

1 **Distributed specific sediment yield estimations in Japan**  
2 **attributed to extreme-rainfall-induced slope failures under a**  
3 **changing climate**

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11  
12 **Abstract**

13 The objective of this study was to estimate the potential sediment yield distribution in Japan  
14 attributed to extreme-rainfall-induced slope failures in the future. For this purpose, a  
15 regression relationship between the slope failure hazard probability and the subsequent  
16 sediment yield was developed by using sediment yield observations from 59 dams throughout  
17 Japan. The slope failure hazard probability accounts for the effects of topography (as relief  
18 energy), geology and hydro-climate variations (hydraulic gradient changes due to extreme  
19 rainfall variations) and determines the potential slope failure occurrence with a 1-km  
20 resolution. The applicability of the developed relationship was then validated by comparing  
21 the simulated and observed sediment yields in another 43 dams. To incorporate the effects of  
22 a changing climate, extreme rainfall variations were estimated by using two climate change  
23 scenarios (the MRI-RCM20 Ver.2 model A2 scenario and the MIROC A1B scenario) for the  
24 future and by accounting for the slope failure hazard probability through the effect of extreme  
25 rainfall on the hydraulic gradient. Finally, the developed slope failure hazard-sediment yield  
26 relationship was employed to estimate the potential sediment yield distribution under a  
27 changing climate in Japan.

28 Time series analyses of annual sediment yields covering 15-20 years in 59 dams reveal that  
29 extreme sedimentation events have a high probability of occurring on average every 5-7

1 years. Therefore, the extreme-rainfall-induced slope failure probability with a five-year return  
2 period has a statistically robust relationship with specific sediment yield observations (with  $r^2$   
3 = 0.65). The verification demonstrated that the model is effective for use in simulating  
4 specific sediment yields with  $r^2 = 0.74$ . The results of the GCM scenarios suggest that the  
5 sediment yield issue will be critical in Japan in the future. When the spatially averaged  
6 sediment yield for all of Japan is considered, both scenarios produced an approximately 17-  
7 18% increase around the first half of the 21<sup>st</sup> century as compared to the present climate. For  
8 the second half of the century, the MIROC and MRI-RCM20 scenarios predict increased  
9 sediment yields of 22% and 14%, respectively, as compared to present climate estimations.  
10 On a regional scale, both scenarios identified several common areas prone to increased  
11 sediment yields in the future. Substantially higher specific sediment yield changes (over 1000  
12  $\text{m}^3/\text{km}^2/\text{year}$ ) were estimated for the Hokuriku, Kinki and Shikoku regions. Out of 105 river  
13 basins in Japan, 96 will have an increasing trend of sediment yield under a changing climate,  
14 according to the predictions. Among them, five river basins will experience an increase of  
15 more than 90% of the present sediment yield in the future. This study is therefore expected to  
16 guide decision-makers in identifying the basins that are prone to sedimentation hazard under a  
17 changing climate in order to prepare and implement appropriate mitigation measures to cope  
18 with the impacts.

19

## 20 **1 Introduction**

21 The latest report from the Intergovernmental Panel on Climate Change (IPCC AR4 by Parry  
22 et al., 2007) along with many other studies predicts increases in the frequency and intensity of  
23 heavy rainfall in high-latitude areas under enhanced greenhouse conditions (Jones and Reid,  
24 2001; Palmer and Raisanen, 2002; Fowler et al., 2005). Several studies in Japan have  
25 supported this conclusion by illustrating long-term increases in rainfall intensity (Iwashima  
26 and Yamamoto, 1993; Kajiwara et al., 2003) and frequency (Suzuki, 2004) in the 20<sup>th</sup> century  
27 and by also predicting increases in the total rainfall amount by the end of the 21<sup>st</sup> century  
28 (Gunawardhana and Kazama, 2010).

29 Slope failures, debris flows, and mass movements might be some of the most devastating  
30 outcomes associated with extreme rainfall (Cheng et al., 2005; Crosta and Frattini, 2008). The  
31 rapid buildup of pore water pressure beyond hydrological thresholds following extreme  
32 rainfall events can induce substantial increases in sediment yields through slope failure,

1 resulting in enormous economical and environmental damage. Economically, extreme rainfall  
2 events bringing large quantities of sediment can push many structures, including hydro-power  
3 plants, sabo dams, urban drainage facilities, flood barriers, and other infrastructure facilities,  
4 to failure (Kunkel et al., 1999; IADB, 2000). Moreover, high sediment concentrations added  
5 to rivers and streams can degrade the drinking water quality, thus increasing the cost of water  
6 treatment, and are harmful to certain species of fish and aquatic organisms (Waters, 1995).

