

1 **HESS-2015-176**
2 **COMMENTS FROM EDITORS AND REVIEWERS**

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5
6 **Editor**

7 I agree with the reviewers comments, in particular reviewer #2 that:

8
9 **Ed.1)** ... the paper is currently hard to follow as it refers to earlier work which needs to be more clearly
10 included.

11 *Answer from authors: The comparison with the previous study gives a great opportunity to answer a*
12 *central question: is it worth the computational effort to include the groundwater as a dynamic reservoir. In*
13 *this updated manuscript we have tried to make clearer links to the previous study and put more effort in*
14 *explaining each comparison.*

15
16 **Ed.2)** ... Other important issues include a justification for the simplifications.

17 *Answer from authors: We have carefully explained and justified the many simplifications throughout the*
18 *manuscript. The Discussion has now more focus on the different types of simplifications: 1) the*
19 *simplifications needed to keep the chosen SDP method computationally feasible and 2) the simplifications*
20 *needed due to poor data availability.*

21
22

23 **Anonymous Referee #1**

24 In this paper, a hydroeconomic modelling approach is used to find cost-optimal sustainable surface water
25 and groundwater allocation strategies for a river basin. A simplified management problem with conjunctive
26 use of scarce surface water and groundwater under inflow and recharge uncertainty is presented. Because
27 of head-dependent groundwater pumping costs the optimization problem is non-linear and non-convex,
28 and a genetic algorithm is used to solve the 1-step-ahead sub-problems with the objective of minimizing
29 the sum of immediate and expected future costs. A real-world application in the Ziya River Basin in
30 northern China is used to demonstrate the model capabilities. It's estimated that the annual cost of ending
31 groundwater overdraft in the basin is estimated to be 5.47 billion CNY/year. Both the methods and the
32 results have reference value.

33
34 **AR1.1)** "Persistent overdraft from the groundwater aquifers on the North China Plain has caused declining
35 groundwater tables, salinization and infiltration of wastewater." Here the expression of "salinization" is
36 confusing. In fact, the lowdown of ground water table has been favorable for the control of salinization and
37 alkalization. Maybe the expression can be changed to "Persistent overdraft from the groundwater aquifers
38 on the North China Plain has caused declining groundwater tables, and infiltration of saline water and
39 wastewater.

40 *Answer from authors: Yes. We have also deleted "saline" since this is outside the focus areas of our study:*
41 *"Persistent overdraft from the groundwater aquifers on the North China Plain has caused declining*
42 *groundwater levels".*

43
44 **AR1.2)** For table 3: What the meaning of SP E? Why it's the same for different scenarios (before and after
45 SNWTP)?

46 *Answer from authors: "SP E" is the average increase in the total costs as a consequence of introducing a*
47 *minimum in-stream flow constraint. To estimate this shadow price, we need comparable scenario runs with*
48 *and without this additional ecosystem constraint. The presented shadow price is only valid for that*
49 *particular scenario and does not cover different situations, e.g. before/after the SNWTP. The white area*
50 *above and below the listed value in Table 3 has been confusing because it may indicate that the shadow*
51 *price covers all scenarios.*

52 *To avoid confusion, we decided to delete these almost blank columns and instead describe them in the*
53 *Results section. Thereby, we also avoid the use of SP E.*

54 **AR1.3)** For table 3: What's the meaning of SP SNWTP? Why for LGW, the shadow price is the lowest?
55 Because the initial condition is more severe, the value of water should be higher. Please give an
56 explanation.

57 *Answer from authors: "SP SNWTP" is the average reduction in the total costs, associated with the*
58 *introduction of the South-to-North Water Transfer Project (SNWTP). The total costs after the SNWTP is put*
59 *in operation are compared to the scenario without the SNWTP (pre-2008) and divided by the allocated*
60 *SNWTP water. While this is meaningful when comparing identical scenarios with and without the SNWTP, it*
61 *is misleading when the increased costs are caused by other factors. We have deleted the SP SNWTP values*
62 *for all scenarios not identical to the baseline setup. The remaining two SP SNWTP-values (row 2 and 5) have*
63 *been presented in the text. The SP SNWTP and SP E columns have been deleted from Table 3.*

64
65

66 **Anonymous Referee #2**

67 The manuscript presents a method consisting in a Stochastic Dynamic Programming (SDP) management
68 model for a system including one reservoir and one aquifer. The aquifer is represented as a box model. The
69 problem is solved with a combination of Genetic Algorithms and Linear Programming (GA-LP) to tackle the
70 non-linearities and non-convexities caused by the head-dependent pumping costs. The framework is
71 applied to the Ziya River system (North China), where groundwater overdraft has led to a significant
72 decrease in the aquifer levels. The results of the SDP are provided in the form of water value tables used as
73 prices in a forward-moving simulation run. The estimated costs given by the model when the aquifer levels
74 reach equilibrium, in comparison with business-as-usual values not considering groundwater
75 overexploitation (previous paper), serve as estimation of the cost associated to a recovery in the aquifer
76 level.

77
78 The provided manuscript refers to a critical problem in many arid and semiarid areas: persistent
79 groundwater overexploitation, which has caused considerable damage in both water quantity and quality
80 across the world. The methodology is well-presented and exposed in the case study. Coupling stochastic
81 programming and groundwater simulation is cumbersome, and new approaches to alleviate its complexity
82 and transform those results into management policies could support the application of those tools in water
83 resources management. For that, this paper has a considerable potential interest for publication in HESS. In
84 addition, it is well-written and well-structured. However, there are some important points that the authors
85 should address in order to enhance the manuscript.

86
87 **AR2.1)** The method strongly simplify the hydrology (just a Budyko model for assessing runoffs, and fixed %
88 of groundwater recharge no justified), as well as the spatial representation of the system (all surface
89 reservoirs lumped into a single one) and the groundwater simulation (a lumped box model with unclear if
90 not missing representation of stream-aquifer interaction). Despite the presentation as a hydroeconomic
91 model, the economics is also highly simplified (constant water demands, constants curtailment cost). These
92 simplifications need to be justified, including an analysis of how realistic these assumptions are. This can be
93 done along the text when the assumptions are presented. Overall, the limitations of the modelling
94 approach should be clearly stated either in the Discussion or the Conclusions.

95 *Answer from authors: There are two main reasons for the high level of simplification: Limited data*
96 *availability and the limitations of the SDP method (curse of dimensionality). All assumptions, simplifications*
97 *and their implications are now carefully discussed in the revised manuscript, e.g. as inserted in the*
98 *Discussion: "the simple system representation needed in SDP required assumptions of inflow and storage*
99 *discretization, generalized estimates of pumping cost and a lumped groundwater model which all contribute*
100 *to the uncertainty. Further, poor data availability for the case study area required some rough estimates of*
101 *the natural water availability, single-point demand curves and perfect correlation between rainfall and*
102 *groundwater recharge. The method-driven assumptions generally limit the decision support to basin-scale*
103 *while the simple estimates caused by poor data availability contribute to raising the general uncertainty of*
104 *the model results. Given the computational challenges and the diverse and significant uncertainties, the*
105 *model results should be seen as a demonstration of the model capabilities rather than precise cost*
106 *estimates. Better estimates will require access to a better case dataset and involve a more comprehensive*
107 *sensitivity analysis."*

108 **AR2.2)** The paper constantly refers to the previous analysis done by the authors, published in another
109 paper, whose results represent the business as usual situation, not shown in this one with the exception of
110 the total annual cost (Discussion). Thus, the presented paper looks like a second part of the one previously
111 referred, since which it is quite hard to fully understand it without the other one. Maybe the authors could
112 briefly include more description of the method and results for the business as usual situation, or update
113 those at the light of the findings of this paper, in order to facilitate the comparison between both
114 alternatives in this paper.

