

Submission to HESSD
“Estimation of flood warning runoff thresholds in ungauged basins with asymmetric error functions”, by Elena Toth

Reply Letter

I warmly thank the Editor, Dr. Stacey Archfield, for her kind encouragement and support and the two Referees for having so carefully read the paper, providing very useful and constructive insights, that prompted me to integrate and improve the work.

In the replies that follow, I explain how and where the required information were added and the suggested modifications implemented, for each comment.

For each remark, the original comment of the referee is copied in bold, the reply follows afterwards and the change made in the revised version are in red.

Please note that the insertion of the new figures and the new table added following the suggestions of the Referees have modified the overall numbering of the existing ones, and in particular:

Figure 2 of the original manuscript is Figure 4 in the revised version

Table 1 of the original manuscript is Table 2 in the revised version

Following the replies, the new manuscript is attached, where all the changes to the original version are highlighted in yellow.

I do hope the revisions adequately address the Referees' comments and I am at your disposal for any clarification that may be needed.

My best regards,

Elena Toth

Author's Reply to Referee #1's Comments

General Comments

Comment 1) • What are the advantages of using a biased estimate of $Q_{T=2yr}$? [...]

I do thank the Referee for having pointed out the importance to better underline in the paper the limitations and advantages of the proposed approach, and in particular with respect to a probabilistic one.

It is certainly true that a probabilistic approach (it is possible to find examples also for neural network models, see for example Khosravi et al., 2011) may be able to add very valuable insights for a more complete evaluation of the prediction model, supplementing the information provided by point-value predictions, and in the future I intend to attempt to investigate the uncertainty of the issued predictions, (as I have added in the concluding session in the revised version) but I do not believe it would be possible performing such complex analysis here. It should in fact be considered that uncertainty assessment methods should take into account all error sources (see for example Montanari, 2007) and not only those related to the choice of their parameters (the majority of the uncertainty methods deal only with a single source of uncertainty, for instance, Monte Carlo-based methods analyze the propagation of uncertainty of parameters only) - and are subject, as well as the prediction model itself, to errors in their underlying assumptions and structure as well as in the determination of their parameters (Xiong et al., 2009) so that it is needed, even if it is far from easy, to correctly evaluate also their quality (in many methodologies it is hypothesised and not verified if the distribution of the forecasts is the real one). As a consequence, implementing a correct, fully comprehensive procedure for a consistent and reliable estimation of the global uncertainty is certainly not straightforward (nor it would be possible to describe it briefly) and this is why it may be subject of a separate, future work.

On the other hand, I do not believe it is necessary to add the implementation of a probabilistic framework here, since the presented methodology is a deterministic one, where an optimal point forecast is obtained by minimizing the conditional expectation of the future loss.

Such framework has not the pros of a probabilistic one in terms of quantification of the uncertainty, but its advantage is the operational value of the forecast in terms of an optimal decision that minimizes the cost; in fact asymmetric loss functions are more appropriate in many types of decision settings, as shown by recent forecasting literature analysing the statistical properties of optimal predictions under asymmetric loss (e.g. Christoffersen and Diebold, 1997, Granger and Pesaran, 2000, Patton and Timmermann, 2004; and in particular Zellner, 1986, 2004, showed that once the symmetric loss function is abandoned, optimal forecasts need not be unbiased) and showing that in many " practical applications, asymmetric loss functions can be critical to effective forecasting" (Elliott et al., 2006).

Minimising the asymmetric error function has the purpose of minimizing the cost, thus optimizing the threshold from an operational point of view. A probabilistic forecasting approach applied to the symmetric error function (provided that the methodology is able to include all sources of uncertainty and its quality may be objectively assessed/verified) would certainly provide awareness on the uncertainties associated with the point forecasts, but identifying the upper (e.g. 95%) uncertainty bound would not allow the decision-maker to choose the optimal value for the threshold in terms of costs/operational utility, since such value (upper bound) would be (if reliable) the one that identifies an assigned risk of underestimation (and, even if this is not the point here, it would, I expect, result in a very high value for a small assigned risk,

given the large uncertainty of the approach, mainly due to the intrinsic limitations and shortcomings of the data set for such an heterogeneous area...) but it would not take into account in any way the overestimation costs resulting from high negative errors, nor it would consider the balance between the costs of positive and negative errors, as it may be done, instead, within an asymmetric loss approach.

In the revised version I have better specified the purpose of the proposed approach, along with considerations on the advantages/disadvantages in respect to a probabilistic framework, adding two new paragraphs in Section 2 and in the Conclusions.

Comment 2) • Regional Flood Frequency analysis is not regression. In a couple of locations in the text, page 6014 line 14–29 and page 6030 lines 10–18, there seems that there is the direct association between Regional Flood Frequency analysis to Regression with catchment attributes (regression or related techniques like ANN's). [...]

In the revised version I will certainly rephrase ll. 24-27 p. 6014 and ll. 13-14 p. 6030, since I definitely did not mean to reduce Regional Flood Frequency Analysis to the application of regression techniques, but only to refer to that thematic area, because the runoff threshold literature generally does not include these issues. I fully agree (as highlighted, as the Referee underlines, also in the chapter on floods prediction that I co-authored of the 2013 book) that regression methodologies are only one of the possible methods (statistical and process-based) to predict floods in ungauged basins and in particular I should better specify in the text that their use is especially frequent only as far as the estimation of the index flood values is concerned.

In the revised version I have rephrased both paragraphs (in Section 1 and in the Conclusions)

Comment 3) • Relative error could be also very valuable. For assessing the performance of several variants of the proposed method, the measures MAE and RMSE are proposed, both functions of the error. Given the large range of discharges considered in the study, it could be also very valuable to report additionally boxplots of the relative errors. [...]

I definitely agree that a more comprehensive description of the errors would be very helpful to interpret the results, especially given the large discharge range, as underlined by the Referee; I am not sure the relative errors would be the fairest way to analyse the results in the presented decision setting framework (see reply to first Comment), given that the costs are weighted in respect to the 'not-relative' errors in the loss functions, so I would prefer, if the Referees agrees, adding in the revised paper the scatterplots of observations/predictions, that I believe allow the most complete visualization of the results over the entire discharge range, showing every single prediction in respect to the corresponding observation. In addition, also OverH and UnderH are already defined in relative terms, since they represent the number of errors greater than 30%.

In the revised version I have added a new figure (Figure 3) showing the scatterplots of the results issued by the 5 models over the independent test set, showing every prediction in the respect to the corresponding observation.

Admittedly the scatterplot highlight that the errors are far from negligible for both the traditional and the asymmetric networks (as already underlined in the original paper); but the test is indeed an exacting one: a single regression model applied over a fully independent test set that includes extremely diverse catchments (from Alpine to Mediterranean) and based on a dataset (the only one currently available at national scale) that unfortunately does not include important influencing factors, the most important one being information on the rainfall extremes.

On the other hand, the objective of the work is to propose a new methodological approach, that may be applied to much better databases and the Italian case study is proposed for a comparative analysis, where all the applied models have the same limitations due to the dataset.

Detailed Comments

Page 6013, Line 27: is “real-world” here “real-time”?

Actually, I mean both: ‘real-time’ warning systems actually implemented by ‘real-world’ organization (that is not only in literature simulation studies): I will rephrase to make it clearer.

In the revised version I have replaced “real-world” with “operational real-time”. (Introduction section)

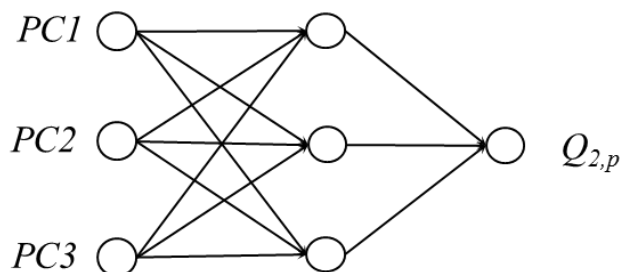
Page 6017, Lines 17–19: The author defines here the error as the observed minus the predicted value. To my knowledge, in runoff prediction in ungauged basins, it is almost a consensus to define error as predicted minus observed. If the author wants to define it here inversely, a stronger warning to the reader should be given, in order to avoid confusions.