7 Japan is a country that is particularly prone to slope failures due to its steep terrains and weak  
8 geological formations. For example, extreme meteorological events in 2004 caused 326  
9 deaths (3.4 times more than the average for 2000-2003) and resulted in damage cost of an  
10 estimated 287.5 billion JPY to agricultural production (2.7 times higher than the average for  
11 2000-2003). The number of annual average slope failure events has doubled to more than  
12 2530 throughout Japan (Climate Change Monitoring Report, 2004, 2005). Consequently,  
13 elevated sediment concentrations have been reported, which have had an impact on river  
14 water quality and associated ecosystems over a considerably long period of time. For  
15 example, following heavy rainfall events in July 2004 in Fukui prefecture, an elevated  
16 turbidity concentration of over 100 degrees was recorded for over five months in the  
17 Managawa dam area (Sakamoto, 2008).

18 In response to the increasing evidence of impacts, sediment transport attributed to extreme  
19 rainfall has become an important issue under changing climate conditions, although the cause  
20 and effect of this phenomenon have still not been practically proven. By identifying and  
21 mapping the areas prone to slope failure, the spatial distribution of this hazard can be  
22 assessed. Moreover, by linking the probability of extreme rainfall events in the past with the  
23 output of General Circulation Models (GCM), we can create climate hazard maps for the  
24 future. There have already been several studies accounting for the effects of climate changes  
25 and land use changes on sediment yields (Borga et al., 2002; Asselman et al., 2003; Philip et  
26 al., 2009). However, thus far, studies have considered only geological and geographical  
27 conditions as the triggering parameters related to the effects of time-averaged climate change  
28 scenarios. There has been no research to assess the regional-scale sediment yield attributed to  
29 extreme rainfall under changing climate conditions. Therefore, the objective of this study is to  
30 estimate the spatially distributed sediment yield attributed to extreme-rainfall-induced slope  
31 failure hazards under changing climate conditions for all of Japan. Given the socio-economic  
32 and ecological importance of this issue, zones must be defined in terms of the probability of

1 occurrence of slope failures and subsequent sediment production for particular return periods.  
2 Such information will be essential for the decision-making process in Japan for hazard  
3 mitigation through proper regional planning and implementation.

## 4 5 **2 Methodology**

6 In general, there are three basic approaches, i.e., qualitative methods, physically based  
7 deterministic methods, and statistical methods, that have long been used for evaluating  
8 rainfall-induced slope failure hazards (Xie et al., 2007; Yilmazer et al., 2007; Westen et al.,  
9 2003; Temesgen et al., 2001) and corresponding sediment yield estimations (Wicks and  
10 Bathurst, 1996; Bemporad et al., 1997; Westen et al., 1999; Bathurst, 2002, Maa et al.,  
11 2009). All of these methods have unique advantages when employed in different conditions  
12 and also suffer from their own drawbacks. The qualitative approaches are simple and easy to  
13 apply but fail to properly represent hydro-geological processes. In contrast, physically based  
14 deterministic methods systematically approximate physical concepts such as the equilibrium  
15 of the slope stability and surface water infiltration by applying a set of mathematical  
16 formulae, and modeling is done by using physically meaningful parameter set. The model  
17 calibration and validation is therefore carried out with comparatively short time series of field  
18 measurements. However, these methods are computationally expensive and demand  
19 reasonably accurate spatially distributed parameters and meteorological inputs at a fine  
20 resolution (Hutchinson, 1995; Guzzetti et al., 2005). Therefore, physically based deterministic  
21 methods are difficult to apply practically over large areas with complex topography and  
22 geological formations. Statistical methods, on the other hand, do not require such large  
23 amounts of detailed input data. They develop a statistical relationship between the impacts  
24 (e.g., sediment yield) and impact-triggering parameters (e.g., extreme rainfall, earthquakes)  
25 based on a series of past observations. The statistical methods therefore possess a unique  
26 advantage over the other methods because they can be applied at the regional scale (Shou et  
27 al., 2009). Moreover, they can be used to predict the susceptibility of the impact, which  
28 enables us to link the model with climate change studies. In this study, we employed  
29 statistical methodology to relate extreme-rainfall-induced slope failures with the subsequent  
30 sediment yield. The working procedure of the study includes:

- 1 1. Use of the probability model developed by Kawagoe et al. (2010) for the relationship  
2 between the slope failure hazard and triggering parameters, including spatially distributed  
3 extreme rainfall,
- 4 2. Development of a regression relationship between the probability of slope failure and  
5 subsequent sediment yield and
- 6 3. Application of the relationships developed with selected GCM scenarios to estimate  
7 the resultant sediment yield under changing climate conditions.

