

Abstract

The objectives of this work are: (a) to statistically test and quantify the decreasing trends of streamflow and sediment discharge in the middle reaches of the Yellow River in China during 1950–2008, (b) to identify change points or transition years of the decreasing trends, and (c) to diagnose whether the decreasing trends were caused by precipitation changes or human intervention, or both. The results show that significant decreasing trends in annual streamflow and sediment discharge have existed since the late 1950s in the middle reaches of the Yellow River ($P=0.01$). Change-point analyses further revealed that transition years existed and that abrupt decline in streamflow and sediment discharge began in 1985 and 1981, respectively, in the middle reaches of the Yellow River ($P=0.05$). Adoption of conservation measures in the 1980s and 1990s corroborates the identified transition years. Double-mass curves of precipitation vs. streamflow (sediment) for the periods before and after the transition year show remarkable decreases in proportionality of streamflow (sediment) generation. Compared with the period before the transition year, cumulative streamflow and cumulative sediment discharge reduced respectively by 17.8% and 28% during 1985–2008, which was caused by human intervention, in the middle reaches of the Yellow River. It is, therefore, concluded that human activities occupied a dominant position and played a major role in the streamflow and sediment discharge reduction in the middle reaches of the Yellow River.

1 Introduction

Streamflow and sediment discharge provide useful information on the processes of soil erosion and sediment delivery occurring in a basin (Siakeu et al., 2004). The middle reaches of the Yellow River flows through the Loess Plateau. Due to improper land use and excessive exploitation, the Loess Plateau is counted among the most severely eroded areas in the world. Meanwhile, rivers in this region transport a large amount

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downward trend exists, it is important and necessary to further check exactly when the change began and what the driving factors are. Understanding the impacts of climate variation and human activity on hydrological regime and sediment dynamics is useful for developing effective conservation strategies in the middle reaches of the Yellow River basin. Therefore, the objectives of this study are: (a) to statistically detect trends and change-points in annual streamflow and sediment discharge in the middle reaches of the Yellow River basin; (b) to analyse possible impacts of precipitation and human activities on annual streamflow and sediment discharge dynamics in relation to change-points or transition years detected in this study; and (c) to further estimate the effects of the identified driving factors on both streamflow and sediment discharge decline by comparing two contrasting periods before and after the transition years.

2 Study area and datasets

The middle reaches of the Yellow River (MRYP), between Toudaoguai and Huayuankou, is the study region of this paper. The region lies between 104° E–113° E and 32° N–42° N, with a drainage area of 362 000 km².

A dataset from 33 meteorological stations with long-term annual precipitation data (1957–2008) in the MRYP basin was analysed in this study (Fig. 1 and Table 1). The precipitation data were provided by The National Meteorological Information Centre (NMIC), China Meteorological Administration (CMA). Two key hydrological stations in the MRYP mainstream (Toudaoguai and Huayuankou) were chosen to calculate the streamflow and sediment discharge in the region for analysis (Fig. 1 and Table 1). Annual streamflow and sediment discharge data at the two stations from 1950 to 2008 were obtained from the Chinese River Streamflow and Sediment Communiques, the Ministry of Water Resources of PRC (MWR). All measured data used in this study are of good quality and were checked for quality control by corresponding agencies.

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3 Analysis methods

3.1 Trend test

The rank-based, non-parametric Mann-Kendall statistical test (Mann, 1945; Kendall, 1975) is commonly used for trend detection due to its robustness for abnormally distributed and censored data, which are frequently encountered in hydroclimatic time series (e.g., Hirsch et al., 1982; Burn and Elnur, 2002; Yue et al., 2003; Yue and Pilon, 2004).