I used the notation by Elliot et al (2005), for consistency with their definition of the loss function in Eq 1.

In the revised version I have modified the definition of error as suggested by both Referees, since it indeed has generated confusion (both in the readers and in the author...). Accordingly, I have modified Fig 1, Equations 1 and 3, all the references to positive/negative errors throughout the paper and also the box-plot (ex-Figure 2).

Pages 6021–6023: Maybe adding an schematic figure with the structure of the selected ANN could help the reader.

I fully agree with your suggestion: I will add a figure showing the ANN architecture:



In the revised version I have added the above figure of the network architecture (Fig. 1c, section 4.2).

Page 6027, line 4: is here “scour” the Q2?

Page 6029, line 7: ... the errors are not negligible...

In the revised version I have amended the mistakes.

Page 6027, line 9: I think “prudence” is not the right word here. Maybe “tendency to over/underestimate”?

Page 6031, line 13: Again, “prudentially” is not the right word here.

In the revised version I have rephrased both sentences as suggested.

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Author's Reply to Referee #2's Comments

Detailed comments

Comment: It is indicated that the annual maximum flow records for some stations are available for as few as only 5 years. How was the quantile of interest in this work estimated and how meaningful is the estimate done using such a short data set?

The only quantile that is estimated in this work is the 2-years one and it was estimated as the median of the available historical records of flood maxima. Even if of course it would be preferable having longer time-series, five years should be sufficient for such a short return period, for example according to the classical guideline by Cunnane (1987), that suggests not to extrapolate statistical inference beyond a return period of 2 times the sample length (and for the shortest records in the present application, it is inferred a quantile with a return period that is less than half the sample length). In addition, the stations with less than 8 years of data are only 9, so that I believe the dataset, in terms of the length of the records, may be considered, overall, sufficiently meaningful for the purpose of estimating the 2-years return period flows.

In the revised version I have added such observation in Section 3.1.

Comment: Why were only three classes of catchment descriptors used to sample representative catchments from for the three groups of catchments? Are they not too few to enable a fair distribution of different ranges of the catchment characteristics evenly across all the three groups?

I fully agree that such choice is subjective and that a different number of classes could have been chosen; given the small number of features characterizing the catchments (the 3 first principal components), I believe that 3 classes should be sufficient for identifying training, cross-validation and test sets that are sufficiently similar. I report below the graphs showing the mean value (red dash) and the 90% and 10% percentiles of the resulting sets, for each of the three input variable (PC1, PC2 and PC3). The graphs seems to highlight a good degree of similarity in the distribution of the values over the three sets. Such graphs might be added in the revised manuscript at the end of Section 3.2.

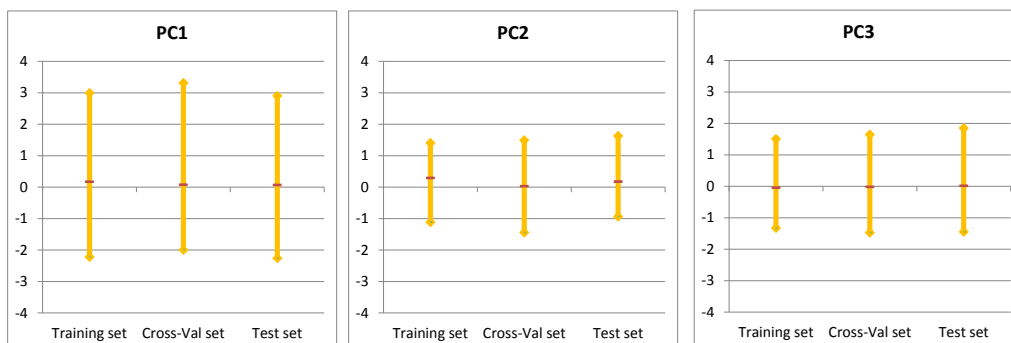


Figure: Mean value (red dash) and the bars comprised between the 90% and 10% percentiles of the resulting training, cross-validation and testing sets, for each of the three input variable (PC1, PC2 and PC3).

In the revised version I have added the above graphs in a new figure (Fig. 2) at the end of section 3.2.

Comment: How were the output values standardized in the range between -1 and 1 (page 6023, line 12). Here I assume the output variables to be the 2-year flood values.

Related to my previous comment, are the error terms in Equations 3-5 estimated from the normalized 2-years flood values or from the actual values? If they are estimated from the actual

values, as it looks is the case by looking at the values of MEA and RMSE in Table 1 and the errors in Figure 2, how was the scale inconsistency at the different stations handled? It is mentioned somewhere that these values range between 10 and 1000m³/s.

Data standardization is generally used in neural network in order to ensure that the data receive equal attention during the training process (Maier and Dandy, 2000) and it is also important for the efficiency of training algorithms (Dawson and Wilby, 2001).

In the present case, the output data are rescaled as a function of the minimum and maximum values to the [-1 1] range (actually to the [-0.95 +0.95] range, to avoid the problem of saturation – I had not explained this issue in order not to further complicate the explanation of the neural networks working). Such rescaled values are those that are simulated by the ANN model: those corresponding to the training and cross-validation sets are used, as ‘target’, for training the neural networks; when the model is successively used for predicting the standardized $Q_{2,p}$ over the independent test set, the ANN output values are then transferred back reversing the function, restoring the values - in a proportional manner - to the original ranges.

Finally, the error statistics presented in section 5 (and in Table 1 and Figure 2) are calculated on the errors between the actual observed values and the re-transformations of the values issued by the neural network.

It is certainly necessary to better explain the procedure in the revised manuscript, both in Section 4.2 and 5.1 (and in the latter it is indeed important to clarify that the prediction $Q_{2,p}$ is not directly the value issued by the model but its ‘de-standardised’ value, since the present wording is, as highlighted by the Referee, confusing).

In the revised version I have better explained the standardization method in Section 4.2 and in Section 5.1 I have explained that the model results are compared after having de-standardized the network outputs.

Comment: I find the whole text on page 6028 messy. Most of the discussion on results is presented on this page, but it is very confusing. The author mentions that negative errors mean overestimation and a couple of lines later a contradictory statement is made (statements on line 6 and 10). Similarly, it is mentioned somewhere that the overestimation error reduces with increasing alpha value and the opposite is mentioned elsewhere. There is even little consistency between what is discussed here and the referred Table 1 and Figure 2.

Thank you indeed for having identified two mistakes in the same sentence (I really have to apologise: I had in my last version reworded the sentence changing the focus from under to overestimation errors and I have inadvertently maintained a part of the original sentence...): lines 10-12 should read:

“At the same time, and more importantly, the number of negative (~~under~~ overestimation) errors larger than 30% substantially decreases with a , with OverH% reaching a value that is much lower than that of the ANN-Symm model when a arrives at ~~0.4~~ 0.1 (18% vs 34%)”

I do hope that amending the wrong sentence, the text will become more clear and it should be consistent with the results shown in Table 1 and Figure 2.

In the revised version, also due to the change in the definition of the errors (as suggested by the Referee in the comment below), I have amended the mistaken sentence and also rephrased the following part of page 6028 that was indeed not clear.

Comment: Why did the author choose to define the error term as the observed minus the simulated values? Defining it in a more conventional way would have helped to avoid such inconsistency.

In the manuscript I used the notation by Elliot et al (2005), for consistency with their definition of the loss function in Eq 1.

In the revised version I have modified the definition of error as suggested by both Referees , since it indeed has generated confusion (both in the readers and in the author...). Accordingly, I have modified Fig 1, Equations 1 and 3, all the references to positive/negative errors throughout the paper and also the box-plot (ex-Figure 2).

Other comments

Comment: Define the variable M in Equation 3.

Thank you for pointing this out:

In the revised version I have added that M is the number of records in the set (either the early-stopping validation set or the test one).

Comment: I suggest that the catchment descriptors be listed in a table. I am a bit astonished to read that data on soils and land cover are missing when there are open data sources on both that are often used in modeling.

