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Climate change and its impacts on river discharge in two climate regions in China

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Abstract

Understanding the heterogeneity of climate change and its impacts on annual and seasonal discharge, and the difference between mean flow and extreme flow in different climate regions is of utmost importance to successful water management. To quantify the spatial and temporal heterogeneity of climate change impacts on hydrological processes, this study simulated river discharge in the River Huangfuchuan in semi-arid northern China and the River Xiangxi in humid southern China. We assessed the uncertainty in projected discharge for three time periods (2020s, 2050s and 2080s) using seven equally weighted GCMs for the SRES A1B scenario.

Climate projections that were applied to semi-distributed hydrological models Soil Water Assessment Tools (SWAT) in both catchments showed trends toward warmer and wetter conditions, particularly for the River Huangfuchuan. Results based on seven GCMs' projections indicated −1.1 to 8.6 and 0.3 to 7.0 °C changes in seasonal temperature and −29 to 139 and −32 to 85 % changes in seasonal precipitation in River Huangfuchuan and River Xiangxi, respectively. The largest increases in temperature and precipitation in both catchments were projected in the spring and winter seasons. The main projected hydrologic impact was a more pronounced increase in annual discharge in the River Huangfuchuan than in the River Xiangxi. Most of the GCMs projected increased discharge in all seasons, especially in spring, although the magnitude of these increases varied between GCMs. Peak flows was projected to appear earlier than usual in River Huangfuchuan and later than usual in River Xiangxi. While the GCMs were fairly consistent in projecting increased extreme flows in both catchments, the increases were of varying magnitude compared to mean flows. For River Huangfuchuan in the 2080s, median flow changed from −2 to 304 %, compared to a −1 to 145 % change in high flow (Q05 exceedence threshold). For River Xiangxi, low flow (Q95 exceedence threshold) changed from −1 to 77 % and high flow changed from −1 to 62 %, while mean flow changed from −4 to 23 %. The uncertainty analysis provided an improved understanding of future hydrologic behavior in the watershed.

Furthermore, this study indicated that the uncertainty constrained by GCMs was critical and should always be considered in analysis of climate change impacts and adaptation.

1 Introduction

The impacts of climate and hydrological changes cover all spatial scales, from local to global (Lahmer et al., 2001; Coulthard et al., 2005). There is now substantial evidence indicating that over the most recent decades, the global hydrological cycle has already been responding to observed global warming (Bates et al., 2008), which includes increasing atmospheric water vapor content and changing precipitation pattern. In many regions, changing precipitation or melting snow and ice are altering hydrological systems, affecting water resources in terms of quantity and quality (Jiménez Cisneros et al., 2014). To some extent, global climate change has also changed the availability of water resource in China. The precipitation, evaporation, and discharge of China's main rivers and lakes have changed to varying degrees. But in general, the measured discharges of northern rivers such as the Haihe, Yellow, and Liaohe have decreased (Second National Assessment Report for Climate Change, 2011). In contrast, the water cycle in southern China is significantly different from that of northern China (Liu et al., 2004). Moreover, there is evidence that water cycle will likely intensify further. So exploring the impacts of projected climate change on river discharge in the south and north China will be an interesting topic.

Global Climate Models (GCMs) are useful tools for simulating climate system and developing climate change research, which generate possible future climate scenarios. Within the most hydrological projection studies use the precipitation and temperature downscaled from GCMs to driven hydrological model. From these studies, it is abundantly clear that climate change has the potential to substantially impact water resource. It emerges the uncertainties in projected changes to river runoff constrained by the uncertainties in regional climate projection. Considering the usefulness of

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climate scenarios for the decision-making and substantial uncertainties in climate projection, better quantifying uncertainties is helpful to reduce the future risk and adopt adaptive water management.

In a previous study, two typical sub-catchments River Huangfuchuan of the Yellow and River Xiangxi of Yangtze Rivers were selected as study areas for quantitative evaluation of the projected impacts of climate change on river discharge for the 2050s. The results indicate a consistent trend toward warmer and wetter conditions and increased river discharge in both catchments. Substantially larger increases in river discharge relative to baseline were consistently projected for the semi-arid River Huangfuchuan catchment in northern China compared to the subtropical humid River Xiangxi catchment in southern China. In this paper, we analyzed the changes in air temperature, precipitation, and river discharge of the two catchments in the early (2020s), middle (2050s), and late 21st century (2080s) using seven GCMs under the SRES A1B emission scenario within the CMIP3 structure. We focused on extreme flow, constraining uncertainty in the projected river discharges, and examining the contrasts between the southern and northern catchments.

2 Study areas and methodology

2.1 Characteristics of study areas

The River Huangfuchuan is a primary catchment in the middle reaches of the Yellow River. The River Xiangxi is the first tributary supplying the Three Gorges Dam (Fig. 1). In addition to their being located in different climate regions, these two catchments also have different climate conditions and historical climate change trends (Xu et al., 2011).

The River Huangfuchuan is located in northern China, has a semi-arid climate, and is primarily a pastoral farming region. The mean annual temperature is 7.5 °C, and increased at a rate of 0.24 °C per decade from 1961 to 2010. The mean annual precipitation is 388 mm, but decreased over the period of 1961–2000, and increased

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over 2001–2010 at a rate of 0.87 mm decade^{-1} (Gao et al., 2005; Sun et al., 2012). The mean annual runoff was 42.4 mm for 1956–2005, with a range of 74 mm in the 1950s to 28 mm in the 1990s (Wang et al., 2009, 2012). Given the impacts of soil erosion, ecological water shortage, land desertification, flooding, and human water use, the River Huangfuchuan is very sensitive to global change (Yang et al., 2004).