The results of the trend test can be used to determine whether or not the observed time series of hydrological variables exhibits a trend that is statistically significant from a trend that could occur by chance; and to do this, it is necessary to test the serial correlation of the data (Jenkins and Watts, 1968). However, the presence of serial correlation can complicate the identification of trends, in that a positive serial correlation can increase the expected number of false-positive outcomes for the Mann-Kendall test (von Storch and Navarra, 1995). Thus, any serial correlation should be removed before conducting the Mann-Kendall trend test. In this work the trend free pre-whitening (TFPW) method of Yue et al. (2003) was used as follows.

Firstly, remove any significant linear trend from the raw time series using:

$$Y_t = X_t - \beta t \quad (1)$$

where X_t is the series value at time t ; β is the linear regression slope of the trend in the raw time series, Y_t is the de-trended series.

Secondly, remove serial correlation if the lag-one serial correlation coefficient (r_1) of the de-trended series is statistically significant at the 5% level, using the pre-whitening method of Kulkarni and von Storch (1995):

$$Y'_t = Y_t - r_1 Y_{t-1} \quad (2)$$

where Y'_t is the de-trended and pre-whitened series, which is referred to as the residual series.

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Thirdly, add the linear trend that was removed in step 1 back to the de-trended or residual series, using:

$$Y''_t = Y'_t + \beta t \quad (3)$$

where Y''_t is the trend free pre-whitened series.

Z statistic was obtained from the Mann-Kendall test on the whitened series from step third. In addition, to confirm the results provided by the Mann-Kendall test, we also performed linear regression analysis.

3.2 Change-point analysis

A number of methods can be applied to determine change points of a time series (Buishand, 1982; Chen and Gupta, 2000; Radziejewski et al., 2000). In this study, we used the non-parametric approach developed by Pettitt (1979) to detect change-points in streamflow and sediment discharge time-series. This method detects a significant change in the mean of a time series when the exact time of the change is unknown. The test uses a version of the Mann-Whitney statistic $U_{t,N}$, that tests whether two sample sets x_1, \dots, x_t and x_{t+1}, \dots, x_N are from the same population. The test statistic $U_{t,N}$ is given by:

$$U_{t,N} = U_{t-1,N} + \sum_{j=1}^N \text{sgn}(X_t - X_j) \quad \text{for } t=2, \dots, N \quad (4)$$

and

$$\begin{aligned} \text{if}(X_t - X_j) > 0, & \quad \text{sgn}(X_t - X_j) = 1 \\ \text{if}(X_t - X_j) = 0, & \quad \text{sgn}(X_t - X_j) = 0 \\ \text{if}(X_t - X_j) < 0, & \quad \text{sgn}(X_t - X_j) = -1 \end{aligned} \quad (5)$$

The test statistic counts the number of times a member of the first sample exceeds a member of the second sample. The null hypothesis of the Pettitt's test is the absence

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of a change point. The test statistic K_N and the associated probability (P) used in the test are given as:

$$K_N = \max_{1 \leq t \leq N} |U_{t,N}| \quad (6)$$

$$P \cong 2 \exp \left\{ -6(K_N)^2 / (N^3 + N^2) \right\} \quad (7)$$

3.3 Double mass curve

The theory of the double-mass curve is based on the fact that a plot of the two cumulative quantities during the same period exhibits a straight line so long as the proportionality between the two remains unchanged, and the slope of the line represents the proportionality. This method can smooth a time series and suppress random elements in the series and, thus, show the main trends of the time series. In this study, double-mass curves of precipitation vs. streamflow and precipitation vs. sediment are plotted for the two contrastive periods to estimate changes in the regression slope (proportionality) to quantify the overall efficiency of soil conservation measures before and after the transition years.

4 Results and discussion

4.1 Trend analysis of observed annual precipitation, streamflow and sediment discharge

The characteristics of precipitation, streamflow and sediment discharge in different decades in the MRYS are given in Table 2. The annual average precipitation, streamflow and sediment discharge is 511.41 mm, $165.21 \times 10^8 \text{ m}^3$ and $8.21 \times 10^8 \text{ t}$ respectively; and the mean coefficient of variation (C_v) is 0.17, 0.53 and 0.72, respectively, in the MRYS during 1950–2008.