I will list, as suggested, the descriptors in a new Table (new Table 1):

1	<i>Long</i> - UTM longitude of catchment centroid
2	<i>Lat</i> - UTM latitude of catchment centroid
3	<i>A</i> - Catchment drainage area
4	<i>P</i> - Catchment perimeter
5	z_{max} - Maximum altitude of the catchment area
6	z_{min} - Elevation of the catchment outlet
7	z_{mean} - Mean altitude of the catchment area
8	<i>L</i> - Length of the Maximum Drainage Path
9	S_L - Average slope along the Maximum Drainage Path
10	S_A - Catchment average slope
11	Φ - Catchment orientation
12	<i>MAP</i> - Mean Annual Precipitation

And I definitely take the Referee's point that it would be extremely helpful to extend the database content, and working on a consistent, comprehensive database of Italian catchments with validated and reliable information on other important features of the catchments' areas.

However, such compilation of an extended database for the Italian country was not the object of the present analysis, that presents a comparison of methodologies applied utilizing the same dataset, and I based the analysis on the data made available by the CUBIST project (the most recent National project of characterization of the Italian basins) and already used in Di Prinzio et al. (2011).

I do hope, in the (hopefully near) future, that the colleagues who prepared the CUBIST database (and who have already developed the analyses for the delineation of the catchment boundaries) will find the time (and I may certainly offer my help, too) to include additional descriptor to the national database.

On the other hand, the objective of the work is to propose a new methodological approach, that may be applied to much better databases and the Italian case study is proposed for a comparative analysis, where all the applied models have the same limitations due to the dataset.

In the revised version I have listed the descriptor in a new table (new table 1) and I have added in section 3.1 that "the CUBIST set is currently the only database available in the Italian hydrologists community at national scale".

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1 Estimation of flood warning runoff thresholds in ungauged 2 basins with asymmetric error functions

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7 Abstract

8 In many real-world flood forecasting systems, the runoff thresholds for activating warnings or
9 mitigation measures correspond to the flow peaks with a given return period (often the 2-year one,
10 that may be associated with the bankfull discharge). At locations where the historical streamflow
11 records are absent or very limited, the threshold can be estimated with regionally-derived empirical
12 relationships between catchment descriptors and the desired flood quantile. Whatever is the
13 function form, such models are generally parameterised by minimising the mean square error, that
14 assigns equal importance to overprediction or underprediction errors.

15 Considering that the consequences of an overestimated warning threshold (leading to the risk of
16 missing alarms) generally have a much lower level of acceptance than those of an underestimated
17 threshold (leading to the issuance of false alarms), the present work proposes to parameterise the
18 regression model through an asymmetric error function, that penalises more the overpredictions.

19 The estimates by models (feedforward neural networks) with increasing degree of asymmetry are
20 compared with those of a traditional, symmetrically-trained network, in a rigorous cross-validation
21 experiment referred to a database of catchments covering the Italian country. The analysis shows
22 that the use of the asymmetric error function can substantially reduce the number and extent of
23 overestimation errors, if compared to the use of the traditional square errors. Of course such
24 reduction is at the expense of increasing underestimation errors, but the overall accurateness is still
25 acceptable and the results illustrate the potential value of choosing an asymmetric error function
26 when the consequences of missed alarms are more severe than those of false alarms.

27 1 Introduction

28 In the operation of flood forecasting systems, it is necessary to determine the values of threshold
29 runoff that trigger the issuance of flood watches and warnings. Such critical values might be used for
30 threshold-based flood alert based on real-time data measurements along the rivers (WMO, 2011) or
31 for identifying in advance, through a rainfall-runoff modelling chain, the rainfall quantities that will
32 lead to surpass such streamflow levels, as in the Flash Flood Guidance Systems framework (Carpenter
33 et al., 1999; Ntelekos et al., 2006; Reed et al., 2007; Norbiato et al., 2009).

34 A runoff threshold should correspond to a 'flooding flow', that is to a value that may produce flood
35 damages, and it is very difficult to determine on a regional or national scale: it may be defined as a
36 flow that just exceeds bankfull conditions, but in practice, both in gauged and in ungauged river

1 sections, such conditions are arduous to quantify due to the lack of local information (Reed et al.,
2 2007; Hapuarachchi et al., 2011).

3 In absence of more sophisticated physically-based approaches, based on detailed information of
4 each specific cross-section that are rarely available due to limited field surveys, the literature
5 suggests to estimate the bankfull flow as the flood having a 1.5 to 2 years return period (Carpenter et
6 al., 1999; Reed et al., 2007; Harman et al., 2008; Wilkerson, 2008; Hapuarachchi et al. 2011; Cunha et
7 al., 2012; Ward et al., 2013) and a flow that is slightly higher than bankfull may be identified with the
8 2-year return period flood (Carpenter et al., 1999; Reed et al., 2007).

9 Many operational systems all around the world adopt a statistically-based definition of the flooding
10 flow and the flows associated with given return periods are used as threshold stages for activating
11 flood warning procedures.

12 The 2-year recurrence is used by many River Forecast Services in the United States, as suggested by
13 Carpenter et al. (1999), also due to the fact that *“the good national coverage of the 2-yr return period
14 flows that the U.S. Geological Survey (USGS) maintains nationwide supports its use”* (Ntelekos et al.,
15 2006), as well as in British Columbia (Canada).

16 However, the floods with different annual exceedance probabilities, associated with different levels
17 of risk, are also frequently adopted in **operational real-time** flood warning systems: for example in
18 the Czech Republic, flood watch usually corresponds to a 1- to 5-year flow return period (Daňhelka
19 and Vlasák, 2013). In Italy, where a national directive issued in 2004 introduces a system articulated
20 on at least two levels of flow thresholds, many Regions have identified the alert levels as flood
21 quantiles with return periods of 2, 5 or 10 years (e.g. the Abruzzo, Lombardia, Puglia Regions). In the
22 South of France, the AIGA flood warning system compares real-time peak discharge estimated along
23 the river network (on the basis of rainfall field estimates and forecasts) to flood frequency estimates
24 of given return periods (with three categories: yellow for values ranging from the 2-year to the 10-
25 year flood, orange for between the 10 and the 50-year flood, and red for peaks exceeding the 50-
26 year flood) in order to provide warnings to the national and regional flood forecasting offices (Javelle
27 et al., 2014).

28 For river sections where the streamflow gauges are newly installed or where historical rating curves
29 are not available, the observations of the annual maxima are absent or very limited and it is not
30 possible obtaining a reliable estimate of flood quantiles on the basis of statistical analyses of series of
31 observed flood peak discharges.

32 For these ungauged or poorly gaged basins, the peak flow of given frequency to be associated with
33 the watch/warning threshold can be estimated **transferring information from data-rich sites to data-
34 poor ones, as it is done in the corpus of methodologies** applied in RFFA (Regional Flood Frequency
35 Analysis) at ungauged sites, that have always received considerable attention in the hydrologic
36 literature (Bloeschl et al., 2013). **Among the possible approaches (statistical and process-based) to
37 predict floods in ungauged basins, many researchers have traditionally applied regression-like
38 regionalisation methods for** i) the estimation of the index flood (Darlymple, 1960), usually defined as
39 either the mean or the median (that is the 2-year return period quantile) of the annual maximum
40 flood series, or for ii) the direct estimate of other quantiles of annual maxima in ungauged basins

1 (Stedinger and Lu, 1995; Salinas et al., 2013). Such **methods** are based on the assumption that there
2 is a relationship between catchment properties and the flood frequency statistics and are
3 implemented through a regression-type model that relates the flood quantile or the index flood to a
4 number of relevant morpho-climatic indexes. Linear or power (often linearized through a log-
5 transformation) forms, with either a multiplicative or additive error term, are the most commonly
6 used functions (see e.g. Stedinger and Tasker, 1985; GREHYS, 1996; Pandey and Nguyen, 1999; Brath
7 et al., 2001; Kjeldsen et al., 2001, 2014; Bocchiola et al., 2003; Merz and Bloeschl, 2005; Griffis and
8 Stedinger, 2007; Archfield et al., 2013; Smith et al., 2015).