The River Xiangxi lies in a subtropical humid climate region. Mean, minimum, and maximum annual air temperature from 1961 to 2008 was 17.0, 12.7, and 22.9 °C, respectively. Minimum and maximum annual temperatures have increased over this period, especially since the 1980s. Mean annual precipitation during this period was 992 mm with a slight decrease in recent years. The mean annual runoff was 688 mm for 1961–2005, with a decrease from 733 mm in the 1960s to 552 mm in the 1990s. Due to land shortage, the natural vegetation has been progressively converted to farmland, exposing large areas of soil and leading to serious erosion and water loss (Jin et al., 1996; Jiang et al., 2002).

2.2 Methodology

2.2.1 Hydrological model calibration and validation

The hydrological model used in this study was the Soil-Water-Assessment-Tool (SWAT) model developed by the US Department of Agriculture (USDA). We used a previously calibrated SWAT model of River Huangfuchuan and River Xiangxi (Xu et al., 2011). A digital elevation model with a scale of 1 : 250 000 was prepared by the China Fundamental Geographic Information Center. Spatial soil data with a scale of 1 : 1 000 000 was derived from Environment and Ecology Scientific Data Center of western China, National Natural Science Foundation of China. Soil properties were generated from the Soil Attribute Data Set which based on “Soil Species of China” and other sources with total information includes 7300 soil profiles collected from all over China. The most recent land-use maps for the River Xiangxi compiled by the Hubei Land Management Bureau in the 1990s and land-use records from the Inner

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Mongolia Autonomous Region Department of Land and Resource in the 1980s were used to represent catchment land use. Monthly climate dataset CRU TS3.0 (Mitchell and Jones, 2005) which cover the two catchments were stochastically disaggregated to daily resolution following the procedures developed by Arnell (2003) and further described by Todd et al. (2010). Station-based daily precipitation and temperature within and around the two catchments obtained from National Climate Information Center, China Meteorological Administration were used for local calibration of the daily disaggregation procedure.

Based on the input digital maps, a total of 10 and 13 sub-watersheds were generated based on dominant soil and land use for each subbasin. SWAT model was calibrated for the monthly river discharge of the Rivers Huangfuchuan and Xiangxi for a baseline period of 1961–1990, with the remaining 1961–1997 and 1961–1994 data of the Rivers Huangfuchuan and Xiangxi for validation. The model performed well against the monthly river discharges observed from Huangfu gauging station of River Huangfuchuan and the Xingshan gauging station of River Xiangxi. The calibration and validation results showed that SWAT model was able to simulate the monthly discharge well, while Nash–Sutcliffe efficiency were 0.64 and 0.67 for River Huangfuchuan, and were 0.44 and 0.57 for River Xiangxi, respectively, for calibration and validation periods. The frequency distributions of simulated river discharge in both sub-catchments closely approximates those of the observed discharge records as indicated by flow duration curves, while peak flows of the River Xiangxi are very slightly underestimated. More details on input datasets, model calibration and validation results can be found in Xu et al. (2011).

2.2.2 Climate scenarios and hydrological projection

The climate change projection data used in this study seven GCMs were from the CMIP-3 dataset under the SRES A1B emission scenario from 2010 to 2099, which included the following: UKMO HadCM3, UKMO HadGEM1, NCAR CCSM3.0, MPI ECHAM5, IPSL CM4, CSIRO MK3.0, and CCCMA CGCM3.1. The period from 1961

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to 1990 was used as the baseline, based on CRU TS3.0 gridded ($0.5^{\circ} \times 0.5^{\circ}$) climate data. This study utilized the monthly temperature and precipitation projections from different GCMs using the ClimGen pattern-scaling technique (Osborn, 2009), which were subsequently downscaled to daily resolution. ClimGen created monthly climate scenarios through a pattern scaling approach in which climate change patterns as simulated by a suite of GCMs (Osborn, 2009), and later downscaled to daily resolution following the procedure outlined above. The baseline 1961–1990 used to represent the "present day" climatology of the study area. Climate scenarios were centered around three time periods: 2020s (2010–2039), 2050s (2040–2069), and 2080s (2070–2099), representing the early, middle and late of 21st century. The annual and seasonal changes for projected temperature and precipitation were compared with baseline period from 1961 to 1990 for two catchments over the three time horizons.

For subsequent hydrological projections, this study adopted downscaled projection data derived from the GCMs and validated SWAT models, and projected the impact of climate change on river discharges from 2010 to 2099. The average hydrograph, annual and monthly discharge changes were calculated using 30 years of projected monthly discharge over each of the three time horizons, and then compared with the discharge simulated discharge based on CRU_TS3.0 climate data for baseline period rather than the actual observed discharge data. This technique was used to avoid systematic errors that the SWAT model would introduce in comparing the projection period with the reference period.

The uncertainty envelope of climate projection was showed as function of each GCMs, with the assuming that each climate projection had an equal probability of occurrence. Using the result from 30 year simulations, empirical probability density functions (PDFs) of the projected annual temperature, annual precipitation, and simulated annual discharge were generated. The PDFs indicated the range of possible values for each variable and for each time horizon.

