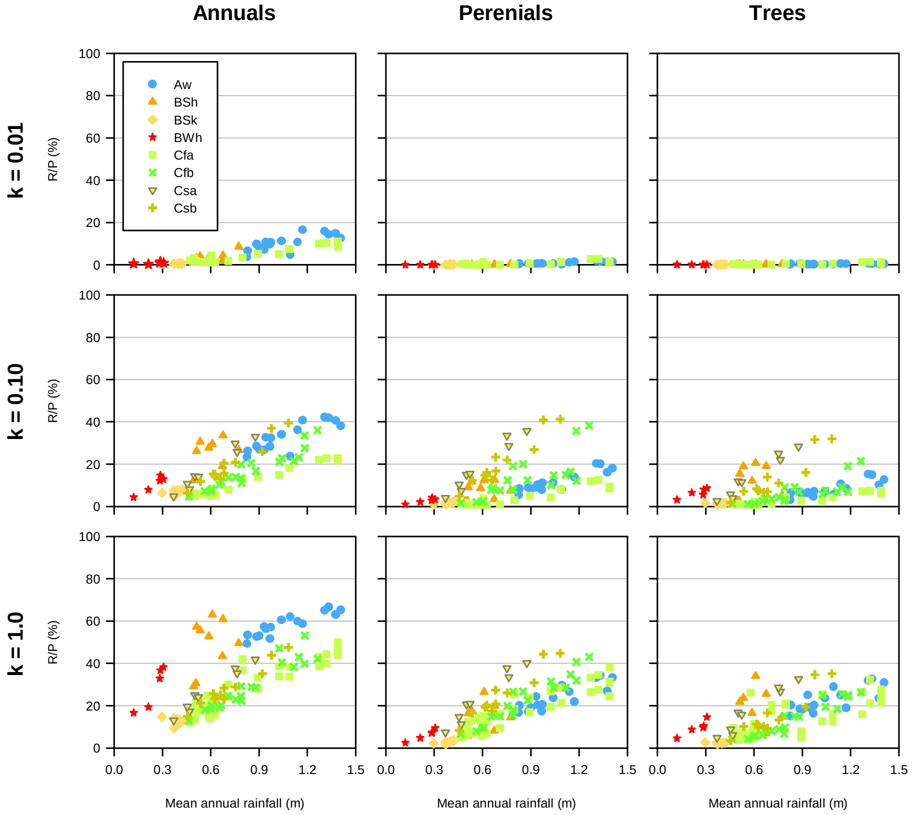


Fig. 7. Relationship between mean annual rainfall (P) and percent P which becomes recharge (R/P) for all modelled locations.