115 *Answer from authors: The previous study was a traditional implementation of SDP on a single-reservoir*
116 *system and shows optimal management while disregarding dynamic groundwater storage and head-*
117 *dependent groundwater pumping costs.*

118 *We have added a brief and clear summary of the previous study in the case study description with focus on*
119 *underlining the differences between the previous and the present study. We have also revised the links to*
120 *the previous study with focus on keeping the explanations short and precise. Together, this facilitates a*
121 *better comparison between the alternatives within the paper.*

122
123 **AR2.3)** Introduction. While being successful in presenting the problem, the Introduction seems a little
124 confusing. At first, one would expect some comments about why is important to jointly manage surface
125 and groundwater prior to enumerating the state-of-the-art on conjunctive use optimization. While the
126 division between deterministic and stochastic programming is adequate, the state-of-the-art presented
127 consists in describing several references rather than explaining briefly both approaches supporting both
128 explanations with references. It is said in the paper that “has been addressed widely in the literature”
129 (which is true) but then only 4 references for deterministic and 2 for stochastic are shown. I would prefer to
130 not explain what has been done in a little number of papers, but to discuss the different approaches
131 employed and then enumerate the references. Besides, the review seems to not have moved prior to the
132 90’s, when the topic appeared in the 60’s and 70’s.

133 *Answer from authors: The introduction has been revised as suggested by the reviewer:*

- 134 • *A line has been added to motivate why it is important to jointly manage surface and groundwater:*
135 *“Optimal allocation of the water resources should address coordinated use of the water resources*
136 *by considering the long term total costs while utilizing the groundwater as a buffer.”*
- 137 • *The literature review has been completely rewritten and now provides a more complete overview of*
138 *the major approaches employed within conjunctive surface water and groundwater management.*
- 139 • *The last paragraph has been added stronger links to the existing methods.*

140
141
142 **AR2.4)** (Case study) p. 5935. It is assumed that the full storage capacity can be managed flexibly without
143 consideration of storage reserved for flood protection or existing management rules. Why ?

144 *Answer from authors: Reservoir rule curves and flood control volumes were not available as such*
145 *information is classified in China. A sentence has been added to clarify this: “While reservoir rule curves and*
146 *flood control volumes can easily be accommodated, the present policies were not available for the case*
147 *area”.*

148
149 **AR2.5)** So how flood protection pools are taken into account? Are you using a realistic useful storage?

150 *Answer from authors: Flood protection is not taken into account in this study. It will, however, be easy to*
151 *implement a volume reserved for flood storage within the proposed framework. This will reduce the*
152 *available storage and increase water scarcity in the long dry season. In the present model setup, we find the*
153 *lower limit on water scarcity costs, assuming that the entire storage capacity is available for storing water.*
154 *Reservoir spills will cause an economic loss, and the model tends to avoid spills by entering the rainy season*
155 *with a low reservoir storage level.*

156 *This has been clarified in the second paragraph of the “2.1 Study area”.*

157
158

159 **AR2.6)** p. 5935. . . . analysis of dynamic interactions between the groundwater and surface water resources.
160 It seems that the box model that you use for groundwater does not account for any dynamic interaction
161 between groundwater and surface water. Is this correct? If that is the case, groundwater discharges
162 (outflow) and stream-aquifer interaction are not considered . . . Please show that it is correct to neglect this
163 groundwater outflow components. Otherwise, we have an incomplete groundwater balance.

164 *Answer from authors: The groundwater model is a simple box model (Infiltration + Storage = Pumping +*
165 *Overflow). The groundwater overflow is only used in extreme cases where the total demands + available*
166 *storage < infiltration. The spills will go to the spill variable and leave the system, practically as baseflow to*
167 *the rivers (unavailable for allocation). The aquifer is so heavily over-exploited that no significant baseflow is*
168 *being created or will be created in any foreseeable future. This has now been clarified in the manuscript.*
169

170 **AR2.7)** A rainfall-runoff model previously used in the paper of the business-as-usual run. It is unclear if you
171 simply took the resulting inflow values of that study or if you update that model. If it is an update, then the
172 calibration results should be presented.

173 *Answer from authors: The exact same hydrological model results were used in both studies. No new*
174 *calibration was performed, and space was therefore not used to repeat details. Note that the hydrological*
175 *model does not represent the actual modified discharge in the rivers today, but is an estimate of the natural*
176 *water availability. We have clarified this in the manuscript.*
177

178 **AR2.8)** In addition, I do not see the point of developing a daily model and then aggregate the results. It
179 would have been easier to directly develop a monthly model.

180 *Answer from authors: We need an estimate of the natural water availability and chose to reuse the*
181 *estimate from our previous peer-reviewed study. In this study, we had access to daily weather data from the*
182 *Chinese Meteorological Services and daily runoff from an almost natural river.*
183

184 **AR2.9)** Besides, it is said that the recharge is estimated upon the precipitation, using the average
185 precipitation value corresponding to the inflow class as characteristic value. That assumes a perfect
186 correlation between precipitation and inflows, which is uncommon. Would have then possible to be
187 included in the Markov chain? . . . although it would suppose an increase in the curse of dimensionality
188 phenomenon . . .

189 *Answer from authors: It would be possible to include another Markov Chain describing the groundwater*
190 *recharge transition probabilities. With 3 flow classes for both runoff and recharge the number of inflow*
191 *scenarios would increase to $3 \times 3 = 9$. However, we do not have any observations of groundwater recharge to*
192 *develop these statistics. In the absence of such data, we decided to assume perfect correlation.*
193

194 **AR2.10)** (2.2. optimization model formulation) There is a variable named “groundwater spill”. Does it refer
195 to “groundwater discharge”. Where does physically go this discharge? Please give an explanation about
196 what means this spill, and how this is modeled.

197 *Answer from authors: The groundwater spill is only used in rare extreme cases where the total demands +*
198 *available storage < infiltration. These spills will go to the spill variable and leave the system, practically as*
199 *base flow to the rivers (unavailable for allocation). As we are discretizing the entire groundwater storage*
200 *(empty to full), we experience this situation occasionally in the backward iteration. The resulting lower*
201 *water values and the large discrete storage intervals will prevent that these spills appear in the forward*
202 *simulation. We have clarified this in the manuscript (see AR2.6).*
203

204 **AR2.11)** (2.4 Solving non-linear and non-convex sub-problems) The non-linearities tackled by your GA-LP
205 algorithm are the decision variables regarding final storages. In an alternative SDP approach, these
206 variables are kept discrete. If you keep them discrete, the problem becomes linear again and there is no
207 need to maintain the timeconsuming GA procedure. In fact, that ability to work out non-linearities is one of
208 the main advantages of Dynamic Programming (DP). Why have you not taken the ending groundwater table
209 $V_{gw,t+1}$ discrete? It would have saved you a huge amount of time, although with less quality in the results,
210 as you point out. I would think it would have been worth it, specially regarding at the steady water values
211 found in the aquifer.