8 Numerous studies have pointed out the importance of various processes, such as geographical,  
9 geological and hydrological processes, in rainfall-induced slope failures and sediment yield  
10 assessments (Hutchinson, 1995; Dai et al., 2001). In this study, we consider three important  
11 triggering parameters; 1) the hydraulic gradient represents the hydro-climate effect, 2) the  
12 relief energy represents the geographical effect and 3) four geological formations represent  
13 the geological effect, attributed to slope failure hazards.

14 Regardless of the lithological structure, slope failures are more potentially causative in steep  
15 terrains than in gentle gradients. Therefore, the relief energy, which is defined as the elevation  
16 difference between the highest and lowest point in each grid cell, was used to represent the  
17 effect of geography on slope failure hazards. The National-Land Information Database  
18 (2001), which has well-detailed, fine-resolution (1 km × 1 km), digital elevation model data  
19 (KS-META-G05-54M) for all of Japan, was used to estimate the representative relief energy  
20 for each grid cell. The same database (KS-META-G05-56M) was used to classify the area for  
21 different geological zones at the same grid resolution. There are four geological formations  
22 that are commonly found in Japan: colluviums, Neogene sedimentary rocks, Paleogene  
23 sedimentary rocks and granites; these were considered based on their likelihood in the  
24 formation of slope failures.

## 25 **2.1 Estimation of the hydraulic gradient**

26 The hydraulic gradient is defined as the rate of hydraulic head change per unit distance in a  
27 particular direction. Temporal changes in hydrological conditions (changes in soil moisture  
28 content from unsaturated to saturated and vice versa) due to variations in extreme rainfall and  
29 the resulting infiltration rate have an intensive impact on slope failure formations. The  
30 unsteady nature of this parameter offers a unique opportunity to combine our assessments  
31 with climate change studies. Nevertheless, it requires a large computational effort as

1 compared to other triggering parameters. To estimate the hydraulic gradient attributed to  
2 extreme rainfall at a 1-km resolution, we followed the method previously developed by  
3 Kawagoe et al. (2010). The two-dimensional form of Richard's equation was employed to  
4 obtain the hydraulic gradient, which was numerically solved by considering soil data, the  
5 slope angle and extreme rainfall as the independent input variables in each grid cell (more  
6 details can be found in Kawagoe et al., 2010). To estimate extreme rainfall events, 24-hour  
7 maximum rainfall data covering 21 years (1980-2000) from the Automated Meteorological  
8 Data Acquisition System (AMeDAS) were employed with the Generalized Extreme Value  
9 (GEV) probability distribution function. For 1024 AMeDAS meteorological stations  
10 throughout Japan, GEV analysis independently generates 1024 extreme rainfall values to each  
11 station. Considering the fact that the rainfall patterns in mountain areas are largely influenced  
12 by irregular topography (Buytaert et al., 2006), to distribute the estimated extreme rainfall at a  
13 resolution of 1-km, we used the "Mesh Climate Data 2000" rainfall database developed by the  
14 Japanese Meteorological Business Support Center (2002). In this database, the rainfall  
15 distribution over Japan was estimated by regression models constructed using independent  
16 variables developed from geographical factors (Lookingbill and Urban, 2003; Ueyama, 2004).  
17 The data set includes the monthly averaged rainfall over 30 years (1971-2000) assembled at a  
18 1-km grid resolution. A relationship between the estimated extreme rainfall and maximum  
19 monthly rainfall from the Mesh Climate Data 2000 was developed for distributing the  
20 extreme rainfall to a 1-km grid resolution. To develop a statistically better relationship, the  
21 AMeDAS stations were categorized into three seasonal classes, winter (December-February),  
22 spring-summer (March-August) and autumn (September-November) based on the probability  
23 of an extreme rainfall event. As an example, the mountain areas at the seaside receive their  
24 maximum rainfall during the winter, while the south islands of Japan receive the most rainfall  
25 in the spring-summer category (Kawagoe et al., 2010). Therefore, three separate regression  
26 analyses were performed to obtain the relationship between the extreme rainfall and  
27 maximum monthly rainfall from the Mesh Climate Data 2000, and later extreme rainfall  
28 values were distributed with a 1-km grid resolution based on the maximum monthly rainfall  
29 values from the Mesh Climate Data 2000 at each grid box. These extreme rainfall values were  
30 then used as the main input in the infiltration analysis to find the hydraulic gradients.