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The observed annual precipitation, streamflow and sediment discharge during 1950–2008 in the MRYR are shown in Fig. 2, and their corresponding Mann-Kendall test results are given in Table 3. The streamflow and sediment discharge decreased significantly, with the average annual decrease coefficient being $-3.21 \times 10^8 \text{ m}^3/\text{a}$ and $-0.21 \times 10^8 \text{ t}/\text{a}$, respectively. The precipitation showed a downward trend, with an average annual decrease coefficient being $-1.69 \text{ mm}/\text{a}$.

4.2 Change-point analysis

Since the Mann-Kendall tests showed significant downward trends in precipitation, streamflow and sediment discharge, the Pettitt's test was further used to detect the change points or transition years (Fig. 2). These results corroborated the Mann-Kendall test results. For annual precipitation, the change-point year was detected in 1990 ($P=0.1$), and the precipitation was 529.7 mm and 476.8 mm, respectively, before and after the transition year. For annual streamflow, the change-point year was detected in 1985 ($P=0.05$), it was $205.9 \times 10^8 \text{ m}^3$ and $101.5 \times 10^8 \text{ m}^3$, respectively, in the periods before and after 1985. For annual sediment discharge, the change-point year was detected in 1981 ($P=0.05$). Sediment discharge was $11.3 \times 10^8 \text{ t}$ and $4.5 \times 10^8 \text{ t}$, respectively, during 1950–1981 and 1981–2008.

4.3 Double mass curve of precipitation-streamflow and precipitation-sediment

To further quantify the streamflow and sediment discharge changes before and after the transition years, double mass curves, along with the linear regression lines, were plotted in Fig. 3. There existed clear breakpoints between the two regression lines for both streamflow and sediment discharge in the basin, suggesting that the transition years, identified by Pettitt's method, are correct and meaningful. The slopes of the regression lines were lower after the breakpoints or transition years (i.e. at higher cumulative precipitation values) than before for both streamflow and sediment discharge in the basin. To estimate the relative reduction of total streamflow and sediment discharge for the

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period after the transition years, the information of these two factors and precipitation before the transition years were used to establish regression equations (Tables 4 and 5), and to further extrapolate the cumulative streamflow and sediment up until 2008. The extrapolated cumulative streamflow (R_c in Table 4) and sediment discharge (S_c in Table 5) were based on the assumption that environmental conditions, including human impacts in the basin in the first period before the transition years, remained unchanged in the second period after the transition years. Compared with the extrapolated cumulative streamflow (R_c), observed cumulative streamflow (R_o in Table 4) reduced by 17.8% in the basin (Table 4). The corresponding reduction for sediment discharge was 28.0% in the basin (Table 5). It should be noted that the percent reductions in cumulative sediment discharge were greater than those in cumulative streamflow. The reasons for this are similar to those given in the end of the section above.

4.4 Impacts of precipitation and human intervention

We can calculate the annual streamflow and sediment discharge for the period after the transition years using the regression equations established from the double mass curve of precipitation-streamflow and precipitation-sediment before the transition years. The difference between the calculated values in different periods is because of the impact of precipitation changes. However, the difference between the calculated values and measured values in the same period is the result of human activities. The results were shown in Tables 6 and 7.

For the streamflow reduction, the impact of human activities showed an increasing trend from 61.6% in the 1980s to 81.1% in the early 21st century (in Table 6), and it had been the dominant factor ever since the transition year. And the average human activities contribution rate is 72.0% from 1986 to 2008, which is significantly stronger than the contribution rate of precipitation (28%). On the other hand, the impact of human activity was also the main factor in the sediment discharge decline after the transition year; although this effect weakened in the 1990s. The average human activities contribution rate is 87.8% from 1982–2008, which is still significantly stronger than the

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contribution rate of precipitation (12.2%). The analysis showed that: human activities played a major role in the streamflow and sediment discharge reduction in the MRYP basin.