212 *Answer from authors: This was also our initial idea. First problem is the discretization. We would need a*
213 *very fine discretization of the groundwater aquifer to allow discrete storage levels and decisions. If not, the*
214 *discrete volumes of the large aquifer become much larger than the combined monthly demands. Storing all*
215 *recharge will therefore not be sufficient to recharge to a higher discrete storage level. Similarly, the*
216 *demands will be smaller than the discrete volumes, and pumping the remaining water to reach a lower*
217 *discrete level would also be infeasible. For this reason, we decided to allow free end storage. Free end*
218 *storage requires interpolation between the discrete storage levels. With free surface water and*
219 *groundwater end storages, the future cost function has three dimensions (surface water storage,*
220 *groundwater storage and expected future costs). With our head-dependent pumping costs and increasing*
221 *electricity price, we observed that the future cost function changes from strictly convex (very low electricity*
222 *price) to strictly concave (very high electricity price). At realistic electricity prices, we observed a mix of*
223 *concave and convex shape. For the use of Benders' decomposition (require strict convexity), this caused a*
224 *problem. Instead, we developed the hybrid LP-GA model which was applied successfully. This model can deal*
225 *with any electricity price (= any groundwater pumping costs) at any storage level.*
226 *We have focused on communicating this better in the manuscript.*

227
228 **AR2.12** (2.4 Solving non-linear and non-convex sub-problems) A misunderstanding regarding piecewise
229 linear interpolation is found in this section. You said that, according to Pereira and Pinto, piecewise linear
230 interpolation requires strict convexity. However, Pereira and Pinto used a Benders decomposition, which
231 employs piecewise linear approximations and requires convexity, but it is different from the regular
232 procedure, which does not need the cost-to-go function to be convex. You can fit a linear function between
233 your point and the neighboring ones, as you did when interpolating the future costs with cubic functions.
234 Please correct that.

235 *Answer from authors: The previous study used Benders' decomposition and not a piecewise linear*
236 *interpolation. We have updated this paragraph in the manuscript and now explaining why Benders'*
237 *decomposition and a traditional linear interpolation are problematic.*

238
239 **AR2.13** (3 Results) In the first paragraph of page 5946, it can be read that, at the equilibrium groundwater
240 storage level, the willingness to pay is equal to 2.3 CNY m⁻³. In Figure 6 user's price for groundwater is
241 always below that threshold if initial groundwater storage is at equilibrium. If the user's price for
242 groundwater is always below the curtailment cost, why is the model curtailing the wheat agriculture? One
243 would expect that pumping would fluctuate according to surface water availability, but without any
244 curtailment, since it is more profitable to pump. Is there any constraint forcing that curtailment? Please
245 elaborate.

246 *Answer from authors: The 2.3 CNY m⁻³ is a mistake. The downstream wheat user has a curtailment cost at*
247 *2.12 CNY m⁻³ (rounded to 2.1 CNY m³ in table 1). The user's price for groundwater reported in Figure 6 is*
248 *~2.15 CNY m⁻³ (groundwater value at ~2.06 CNY m⁻³ and a pumping cost at 0.09 CNY m⁻³). This exceeds*
249 *the curtailment cost of wheat agriculture (2.12 CNY m⁻³), and this user is therefore curtailed. These values*
250 *have been updated and the conclusion (curtailment of wheat agriculture) underlined.*

251
252 **AR2.14** (3 Results) Why a reservoir storage evolution plot does not appear in the manuscript? It would be
253 important to see the surface and the groundwater storage in order to identify possible conjunctive use
254 patterns. Please include the surface reservoir storage evolution or explain why it is not necessary.

255 *Answer from authors: The reservoir storage plot was not included in an attempt to reduce the length of the*
256 *manuscript. We have now prepared a figure with a comparison of groundwater and surface water storage*
257 *(see Figure 6).*

258
259 **AR2.15** (4 Discussion) In the first paragraph of page 5948, you say that SDDP only samples around the
260 optimal decisions and, consequently, you will not be able to get the complete set of shadow prices for all
261 state combinations. However, the SDDP sampling procedure actually employs samples that are not
262 subjected to a pre-defined grid and, therefore, the samples are not evenly distributed across space,
263 concentrating in the region located near the optimal decisions. The extrapolation process applied in SDDP
264 covers the whole space but with different levels of accuracy depending in which region you look at. The
265 difference between SDP and SDDP regards to the fact that the SDP results have the same accuracy for the

266 whole space, while the SDDP results' accuracy varies across the space, focusing near the optimal decisions
267 while usually decreasing when moving far from them. With SDDP you will get a complete set of shadow
268 prices as well, but with different accuracy levels: some of them better than SDP and some of them worse.
269 Choosing between them does not regard to having or not shadow prices, but to the degree of accuracy that
270 you can accept on them. Please re-elaborate the comparison between SDP and SDDP.

271 *Answer from authors: Thanks for clarifying this. We have revised the SDDP-SDP comparison.*

272
273 **AR2.16)** (3 Results and 4 Discussion) Although a sensitivity analysis was made with regard to the water
274 demands, the curtailment costs and the transmissivity; there are other sources of uncertainty that must be
275 taken into account. Factors like inflow and storage discretization, assumption of perfect correlation
276 between rainfall and in- flow, pumping costs estimation, usage of a lumped model for the aquifer and so
277 on, add a considerable amount of uncertainty to the problem. An explanation about the implications of
278 those sources of uncertainty in the results should be added to the manuscript.

279 *Answer from authors: We have expanded the section on uncertainty and elaborated on the factors that are*
280 *presently not mentioned. We also highlight that "Given the computational challenges and the diverse and*
281 *significant uncertainties, the model results should be seen as a demonstration of the model capabilities*
282 *rather than precise cost estimates. Better estimates will require access to a better case dataset and involve*
283 *a more comprehensive sensitivity analysis."*

284
285 **AR2.17)** (5 Conclusion) As presented, the conclusions would not attract the reader. They seem to appear as
286 part of the discussion rather than a separate section. It should be re-organized in order to clearly highlight
287 what are the novelties of the study and what conclusions can be extracted from the methodology applied
288 and the results obtained in the case study.

289 *Answer from authors: We have reorganized/rewritten the conclusions and put focus on a brief presentation*
290 *of the clear conclusions related to the method and the results.*

291 292 **Detailed comments**

293 **AR2.18)** (page 5934, line 11) One would expect here references about the water value method, not about
294 the SDP one. In addition, Pereira and Pinto (1991) did not used SDP, but SDDP.

295 *Answer from authors: Yes, this is indeed confusing. We have removed Pereira and Pinto (1991) and left the*
296 *reader with Stage and Larsson (1961) (water value method) and Stedinger et al. (1984) (SDP in reservoir*
297 *operation).*

298
299 **AR2.19)** (page 5935) Line 11: upper storage capacity ?. This is storage capacity, what it is represented
300 through a upper bound constraint, but the combination of terms here is unclear. I suggest to remove
301 "upper". Please correct it in all the times this appears in the text.

302 *Answer from authors: Yes, the "upper" has been removed as suggested.*

303 **AR2.20)** (page 5935) Line 24: Why only the upstream users have a pumping limit?

304 *Answer from authors: The river basin has two aquifers (upstream and downstream), which are only*
305 *connected by the river. Ideally, each aquifer should be modelled as a box model, but this extra state variable*
306 *would be computationally challenging within the SDP framework. We therefore set up the box model for the*
307 *downstream and most important aquifer. The upstream aquifer is only bound by an upper pumping limit*
308 *corresponding to the average monthly recharge. This has been clarified.*

309
310 **AR2.21)** (page 5940, line 21) Replace "the thickness of the aquifer" by "groundwater pumping"

311 *Answer from authors: Yes, this has been replaced*

312
313 **AR2.22)** (page 5941, line 1) Is it realistic to assume an even distribution of total pumping across all the
314 wells?

315 *Answer from authors: We have added the following sentence to justify this assumption: "The even pumping*
316 *distribution is a fair assumption, as field investigations showed that 1) the majority of the groundwater*
317 *wells are for irrigation, 2) the timing of irrigation, crop types and climate is homogeneous and 3) the*
318 *groundwater wells have comparable capacities."*

319

320 **AR2.23)** (page 5943, line 18) Replace “program” by “programming”.