## 2.2 Probability model for slope failure

By following the above procedures, the probability of slope failure occurrence was determined by accounting for past events of slope failures at each grid cell. A stepwise logistic regression method was then employed to find the relationship between the triggering parameters and slope failure probability (Eq. 1). Instead of considering the geological type as an independent variable with the appropriate weighting factor in the model, four different models were developed for each geological type.

$$\log\left(\frac{P}{1-P}\right) = \sigma_0 + \sigma_h \times hyd + \sigma_r \times relief \quad (1)$$
$$P = \frac{1}{1 + \exp[-(\sigma_0 + \sigma_h \times hyd + \sigma_r \times relief)]}$$

where  $P$  is the probability of slope failure occurrence,  $\sigma_0$  is the intercept,  $\sigma_h$  is the coefficient of the hydraulic gradient,  $\sigma_r$  is the coefficient of relief energy,  $hyd$  is the hydraulic gradient, and  $relief$  is the relief energy.

## 2.3 Probability model for sediment yield

Many studies have derived magnitude-frequency relationships for sediment yields in hazard assessments (Helsen et al., 2002; Marchi et al., 2002; Hunger et al., 2008). Differences among these relationships reflect the influence of triggering parameters, such as rainfall, and the geomorphologic setting of the catchments. In this study, we developed a relationship between the annual average specific sediment yield and the average probability of slope failure in the representative catchment. Altogether, 59 dams were selected throughout Japan. For each of these dams, the catchment areas are larger than 185 km<sup>2</sup> and more than 15 years of annual sediment yield records are available. The same relationship was developed for various return periods of extreme rainfall, and the goodness of the fit was evaluated against the coefficient of determination to select the best-fit relationship for climate predictions.

## 2.4 GCMs for climate predictions

The Special Report on Emission Scenarios (SRES) along with the IPCC AR4 report has given widely recognized GCMs for climate predictions. In this study, we used two climate scenarios from two GCMs developed in Japan: the Meteorological Research Institute Regional Climate model (MRI-RCM20-Ver.2) embedded with the SRES A2 scenario and the

1 MIROC3.2\_HIRES (high-resolution version of the Model for Interdisciplinary Research on  
2 Climate) embedded with the SRES A1B scenario. The GCMs above have demonstrated good  
3 performance in simulating large-scale circulations and climate features that affect regional  
4 climates (Salathe et al., 2007). However, the resolutions are still far too coarse to use in site-  
5 specific assessments, especially in mountainous areas. Therefore, a statistical downscaling  
6 technique was employed to link the spatial gap between the local scale grid resolution (1-km  
7 in this study) and the GCM grid scale (Iizumi et al., 2008). Climates for three time periods,  
8 the present climate (1980-2000), an intermediate climate (2036-2065 for MIROC and 2031-  
9 2050 for MRI-RCM20-Ver.2) and the future climate (2065-2095 for MIROC and 2081-2100  
10 for MRI-RCM20-Ver.2), were selected to show the transition of the impact in the future. The  
11 two selected GCMs have some advantages as compared to other models presented in IPCC  
12 AR4. Firstly, the model outputs produced finer resolutions, which is particularly useful for  
13 use with mountainous topography. For example, the HADCM3 model has a very coarse  
14 resolution that is approximately equal to 90,465 km<sup>2</sup> of the grid boxes in Japan, while the  
15 MRI-RCM20-Ver.2 model resolution is only 400 km<sup>2</sup>. Secondly, MIROC and MRI-RCM20-  
16 Ver.2 have been proven to be very effective in simulating the climate variables that eventually  
17 produced the impacts for extreme cases over wider ranges (GERF S-4 project document,  
18 2008). Therefore, they avoid the extensive downscaling efforts that are necessary for many  
19 GCM scenarios for predicting the impacts in a reliable range.