9 In order to allow more flexibility to the model structure (whose 'true' form is of course not known),
10 the international literature has recently proposed methods based on the use of artificial neural
11 networks (ANN), providing a non-linear relationship between the input and output variables without
12 having to define its functional form a priori. Successful applications of ANN for the estimation of
13 index floods or flood quantiles at ungauged sites are reported in Muttiah et al., 1997; Hall et al.,
14 2002; Dawson et al., 2006; Shu and Burn, 2004; Shu and Ouarda, 2008; Singh et al., 2010; Simor at
15 al., 2012; Aziz et al., 2013.

16 Both the traditional power form or linear regression methods and the neural networks models are
17 generally parameterized by minimizing the mean or root mean of the squared errors, that is a
18 symmetric function assigning the same importance to overestimation and underestimation errors.

19 Nevertheless, the consequences of under or overestimating the runoff threshold when used for early
20 warning are extremely different.

21 Adopting a watch threshold that is higher than the runoff/stage that actually produces flooding
22 damages would in fact lead to missing such events, failing to issue an alarm. Underestimating the
23 runoff threshold may instead determine the issue of false alarms.

24 False alarms may certainly lead to money losses and also "undermine the credibility of the warning
25 organisation but are generally much less costly than an unwarned event." (UCAR, 2010): in fact the
26 costs of failing to issue an alarm grow rapidly in a real emergency, since a totally missed event has
27 strongly adverse effects on preparedness. Not only the costs of false warnings are commonly much
28 smaller than the avoidable losses of a flood, but they cannot match up to indirect and/or intangible
29 flood damages such as loss of lives or serious injuries (Pappenberger et al., 2008; Verkade and
30 Werner, 2011).

31 Furthermore, regarding the effects of false alarms, "in opposition to 'cry wolf' effect, for some they
32 may provide an opportunity to check procedures and raise awareness, much like a fire practice drill."
33 (Sene, 2013)

34 Overall, false alarms have usually a higher level of acceptance than misses and this entails that the
35 estimate of flood warning thresholds should be cautionary, so as to reduce, conservatively, the
36 number of missed alarms.

37 For the development of watches and warnings it is therefore important to obtain estimates as
38 accurate as possible, minimising both positive and negative errors, but, considering that an error will

1 always be present, it is better underpredicting rather than overpredicting the threshold estimate, for
2 safety reasons.

3 To obtain a conservative estimate of the thresholds, penalising more the predictions that exceed the
4 “observed” values (in the present case represented by the quantile estimate based on the statistical
5 analysis of observed flow peaks) than those that underestimate them, in the present work it is
6 proposed, for the first time to the Author’s knowledge, a parameterisation algorithm that weights
7 asymmetrically the positive or negative errors, in order to decrease the consistency of
8 overestimation and therefore the risk of missing a flooding occurrence.

9 It is important to underline that the proposed asymmetric error function is here applied for
10 optimising a neural network model for predicting the 2-year return period flood (due to its
11 association with the bankfull conditions) but it might be used to improve any other kind of
12 methodology for the estimate of flood warning thresholds associated to any return period.

13 Section 2 presents the asymmetric error functions; the next one describes the information available
14 in a database covering the entire Italian country and the identification of the subsets to be used for a
15 rigorous cross-validation approach. Section 4 presents the implementation of the models for
16 estimating the 2-year return period flood in ungauged catchments, consisting in artificial neural
17 networks calibrated using respectively the symmetric square error and the asymmetric error
18 functions. The results are presented and then discussed in section 5 and section 6 concludes.

19 **2 The asymmetric error function**

20 The scientific literature on forecasting applications, in any scientific area, adopts almost exclusively
21 an objective function based on the sum or mean of the squared discrepancies, that is a symmetric
22 quadratic function, due to the well-established good statistical properties of the minimum mean
23 square error estimator.

24 On the other hand, in economics as well as in engineering and other many fields, there are cases
25 where the forecasting problem is inherently non-symmetric and, in the financial forecasting
26 literature, the use of mean squared error, even if still widely applied, is nowadays not always
27 accepted.

28 Error (or loss) functions devised to keep into account an asymmetric behaviour have been proposed,
29 such as the linear-exponential, the double linear and the double quadratic (Christoffersen and
30 Diebold 1996; Diebold and Lopez 1996; Granger 1999; Granger and Pesaran 2000; Elliot et al. 2005;
31 Patton and Timmerman, 2006). In particular, Elliot et al. (2005) recently presented a family of
32 parsimoniously parameterized error functions that nests mean squared error loss as a special case
33 (Patton and Timmerman, 2006).

34 Such function, adapted from Elliot et al. (2005) and defining the error ε as the prediction minus the
35 observed value (that is, a negative error corresponds to underestimation, a positive one to
36 overestimation), reads:

$$37 \quad L(p, \alpha) = 2 \cdot [\alpha + (1 - 2\alpha) \cdot \mathbf{1}\{\varepsilon > 0\}] \cdot |\varepsilon|^p, \quad (1)$$

1 where $\mathbf{1}(\cdot)$ is a unit indicator, equal to one when $\varepsilon > 0$ and zero otherwise; p is a positive integer that
2 amplifies the larger errors (corresponding to a quadratic error when equal to 2) and $\alpha \in (0,1)$ is a
3 parameter representing the degree of asymmetry.

4 For $\alpha < 0.5$ the function penalises more the overestimation errors ($\varepsilon > 0$), while for $\alpha > 0.5$ more
5 weight is given to **negative** forecast errors (under-predictions); for $\alpha = 0.5$ the loss weights
6 symmetrically positive and negative errors.

7 When $p = 2$ and $\alpha \neq 0.5$, the error becomes the asymmetric double quadratic (Quad-Quad) loss
8 function (see Christoffersen and Diebold 1996), that is used in the present work for a fair comparison
9 with the traditional mean square error estimator. When $p = 2$ and $\alpha = 0.5$, Eq. (1) corresponds in fact
10 to the 'traditional', symmetric, square error:

$$11 \quad L(2,0.5) = \varepsilon^2 \quad (2)$$

12 Figure 1 shows the asymmetric Quad-Quad loss function (with α varying from 0.1 to 0.9) compared
13 with the squared error (SE).

14 In the water engineering field, the asymmetric Elliot error function with quadratic amplification ($p =$
15 2) has been recently applied to parameterise a model for estimating the expected maximum scour at
16 bridge piers, in order to obtain safer design predictions through the reduction of underestimation
17 errors by Toth (2015).

18 **It should be noted that the proposed methodology is a deterministic one, where an optimal point**
19 **forecast is obtained by minimizing the conditional expectation of the future loss; such framework has**
20 **not the pros of a probabilistic one in terms of quantification of the uncertainties of the prediction,**
21 **but it aims at identifying the optimal value for the threshold in terms of operational utility.**

22 In Section 4, the asymmetric quadratic error function is proposed for optimizing the parameters of an
23 input-output model, based on artificial neural networks, between the input variables summarising a
24 set of catchment descriptors (obtainable also for ungauged river sections) and the 2-year return
25 period flood, thus warranting that overestimation errors, that would increase the risk of missing
26 flood warnings, are weighted more than underestimation ones.

27 **3 Available information: the national data set of Italian catchments**

28 The case study refers to a database of almost 300 catchments scattered all over the Italian peninsula,
29 compiled within the national research project "CUBIST – Characterisation of Ungauged Basins by
30 Integrated uSe of hydrological Techniques" (Claps et al., 2008).

31 **3.1 Input and output variables**

32 The 12 geomorphological and climatic descriptors **are listed in Table 2**. The dataset unfortunately
33 lacks information on other hydrological properties (e.g. on soils, land-cover, vegetation) and the
34 climatic characterisation is very limited (for example information on extreme rainfall would be
35 extremely important), but **the CUBIST set is currently the only database available in the Italian**
36 **hydrologists community at national scale.**

1 The dataset is described in Di Prinzio et al. (2011), where, following a catchment classification
2 procedure based on multivariate techniques, the descriptors were used to infer regional predictions
3 of mean annual runoff, mean maximum annual flood and flood quantiles through a linear
4 multiregression model.