321 *Answer from authors: Yes.*

322

323 **AR2.24)** (page 5944, line 24) I think that, besides the larger storage, one important reason beyond the
324 stability shown by the groundwater values is the fact that the interaction between surface water and
325 groundwater is not represented. If some sort of stream-aquifer interaction had been found, the
326 groundwater values would have been affected by surface waters and vice versa.

327 *Answer from authors: Yes, for large permanent rivers this would probably be an important factor. This has
328 now been clarified in the following sentence: “Addition of stream-aquifer interactions to the model is
329 generally expected to affect this stability, but since the flow in rivers/canals in the case study area is small
330 most of the year, and since most areas are far from a river, it is a reasonable assumption to ignore these
331 dynamics.”*

332

333 **AR2.25)** (page 5945, line 1) Rather than decision rules, the water values tables act as pricing policies. In
334 fact, you do that in the Discussion and the Conclusions sections.

335 *Answer from authors: The water value tables are the main drivers behind the release decisions and, if fully
336 implemented in the decision process, should be referred to as decision rules. For consistency, we have now
337 used “pricing policy” throughout the manuscript.*

338

339 **AR2.26)** (page 5947, line 17) You should add “with SDP” after “feasible today”. Other alternatives are able
340 to handle large water resources systems.

341 *Answer from authors: Yes, indeed. This has been added.*

342

343 **AR2.27)** (page 5947, line 24) Has a simulation model with higher spatial resolution been used? If not, please
344 clearly indicate in the results section (page 5945, line 1) that the forward-moving simulation uses the same
345 system scheme.

346 *Answer from authors: No, we have only used a simulation model with the same system scheme. This has
347 now been clarified.*

348

349 **AR2.28)** (page 5949, line 24) I think that the reason beyond the small differences between SDP and DP
350 regard to the inclusion of the aquifer rather than a very good performance of the SDP algorithm (although
351 it is good). If you consider groundwaters in the analysis, their buffer value gives a high robustness to the
352 surface system. This is reflected in the fact that the SDP empties the reservoir almost every year while not
353 doing that if groundwater was not considered: it can always pump so it hedges the reservoir in an
354 aggressive way.

355 *Answer from authors: Yes, we have added this point more clearly in the discussion.*

356

357 **AR2.29)** (page 5950, line 15) The groundwater results are independent in the recharge as well. It should be
358 added to the list.

359 *Answer from authors: Yes, this has now been added here and in the conclusion.*

360

361 **AR2.30)** (page 5951, line 4) I do not understand how the opportunity costs are reduced if electricity prices
362 grow. This would apply exclusively if all the demands could freely pump and all of them had the same
363 pumping head, which is not the case (you have demands that are subjected to pumping quotas while other
364 cannot pump). However, the fact that electricity prices can be used to internalize the groundwater prices is
365 valuable regardless of that.

366 *Answer from authors: This is true. The electricity price statement has been deleted and focus put on the
367 internalization of the groundwater price.*

368

369 **AR2.31)** (page 5951, line 7) Rather than opportunity cost pricing (OCP), the name should be marginal cost
370 pricing (MCP). Please replace this definition here and in the rest of the document.

371 *Answer from authors: Yes, this has been updated throughout the manuscript.*

372

373

374 **AR2.32)** (page 5951, line 10) The title of the section should be “Conclusions”.

375 *Answer from authors: Yes, this has been corrected.*

376

377 **AR2.33)** (page 5951, line 20) The non-convexity is caused by the headdependent pumping costs rather than
378 the inclusion of the groundwater reservoir.

379 *Answer from authors: Yes, this has been clarified in the conclusion: “Non-convexity caused by head and rate*
380 *dependent groundwater pumping costs was accommodated with the use of a GA and was further extended*
381 *to include stationary Thiem local drawdown cones”*

382

383 **AR2.34)** (page 5958, Table 2) This table has not been cited in the text. Remove it or cite it.

384 *Answer from authors: An error has happened in the layout version. The reference is wrongly listed as “Table*
385 *1” on page 5945 in line 15 and 26. We will make sure that the table references are corrected in the final*
386 *version.*

387

388 **AR2.35)** (page 5963, Figure 4) In the surface water values part of the Figure, V_{gw} must be 50% rather than
389 80%.

390 *Answer from authors: We have plotted for 80% (SW) and 50% (GW) to better represent the changes. The*
391 *surface water values are changing mostly at higher storage levels, while the groundwater values are not*
392 *depending on the SW values. The figure caption wrongly states 50% - this has been corrected.*

393

394 **AR2.36)** (page 5965, Figure 6) Do you mean Davidsen et al (2015) rather than Davidsen et al (2014)? If not,
395 please add Davidsen et al (2014) to the reference list.

396 *Answer from authors: Yes, Davidsen et al (2015) is the correct citation. The paper was only published online*
397 *(2014) when this manuscript was submitted. The reference has been corrected throughout the manuscript.*

398

399

.....

400

Other minor changes

401

- The abstract have been updated to match the revised conclusion.

402

- Eq. (3) has been updated to match the style of the other equations (location of t).

403

- The units of P in Eq. (11)-(13) have been updated to avoid confusion (before P used both J/m^3 and kWh/m^3 as unit).

404

- The explanation and unit of the hydraulic conductivity has been updated.

405

- The result and discussion sections have been reorganized so that no results are presented in the discussion.

406

- Reference added in the caption to Table 1.

407

- In Figure 5, $V_{max,gw}$ have been updated to match the used nomenclature.

408

- Grammar changes throughout the manuscript.

409

410

1 The cost of ending groundwater overdraft on the North 2 China Plain

3
4 C. Davidsen^{1,2,3,*}, S. Liu^{2,*}, X. Mo², D. Rosbjerg¹ and P. Bauer-Gottwein^{1,*}

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13 Abstract

14 Over-exploitation of groundwater reserves is a major environmental problem around the
15 world. In many river basins, groundwater and surface water are used conjunctively and joint
16 optimization strategies are required. A hydroeconomic modelling approach is used to find
17 cost-optimal sustainable surface water and groundwater allocation strategies for a river basin,
18 given an arbitrary initial groundwater level in the aquifer. A simplified management problem
19 with conjunctive use of scarce surface water and groundwater under inflow and recharge
20 uncertainty is presented. Because of head-dependent groundwater pumping costs the
21 optimization problem is non-linear and non-convex, and a genetic algorithm is used to solve
22 the 1-step-ahead sub-problems with the objective of minimizing the sum of immediate and
23 expected future costs. A real-world application in the water-scarce Ziya River Basin in
24 northern China is used to demonstrate the model capabilities. Persistent overdraft from the
25 groundwater aquifers on the North China Plain has caused declining groundwater ~~tables~~levels,
26 ~~salinization and infiltration of wastewater~~. The model maps the opportunity marginal cost of
27 water in different scenarios, and the minimum cost of ending groundwater overdraft in the
28 basin is estimated to be 5.5847 billion CNY/year. The study shows that it is cost-effective to
29 slowly recover the groundwater aquifer to a level close to the surface while gradually
30 lowering the groundwater value to the equilibrium at 2.15 CNY/m³. The model can -be used

31 to guide decision makers to economic efficient ~~ensure~~ long-term sustainability of
32 groundwater and surface water resources management ~~in the basin in an economically optimal~~
33 ~~way~~.