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3 Results

3.1 Projected climate change

3.1.1 Annual temperature and precipitation changes

The mean annual temperature and precipitation projections from the seven GCMs for River Huangfuchuan and River Xiangxi are shown in Fig. 2. All seven GCM projections for the River Huangfuchuan indicated warming and wetting, with air temperature rose from 1.0 to 1.8, 2.1 to 3.8 and 3.0 to 5.5 °C in 2020s, 2050s and 2080s, while precipitation increased by 1 to 13, 1 to 27, and 2 to 39 % respectively for the same slices. For the River Xiangxi, the GCM projections consistently showed rising temperature, with temperature rose from 0.9 to 1.7, 1.9 to 3.4, and 2.7 to 4.9 °C in 2020s, 2050s and 2080s, but two GCMs projected precipitation decreases (CCSM3.0 and ECHAM5), while precipitation changed from −1 to 6, −2 to 13, and −2 to 18 % respectively for the same slices.

The projected ratio of precipitation changing with temperature ranged about 7.8 % $°C^{-1}$ (CCSM3.0) to 0.3 % $°C^{-1}$ (ECHAM5) for River Huangfuchan, with 4 GCMs projected the ratio greater than 5.8 % $°C^{-1}$, and 3 GCMs' projection less than 3.4 % $°C^{-1}$. There were 3 GCMs projected the ratio greater than 3.5 % $°C^{-1}$ and 4 GCMs' projection less than 2.3 % $°C^{-1}$ with 2 GCMs' projection showed the precipitation decreasing with warming as mentioned before for River Xiangxi.

3.1.2 Seasonal temperature and precipitation changes

The projected seasonal temperature and precipitation (Fig. 3) indicated that there was consistent warmer in winter, spring, and summer in the River Huangfuchuan, with temperature increases ranging from 0.7 to 5.3 °C in the 2020s and from 2.5 to 8.6 °C in the 2080s. Winter showed the greatest temperature rise, while several GCMs projected temperature decreasing in autumn. For the River Xiangxi, all seven GCMs projected

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temperature increases in all seasons, ranging from 0.3 to 2.1 °C in the 2020s and 1.9 to 7.0 °C in the 2080s. The temperature increases appeared to be greatest in autumn and least in spring. Both the estimated emissions of greenhouse gases and the total radiative forcing increase are greater at the end of 21st century than earlier of 21st century, which cause the projected temperature are larger in 2080s compare to 2020s. The projected seasonal changes are generally consistent with the projected seasonal changes in eastern Asia. The projected temperature increase are generally greater in winter and autumn compared to summer and spring in eastern Asia support by regional averages of temperature projections from a set of 21 global models for A1B scenario by CMIP3 (Christensen, 2007), and CMIP5 results support this assessment (Christensen, 2013).

There were consistent increases in projected precipitation for winter and spring across the seven GCMs, but the consistency was poorer for summer and autumn precipitation changes. The ratio of percentile precipitation changing with temperature was the highest in winter and the least in summer in both catchments. The projected seasonal precipitation increased more for the River Huangfuchuan than that of the River Xiangxi. For the River Xiangxi, seasonal mean precipitation increases in the 2020s and 2080s were 1.3 and 8.6 %, respectively, while, seasonal mean precipitation increases were 8.6 and 33.6 % for the River Huangfuchuan during the same periods. The difference in projected precipitation among GCMs increased over time horizons in each season, with the maximum range in winter and minimum range in summer among GCMs.

3.1.3 Uncertainties in temperature and precipitation projections

Based on the climate change projections, the calculated Probability Density Functions (PDFs) showed the possible ranges of temperature and precipitation changes during all of the three time horizons (Fig. 4). The most important findings were the increased uncertainties in projected mean annual temperature and precipitation toward the end of the 21st century. Further, the projected mean annual temperature increased in all

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of the GCMs, while precipitation projections showed relatively consistent increases and shifts toward extreme conditions, with the exception of ECHAM5 and CCSM3.0, which showed a decrease in the River Xiangxi. However, while the GCMs showed a consistent direction of changes in temperature and precipitation, there were large differences in the magnitudes of increase. Finally, the magnitudes of the temperature and precipitation changes in the River Huangfuchuan were more than that of the River Xiangxi, indicating that the climate change uncertainty was greater for the River Huangfuchuan.

For River Huangfuchuan, the projected temperature increase from CCSM3.0 and CSIRO showed the smallest magnitude with about 1.0 and 3.0° for 2020s and 2080s, while the ECHMM5 and HadCM3 were at the other end of the spectrum, with increase of 1.8 and 5.3° for the same horizons. The middle-ground model was CCCMA, with projected increase of 1.3 and 3.8° for 2020s and 2080s horizons. For River Xiangxi, the projected temperature from CSIRO and CCSM3.0 also showed the smallest increase with the magnitude about 0.9 and 2.8° for 2020s and 2080s, while the HadCM3, ECHAM5 and IPSL models were at the other end of the spectrum, with increase of 1.7 and 4.9° for the same time horizons. The middle-ground model was also CCCMA, with the projected increase of 1.4 and 4.2° for 2020s and 2080s horizons. The projected temperature from 7 GCMs showed substantial consistency between the relative magnitudes of change associated with the different GCMs for the different time slices for an individual catchment.