4.5 Influence of human activities on streamflow and sediment discharge decline

Many studies have documented that human and economic activities (especially after the 1980s) might play an important role in streamflow and sediment discharge reduction in the Yellow River basin (van den Elsen et al., 2003; Xu, 2003; Huang and Zhang, 2004; Mu et al., 2007; Gao et al., 2010). Those activities can be summarized as follows:

(a) Increased demand for water resources in the Yellow River due to national economic development. With the rapid development of China's national economy, water extraction and diversion has dramatically increased for agricultural irrigation and urban and industrial use, especially after the 1980s (Liu and Zhang, 2004). The average annual water extraction and diversion was $64.01 \times 10^8 \text{ m}^3$ during 1998–2007, accounting for 38.75% of average annual streamflow in the MRYP basin.

(b) Impact of soil and water conservation programs and ecological environment rehabilitation campaign in the Loess Plateau. The Loess Plateau, situated in the MRYP basin, is the major sediment source area for the river. To reduce water and soil erosion, some water and soil conservation measures were implemented between 1950 and 1978 (Mu et al., 2007). However, a large-scale of conservation measures were carried out between 1979 and 1997 due to various government-sponsored conservation programmes and environmental rehabilitation campaigns in the MRYP basin. Until 1978, different types of conservation measures covered 13.2% of the basin, and the percentage increased to 41.1% in 1997. Based on the average area that received conservation measures per year, the increasing rates of terrace building, check dam construction, afforestation, and grass planting during 1979–1997 were 2.1, 1.3, 4.1 and 5.1 times the rates during 1950–1978 (Table 8) (Gao, 2010). By 2006, about 49%

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of eroded land had been under protection, with sorts of soil and water conservation measures (including 52 729 km² of prime farmlands, 94 613 km² of soil and water conservation forest and 34 938 km² of grass planting), more than 2700 structures of key projects for gully erosion control and more than 4 300 000 structures of assisted small-scale projects in the Loess Plateau. Undoubtedly, the rapid adoption of soil and water conservation measures and engineering structures in the 1980s and 1990s played a significant role in streamflow and sediment discharge reduction in the Loess Plateau in the MRYP. The results of sediment discharge and streamflow reduction through soil and water conservation measures in different periods in the MRYP were shown in Table 9, and the effects of those measures showed an increasing trend over time (Ran, 2006). This rapid adoption period is in good agreement with the transition year of sediment discharge (1981) identified by the change-point analysis, suggesting that the conservation effects on streamflow reduction may have a time lag in such a large basin.

(c) Impact of the construction of water control projects. Construction of large/medium-sized multi-purpose water control projects has some effect on streamflow and sediment discharge decline in the Yellow River (Tian et al., 2005; Wang et al., 2005). The evaporation and leakage losses from the reservoir was about $10.04 \times 10^8 \text{ m}^3/\text{a}$, accounting for 6.08% of average annual streamflow in the MRYP basin. Reservoir siltation, though undesirable, has reduced sediment discharge in the MRYP basin. There were four large/medium-sized multi-purpose water control projects including Wanjiashai, Tianqiao, Sanmenxia and Xiaodangdi in the mainstream. These reservoirs, facing various degrees of siltation problems, had reduced sediment discharge in downstream region. In almost 50 years, these four large reservoirs had totally deposited $89.44 \times 10^8 \text{ m}^3$, approximately $134.16 \times 10^8 \text{ t}$ sediment, and the sediment deposition rate was about $2.80 \times 10^8 \text{ t/a}$.