34 **1 Introduction**

35 Groundwater aquifers are of high economic importance around the world and often act as
36 buffers in the water supply system during droughts (Tsur and Graham-Tomasi, 1991; Tsur,
37 1990). On the North China Plain, persistent groundwater overexploitation over the past
38 decades has caused decline of the shallow and deep groundwater tables (Liu et al., 2001). The
39 immediate benefits of satisfying the water demands greatly exceed the costs of pumping,
40 which highlights the problem of the present self-regulating management. As the groundwater
41 resource is overexploited, the immediate benefits of the increased unsustainable supply have
42 to be traded off against the long term increase in pumping costs and reduced buffering
43 capacity. Optimal allocation of the water resources should address coordinated use of the
44 water resources by considering the long term total costs while utilizing the groundwater as a
45 buffer. This is in line with the 2011 Chinese No. 1 Policy Document, which targets
46 improvement of the water use efficiency and reduction of water scarcity (CPC Central
47 Committee and State Council, 2010).

48 ~~Optimal management of conjunctive use of surface water and groundwater has been~~
49 ~~addressed widely in the literature. Harou and Lund (2008) used a deterministic~~
50 ~~hydroeconomic optimization approach to examine the economic effects of ending long term~~
51 ~~groundwater overdraft in California. The linear model was run under different scenarios and~~
52 ~~used to estimate the water users' willingness to pay, water scarcity costs and the benefits of~~
53 ~~conjunctive use facilities. Andreu et al. (1996) developed the deterministic AQUATOOL~~
54 ~~simulation software based on the Out-of-Kilter Algorithm to minimize deficits in demand and~~
55 ~~minimum flows in a coupled surface water-groundwater environment. This model was later~~
56 ~~applied in a hydroeconomic context by Pulido-Velázquez et al. (2006) to minimize the sum of~~
57 ~~scarcity costs and variable operating costs for a coupled setup with a distributed parameter~~
58 ~~groundwater simulation. The integrated aquifer model allowed variable pumping costs in a~~
59 ~~forward moving, scenario-based framework but lacked the ability to give predictions in an~~
60 ~~uncertain real-time management environment. An alternative optimization approach was~~
61 ~~demonstrated by Riegels et al. (2013), who maximized welfare subject to ecosystem~~
62 ~~constraints by adjusting time-constant water prices.~~

63 ~~While a high level of complexity can be accommodated in deterministic simulation models,~~
64 ~~the objective functions of stochastic optimization models are kept simpler to remain~~
65 ~~computationally feasible. Philbrick and Kitanidis (1998) applied Stochastic Dynamic~~
66 ~~Programming (SDP) to a multi-reservoir system to optimize conjunctive use of surface water~~
67 ~~and groundwater given stochastic inflow. The second order gradient dynamic programming~~
68 ~~method, a modification of the classical recursive SDP, was used to mitigate the well-known~~
69 ~~*curse of dimensionality*. Pumping costs were linked linearly to pumping rates but changes in~~
70 ~~pumping costs due to long term depletion of the aquifer were not considered. Head-dependent~~
71 ~~pumping costs were included in the SDP model by Knapp and Olson (1995), who analyzed~~
72 ~~conjunctive use of groundwater with randomly generated runoff. Non-linearity arising from~~
73 ~~the head-dependent pumping costs was overcome with lattice programming techniques in this~~
74 ~~qualitative model setup.~~Optimal management of conjunctive use of surface water and
75 groundwater has been addressed widely in the literature (e.g. Booker et al., 2012; Burt, 1964;
76 Knapp and Olson, 1995; Labadie, 2004; Noel and Howitt, 1982). While control-based
77 methods, such as Model Predictive Control (MPC, e.g. Morari and Lee, 1999; Mayne et al.,
78 2000) and Reinforcement Learning (RL, Lee and Labadie 2007), focus on deriving real-time
79 optimal control policies, this study will focus on planning oriented optimization techniques.
80 Deterministic optimization problems for a given time horizon allow a detailed representation
81 of the groundwater system using spatially distributed groundwater models (Andreu et al.,
82 1996; Harou and Lund, 2008; Marques et al., 2006; Pulido-Velázquez et al., 2006).
83 Stochasticity is commonly represented in scenarios where a regression analysis is used to
84 formulate operation rules, see e.g. the Implicit Stochastic Optimization (ISO) approaches
85 reviewed by Labadie (2004). Singh (2014) reviewed the use of simulation-optimization (SO)
86 modeling for conjunctive groundwater and surface water use. In SO-based studies, efficient
87 groundwater simulation models are used to answer “what if”-questions while an optimization
88 model is wrapped around the simulation model to find “what is best”. Groundwater aquifers
89 have been represented as simple deterministic box or “bathtub” models (e.g. Cai et al., 2001;
90 Riegels et al., 2013) and as spatially distributed models (e.g. Maddock, 1972; Siegfried et al.,
91 2009) with stochasticity (Reichard, 1995; Siegfried and Kinzelbach, 2006). While the results
92 obtained from these methods are rich in detail, they yield only a single solution to the
93 optimization problem.

94 Dynamic Programming (DP, Bellman 1957) based methods have been used extensively
95 to demonstrate the dynamics of conjunctive groundwater – surface water use for both

96 deterministic (e.g. Buras, 1963; Provencher and Burt, 1994; Yang et al., 2008) and stochastic
97 (SDP, e.g. Burt, 1964; Philbrick and Kitanidis, 1998; Provencher and Burt, 1994; Tsur and
98 Graham-Tomasi, 1991) optimization problems. In DP-based methods, the original
99 optimization problem is decomposed into subproblems which are solved sequentially over
100 time. The entire decision space is thereby mapped, enabling use of the results as dynamic
101 decision rules. However, the number of subproblems grows exponentially with the number of
102 state variables and this *curse of dimensionality* has frequently limited the use of DP and SDP
103 (Labadie, 2004; Provencher and Burt, 1994; Saad and Turgeon, 1988). Although it causes loss
104 of detail and inability to disaggregate the results, reservoir aggregation has been suggested as
105 one solution strategy (Saad and Turgeon, 1988).

106

107 This study aims to answer the following two macro-scale decision support questions for
108 conjunctive groundwater and surface water management for the Ziya River Basin in North
109 China: 1) *what are the minimum costs of ending groundwater overdraft?* and 2) *what is the*
110 *cost-efficient recovery strategy of the over-pumped aquifer?* ~~Ademonstrates how a~~
111 hydroeconomic modeling approach ~~can be~~ used to identify the least-cost strategy to achieve
112 sustainable groundwater abstraction, ~~defined as the. In this context, “sustainable” means that~~
113 ~~the~~ long term average abstraction ~~does not exceeding~~ the long term average recharge. To
114 ~~overcome. At the water management problem with conjunctive use of surface water and~~
115 ~~groundwater~~ similar to Harou and Lund (2008) ~~with is addressed. I~~ increased complexity ~~is~~
116 caused by uncertain surface water runoff and groundwater recharge, ~~and non-linearity arising~~
117 ~~from head and rate dependent groundwater pumping costs. the surface water reservoirs are~~
118 ~~aggregated. This is adequate at macro-scale (Davidsen et al., 2015) and allow use of dynamic~~
119 ~~programming based approaches.~~ The cost minimization problem is solved with the water
120 value method, a variant of SDP (Stage and Larsson, 1961; Stedinger et al., 1984) which
121 ~~produces dynamic tables of marginal costs linked to states, stages and water source. Head and~~
122 ~~rate dependent pumping costs introduce non-linearity in the. The non-linear~~ discrete sub
123 problems. ~~This nonlinearity is are solved/handled~~ with a ~~combined-hybrid G~~ genetic
124 ~~A~~ algorithm (GA) -and ~~L~~ linear ~~P~~ programming- (LP) method similar to that used by Cai et al.
125 (2001), ~~but here~~ applied ~~to in~~ a coupled groundwater-surface water management problem
126 ~~with~~in an SDP framework.