For River Huangfuchuan, the projected precipitation from ECHAM5 and CSIRO showed the smallest increase, of less than 5.0 and 15.0 mm for 2020s and 2080s, while the HadCM3 showed the largest increase, of about 50 and 150 mm for 2020s and 2080s. For River Xiangxi, the projected precipitation from ECHAM5 and CCSM3.0 showed decrease, of about maximum decrease 10 and 20 mm for 2020s and 2080s; while the CCCMA, HadCM3 models showed the largest increase, of about 60 and 190 mm for 2020s and 2080s. The projected precipitation change associated with the different GCMs showed consistent for an individual catchment in three time

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horizons, but various not only the magnitude but also the direction between the different catchments. More precisely, the projected precipitation from CCSM3.0 and ECHAM5 showed decrease for River Xiangxi while increase for River Huangfuchuan for the three horizons. Among all GCMs, HadCM3 showed substantial increase for projected precipitation for both catchments.

PDFs also showed that the mean annual temperature of the River Huangfuchuan in the 2080s was outside the natural temperature variation of the baseline with the cold years in 2080s were warmer than the warmest years for baseline. For River Xiangxi, a similar pattern was simulated in the 2050s. The annual precipitation projections were very different from that of temperature. Compared to the baseline, all 7 GCMs' projections indicated that the Rivers Huangfuchuan and Xiangxi will become wetter in the future and the frequency of extreme wet and dry years will also increase.

3.2 Projected climate change impact on discharge

3.2.1 Average hydrograph

The projected average hydrographs of the Rivers Huangfuchuan and River Xiangxi over each of the three time horizons are presented in Fig. 5. Average hydrograph shows a general increase in discharge for the Rivers Huangfuchuan and Xiangxi, with the exception of HadGEM1 and ECHAM5, which project a decrease in the River Xiangxi in summer (June to August). The projected peak discharge showed great increase and appeared earlier during the flood season in the River Huangfuchuan, while that of the River Xiangxi appeared later.

3.2.2 Annual and seasonal river discharge changes

The changes in projected annual and seasonal river discharges are presented in Fig. 6. The projected annual river discharge decreased for River Xiangxi in the ECHAM5 with the magnitude ranged from −1 to −1.7 % during the three time horizons, and the

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projections from other GCMs showed an increase with the magnitude ranged from 0.3 to 7 % in the 2020s, 2 to 18 % in the 2050s, and 3 to 25 % in the 2080s. The projected annual river discharges in the River Huangfuchuan showed consistent increase across all of the GCMs with the magnitude ranged from 5 to 29 % in the 2020s, 12 to 73 % in the 2050s, and 17 to 142 % in the 2080s. The comparison between the two catchments showed that the River Huangfuchuan had substantial increase in annual river discharge than the River Xiangxi.

The changes in projected seasonal river discharge indicated the larger difference t for both of the magnitude and direction comparing with the changes in projected annual river discharge across the GCMs with the three time horizons, especially in the River Xiangxi. The changes in projected discharge increased the most in spring for the River Huangfuchuan.

3.2.3 Extreme discharge changes

Figure 7 shows the projected extreme discharges for both catchments. There was general increase in the extreme discharges (Q05 for high flow and Q95 for low flow) for both catchments in the three time horizons, and the increase in the River Huangfuchuan was more substantial than River Xiangxi. However, the changes in extreme discharge were totally different comparing that of mean flow. The increase in projected high flow for the River Huangfuchuan was less than the increase in mean flow; with substantial uncertainty was projected for mean flow. However, the River Xiangxi showed an increase in projected extreme discharge that was more substantial than that of mean flow, with a larger range.

For the River Huangfuchuan, CSIRO was the only model that projected decreases of Q05 and mean flow (Q50) in the 2020s; the projected Q05 under all of the other GCMs increased over the three time horizons. The projected maximum changes in Q05 and Q50 in the River Huangfuchuan during the 2020s were 39 % (CGCM3.1) and 38 % (IPSL), while the projected maximum changes in Q05 were 70 and 146 % (HadCM3),

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and the projected maximum changes in Q50 were 119 and 304 % (HadCM3) in the 2050s and 2080s.

For the River Xiangxi, HadGEM1, ECHAM5, and CISRO projected a slight decrease in Q50, whereas most GCMs projected an increase in extreme discharge over the three time horizons. The projected Q05 from HadGEM1 decreased during the 2050s and 2080s, while increased for the other GCMs. The maximum increase for Q05 was from the CCSM3.0, with the magnitude of 17 % (2020s), 41 % (2050s), and 63 % (2080s). The projected Q95 decreased in ECHAM5 and IPSL during the 2080s, while increased fort other GCMs, and the maximum increase from the HadCM3 with the magnitude of 27 % (2020s), 38 % (2050s), and 77 % (2080s). The consistent and large increases in Q05 from CCSM3.0 and Q95 from HadCM3 for the three time horizons should be considered when using these information for decision making.

3.2.4 Uncertainty in river discharge projections

The PDFs of the mean annual river discharges are shown in Fig. 8. There were large uncertainties in the model projections, especially toward the end of the 21st century. Some models behaved very differently than others, such as the results of HadCM3 for River Huangfuchuan and HadGEM1 for River Xiangxi. Besides the model uncertainties, projections indicated that the entire discharge distribution shifted toward more extreme events compared to the baseline period from 1961 to 1990. Both the mean and extreme events increased for the two river basins in the future.