Through the above analysis, we can roughly quantitatively estimate the impact of human activities on streamflow and sediment discharge reduction in the MRYP basin. Reduced cumulative streamflow was about $1740 \times 10^8 \text{ m}^3$, approximately $72.5 \times 10^8 \text{ m}^3/\text{a}$

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during 1986–2008; while reduced cumulative sediment discharge was about 154×10^8 t, approximately 5.7×10^8 t/a during 1982–2008. Reduced water and sediment affected by various human activities were estimated as follows: Reduced water was about 45×10^8 m³/a, 10×10^8 m³/a and 6.0×10^8 m³/a, respectively, due to water extraction and diversion, soil and water conservation measures, and evaporation and leakage losses (Table 10) (Gao, 2010). Reduced sediment was about 0.6×10^8 t/a, 2.7×10^8 t/a and 1.8×10^8 t/a, respectively, due to water extraction and diversion, soil and water conservation measures, and reservoir siltation (Table 11) (Gao, 2010).

5 Summary and conclusion

Precipitation, streamflow and sediment discharge during 1950–2008 were analysed by the Mann-Kendall trend test and the Pettitt's change-point analysis in the MRYSR basin. Significant downward trends in annual streamflow and annual sediment discharge were detected in the MRYSR basin, while the annual precipitation showed a downward trend in the same period but not significant. The change-point or transition year was 1985 ($P=0.05$) for streamflow, 1981 ($P=0.05$) for sediment discharge, and 1990 ($P=0.1$) for precipitation in the MRYSR basin.

Human intervention was largely responsible for the downward trends of streamflow and sediment discharge after the transition years in the MRYSR basin. The effects of human intervention on streamflow and sediment discharge decline could be quantified by comparing the two periods using the double-mass curves. Compared to the period before the transition years, measured cumulative streamflow and sediment discharge decreased by 17.8% and 28%, respectively, from the transition years to 2008. Human activities played a major role in the streamflow and sediment discharge reduction. The average human activities contribution rate is 72.0% and 87.8% for the streamflow and sediment discharge reduction, which are significantly stronger than the precipitation contribution rate (28.0% and 12.2%).

Soil and water conservation in the MRYSR basin began in the late 1950s, and the

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pace was more than triple after the 1980s; the calculation was based on the areas conservation measures covered every year. The extensive adoption of conservation measures in the 1980s and 1990s altered the natural regimes of streamflow and led to an abrupt decline in streamflow in 1985, and the transition year of sediment discharge (1981) was also in good agreement with the period that conservation measures developed. The rates of decreases in streamflow and sediment discharge coincided well with the intensity and extent of human intervention and activities. The overall results showed that human activities, such as soil and water conservation programmes, eco-environmental rehabilitation campaign, construction of key water control projects and so on, appear to be the major factor of a significant decrease in annual streamflow and sediment discharge in the recent 50 years in the MRYS basin.

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Table 1. The location and data series of the hydrologic and rain station.

Station	Longitude(E)	Latitude(N)	Series(Year)
Toudaoguai	111° 02'	40° 17'	1950–2008
Huayuankou	113° 40'	35° 54'	1950–2008
Xiji	105° 43'	35° 58'	1958–2008
Tianshui	105° 45'	34° 35'	1957–2008
Guyuan	106° 16'	36° 00'	1957–2008
Baoji	107° 08'	34° 21'	1957–2008
Huanxian	107° 18'	36° 35'	1957–2008
Xifeng	107° 38'	35° 44'	1957–2008
Changwu	107° 48'	35° 12'	1957–2008
Wuqi	108° 11'	36° 50'	1957–2008
Wugong	108° 13'	34° 15'	1957–2008
Xi'an	108° 56'	34° 18'	1957–2008
Tongchuan	109° 04'	35° 05'	1957–2008
Hengshan	109° 14'	37° 56'	1957–2008
Luochuan	109° 30'	35° 49'	1957–2008
Yan'an	109° 30'	36° 36'	1957–2008
Yulin	109° 42'	38° 14'	1957–2008

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Table 1. Continued.