20 In the first step of downscaling, the procedure explained in Sect. 2.1 was applied with the  
21 GCM-produced daily rainfall to obtain the extreme rainfall distribution at the GCM grid scale  
22 (hereafter referred to as ERF<sub>GCM</sub>). Because the bias correction was to be performed at the  
23 GCM grid scale, the extreme rainfall in the present climate (hereafter referred to as ERF<sub>PC</sub>) as  
24 derived in Sect. 2.1 at a 1-km grid resolution was aggregated with the grid scale of the  
25 climate model (e.g., 20 km for MRI-RCM20-Ver.2). The ERF<sub>GCM</sub> data for each time period  
26 (intermediate and future climate) were then separately matched with the ERF<sub>PC</sub> data  
27 belonging to the three seasonal categories to form transfer functions (six transfer functions for  
28 each GCM scenario separated into three seasonal classes and two time resolutions). In the  
29 final step, the original extreme rainfall values at a 1-km grid resolution were used in the  
30 corresponding transfer function to obtain the future (intermediate and future climate) extreme  
31 rainfall values. This process was repeated for each grid cell in the domain to obtain the  
32 extreme rainfall distribution for two scenarios in two future time periods.



## 1 **3 Results and discussion**

### 2 **3.1 Slope failure probability and sediment yield**

3 The spatial distribution of sediment yield in Japan attributed to extreme-rainfall-induced slope  
4 failure probability was estimated. In the first stage of the research, the spatial distribution of  
5 the slope failure probability was estimated by considering the extreme-rainfall-induced  
6 hydraulic gradient, relief energy and geological formation as the triggering parameters. The  
7 results portrayed two distinct aspects of the slope failure probability (Table 3 in Kawagoe et  
8 al., 2010). Firstly, the calculated standardized partial regression coefficient produced  
9 noticeably different values for the two triggering parameters. This coefficient explains the  
10 change in slope failure hazard probability in the model when one triggering parameter  
11 (hydraulic gradient or relief energy) is changed by one unit while the other parameter is held  
12 constant. The standardized partial regression coefficient was higher for the hydraulic gradient  
13 than for the relief energy for all four geological formations. This suggests that the hydraulic  
14 gradient is more influential than the relief energy in terms of triggering slope failures.  
15 Secondly, differences in magnitude of the coefficient of each parameter in different geological  
16 settings indicate variations in their resistance to slope failures. The probability of slope  
17 failure occurrence varied from the highest in colluviums to the lowest in granites. The  
18 generally loose, non-consolidated nature of the colluviums has been proven to be more  
19 significant in the occurrence of slope failures than hard compact formations such as granite  
20 (Restrepo et al., 2006).

21 In the next step, a regression model for the slope failure hazard probability and subsequent  
22 sediment yield was developed. Among the various return periods of extreme rainfall  
23 considered, the five-year return period gave the best fit with a determinacy coefficient of 0.65  
24 (Fig. 1a). The underlying reason for the best fit with respect to the five-year return period was  
25 tested by examining the number of years of extreme sediment yield in each dam. Annual  
26 sediment yield records covering 15-20 years at each dam were examined. Assuming that the  
27 annual sediment yield averaged over the catchment is normally distributed throughout the  
28 recording period (15-20 years), the lower bound of the extreme sediment yield ( $SY_{LB}$  in  
29  $m^3/km^2/year$ ) in each catchment is defined as in Eq. (2).

$$30 \quad SY_{LB} = SY_{Avg} + SD \quad (2)$$

1 where  $SY_{Avg}$  (in  $m^3/km^2/year$ ) is the annual average sediment yield in the catchment and  $SD$  is  
2 the standard deviation of the annual sediment yield data series. Sediment yields exceeding the  
3 threshold of  $SY_{LB}$  are defined as extreme sediment yield events. The threshold of  $SY_{LB}$  is then  
4 used to separate the years with extreme sediment yield events at each dam site. Figure 1b  
5 indicates the average recurrence interval (Chow et al., 1988) of extreme sediment yield events  
6 at all selected dam sites throughout Japan. According to this, over 55% of the dams studied,  
7 experience an extreme sediment yield event every 5-7 years, and over 80% of the dams  
8 experience one every 5-10 years, on average. These figures clearly explain the reason for the  
9 statistically better relationship obtained between the extreme rainfall and extreme sediment  
10 yield over a five-year return period.

11 Figure 2 depicts the relationship between the catchment-averaged probability of the slope  
12 failure occurrence and the annual-averaged sediment yield for a five-year extreme rainfall  
13 return period, and Eq. (3) shows the representative regression relationship.

$$14 \quad SY = 0.1051 \exp(0.0301P) \quad (3)$$

15 where  $SY$  is the annual average sediment yield ( $m^3/km^2/year$ ) in a particular dam and  $P$  is the  
16 spatially averaged probability of slope failure occurrence in a specific dam catchment. The  
17 exponential shape of the relationship indicates that the sediment yield may substantially  
18 increase with increasing probability of slope failure.