5 As described in such work, in order to reduce the high-dimensionality of the geomorphological and
6 climatic descriptors set, a Principal Components (PC) analysis was applied, obtaining a set of derived
7 uncorrelated variables. The PC variables are as many as the original variables, but they are ordered in
8 such a way that the first component has the greatest variability, the second accounts for the second
9 largest amount of variance in the data and is uncorrelated with the first and so forth. In the present
10 data set, the first three principal components explain more than three quarters of the total variance
11 (see Di Prinzio et al., 2011) and such three first PCs are here chosen as input variables to the models
12 described in the following, assuming that they may adequately represent, in a parsimonious manner,
13 the main features of the study catchments.

14 The data base, in addition to the morpho-pluviometric information, includes the annual maxima flow
15 records for periods ranging from 5 to 63 years, whose median values, corresponding to the 2-year
16 return period, represent the output variable to be simulated by the models. Even the shortest
17 records (and actually only 9 of the locations have less than 8 years of data) should be sufficient for
18 such a short return period, for example according to the classical guideline by Cunnane (1987), that
19 suggests not to extrapolate statistical inference beyond a return period of 2 times the sample length.

20 The data set covers a great diverseness of hydrological, physiographic and climatic properties and in
21 order to partially reduce such heterogeneity, it was decided to limit the analysis to catchments
22 having a 2-year flood included in the range 10-1000 m³/s, that is 267 over the original 296 basins.

23 **3.2 Identification of balanced cross-validation subsets with SOM clustering of** 24 **input data**

25 As will be detailed in Section 4, the database is to be divided in three disjoint subsets (called training,
26 cross-validation and test sets) in order to allow a rigorous independent validation and also to
27 increase the generalization abilities of the model when encountering records different from those
28 used in the calibration (or 'training') phase, following an 'early stopping' parameterisation procedure.

29 The way in which the data are divided may have a strong influence on the performance of the model
30 and it is important that each one of the three sets contains all representative patterns that are
31 included in the dataset. As proposed in the recent literature (Kocjancic and Zupan, 2001; Bowden et
32 al., 2002; Shahin et al., 2004) a self-organising map (SOM) may be applied to this aim. The SOM is a
33 data-driven classification method based on unsupervised artificial neural networks that may be
34 applied for several clustering purposes (for hydrological applications see, for example, Minns and
35 Hall, 2005; Kalteh et al, 2008).

36 In the recent years, SOMs were also successfully applied for catchments classification either based on
37 geo-morpho-climatic descriptors (Hall and Minns, 1999; Hall et al., 2002; Srinivas et al., 2008; Di
38 Prinzio et al., 2011) or based on hydrological signatures (Chang et al., 2008; Ley et al., 2011; Toth,
39 2013); however, it is important to underline that the clustering is not carried out here in order to

1 identify a pooling group of similar catchments for developing a region-specific model, but for the
2 optimal division of the available data for the parameterization and independent testing of a single
3 model to be applied over the entire study area.

4 The SOM is in fact used to cluster similar data records together: an equal number of data records is
5 then sampled from each cluster, ensuring that records from each class (that is catchments with
6 different features) are represented in the training, validation and test sets, that, as a result, have
7 similar statistical properties (Bowden et al., 2002; Shahin et al., 2004).

8 A SOM (Kohonen, 1997) organizes input data through non-linear techniques depending on their
9 similarity. It is formed by two layers: the input layer contains one node (neuron) for each variable in
10 the data set. The output-layer nodes are connected to every input through adjustable weights,
11 whose values are identified with an iterative training procedure. The relation is of the competitive
12 type, matching each input vector with only one neuron in the output layer, through the comparison
13 of the presented input pattern with each of the SOM neuron weight vectors, on the basis of a
14 distance measure (here the Euclidean one). In the trained (calibrated) SOM, all input vectors that
15 activate the same output node belong to the same class.

16 In the present application, the dimension of the input layer is equal to three (that is, the first three
17 principal components of the catchments descriptors); as far as the output layer is concerned, there is
18 not a predefined number of classes and, given the small dimension of the input variables, it was here
19 chosen a parsimonious output layer formed by three nodes in a row, each one corresponding to a
20 class.

21 The three resulting clusters are formed respectively by 121, 70 and 76 catchments; each cluster is
22 then divided into three parts, and one third is assigned to the training, validation and test sets
23 respectively. Overall, the training, validation and test sets are therefore equally numerous (91, 88
24 and 88 records respectively) and formed by the same proportion of catchments belonging to each of
25 the clusters, having eventually a similar information content, as shown by the similar statistics of the
26 three variables in the three sets represented in Figure 2.

27

28 **4 Development of symmetric and asymmetric artificial neural networks** 29 **models for estimating the 2-year return period flows at ungauged sites**

30 **4.1 Feedforward Artificial Neural Networks**

31 Artificial neural networks are massively parallel and distributed information processing systems,
32 composed by nodes, arranged in layers, that are able to infer a non-linear input-output relationship.
33 ANN, and in particular feedforward networks have been widely used in many hydrological
34 applications (see for example the recent review papers by Maier et al., 2010 and by Abrahart et al.,
35 2012) and the readers may refer to the abundant literature for details on their characteristics and
36 implementation.

1 Three different layer types can be distinguished: input layer, connecting the input information, one
2 or more hidden layers, for intermediate computations, and an output layer, producing the final
3 output; adjacent layers are connected through multiplicative weights and, in each node, the sum of
4 weighted inputs and a threshold (called bias) is passed through a non-linear function known as an
5 activation.

6 The models here applied are networks formed by one hidden layer, with tan-sigmoid activation
7 functions, and a single output node (corresponding to the estimated flood with 2-year return period),
8 with a linear activation function.

9 The identification of the network's weights and biases (called training procedure) is carried out with
10 a non-linear optimization, searching the minimum of an error (or learning) function measuring the
11 discrepancy between predicted and observed values, and feedforward networks are generally
12 trained with a learning algorithm known as BackPropagation (Rumelhart et al. 1994) based on
13 steepest descent or on more efficient quasi-newton methods.

14 In order to avoid overfitting, that degrades the generalisation ability of the model, the Early Stopping
15 or Optimal Stopping procedure was applied (see, for example, Coulibaly et al., 2000). For applying
16 Early Stopping, the available data have been divided into three disjoint subsets with a similar
17 information content, as described in Section 3.2: a training set, an early-stopping validation set and a
18 test set. While the network is parameterised minimising the error function on the training set, the
19 error function on the early-stopping validation set is also monitored; if the error function on such
20 second set increases continuously for a specific number of iterations, this is a sign of overfitting of
21 the training set: the training is then stopped and network parameters at the lowest validation error
22 are returned. The third set (test set) is not used in any way during the parameterization phase, but it
23 is used for out-of-sample, independent evaluation of the resulting models.

24 **4.2 Implementation of the symmetric model**

25 Neural networks, including those applied in the recent hydrological literature for the estimation of
26 index floods or flood quantiles at ungauged sites, are traditionally trained minimizing the square
27 error function, which is symmetrical about the y-axis and negative or positive discrepancies of the
28 same magnitude result in the same function value.

29 In the present work, the results obtained by a network trained with a 'conventional' square error
30 function are compared with those obtained when parameterising the network through the
31 minimisation of an asymmetric loss function, that takes into account both over and underestimation
32 discrepancies but penalizes more the overprediction errors, since the consequences of missing
33 alarms are more severe than those of false alarms.

34 For both type of models, the output values (2-year flood values) are rescaled as a function of the
35 overall minimum and maximum values to the $[-0.95,+0.95]$ range, to facilitate the optimization
36 algorithms and also avoid saturation problems by accommodating possible extreme values occurring
37 outside the range of available data (Dawson and Wilby, 2001). Each implemented architecture is
38 randomly initialized for ten times to help avoiding local optima: the parameter set that results in the

1 minimum error function on the early stopping validation data (second set) is chosen as the trained
2 network.

3 The first implemented model is obtained through the minimization of the traditional, symmetric
4 mean squared error, applying the quasi-Newton Levenberg-Marquardt BackPropagation algorithm
5 (Hagan and Menhaj 1994), widely applied and regarded as one of the most efficient neural network
6 training algorithms.