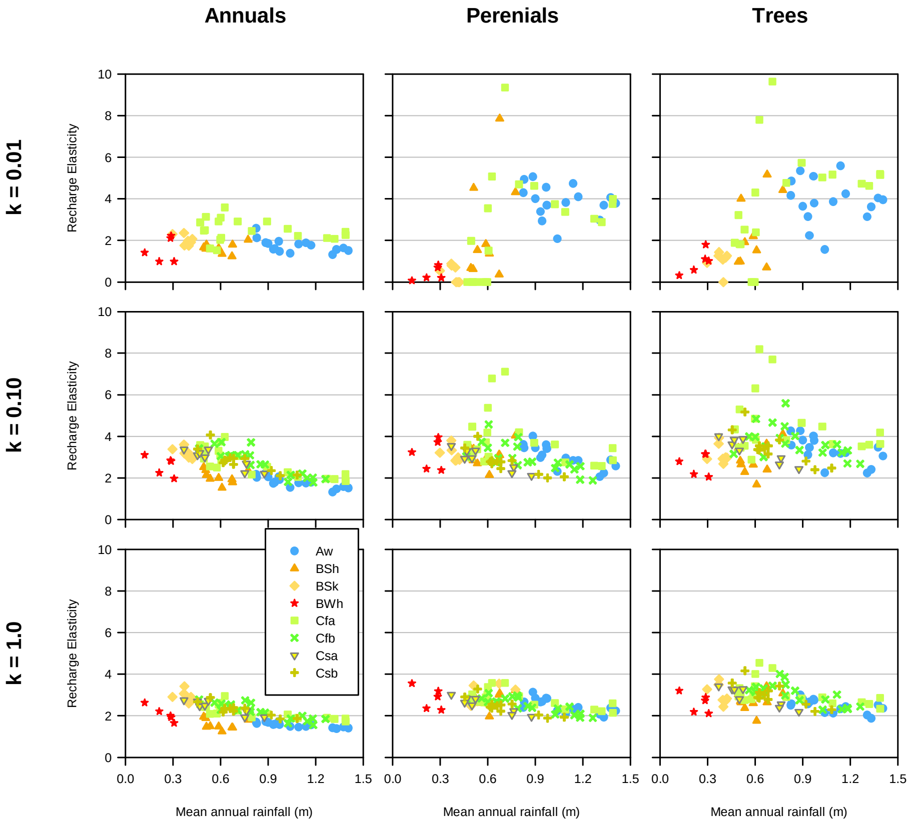


Fig. 8. Relationship between mean annual rainfall and recharge elasticity for all modelled locations.

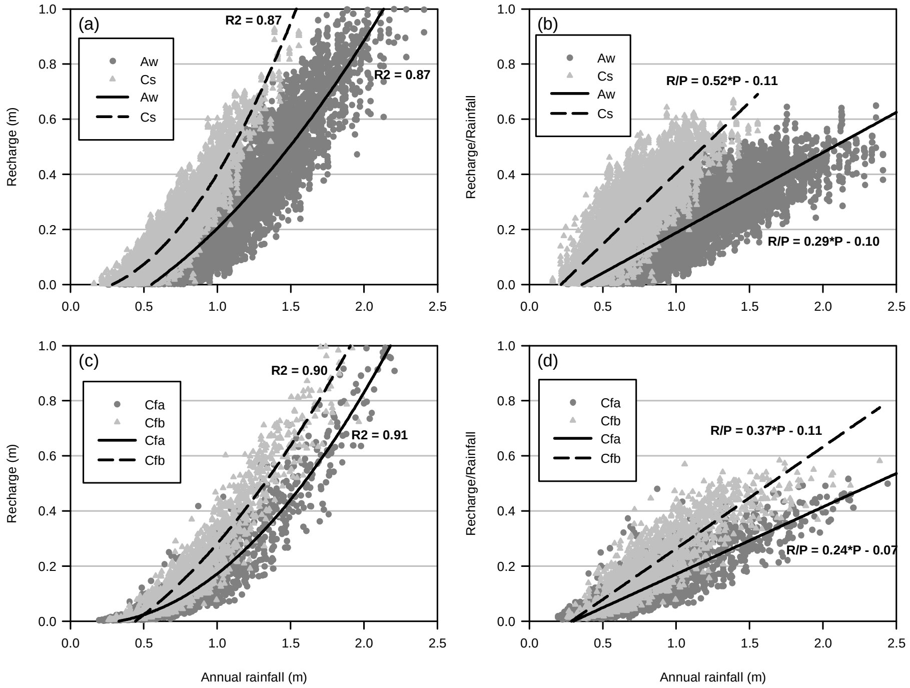
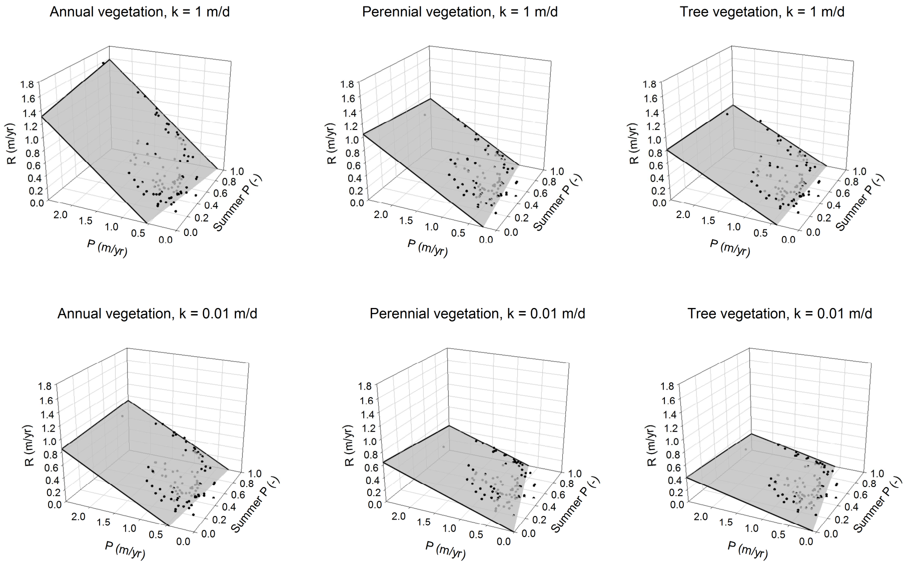