19 The validity of the developed relationship was tested prior to its use in climate impact  
20 predictions. Another 43 dams, which were not considered in developing the original  
21 regression relationship, were selected, covering all of Japan. Figure 3a shows the locations of  
22 these dams in Japan based on their annual average sediment productions. Figure 3b depicts a  
23 comparison of the simulated and observed sediment yields. Out of the 43 dams, only 5 dams  
24 whose observed annual average sediment accumulations were greater than  $1.2 \times 10^3$   
25  $m^3/km^2/year$  were not well-predicted in terms of sediment yield. It is noted that the catchment  
26 areas of the dams with low accuracy predictions were comparatively small, and large-scale  
27 slope failures have occurred following recent extreme rainfall events. The relationship  
28 developed in Eq. (3) considered long-term sediment accumulations in the dams (averaged  
29 over 15-20 years). Therefore, when the sediment yield records were averaged over time, dams  
30 with recently outsized sediment accumulations eventually produced atypical average sediment  
31 yield values. Therefore, by averaging the values with a set of future sediment yield records  
32 over a wider time scale would match the sediment yield records of the five dams with the

1 developed regression relationship. When the results of the 5 dams with excessive sediment  
2 accumulation were disregarded, the sediment yields calculated from the model indicated good  
3 agreement with the observed sediment loads in the other 39 dams, with a determinacy  
4 coefficient of 0.74. This confirms that the relationship developed between the probability of  
5 slope failure and sediment yield can be successfully applied in long-term studies of climate  
6 change impact predictions.

7 The sensitivity of the sediment yield model to the triggering parameters was also tested.  
8 Figure 4 shows the variations in sediment yield with relief energy and hydraulic gradient for  
9 four selected geological formations. Similar to the slope failure hazard probability, the  
10 sediment yield potential is highest in colluviums and decreases in the order of Neogene  
11 sedimentary rocks and Paleogene sedimentary rocks to the lowest potential in granites. As an  
12 example, for a unit change in the hydraulic gradient, colluviums formations produce  $12.2 \times$   
13  $10^3 \text{ m}^3/\text{km}^2/\text{year}$  of sediment yield, which is 94% higher than the sediment yield production of  
14 granites under the same conditions. Similarly, Neogene sedimentary rock and Paleogene  
15 sedimentary rock produce  $8.6 \times 10^3 \text{ m}^3/\text{km}^2/\text{year}$  and  $6.5 \times 10^3 \text{ m}^3/\text{km}^2/\text{year}$  of sediment load  
16 for a unit change in the hydraulic gradient, respectively.

17 Figure 4 also reveals an important aspect that would be critical under changing climate  
18 conditions. The hydraulic gradient is a rainfall-sensitive parameter that can be significantly  
19 elevated with an increase in the intensity and frequency of rainfall with climate change  
20 effects. According to Fig. 4, the rate of change of the sediment yield (gradient of the curve) is  
21 more sensitive to a small change in the hydraulic gradient, especially within the rising limb of  
22 the curve (e.g.,  $12.2 \times 10^3 \text{ m}^3/\text{km}^2/\text{year}$  per unit change of hydraulic gradient for colluvium  
23 formations). Therefore, areas that will cross the lower edge of the rising limb in the future  
24 may have a critical impact on the sediment yield under changing climate conditions.

### 25 **3.2 Spatial variability of sediment yield**

26 By applying the developed sediment yield model with the distributed slope failure probability,  
27 the spatial variability of the sediment yield can be estimated. Figure 5a shows the spatial  
28 distribution of the sediment yield estimated at a 1-km grid resolution. Moreover, the model-  
29 predicted sediment yields were further aggregated to the major river basins of Japan. These  
30 river basins were categorized based on the existence of first-order rivers in Japan. Figure 5b  
31 depicts the average sediment yield based on the different basins. Areas with significantly

1 higher specific sediment yields (over 2000 m<sup>3</sup>/km<sup>2</sup>/year) are distributed throughout the  
2 Tenryu, Ooi and Kiso river basins in the Hokuriku and Tokai regions and the Shimanto and  
3 Naka river basins in the Shikoku region. The lithology and relief energy differences between  
4 the various regions may play an important role in producing sediment yield (Fig. 4). As an  
5 example, the greater yields corresponding to the Yoshino river basin in the Tokai region  
6 consists of 40% colluviums and 27% Neogene sedimentary rock formations, whose soils have  
7 a low resistance to the sediment yield, while granites with high resistance to the sediment  
8 yield cover only 1% of the area. Moreover, the Tohoku and Hokuriku mountain seaside  
9 regions with comparatively high relief energy have a significantly higher specific sediment  
10 yield (spatial average of 600-800 m<sup>3</sup>/km<sup>2</sup>/year in Fig. 5b) as compared to areas with low relief  
11 energy, such as the eastern side of the Kanto region.