7 The input variables are the first three principal components of the catchment descriptors, so the
8 input layer is formed by three nodes; the output node corresponds to the estimated flood with 2-
9 year return period; as far as the dimension of the hidden layer is concerned, there is, unfortunately,
10 no definitive established methodology for its determination, because the optimal network
11 architecture is highly problem-dependent: different architectures with a number of hidden nodes
12 varying from 2 to 6 were set up and the mean squared error of the estimates issued for the third,
13 independent set resulted the lowest with the hidden layer formed by 3 nodes.

14 The architecture with three input nodes, three hidden nodes and 1 output node, **represented in**
15 **Figure 3**, is therefore the network finally chosen; the network parameterized minimising the
16 symmetric mean square error function will be denoted as ANN-Symm, and its results will be in
17 Section 5 compared with those of the asymmetric models having the same architecture but a
18 different error function.

19 **4.3 Implementation of asymmetric models with varying degree of asymmetry**

20 The Quad-Quad loss function described in Section 2 is here applied for calibrating the network
21 parameters of the asymmetric models. The learning function to be minimized is therefore the
22 average value of the double quadratic errors (Mean Quad-Quad Error, *MQQE*), obtainable averaging
23 **the M (number of records in the set)** errors given by Eq. (1) when $p=2$:

$$24 \quad MQQE = \frac{2}{M} \sum_{j=1}^M [\alpha - (1 - 2\alpha) \cdot \mathbf{1}_{\{\varepsilon > 0\}}] \cdot |\varepsilon_j|^2 \quad (3)$$

25 The value of α , corresponding to the degree of asymmetry of the loss function, cannot be fixed a
26 priori, since such choice should be based on a location-specific cost-benefit analysis, keeping into
27 account the avoidable losses (that is the direct and indirect losses, provided they may be
28 quantifiable, that may be prevented by mitigation actions following an alarm issue) and the cost of
29 the mitigation measures themselves. Such analysis is acknowledged to be extremely difficult,
30 especially since it involves also intangible costs such as life losses, but also warning credibility issues;
31 furthermore, the costs may change over time and are also dependent on the warning lead-time (see
32 e.g. Martina et al., 2006; Verkade and Werner, 2011, Montesarchio et al., 2011/2014).

33 For this reason, in the present application, different asymmetric networks, with α varying from 0.4 to
34 0.1, are implemented, in order to compare the results obtainable with a different asymmetry degree,
35 that is a different extent of importance given to over/underestimation errors. Such asymmetrically
36 trained network are in the following denoted as “Asymm- 0.4”, “Asymm- 0.3”, “Asymm-
37 0.2”, “Asymm- 0.1”.

1 The training of the four asymmetric networks, based on the minimisation of the Mean Quad-Quad
2 Error, is carried out through the generalization of the backpropagation algorithm proposed by Crone
3 (2002) and applied by Silva et al. (2010), that may be used for parameterising artificial neural
4 networks with any differentiable (analytically or numerically) error function.

5 **5 Results and discussion**

6 **5.1 Goodness-of-fit measures and plots**

7 As described in section 4.2, the neural networks are trained over the standardized (rescaled) output
8 values of the training and cross-validation sets and they are successively used for predicting the
9 output over the independent test set: such ANN output values are then scaled back, obtaining the
10 predictions $Q_{2,p}$.

11 The performances of the **models** are evaluated through a set of indexes that describe the prediction
12 error, ε , that is the difference between the de-standardised predictions, $Q_{2,p}$ issued by the models (as
13 a function of morpho-climatic attributes only) and the 'observed' 2-year flood values (the median of
14 historical annual maxima), $Q_{2,o}$, on the third set (test set), formed by $N=91$ catchments distributed all
15 over the country, whose data have not been used in any capacity in the models' development.

16 The following error statistics have been computed:

17 *MAE (mean absolute error)*

$$18 \quad MAE = \frac{\sum_{i=1}^N |\varepsilon(i)|}{N} \quad (4)$$

19 *RMSE (root mean square error)*

$$20 \quad RMSE = \sqrt{\frac{\sum_{i=1}^N (\varepsilon(i))^2}{N}} \quad (5)$$

21 MAE and RMSE both represent a symmetric accuracy, corresponding to the distance of the
22 predictions from the observations independently of the error sign (and the RMSE, being quadratic,
23 emphasizes more the larger errors).

24 In order to keep into account the differences in sign of the errors, representing the extent of
25 overpredictions as compared to underpredictions, the overall percentage of **positive** errors (Over%),
26 is computed:

27 *Over% (percentage of overestimates)*

$$1 \quad \text{Over}\% = \frac{\{i=1, \dots, N | Q_{2,p}(i) > Q_{2,o}(i)\}}{N} \quad (6)$$

2 Such metric shows the general tendency of the model to overestimate (or to underestimate, since
 3 $100 - \text{Over}\%$ represents, conversely, the proportion of underpredictions), but these indexes do not
 4 distinguish among errors of different magnitude, since they count also predictions that may be only
 5 barely above (or below) the targets, that is very good predictions, with minimum errors.

6 It is therefore computed also the number of the ‘high’ overestimation errors, keeping into account
 7 only the more relevant, and therefore potentially more dangerous, overpredictions. It was here
 8 considered as ‘high overprediction’ an estimate that is more than 30% higher than the corresponding
 9 target value:

10 *OverH% (percentage of high overprediction errors)*

$$11 \quad \text{OverH}\% = \frac{\{i=1, \dots, N | Q_{2,p}(i) > 1.3 \cdot Q_{2,o}(i)\}}{N} \quad (7)$$

12 The more conservative is the threshold estimate, the lower is the value of OverH%.

13 On the other hand, even if - as discussed – generally less crucial in terms of consequences, also the
 14 number of high underestimation errors should be monitored, since excessively low values imply the
 15 tendency of the model to establish thresholds leading to the issuance of too many false alarms.

16 *UnderH% (percentage of high underprediction errors):*

$$17 \quad \text{UnderH}\% = \frac{\{i=1, \dots, N | Q_{2,p}(i) < 0.7 \cdot Q_{2,o}(i)\}}{N} \quad (8)$$

18 In addition to the goodness-of-fit measures (reported in Table 2), the boxplot of the errors (predicted
 19 minus observed quantiles) is shown in Figure 4: the bottoms and tops of the rectangular boxes are
 20 respectively the lower and the upper quartiles, the horizontal segment inside the box is the median
 21 and the whiskers represent the 5th and 95th percentiles.

22 The results may be evaluated also through the scatterplots of predicted (y-axis) vs observed (x-axis)
 23 quantiles, presented in Figure 5 that show every prediction $Q_{2,p}$ in respect to the corresponding
 24 ‘observation’ $Q_{2,o}$.

25

26 5.2 Discussion of the results

27 The boxplot (Fig. 4) allows to visually assess both the accuracy and the tendency to
 28 over/underestimate of the models: the boxes should be compact and close to the dotted line

1 representing zero error but at the same time it is better if the data lie **below** such line, thus indicating
2 that the method do not tend to overpredict the thresholds and the warning system is therefore less
3 subject to miss a potentially dangerous flood.

4 It may be seen that for the network that was trained minimising the traditional Square Error (ANN-
5 Symm) the box and whiskers are centred on the zero-error line and the quantiles (top/bottom of the
6 box, top/bottom whiskers) are at a similar distance from such line, showing that the errors are
7 equally distributed among overestimation and underestimations. The box is compact, demonstrating
8 the good accurateness of the method for a substantial part of the test set, but, due to the symmetric
9 disposition of the errors, many overestimation errors, also remarkably high, are issued, as shown by
10 the position of the **upper** whisker.

11 Analysing Table 2, the relatively good accuracy of the ANN-Symm model is demonstrated by the
12 values of the MAE and RMSE, that are the lowest among the implemented models. The symmetric
13 distribution of the overall errors is shown by an Over% close to 50% and the similar values of the
14 OverH% (34%) and UnderH% (32%) confirm that also the high relative errors are equally split among
15 over and underestimates.

16 Such results were expected since the training is based on a symmetric loss function, but the
17 consequence is that the ANN-Symm model issues a remarkable number of significant overprediction
18 errors, in fact for about one third of the test catchments the estimates are more than 30% higher
19 than the observations.