Fig. 9. Relationship between recharge and annual rainfall **(a, c)** and between percent recharge in rainfall and annual rainfall **(b, d)** for perennial vegetation and soil with $K \sim 1\,\mathrm{m\,d^{-1}}$; (**a** and **b**) summer dominative rainfall (Aw – tropical savannah) and winter dominated rainfall (Cs – temperate climate with dry summer); (**c** and **d**) temperate climate without dry summer (Cfa – hot summer; Cfb – warm summer).



Annual vegetation, k = 1 m/d

Perennial vegetation, k = 1 m/d

Tree vegetation, k = 1 m/d

Annual vegetation, k = 0.01 m/d

Perennial vegetation, k = 0.01 m/d

Tree vegetation, k = 0.01 m/d

Fig. 10. Relationship between annual mean recharge (R), annual mean rainfall (P) and proportion of annual precipitation which falls during summer.

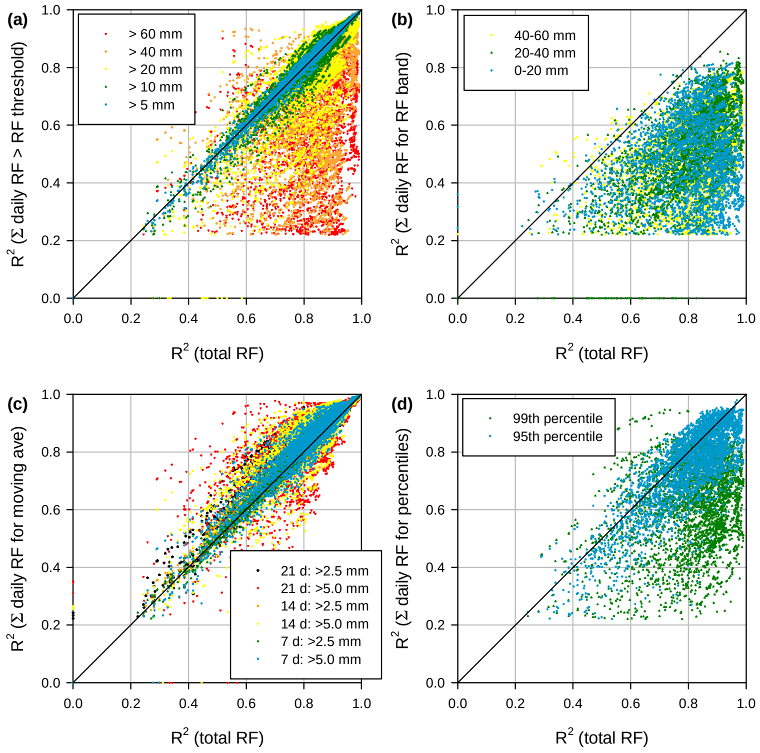


Fig. 11. Comparison of the correlation coefficients (R^2) for the recharge and rainfall relationship: (x-axis) with total annual rainfall and (y-axis) derived annual rainfall characteristics: **(a)** sum of the daily rainfall above the identified thresholds; **(b)** sum of the daily rainfall within the identified bands; **(c)** sum of the daily rainfall as moving average with identified intervals and daily thresholds; and **(d)** sum of 95th and 99th percentile daily rainfall.