127 2 Methods

128 2.1 Study area

129 Northern China and particularly the North China Plain (NCP) have experienced increasing
130 water scarcity problems over the past 50 years due to population growth, economic
131 development and reduced precipitation (Liu and Xia, 2004). The deficit in the water balance
132 has historically been covered by overexploitation of the groundwater aquifer, causing a
133 regional lowering of the groundwater table by up to 1 m/year (Zheng et al., 2010).

134 ~~The case study area is the Ziya River Basin, a part of the Hai River Basin, which is located~~
135 ~~in the Hebei Province on the NCP and with the upper catchment stretching through the~~
136 ~~Taihang Mountains into the Shanxi Province~~ was selected as case study area (see Figure 1).
137 ~~The upper basin is located in the Shanxi Province, while the lower basin is located in the~~
138 ~~Hebei Province on the NCP. The 52,300 km² basin has approximately 25 million inhabitants~~
139 ~~(data from 2007, Bright et al., 2008), and is subject to severe water scarcity is causing, which~~
140 ~~causes multiple conflicts. The 52,300 km² basin, shown in , is home to approximately 25~~
141 ~~million people. A hydroeconomic study of the Ziya River by focused primarily on optimal~~
142 ~~management of the surface water resources.~~ Five major reservoirs with a combined storage
143 capacity of 3.5 km³ are located in the basin. While reservoir rule curves and flood control
144 volumes can easily be accommodated, policies applied in practical management today were
145 not accessible for the case area. Instead it is assumed that the full storage capacity can be
146 managed flexibly without consideration of storage reserved for flood protection or existing
147 management rules. Incorporating flood storage volumes will reduce the available storage and
148 increase water scarcity in the long dry season. In the present model setup, we therefore find
149 the lower limit on water scarcity costs, assuming that the entire storage capacity is available
150 for storing water. Reservoir spills will cause an economic loss, and the model tends to avoid
151 spills by entering the rainy season with a low reservoir storage level.

152 ~~In this study it is assumed that the full storage capacity can be managed flexibly without~~
153 ~~consideration of storage reserved for flood protection or existing management rules.~~

154 A previous hydroeconomic study of the Ziya River Basin was a traditional implementation of
155 SDP on a single-reservoir system (surface water reservoir) and showed optimal water
156 management, while disregarding dynamic groundwater storage and head-dependent
157 groundwater pumping costs (Davidsen et al., 2015). Instead, tIn Davidsen et al., 2014, the

158 groundwater resource was included as a simple monthly upper allocation constraint, ~~which~~
159 ~~prevents analysis of dynamic interactions between the groundwater and surface water~~
160 ~~resources and limits the decision space.~~

161 In the present ~~model setup~~ study, the groundwater resource is included as a ~~simple~~ dynamic
162 aquifer box model with a ~~n upper~~ storage capacity of 275 km³. The river basin has two
163 aquifers (upstream and downstream) which are only connected by the river. Ideally, each
164 aquifer should be modelled as a box model, but this extra state variable would be
165 computationally challenging within the SDP framework. We therefore set up a box model for
166 the downstream and most important aquifer only and abstraction from the upstream aquifer is
167 only bounded by an upper pumping limit corresponding to the average monthly recharge. The
168 box model for the downstream aquifer is formulated as *Infiltration + Storage = Pumping +*
169 *Overflow*. The groundwater overflow is only used in extreme cases, where the total pumping
170 and available storage is less than the infiltration. The spills will go to the spill variable and
171 leave the system, as baseflow to the rivers (unavailable for allocation). The aquifer is so
172 heavily over-exploited that no significant baseflow is being created or will be created in any
173 foreseeable future. This box model allows for more flexible management with larger
174 abstractions in dry years and increased recharge in wet years. The groundwater aquifer can
175 thereby be used to bridge longer drought periods. Except from the groundwater box model,
176 the conceptual model is identical to the one used by (Davidsen et al., (2015)).

177 A conceptual sketch of the management problem is shown in Figure 2 ~~Figure 2~~. The water
178 users are divided into groups of economic activities; irrigation agriculture, industrial and
179 domestic water users. Ideally, each water user group should be characterized by flexible
180 demand curves, but due to poor data availability ~~Each water user group is characterized by a~~
181 ~~constant water demands (m³) and a constant curtailment costs of not meeting the demand~~
182 were used for each group (see Table 1 ~~Table 1~~), ~~as also applied by~~. The water demands are
183 assumed to be deterministic and decoupled from the stochastic runoff. This is a reasonable
184 assumption because the rainfall on the NCP normally occurs in the summer months, while
185 irrigation water demands are concentrated in the dry spring. The irrigation schedule is
186 centrally planned and typically unchanged from year to the same every year. The upstream (u)
187 users have access to runoff and are restricted to an upper pumping limit X_{gw} corresponding to
188 the average monthly upstream recharge, while ~~The water users upstream the surface water~~
189 ~~reservoir (u) have access to the runoff and a monthly limited volume of groundwater, . T~~ the

190 | ~~water users located~~ downstream ~~users the reservoir~~-(d) have access to reservoir releases, water
 191 | delivered through the South-to-North Water Transfer Project (SNWTP) and groundwater
 192 | from the ~~dynamic downstream~~ aquifer.

193 2.2 Optimization model formulation

194 An SDP formulation is used to find the expected value of storing an incremental amount of
 195 surface water or groundwater, given the month of the year, the available storage in surface
 196 and groundwater reservoirs and the inflow scenarios. The backward recursive equation
 197 calculates the sum of immediate and expected future costs for all combinations of discrete
 198 reservoir storage levels (states) and monthly time steps (stages). The immediate management
 199 costs (IC) arise from water supply and water curtailment, whereas the expected future costs
 200 (EFC) are the optimal value function in $t+1$ weighed by the corresponding transition
 201 probabilities. In the present setup, we decided to weigh the IC and EFC equally, but inclusion
 202 of discount rates other than zero is possible. Because of the head and rate dependent
 203 groundwater pumping costs, which will be described in detail later, the immediate cost
 204 depends non-linearly on the decision variables. The objective is to minimize the total costs
 205 over the planning period, given by the optimal value function $F_t^*(V_{gw,t}, V_{sw,t}, Q_{sw,t}^k)$ based on
 206 the classical Bellman formulation:

$$207 \quad F_t^*(V_{gw,t}, V_{sw,t}, Q_{sw,t}^k) = \min \left(IC(V_{gw,t}, V_{sw,t}, Q_{sw,t}^k) + \sum_{l=1}^L (p_{kl} F_{t+1}^*(V_{gw,t+1}, V_{sw,t+1}, Q_{sw,t+1}^l)) \right) \quad (1)$$

208 with IC being the immediate costs:

$$209 \quad IC(V_{gw,t}, V_{sw,t}, Q_{sw,t}^k) = \sum_{m=1}^M (c_{sw} x_{sw} + c_{gw} x_{gw} + c_{SNWTP} x_{SNWTP} + c_{ct} x_{ct})_{m,t} - r_{sw,t} b_{hp} \quad (2)$$