Station	Longitude(E)	Latitude(N)	Series(Year)
Dongsheng	109° 59'	39° 50'	1957–2008
Huashan	110° 05'	34° 29'	1957–2008
Suide	110° 13'	37° 30'	1957–2008
Xixian	110° 57'	36° 42'	1957–2008
Yuncheng	111° 01'	35° 02'	1957–2008
Lushi	111° 02'	34° 03'	1957–2008
Lishi	111° 06'	37° 30'	1957–2008
Xingxian	111° 08'	38° 28'	1957–2008
Hequ	111° 09'	39° 23'	1957–2008
Sanmenxia	111° 12'	34° 48'	1957–2008
Linfen	111° 30'	36° 04'	1957–2008
Wuzhai	111° 49'	38° 55'	1957–2008
Jiexiu	111° 55'	37° 02'	1957–2008
Yangcheng	112° 24'	35° 29'	1957–2008
Youyu	112° 27'	40° 00'	1957–2008
Luoyang	112° 28'	34° 38'	1957–2008
Taiyuan	112° 33'	37° 47'	1957–2008
Zhengzhou	113° 39'	34° 43'	1957–2008

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Table 2. Characteristics of precipitation, streamflow and sediment discharge in different decades in the MRYS basin.

Series	Precipitation (mm)			Streamflow (10^8m^3)			Sediment discharge (10^8t)		
	Average	Extremum ratio	C_v	Average	Extremum ratio	C_v	Average	Extremum ratio	C_v
1950s	564.21	1.48	0.20	240.06	2.40	0.29	14.08	3.76	0.47
1960s	561.28	2.05	0.21	234.97	36.98	0.50	9.31	10.47	0.61
1970s	510.12	1.45	0.11	148.38	2.43	0.28	11.21	3.25	0.39
1980s	515.55	1.64	0.14	172.62	3.01	0.36	6.77	5.39	0.49
1990s	474.52	1.64	0.15	100.14	3.95	0.38	6.43	2.87	0.37
2000s	488.87	1.60	0.17	87.28	6.26	0.44	0.72	14.52	0.88
1950–2008	511.41	2.05	0.17	165.21	36.98	0.53	8.21	212.66	0.72

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Table 3. Results of trend analysis for annual precipitation, streamflow and sediment discharge in the MRYR basin during 1950–2008.

Variable	Mann-Kendall		Linear regression	
	Z statistic	Sig. level	t statistic	Sig. level
Precipitation	−2.028	0.05	−1.595	0.1
Streamflow	−5.617	0.01	−6.082	0.01
Sediment discharge	−5.369	0.01	−5.871	0.01

A negative sign indicates a decreasing trend.

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Table 4. Linear regression equations between cumulative streamflow and cumulative precipitation for the period before the transition years in the MRYS basin.

Regression equation	R_c (10^8m^3)	R_o (10^8m^3)	$R_c - R_o$ (10^8m^3)	$100 \times (R_c - R_o) / R_c$ (%)
$\sum R = 0.3601 \sum P + 203.21 (R^2 = 0.9891 N = 29)$	9779.41	8040.30	1739.11	17.8

R : streamflow; P : precipitation; R_c : extrapolated cumulative streamflow until 2008; R_o : observed cumulative streamflow until 2008.

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Table 5. Linear regression equations between cumulative sediment discharge and cumulative precipitation for the period before the transition years in the MRYS basin.

Regression equation	S_c (10^8 t)	S_o (10^8 t)	$S_c - S_o$ (10^8 t)	$100 \times (S_c - S_o) / S_c$ (%)
$\sum S = 0.0207 \sum P + 1.2069 (R^2 = 0.9915 N = 25)$	551.69	397.23	154.46	28.0

S : sediment discharge; P : precipitation; S_c : extrapolated cumulative sediment discharge until 2008; S_o : observed cumulative sediment discharge until 2008.

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Table 7. The impact of precipitation and human intervention on annual sediment discharge decline in the MRYP basin.