4 Discussions

The projected mean annual temperature of the two catchments showed a consistent increasing trend, and the magnitude increased from the 2020s to the 2080s, while the semi-arid River Huangfuchuan had a more consistent and substantial patterns of warming and wetting in the future. Most GCMs revealed near linear increases in annual

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precipitation and discharge in the two catchments, with the exception of projected decrease in precipitation for the River Xiangxi (CCSM3.0 and ECHAM5) (Fig. 2). Figure 6 shows that ECHAM5 is the only mode that projects decreasing mean annual discharge in the River Xiangxi. ECHAM5 is a warmer model comparing with CCSM3.0 that projects slight decrease in annual precipitation. Even if annual precipitation is on the rise, annual discharges nevertheless decrease due to rising evapotranspiration resulting in a net water loss. However, the changes in projected seasonal temperatures, precipitation and discharges in the two catchments are not univocal. The projected temperatures of the River Huangfuchuan show increasing in autumn for all GCMs in the 2020s, but projections from four GCMs show decreasing in the 2080s. The projected seasonal precipitations vary depending on the GCM, time horizon, and on the season (Fig. 3). The seasonal discharge is affected by the combination of these variables. The projected seasonal discharges in River Huangfuchuan appear increasing consistently, and with the exception of a projected decrease in summer from ECHAM5 and a decrease in autumn from CSIRO. The changes in seasonal discharge in the River Xiangxi are however much variability between GCMs. For this reason, to quantify the climate change impacts and assess the uncertainties, multiple GCMs should be used to capture the probability of future change. It has been suggested that the use of two carefully chosen climate projections (dry/hot and wet/cold projections as an example) may be sufficient (e.g. Brekke et al., 2004; Singh et al., 2006).

The projected climate changes show obvious differences in the two catchment of different climate region. The River Huangfuchuan in semi-arid northern China shows more substantial warming and wetting, with larger magnitudes of change in both temperature and precipitation. The River Xiangxi in humid southern China also shows warming, but the increase in precipitation is very slight. These results coincide with increased total annual precipitation, precipitation intensity and extreme precipitation projected for two catchments in eastern China in a future warming scenario (Feng et al., 2011). The increase in precipitation intensity and extreme precipitation is expected to be larger in the middle reaches of the Yellow River basin than that in

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the middle reaches of the Yangtze River basin (Jiang et al., 2011; Xu et al., 2011). Warmer and wetter scenarios for the River Huangfuchuan are projected to increase river discharge substantially, and if managed properly, this could serve to alleviate current local water shortages. However, on the basis of increasing mean discharge, the projected increase in peak flows may also exacerbate soil erosion in the area of loess plateau. The projected increase in extreme flow in River Xiangxi may be expected to increase the fluxes of nonpoint source pollution and sediment to the river channel. However, increased river discharge could also serve to dilute point source pollution and increase the likelihood that target environmental flows are achieved. Increasing river discharge has important implications for the management of water resources in both catchments. Increases in mean flow expand available water resources but the rise in peak flow (Q05) in both basins could increase flood frequency and flood risk. Therefore, adaptation measures need to consider projected changes in mean and extreme flows, as well as the associated uncertainties.

The results highlight the large uncertainty in climate change impacts due to choice of GCM. During the assessment of the climate change impacts, there are considerable difficulties in choosing appropriate GCMs considering each GCM should be treated equally in the assessment. This study use all seven of the GCMs to quantify the uncertainties and ranges of impacts on river discharge, and provide the basis for water management and further adaptations to climate change. However, considering time and calculation limitations, we chose different GCMs in different regions according to pertinent projections. For example, a single variable (temperature) could be chosen for the River Xiangxi based on the examined the maximum and minimum temperature increases. Nevertheless, in the case of the River Huangfuchuan, the temperature-precipitation combination was more appropriate, and we examined the most cold-dry and the most hot-wet extremes.

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5 Conclusions

This study assesse the climate change and impacts of climate change on river discharge in catchments in two different climate regions in China. The projections are carried out using seven GCMs within the CMIP3 structure and uncertainties in the hydrological changes are also quantified.

There are obvious differences in the climate changes and in the impacts of those climate changes on river discharge in the two catchments. Compared to the catchment in the southern subtropical humid area, the catchment in the northern semi-arid area had more apparent warming and wetting, with a greater increase in river discharge. However, the seasonal changes in temperature, precipitation, and river discharge were more complicated than the annual changes, and the uncertainties were greater among the different models. Moreover, the changes in extreme flows (Q05 and Q95) were different than that of mean flows. For example, in the River Huangfuchuan, the mean river discharge increased greatly, but the changes in extreme flow were less than that of mean discharge, so the uncertainty was relatively small. In contrast, in the River Xiangxi, the changes in extreme flows were much larger, leading to larger uncertainties. Thus, changes in extreme flows are far more critical for water managers.

This study revealed the differences between annual and seasonal river discharges in different climate regions and showed the differences between changes in extreme flows and in mean river discharge. These findings have important implications for the basin-scale management of water resources in these catchments and for adaptation measures. It is insufficient to examine the impacts of climate change or evaluate adaptations based on a single global model. The uncertainties between projections from multiple GCMs must be taken into consideration.