210 subject to:

$$211 \quad x_{sw,m,t} + x_{gw,m,t} + x_{SNWTP,m,t} + x_{ct,m,t} = dm_{m,t} \quad (3)$$

$$212 \quad V_{sw,t} + Q_{sw,t} - \sum_{u=1}^U x_{sw,u,t} - r_{sw,t} - s_{gw,t} = V_{gw,t+1} \quad (4)$$

$$213 \quad r_{sw,t} + s_{sw,t} = \sum_{d=1}^D x_{sw,d,t} + q_{E,t} \quad (5)$$

$$214 \quad V_{gw,t} + Q_{gw,t} - \sum_{d=1}^D x_{gw,d,t} - s_{gw,t} = V_{gw,t+1} \quad (6)$$

$$215 \quad \sum_{u=1}^U x_{sw,u,t} \leq Q_{sw,t} \quad (7)$$

$$216 \quad \sum_{u=1}^U x_{gw,u,t} \leq X_{gw,t} \quad (8)$$

$$217 \quad r_t \leq R, \quad x_{sw,Bei} + x_{SNWTP,Bei} \leq Q_{SNWTP}, \quad q_{E,t} \geq Q_E, \quad V_{sw,t} \leq V_{\max,sw}, \quad V_{gw,t} \leq V_{\max,gw} \quad (9)$$

$$218 \quad c_{gw} = f\left(V_{gw}, \sum_{d=1}^D x_{sw,d}\right) \quad (10)$$

219 [See Table 2 wherefor nomenclature.](#)

220 Eq. (3) is the water demand fulfillment constraint, i.e. the sum of water allocation and water
 221 curtailments equals the water demand of each user. Eq. (4) is the water balance of the
 222 combined surface water reservoir, while Eq. (5) is the water balance of the reservoir releases.
 223 A similar water balance for the dynamic groundwater aquifer follows in Eq. (6). The upstream
 224 surface water allocations are constrained ~~to~~ [by](#) the upstream runoff [as shown in](#) (Eq. (7)), while
 225 the upstream groundwater allocations are constrained to a fixed sustainable monthly average
 226 [as shown in](#) (Eq. (8)). In Eq. (9), the upper and lower hard constraints on the decision variables
 227 are shown. Last, Eq. (10) is the marginal groundwater pumping cost, which depends on the
 228 combined downstream groundwater allocations as described later.

229

230 A rainfall-runoff model based on the Budyko Framework (Budyko, 1958; Zhang et al., 2008)
 231 ~~and has in a~~ previous [study been applied by](#) (Davidsen et al., 2015) is used to estimate the
 232 near-natural daily surface water runoff into reservoirs (Davidsen et al., 2015). The [resulting](#)
 233 51 years (1958-2008) of simulated daily runoff ~~are was~~ aggregated to monthly runoff and
 234 normalized. A Markov chain, which describes the runoff serial correlation between three flow
 235 classes defined as dry (0 – 20th percentile), normal (20th – 80th percentile), and wet (80th –
 236 100th percentile), ~~is was~~ established and validated ~~to~~ ensure second order stationarity
 237 (Davidsen et al., 2015; Loucks and van Beek, 2005). The groundwater recharge is estimated
 238 from the precipitation data also used in the rainfall-runoff model. The average monthly
 239 precipitation (mm/month) for each runoff class is calculated, and a simple groundwater
 240 recharge coefficient of 17.5% of the precipitation (Wang et al., 2008) is used.

241 The SDP loop is initiated with EFC set to zero and will propagate backward in time through
 242 all the discrete system states as described in the objective function. For each discrete

243 combination of states, a cost minimization sub-problem will be solved. A sub-problem will
 244 have the discrete reservoir storage levels ($V_{gw,t}$ and $V_{sw,t}$) as initial conditions and reservoir
 245 inflow is given by the present inflow class in the Markov chain. The optimization algorithm
 246 will search for the optimal solution, given the costs of the immediate management (water
 247 allocations and water curtailments, including reservoir releases and groundwater pumping),
 248 which have to be balanced against the expected future costs. As the SDP algorithm is
 249 propagating backward in time, the future costs will be equal to the minimum total costs from
 250 $t+1$, weighted by the Markov chain transition probabilities. The algorithm will continue
 251 backward in time until equilibrium is reached, i.e. until the shadow prices (marginal value of
 252 storing water for future use) in two successive years remain constant. The SDP model is
 253 developed in MATLAB (MathWorks Inc., 2013) and uses the fast *cplexlp* (IBM, 2013) to
 254 solve the linear sub-problems.

255 The sets of equilibrium shadow prices, referred to as the water value tables, can subsequently
 256 be used to guide optimal water resources management forward in time with unknown future
 257 runoff. In this study, the available historic runoff time series ~~are~~ is used to demonstrate how
 258 the derived water value tables should be used in real time operation. The simulation will be
 259 initiated from different initial groundwater aquifer storage levels, thereby demonstrating
 260 which pricing policy should be used to bring the NCP back into a sustainable state.

261 2.3 Dynamic groundwater aquifer

262 The groundwater aquifer is represented as a simple box model (see ~~Figure 2~~ Figure 2) with
 263 recharge and groundwater pumping determining the change in the stored volume of the
 264 aquifer (Eq. (6)). The pumping is associated with a pumping cost determined by the energy
 265 needed to lift the water from the groundwater table to the land surface (Eq. (10)):

$$266 \quad P = (\rho g \Delta h) / \varepsilon \quad (11)$$

267 where P is the specific pump energy (J/m^3), ρ is the density of water (kg/m^3), g is the
 268 gravitational acceleration (m/s^2), Δh is the head difference between groundwater table and
 269 land surface (m) and ε is the pump efficiency (-). The marginal pumping cost c_{gw} (CNY/m^3)
 270 is found from the average electricity price c_{el} (CNY/Wh) in Northern China:

$$271 \quad c_{gw} = c_{el} P \quad (12)$$

272 Hence this cost will vary with the stored volume in the groundwater aquifer. The present
 273 electricity price structure in China is quite complex, with the users typically paying between
 274 0.4 and 1 CNY/kWh depending on power source, province and consumer type (Li, 2012; Yu,
 275 2011). In this study a fixed electricity price of 1 CNY/kWh is used. The immediate costs of
 276 supplying groundwater to a single user follow:

$$277 \quad c_{gw,t} x_{gw,t} = \rho g \Delta h \varepsilon^{-1} c_{el} x_{gw,t} \quad (13)$$

278 where Δh is found as the mean depth from the land surface to the groundwater table (see
 279 [Figure 2](#)) between t and $t+1$:

$$280 \quad \Delta h = \Delta h_{top} + \left(V_{max,gw} - \frac{V_{gw,t} + V_{gw,t+1}}{2} \right) S_y^{-1} A^{-1} \quad (14)$$

281 where Δh_{top} is the distance from the land surface to the top of the aquifer at full storage (m),
 282 S_y is the specific yield (-) of the aquifer, and A is the area of the aquifer (m²). Here $V_{gw,t+1}$ is
 283 a decision variable, and once substituted into [Eq. \(13\)](#), it is clear that the problem becomes
 284 non-linear.