Period	S_{ao} (10^8t)	S_{co} (10^8t)	ΔS		Impact of precipitation		Impact of human intervention	
			Amount (10^8t)	Percentage (%)	Amount (10^8t)	Percentage (%)	Amount (10^8t)	Percentage (%)
Before 1981	11.00	11.05						
1982–1989	6.42	10.76	4.59	41.7	0.24	5.3	4.34	94.7
1990–1999	6.43	9.82	4.58	41.6	1.18	25.8	3.39	74.2
2000–2008	0.72	10.12	10.28	93.4	0.88	8.6	9.40	91.4

S_{ao} : observed annual average sediment discharge; S_{co} : calculated annual average sediment discharge; ΔS : reduction in observed sediment discharge comparing with the period of 1950–1981.

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Table 9. The results of sediment discharge and streamflow reduction through soil and water conservation measures in different periods in the MRYP.

Period	Reduced water ($10^8\text{m}^3/\text{a}$)	Reduced sediment (10^8t/a)
1970–1979	4.54	1.99
1980–1989	5.70	2.23
1990–1996	6.41	2.61
1970–1996	5.46	2.24

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Table 10. The results of reduced water estimated in the MRYP basin.

Calculated reduced streamflow ($10^8\text{m}^3/\text{a}$)	Human activities impact on reduced water			
	Total ($10^8\text{m}^3/\text{a}$)	Water extraction and diversion ($10^8\text{m}^3/\text{a}$)	Soil and water conservation measures ($10^8\text{m}^3/\text{a}$)	Evaporation and leakage losses ($10^8\text{m}^3/\text{a}$)
72.5	61.0	45.0	10.0	6.0



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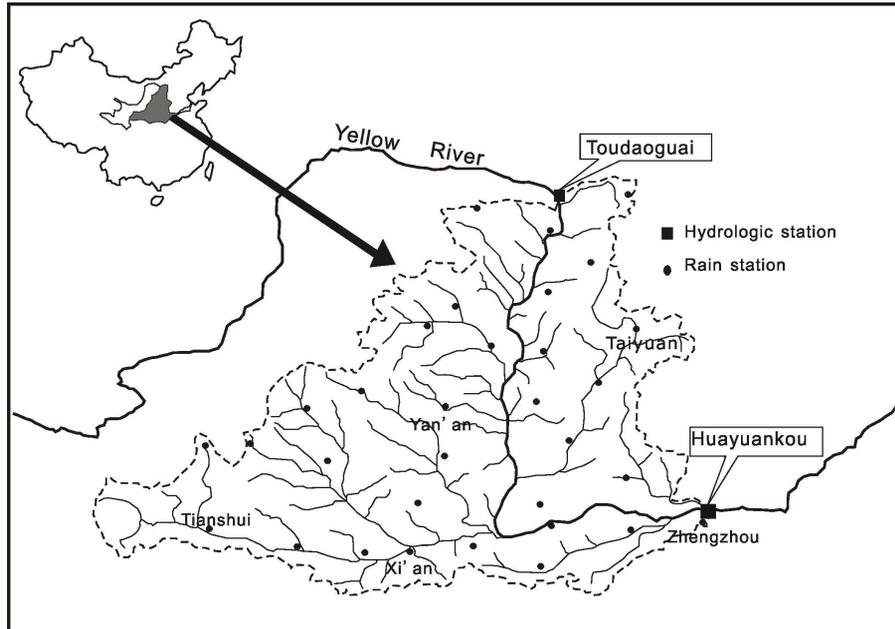


Table 11. The results of sediment reduction estimated in the MRYR basin 10t/a.