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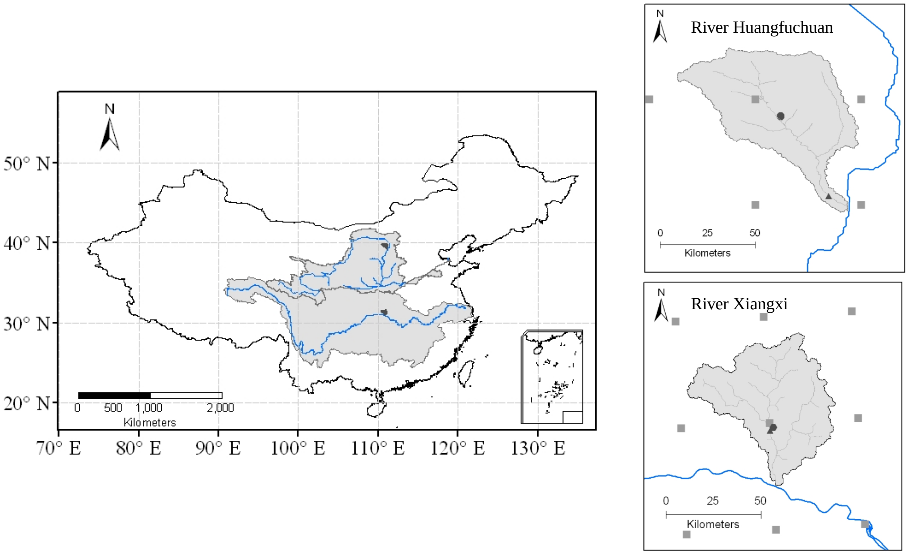

Figure 1. Location of selected sub-catchments in the Yellow River and Yangtze Basins and climate stations (black circle), discharge stations (black triangle), and Climate Research Unit (CRU) grid nodes (grey square).

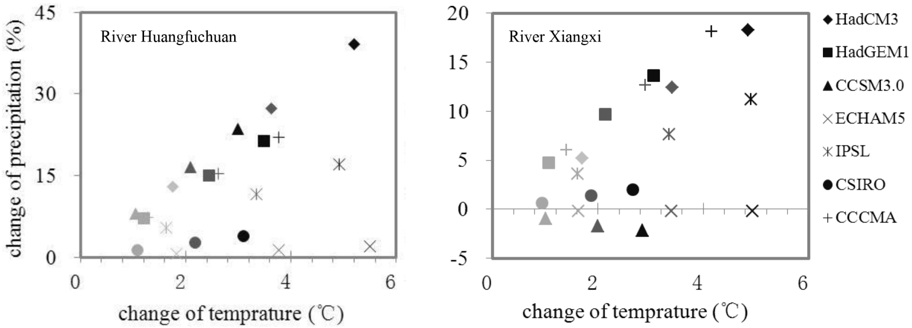

Figure 2. Scatter plots of temperature and precipitation annual changes for seven GCMs projection under 2020s, 2050s and 2080s time horizons for River Huangfuchuan and River Xiangxi (Comparing with 1961–1990 baseline; 2020s: light grey; 2050s: dark grey; 2080s: black).

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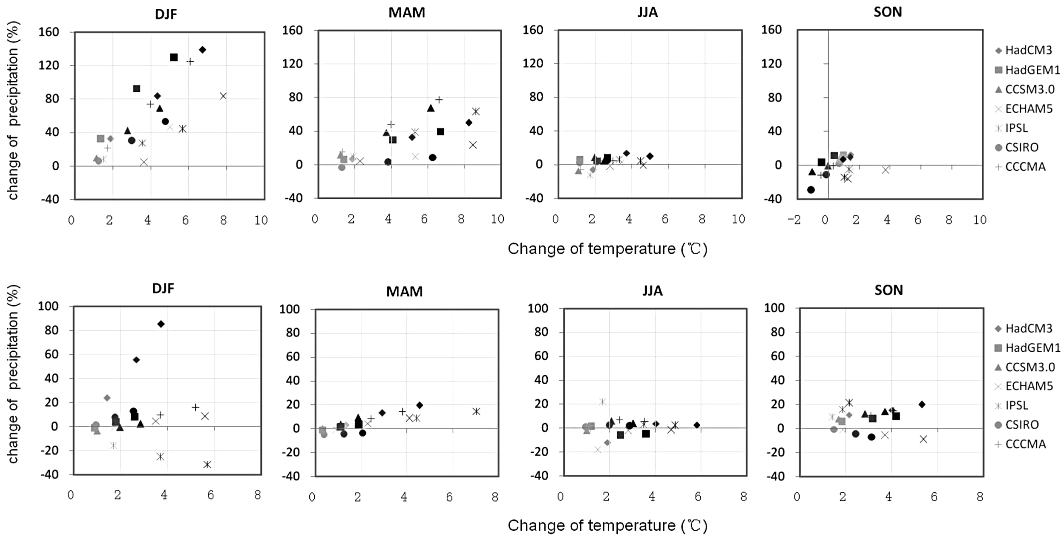

Figure 3. Scatter plots of temperature and precipitation seasonal changes for seven GCMs projection under 2020s, 2050s and 2080s time horizons for River Huangfuchuan (upper row) and River Xiangxi (lower row) (comparing with 1961–1990 baseline; 2020s: light grey; 2050s: dark grey; 2080s: black).