20 The analysis of Table 2 shows that the asymmetrically trained networks tend, for decreasing α values,
21 to reduce the number of overestimations (**positive** errors). For the overall errors this is shown by the
22 different proportion of over/underestimations, that moves from a value that corresponds,
23 approximately, to a balance, to a much more skewed distribution of **negative vs positive** errors, with
24 Over% decreasing up to 31%.

25 At the same time, and more importantly, the number of **positive (overestimation)** errors larger than
26 30% substantially decreases with α , with OverH% reaching a value that is much lower than that of
27 the ANN-Symm model when α arrives at **0.1** (18% vs 34%).

28 Conversely, as expected, the more asymmetric is the network, the higher are the underprediction
29 errors, as shown by the values of UnderH%: the number of significant **negative** errors gradually
30 increases from one third up to 47% of the total.

31 Also the accuracy (given by the total amount of the discrepancies independently of their sign)
32 deteriorates when the asymmetry is more pronounced, but the drop is moderate and the RMSE and
33 MAE values are not so far from those of the ANN-Symm network.

34 **Looking at the parallel boxplots (Fig. 4), it may be seen that with increasing asymmetry the boxes**
35 **become less compact and, as expected, their position shifts downwards. The length of the upper**
36 **whiskers substantially decrease with α but the length of the lower whiskers does not increase at the**
37 **same rate, thus compensating for the fact that the boxes are taller for the more asymmetric models.**
38 **It follows that the global distances from the 5% to the 95% percentiles (given by the distance**
39 **between the ends of the top and bottom whiskers) are very close for the symmetric (ANN-Symm)**

1 and for the two most asymmetric, thus showing that the variability of the errors for the vast majority
2 (middle 90%) of the data is similar. On the other hand, overall, the errors are moving towards the
3 underestimation side for increasing asymmetry (as confirmed also by the corresponding median
4 values) and for Asymm-01, the upper part of the box indicates that only about one quarter of the
5 errors are overestimations.

6 It may be noted, in particular from the scatterplots (Fig. 5), that, for both symmetric and asymmetric
7 models, the errors are not negligible: this is due to the shortcomings of the available data set but
8 mainly to the intrinsic limitations of a regional approach applied to the extreme variability of the
9 study area. As already underlined in Section 3.1, the national data set lacks important information
10 that may help to characterise the hydrological behaviour and the phenomena governing formation of
11 extreme flows. In addition to the unavoidable risk of erroneous data, the absence in the database of
12 additional influences certainly further hampers the possibility to obtain a reliable relationship with
13 the flood quantiles. Most importantly, the data set covers the entire Italian peninsula, characterised
14 by extremely different hydro-climatic settings (from Alpine to Mediterranean ones) and this high
15 heterogeneity is certainly an additional reason that limits the performance.

16 Notwithstanding the limitations of the dataset, that affect equally all the proposed models, the
17 results demonstrate that the use of the double quadratic error function, even if at the expense of
18 more substantial underestimation errors, can substantially decrease the number and extent of
19 overestimation errors, if compared to the use of the traditional square errors.

20 In the application to a specific cross section, the degree of asymmetry might be identified as
21 proportional to the “risk averseness” of the situation: the more the impact of false alarms is,
22 comparatively, small, the more the decision-makers are reluctant to the consequences (economic
23 and social) of a flood and, rather than risking a missed alarm, can accept many cases of false alarm
24 with the associated costs.

25 **6 Conclusions**

26 A crucial issue in the operation of flood forecasting/warning systems at ungauged locations is how to
27 assess the possible impacts of the forecasted flows, that is the identification of streamflow values
28 that may actually cause flooding, to be associated to thresholds that trigger the issuance of flood
29 watches and warnings. The values that may produce damaging conditions (or “flooding flows”), when
30 in absence of detailed local information on each cross-section, are in many parts of the world
31 estimated as the peak floods having a certain return period, often the 2-year one, that is generally
32 associated with the bankfull discharge.

33 For locations where the gauges are new or where historical rating curves are not available, the series
34 of past annual flow maxima are absent or very limited, and the peak flow of given frequency to be
35 associated with the watch/warning threshold can be estimated with regionally-derived empirical
36 relationships, such as those that may be applied for the estimation of the index flood at ungauged
37 sites. Such regression-like methods consist in a relation between a set of catchment descriptors that
38 may be obtained also for ungauged sites and the desired flood quantile; linear or power forms are
39 the most commonly used functions, but recent studies have successfully applied artificial neural
40 network models, due to their flexibility, to flood quantile and index flood estimation.

1 Whatever is the function form, such models are generally parameterised by minimising the mean
2 square error, that assigns equal weight to overprediction or underprediction errors, whereas,
3 instead, the consequences of such errors are extremely different when the estimates are to be used
4 as warning threshold. In fact, false alarms (due to an underprediction of the warning threshold)
5 generally have a much higher level of acceptance than misses (that would derive from an
6 overestimated threshold).

7 For this reason, in the present work, the regression model (a feed-forward neural network) is
8 parameterised minimising an asymmetric error function (of the double quadratic type), that
9 penalizes more the overestimation than the underestimation discrepancies. The predictions of
10 models with increasing degree of asymmetry are compared with those of a traditional (trained on
11 the symmetric mean of square errors) neural network, in a rigorous cross-validation experiment
12 referred to a database of catchments covering all the Italian country.

13 The results confirm, as expected, that the more asymmetric is the network, the more numerous and
14 higher are the underprediction errors, and the less numerous and less severe are the overestimation
15 errors. As also expectable, the symmetric accuracy decreases when the asymmetry is more
16 pronounced, but the drop is moderate and the RMSE and MAE values are not so far from those of
17 the traditionally trained network.

18 Undoubtedly, the nature of the regional approach, as well as the shortcomings of the dataset and the
19 extreme heterogeneity of the study area, generate errors much greater than those obtainable with
20 detailed local studies. On the other hand, where no alternatives exist, the proposed methodology
21 may provide a preliminary estimate of the threshold runoff that do not, prudentially, overestimate
22 the actual flooding flow.

23 Notwithstanding the acknowledged limitations of the dataset, that affect equally all the proposed
24 models, the analysis shows that the use of the asymmetric error function substantially reduces the
25 number and extent of overestimation errors, if compared to the use of the traditional square errors.
26 Of course such reduction is at the expense of increasing underestimation errors, but the overall
27 precision is still acceptable and the study highlights the potential benefit of choosing an asymmetric
28 error function when the consequences of missed alarms are more severe than those of false alarms.

29 Minimising the asymmetric error function has the purpose of optimizing the threshold from an
30 operational point of view, in a deterministic framework: future analyses may be devoted to
31 investigate the uncertainty of the issued predictions, since a probabilistic approach (provided that
32 the methodology is able to include all sources of uncertainty and its quality may be objectively
33 assessed) may provide very valuable insights for a more complete evaluation of the model,
34 supplementing the information provided by point-value predictions.

35 It is important to highlight that the asymmetric error function is used, in this study, to parameterise a
36 neural network, but of course it might be used to optimize any other model or equation, when
37 aiming at obtaining conservative estimates, for safety reasons.

38 The appropriate degree of asymmetry might be identified depending on the risk-averseness of the
39 specific flood-prone context. The quantification of risk aversion is extremely difficult and case-
40 specific: it should keep into account that the perception of society may be very different from a

1 technical appraisal of the involved costs and it should include also indirect, intangible and long-term
2 impacts. More research on the societal perception in different contexts would greatly improve the
3 process of risk-based decision-making (Merz et al., 2009), including the choices concerning flood-
4 warning thresholds. Hopefully, in the next years, a more direct collaboration between the hydrologic
5 and socio-economic research communities, as auspicated in the new Panta Rhei science initiative
6 (Montanari et al., 2013; Javelle et al., 2014), will provide a progress in this direction.