### 12 **3.3 Sediment yield distribution under changing climate conditions**

13 Two climate change scenarios applied to two time periods in the future produced four sets of  
14 results to demonstrate the transition of the sediment yield with climate change effects (Fig. 6).  
15 In general, all of the figures indicate several common areas with significant sediment yield  
16 changes in the future, even though the magnitudes are somewhat different. Above all,  
17 substantial specific sediment yield changes (over 1000 m<sup>3</sup>/km<sup>2</sup>/year) were estimated along the  
18 mountain areas in the Hokuriku and Kinki regions. The southern Shikoku region was also  
19 predicted to have a significant specific sediment yield change, although it varies in the  
20 different figures (250 to over 1000 m<sup>3</sup>/km<sup>2</sup>/year). Second only to the above areas, the southern  
21 Hokkaido region was predicted to have a specific sediment yield increase in the interval  
22 between 250 and 500 m<sup>3</sup>/km<sup>2</sup>/year in three estimations out of four: for MIROC in  
23 intermediate and future climate and MRI-RCM20 in the intermediate climate, while MRI-  
24 RCM20 predicted an increase in the interval between 0 and 250 m<sup>3</sup>/km<sup>2</sup>/year. Moreover,  
25 some areas in the Tohoku region may move from the 0- to 250-m<sup>3</sup>/km<sup>2</sup>/year category to the  
26 250- to 500-m<sup>3</sup>/km<sup>2</sup>/year specific sediment yield category. In contrast, for the northern  
27 Hokkaido, northern Kyushu and Kanto regions, the model does not predict a significant  
28 sediment yield difference in the future.

29 Despite the approximately similar patterns of sediment yield in the above regions in our four  
30 estimations, there were marked differences at the local basin scales. These differences can be  
31 attributed to changes in extreme rainfall in the different scenarios for the different time  
32 periods. As an example, the model predictions for the Toyo river basin in the Tokai region

1 indicate an over 70% sediment yield increment for the MIROC intermediate and future  
2 climates and for the MRI-RCM20 intermediate climate, while the MRI-RCM20 future  
3 estimations predict only a 40% increase. The same phenomenon can be observed in extreme  
4 rainfall changes in the future, where the MIROC intermediate and future climates and the  
5 MRI-RCM20 intermediate climate predicted an average extreme rainfall increase of 28%,  
6 while the MRI-RCM20 future estimations gave an only 8% extreme rainfall increase. Out of  
7 105 river basins that cover the whole area of Japan, the model predicted an approximately  
8 constant or decreasing trend of sediment yield for only 9 river basins in future as compared to  
9 the present estimations. The average percentage of the sediment yield reduction in these nine  
10 river basins was less than 10%, suggesting that almost all of the river basins in Japan will  
11 suffer from an increasing sediment yield risk in the future. For 15 river basins, the model  
12 predicted a more than 50% sediment yield increment in the future (for at least three out of  
13 four estimations in the future), and among them, 5 river basins will experience a more than  
14 90% change as compared to the present sediment yield.

15 When looking at the spatially averaged sediment yield over the whole country, both model  
16 scenarios predicted an increasing trend for the intermediate climate (Fig. 7), implying a  
17 potential impact in the first half of the 21<sup>st</sup> century. With respect to four future estimations for  
18 105 river basins, the MRI-RCM20 future climate, however, predicted a higher sediment yield  
19 than the other three estimations for only 8 river basins. Therefore, for the future climate, the  
20 MRI-RCM20 scenario predicted a small decreasing trend as compared to the intermediate  
21 climate, while the MIROC scenario predicted a continuously increasing trend. These changes  
22 were mainly attributed to variations in extreme rainfall events in the future and were also  
23 influenced by the geology and relief energy of each individual basin.