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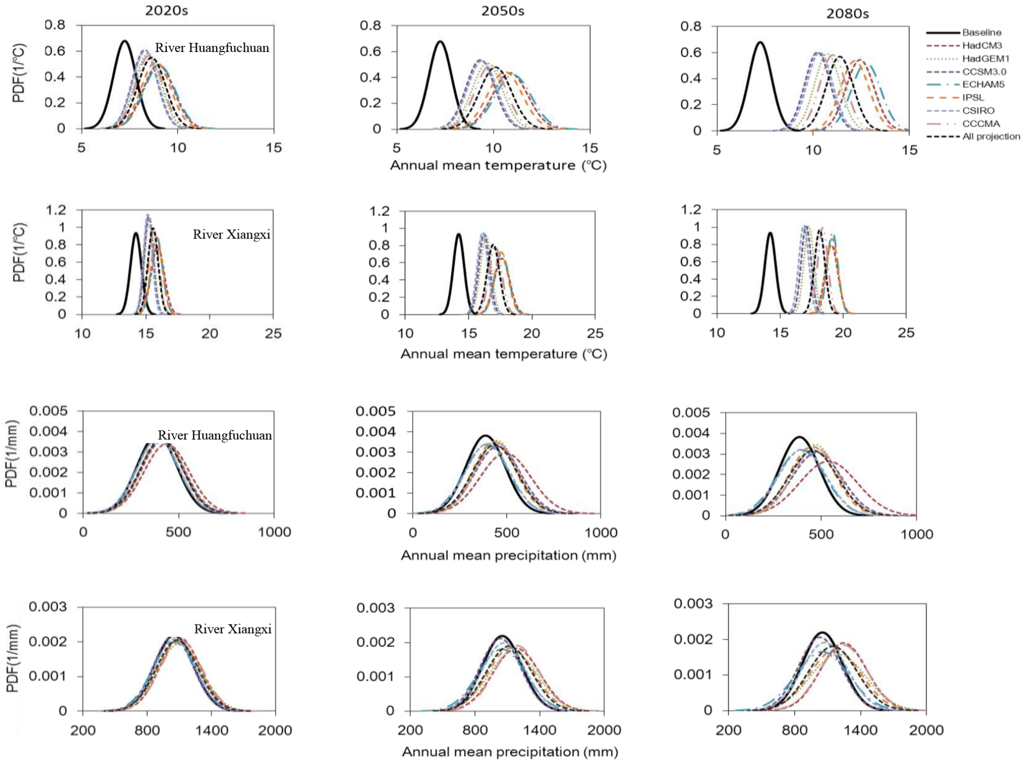

Figure 4. Probability density functions of annual mean temperature and annual precipitation for seven GCMs projection under 2020s, 2050s and 2080s time horizons, and for the 1961–1990 baseline for River Huangfuchuan and River Xiangxi.

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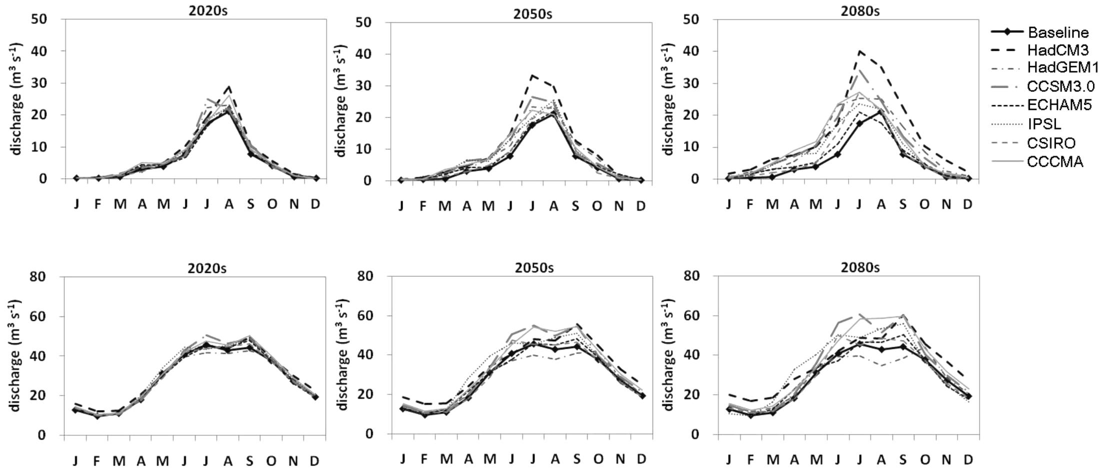

Figure 5. Average hydrographs for seven GCMs projection under 2020s, 2050s and 2080s time horizons, and 1961–1990 baseline for River Huangfuchuan (upper row) and River Xiangxi (lower row).