7

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22

1 **Tables**

2

3 **Table 1. Geomorphological and climatic descriptors of the CUBIST database of Italian catchments**

1	Long - UTM longitude of catchment centroid
2	Lat - UTM latitude of catchment centroid
3	A - Catchment drainage area
4	P - Catchment perimeter
5	zmax - Maximum elevation of the catchment area
6	zmin - Elevation of the catchment outlet
7	zmean - Mean elevation of the catchment area
8	L - Length of the Maximum Drainage Path
9	SL - Average slope along the Maximum Drainage Path
10	SA - Catchment average slope
11	Φ - Catchment orientation
12	MAP - Mean Annual Precipitation

4

5

6 **Table 2. Goodness-of-fit criteria of the 2-year floods estimates obtained by the symmetric and**
 7 **asymmetric networks on the independent test set of catchments.**

Model\Index	MAE (m^3/s)	RMSE(m^3/s)	Over%	OverH%	UnderH%
<i>Symm</i>	98	133	46%	34%	32%
Asymm-04	104	147	42%	32%	35%
Asymm-03	105	152	41%	30%	37%
Asymm-02	108	162	36%	27%	41%
Asymm-01	115	178	31%	18%	47%

8

9

1 **Figure Captions**

2 Figure 1. Asymmetric Quad-Quad loss function (with α varying from 0.1 to 0.9) compared with
3 the Squared Error (SE).

4

5 Figure 2: Mean value (red dash) and the bars comprised between the 90% and 10% percentiles of
6 the resulting training, cross-validation and testing sets, for each of the three input variable (PC1,
7 PC2 and PC3).

8

9 Figure 3. Architecture of the chosen network, with three input nodes, three hidden nodes and 1
10 output node.

11

12 Figure 4. Parallel box-plots of the errors ($\epsilon=Q_{2,o}-Q_{2,p}$) of the 2-year floods estimates obtained
13 by symmetric and asymmetric networks on the independent test set of catchments.

14

15 Figure 5. Scatterplots of the predicted (y-axis) vs observed (x-axis) 2-year floods estimates on the
16 independent test set of catchments, for the symmetric and asymmetric models.

17

Figure 1. Asymmetric Quad-Quad loss function (with α varying from 0.1 to 0.9) compared with the Squared Error (SE).

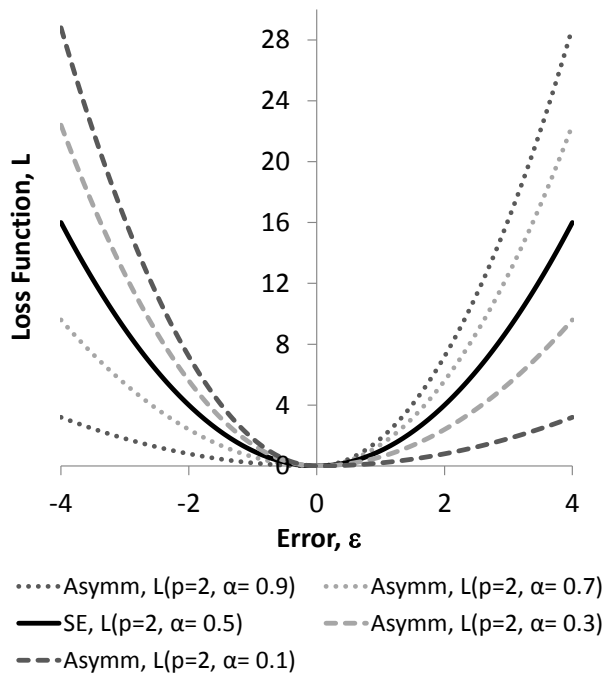


Figure 2: Mean value (red dash) and the bars comprised between the 90% and 10% percentiles of the resulting training, cross-validation and testing sets, for each of the three input variable (PC1, PC2 and PC3).

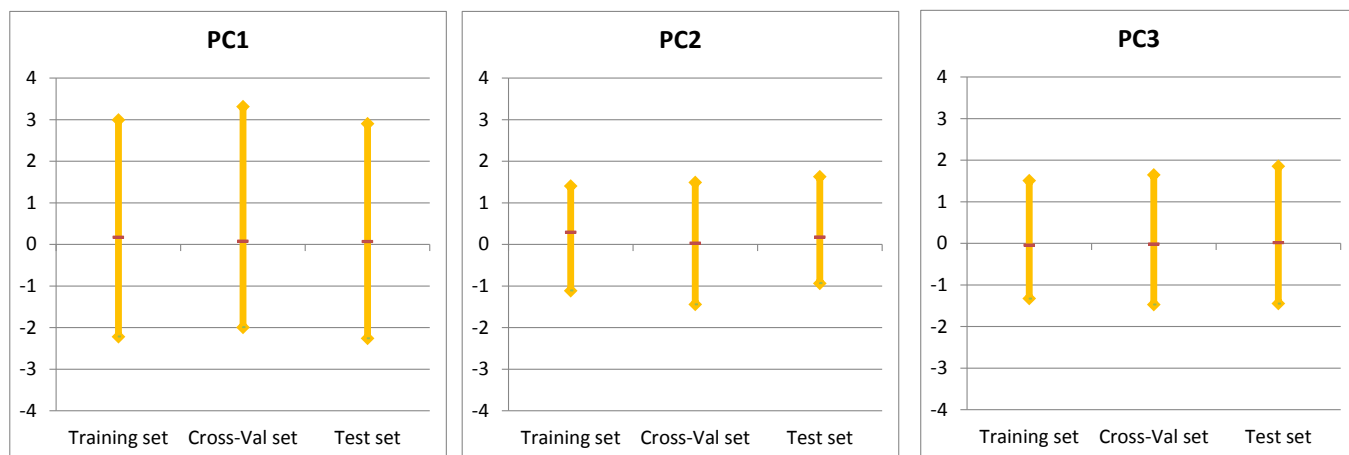


Figure 3. Architecture of the chosen network, with three input nodes, three hidden nodes and 1 output node.

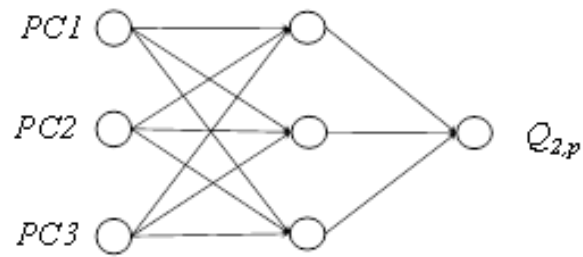


Figure 4. Parallel box-plots of the errors ($\varepsilon=Q_{2,o}-Q_{2,p}$) of the 2-year floods estimates obtained by symmetric and asymmetric networks on the independent test set of catchments.

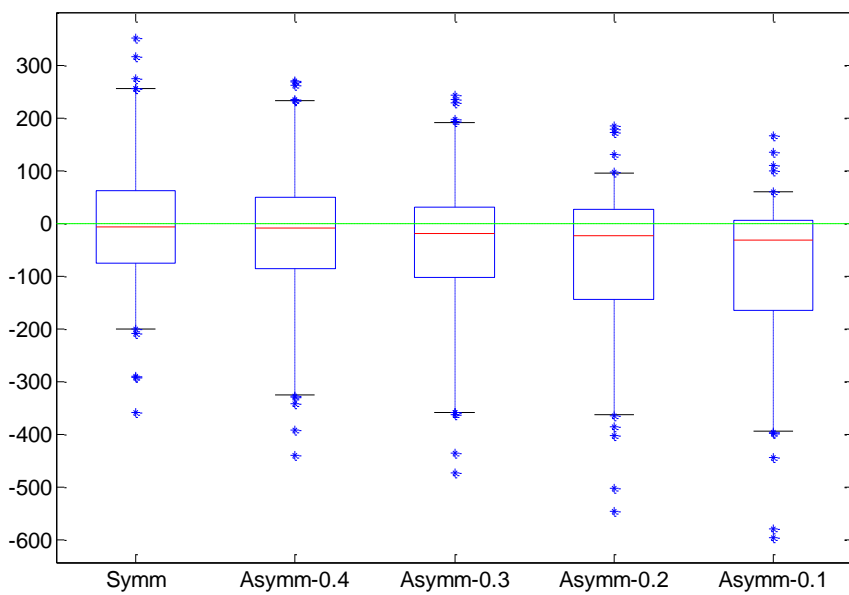


Figure 5. Scatterplots of the predicted (y-axis) vs observed (x-axis) 2-year floods estimates on the independent test set of catchments, for the symmetric and asymmetric models.