24 The estimates for the total sediment yield at the river basin scale during extreme rainfall  
25 events in the future make it easy to identify the hazard-prone areas under a changing climate  
26 conditions. By referring to our results, decision-makers can narrow down the area of interest  
27 to the specific local scales, and proper mitigation measures can be implemented with support  
28 of the respective local authorities. The errors of the model predictions that could not explain  
29 the variations of the observed sediment yield can also be attributed to the land-use-change-  
30 induced sediment yield at the very local scale. Although many studies have documented  
31 sediment yields caused by anthropogenic influences (Asselman et al., 2003; Philip et al.,  
32 2009), it is quite difficult to incorporate them into a probabilistic model in regional-scale

1 analysis. Therefore, the inclusion of more detailed information on land use and sub-basin  
2 watershed characteristics in site-specific approaches should provide more accurate  
3 predictions. After identifying the hazard-prone basins as done in this study, such a detailed  
4 analysis would be appropriate for designing infrastructure facilities for mitigating future  
5 climate change impacts.

#### 6 7 **4 Conclusions**

8 To facilitate the decision-making process by identifying hazard-prone areas under changing  
9 climate conditions, this study developed a probabilistic model for the relationship between the  
10 slope failure probability induced by extreme rainfall and sediment yield. There are three  
11 triggering parameters; the hydraulic gradient, the relief energy and the geology type  
12 representing the hydro-climate (hydrology and extreme rainfall), topography and geological  
13 effects, respectively, were considered in developing the probabilistic model for slope failure.  
14 The relationship between the slope failure and subsequent sediment yield was developed by  
15 matching the annual average sediment yield observations at 59 dams throughout Japan with  
16 the average probability of slope failure hazard in the representative catchment areas. For the  
17 predictions of climate change impacts, two climate model scenarios, MRI-RCM20 A2 and  
18 MIROC A1B, in two time periods (intermediate climate and future climate) in the future were  
19 incorporated.

20 The results show that extreme sedimentation events have a high probability of occurring  
21 every 5-7 years. The verified results of the developed slope failure-sediment yield relationship  
22 demonstrated that the model is effective and useful in estimating the sediment yield attributed  
23 to extreme-rainfall-induced slope failure (with a determinacy coefficient equal to 0.74). The  
24 sensitivity analysis of the model showed that the sediment yield potential was highest in  
25 colluviums and decreased in the order of Neogene sedimentary rocks and Paleogene  
26 sedimentary rocks to the lowest potential in granites. Moreover, it is known that the hydraulic  
27 gradient is more influential than the relief energy.

28 The results of the GCM scenarios predict that the sediment yield impact will increase in the  
29 future. When the spatial average sediment yield for all of Japan is considered, both scenarios  
30 produced an approximately 16-17% and 14-21% increase around the first half and second half  
31 of the 21<sup>st</sup> century, respectively as compared to the present climate. On the regional scale,  
32 substantially higher sediment yield changes (over 1000 m<sup>3</sup>/km<sup>2</sup>/year) were estimated in the

1 Hokuriku, Kinki and Shikoku regions. The southern Hokkaido region is predicted to  
2 experience a moderate sediment yield increase (250-500 m<sup>3</sup>/km<sup>2</sup>/year), while the Tohoku  
3 region is predicted to have a 0- to 250-m<sup>3</sup>/km<sup>2</sup>/year increase in sediment yield. Due to  
4 variations in extreme rainfall events, the sediment yield estimations at the basin scale  
5 predicted changes of different magnitudes. Out of 105 basins in Japan, 96 showed an  
6 increasing trend of sediment yield under changing climate conditions. Among them, five river  
7 basins will experience a more than 90% change as compared to the present sediment yield.

8 Following the increasing trend of extreme meteorological events and the resulting vast impact  
9 on socio-economic and environmental sectors, decision-makers in Japan faced new challenge  
10 to implement mitigation measures under a changing climate. From the results of our study,  
11 proper identification of the basins that are prone to sedimentation hazards under changing  
12 climate conditions can guide decision-makers in preparing and implementing appropriate  
13 mitigation measures to cope with the impacts.

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1 **Figure caption**

2 Figure 1. a) Determination coefficient of probability of the slope failure and specific sediment  
3 yield with respect to different return periods, b) return period of extreme sediment yield  
4 events.

5 Figure 2. Relationship between the probability of slope failure and specific sediment yield.

6 Figure 3. Model validation: a) locations of the selected dams, b) observed and simulated  
7 specific sediment yields.

8 Figure 4. Sensitivity of specific sediment yield to the triggering parameters.

9 Figure 5. Spatial distribution of specific sediment yield: a) at a 1-km grid resolution, b)  
10 averaged to the basin scale.

11 Figure 6. Specific sediment yield distribution under a changing climate: a) specific sediment  
12 yield for different climate change scenarios, b) specific sediment yield change as compared to  
13 the present climate.

14 Figure 7. Spatially averaged specific sediment yield in Japan according to different climate  
15 change scenarios.