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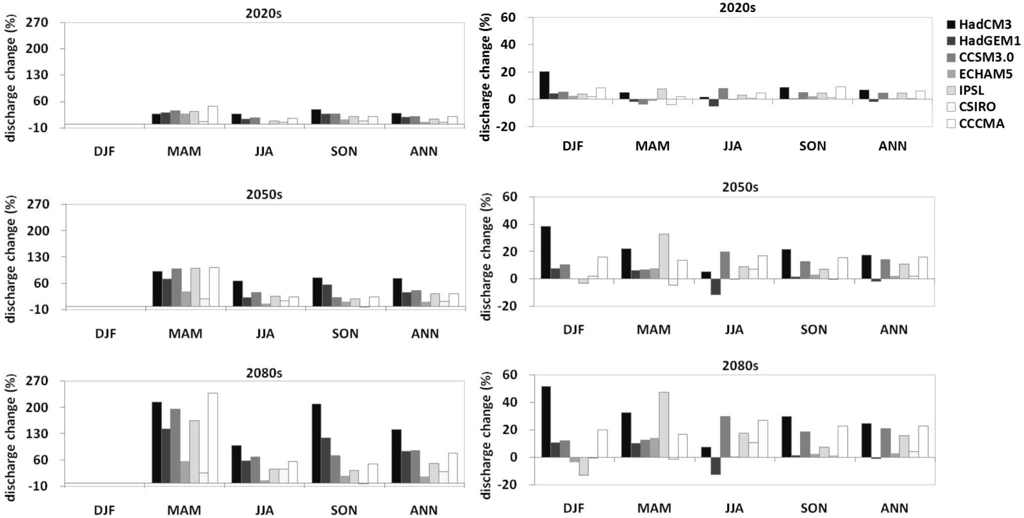

Figure 6. Seasonal and annual discharge changes for 7 GCMs projections under 2020s, 2050s and 2080s time horizons for River Huangfuchuan (left) and River Xiangxi (right) (comparing with 1961–1990 baseline).

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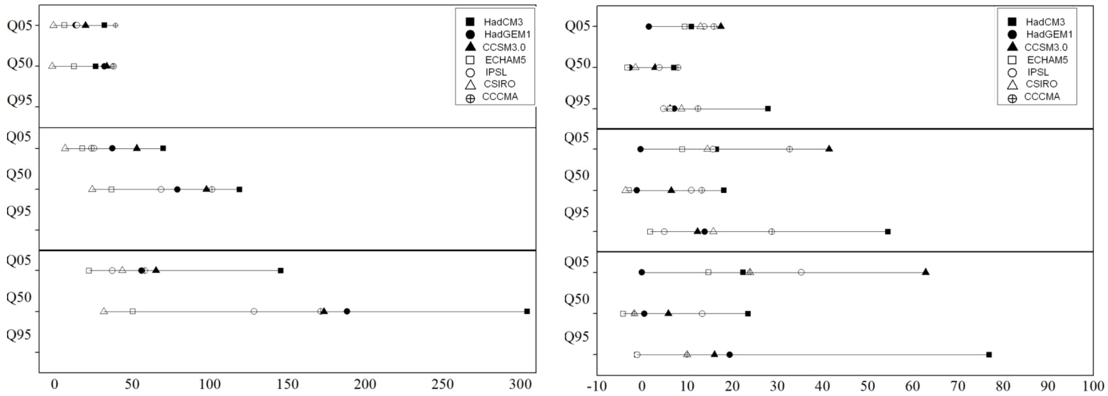

Figure 7. Extreme flows changes for 7 GCMs projections under 2020s, 2050s and 2080s time horizons for River Huangfuchuan (left) and River Xiangxi (right) (Comparing with 1961–1990 baseline; Q05: high flows; Q50: mean flows; Q95: low flows).

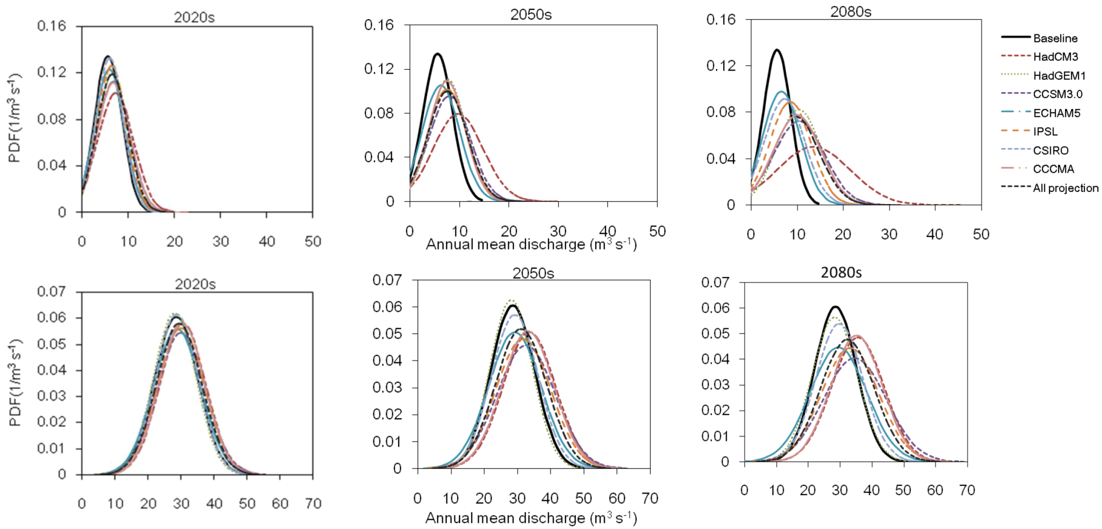

Figure 8. Probability density functions of annual mean discharge for seven GCMs projections under 2020s, 2050s and 2080s time horizons, and for the 1961–1990 baseline for River Huangfuchuan (upper row) and River Xiangxi (lower row).

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