Calculated reduced sediment (10^8 t/a)	Human activities impact on reduced sediment			
	Total (10^8 t/a)	Water extraction and diversion (10^8 t/a)	Soil and water conservation measures (10^8 t/a)	Reservoir siltation (10^8 t/a)
5.7	5.1	0.6	2.7	1.8

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**Fig. 1.** Location of the study region and stations in the middle reaches of the Yellow River basin.[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

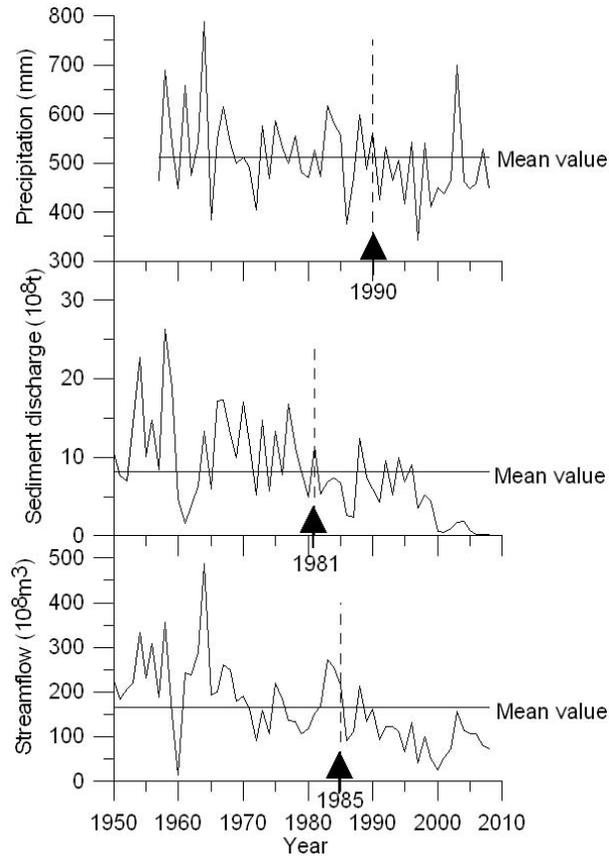


Fig. 2. Observed annual precipitation, streamflow and sediment discharge during 1950–2008 in the MRYR basin.

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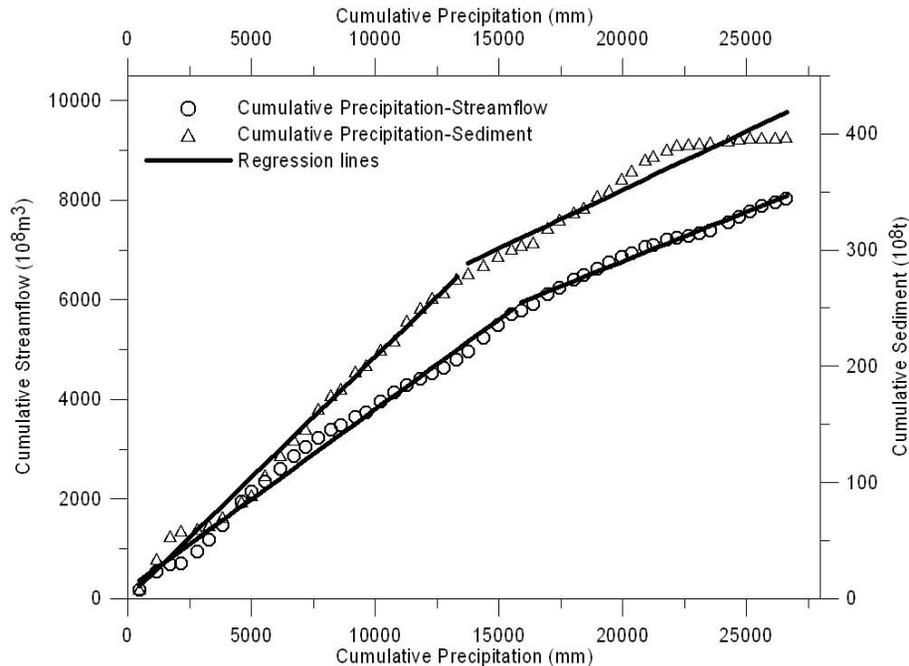


Fig. 3. Double mass curves of precipitation-streamflow and precipitation-sediment during 1957–2008 in the MRYR basin.

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