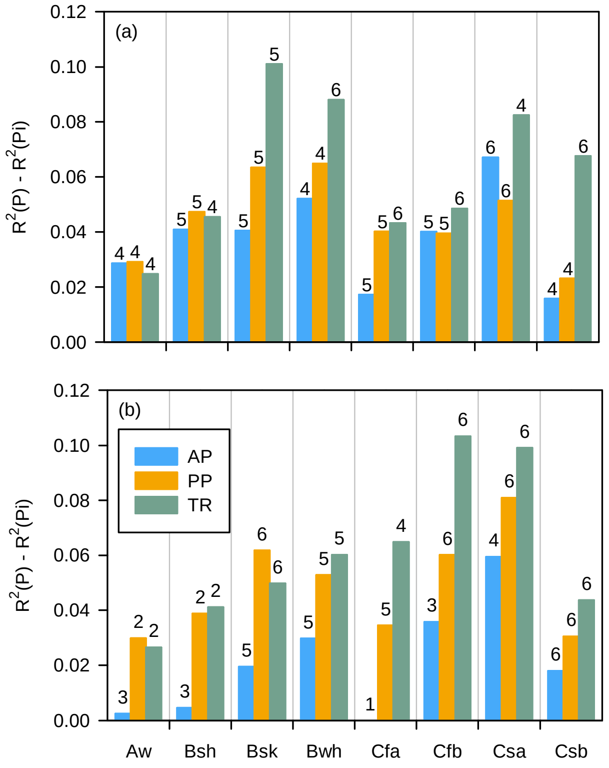


Fig. 12. Differences between R_P^2 and R_{Pi}^2 for all climate, vegetation types and two soil types: **(a)** $K = 1\,\mathrm{m\,d^{-1}}$ and **(b)** $K = 1\,\mathrm{m\,d^{-1}}$. The data labels indicate the method of annual rainfall aggregation resulting in a highest differences between R_P^2 and R_{Pi}^2: 1 – annual rainfall, 2 – a daily rainfall threshold greater than 20 mm, and 3–6 all above a threshold using a moving-average: 3 – 7 day window and a threshold of $2.5\,\mathrm{mm\,d^{-1}}$, 4 – 14 day window and a threshold of $5\,\mathrm{mm\,d^{-1}}$, 5 – 21 day window and a threshold of $2.5\,\mathrm{mm\,d^{-1}}$ and 6 – 21 day window and a threshold of $5\,\mathrm{mm\,d^{-1}}$.

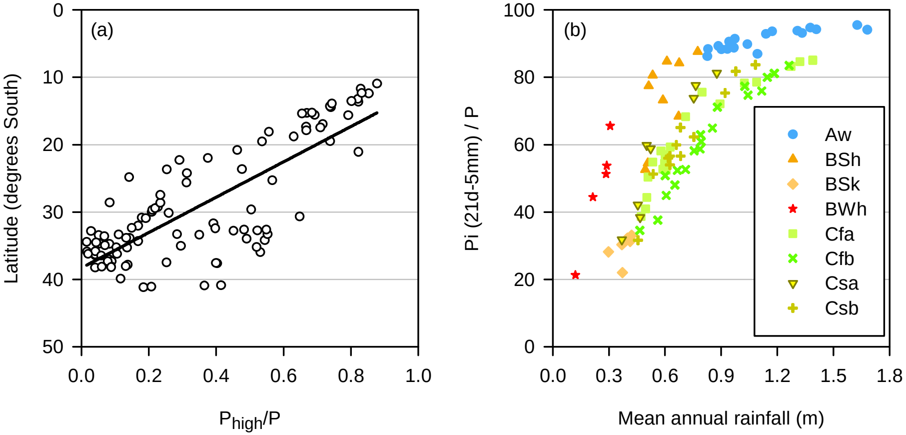


Fig. 13. Relationship between proportion of annual rainfall with high intensity in total annual rainfall **(a)** and **(b)** changes to proportion of high intensity rainfall in total annul rainfall from north to south of the continent, both for moving average over 21 days with daily 5 mm threshold.