285 In [Eq. \(14\)](#) the drawdown is assumed uniform over the entire aquifer. This simplification
 286 might be problematic as the local cone of depression around each well could contribute
 287 significantly to the pumping cost and thereby the optimal policy. Therefore, the steady state
 288 Thiem drawdown (Thiem, 1906) solution is used to estimate local drawdown at the pumping
 289 wells. Local drawdown is then added to [Eq. \(15\)](#) to estimate total required lift:

$$290 \quad \Delta h_{Thiem} = \frac{Q_w}{2\pi T} \ln \left(\frac{r_{in}}{r_w} \right) \quad (15)$$

291 where Q_w is the pumping rate of each well (m³/month), T is the transmissivity (m²/month),
 292 r_{in} is the radius of influence (m), and r_w is the distance from origin to the point of interest
 293 (m), here the radius of the well. The transmissivity is based on a hydraulic conductivity of
 294 $1.3 \cdot 10^{-6}$ m²/month for silty loam (Qin et al., 2013). ~~The hydraulic conductivity which is~~
 295 ~~lower than the expected average for the NCP to provide a conservative estimate of the effect~~
 296 ~~of drawdown. was tested to be realistic in a MIKE SHE model of the ZRB (Marker, 2013).~~
 297 Field interviews revealed that the wells typically reach no deeper than 200 m below surface,
 298 which results in a specific yield of 5%. The ~~thickness of the aquifer~~ groundwater pumping Q_w

299 is defined as the total allocated groundwater within the stage (m^3/month) and, assumed evenly
 300 distributed evenly to the number of wells in the catchment:

$$301 \quad Q_{w,t} = \frac{\sum_{d=1}^D x_{gw,d,t}}{n_w} = \frac{V_{gw,t} - V_{gw,t+1} + Q_{gw,t} - S_{gw,t}}{n_w} \quad (16)$$

302 where n_w is the number of wells in the downstream basin. The even pumping distribution is a
 303 fair assumption, as field investigations showed that 1) the majority of the groundwater wells
 304 are for irrigation, 2) the timing of irrigation, crop types and climate is homogeneous and 3)
 305 the groundwater wells have comparable capacities. –Erlendsson (2014) estimates^d the well
 306 density in the Ziya River Basin from Google Earth to be 16 wells/ km^2 . Assuming that the
 307 wells are distributed evenly on a regular grid and that the radius of influence r_{in} is 500 m,
 308 overlapping cones of depression from 8 surrounding wells are included in the calculation of
 309 the local drawdown. This additional drawdown is included using the principle of
 310 superposition as also applied by Erlendsson (2014).

311

312 **2.4 Solving non-linear and non-convex sub-problems**

313 With two reservoir state variables and a climate state variable, the number of discrete states is
 314 quickly limited by the *curse of dimensionality*. A very fine discretization of the groundwater
 315 aquifer to allow discrete storage levels and decisions is computationally infeasible. A low
 316 number of discrete states increases the discretization error, particularly if both the initial and
 317 the end storages $V_{gw,t+1}$ and $V_{sw,t+1}$ are kept discrete. The discrete volumes of the large aquifer
 318 become much larger than the combined monthly demands, and storing all recharge will
 319 therefore not be sufficient to recharge to a higher discrete storage level. Similarly, the
 320 demands will be smaller than the discrete volumes, and pumping the remaining water to reach
 321 a lower discrete level would also be infeasible. Allowing free end storage in each subproblem
 322 will allow the model to pick e.g. the optimal groundwater recharge and pumping without a
 323 requirement of meeting an exact discrete end state. With free surface water and groundwater
 324 end storages, the future cost function has three dimensions (surface water storage,
 325 groundwater storage and expected future costs). Pereira and Pinto (1991) used Benders'
 326 decomposition approach, which employs piecewise linear approximations and requires
 327 convexity. With head and rate dependent pumping costs and increasing electricity price, we

328 observed that the future cost function changes from strictly convex (very low electricity price)
329 to strictly concave (very high electricity price). At realistic electricity prices, we observed a
330 mix of concave and convex shapes. An alternative is to use linear interpolation with defined
331 upper and lower bounds. However, with two state variables, interpolation between the future
332 cost points will yield a hyperplane in three dimensions, which complicates establishment of
333 boundary conditions for each plane.

334 ~~In the previous study by (Davidsen et al., 2015), the optimization problem was strictly linear~~
335 ~~and strictly convex. The individual sub-problems of the SDP scheme could therefore be~~
336 ~~solved with a fast linear programming algorithm. In this study however, with non-linearity~~
337 ~~from the head-dependent groundwater pumping costs, the expected future cost function is no~~
338 ~~longer strictly convex.~~

339 Non-linear optimization problems can be solved with evolutionary search methods, a sub
340 division of global optimizers. A widely used group of evolutionary search methods ~~is-are~~
341 genetic algorithms (genetic algorithms-(GAs)), which ~~have-arebeen~~ found to be efficient tools
342 ~~to-for~~ getting the-approximate solutions to complex non-linear optimization problems (see,
343 e.g., Goldberg, 1989; Reeves, 1997). GAs use a random search approach inspired by natural
344 evolution and have been applied to the field of water resources management by, e.g., Cai et al.
345 (2001), McKinney and Lin (1994) and Nicklow et al. (2010). Cai et al. (2001) used a
346 combined genetic algorithmGA and linear programming-(LP) approach to solve a highly non-
347 linear surface water management problem. By fixing some of the complicating decision
348 variables, the remaining objective function became linear and thereby solvable with LP. The
349 GA was used to test combinations of the fixed parameters while looking for the optimal
350 solution. The combination yielded faster computation time than if the GA was used to
351 estimate all the parameters.

352
353 ~~A~~This study uses a genetic algorithmGA implemented in MATLAB is used to solve the cost
354 minimization sub-problems. This GA function will initially generate a set of candidate
355 solutions known as the *population*. Each of the candidate solutions contains a set of decision
356 variables (sampled within the decision space), which will yield a feasible solution to the
357 optimization problem. In MATLAB, a set of options specifies: the *population size*, the
358 stopping criteria (*fitness limit, stall limit, function tolerance* and others), the *crossover*
359 *fraction*, the *elite count* (number of top parents to be guaranteed survival) and the *generation*

360 *function* (how the initial population is generated). The options were adjusted to achieve
361 maximum efficiency of the GA for the present optimization problem.

362

363 The computation time for one single sub-problem is orders of magnitude larger than solving a
364 simple LP. As the optimization problem ~~became~~*comes* computationally heavier with
365 increasing number of decision variables, a hybrid version of GA and LP, similar to the
366 method used by Cai et al. (2001), ~~is~~*was* developed (see ~~Figure 3~~*Figure 3*). Decision variables
367 that cause non-linearity are identified and chosen by the GA. Once these complicating
368 decision variables are chosen, the remaining objective function becomes linear and thereby
369 solvable with LP. In the ~~present~~*presented* (in Eq. (1)), the non-linearity
370 is caused by the head-dependent pumping costs as explained in Eq. (13)-(14). Both the
371 regional lowering of the groundwater table and the Thiem local drawdown cones depend on
372 the decision variable for the stored volume in $t+1$, $V_{gw,t+1}$. If $V_{gw,t+1}$ is pre-selected, the
373 regional drawdown is given, and the resulting groundwater pumping rate Q_w can be
374 calculated from the water balance. The groundwater pumping price is thereby also given, and
375 the remaining optimization problem becomes linear.

376 ~~The SDP framework is subject to the curse of dimensionality. With two state variables and~~
377 ~~non-linearity, the computation time is significant and is a limiting factor when choosing~~
378 ~~model discretization. With a low number of discrete states, the discretization error increases,~~
379 ~~particularly if the end storages $V_{gw,t+1}$ and $V_{sw,t+1}$ are kept discrete. Piecewise linear~~
380 ~~interpolation of the future cost function (Pereira and Pinto, 1991) allows for free end storages~~
381 ~~but requires strict convexity. With two state variables, interpolation between the future cost~~
382 ~~points will yield a hyper-plane in three dimensions. In our problem, the EFC is no longer~~
383 ~~strictly convex and therefore both $V_{gw,t+1}$ and $V_{sw,t+1}$ are chosen by the GA.~~

384 For a given combination of stages, discrete states and flow classes, the objective of the GA is
385 to minimize the total costs, TC , with the free states $V_{gw,t+1}$, $V_{sw,t+1}$ being the decisions:

$$386 \quad TC(V_{gw,t+1}, V_{sw,t+1}) = \min IC(V_{gw,t+1}, V_{sw,t+1}) + EFC(V_{gw,t+1}, V_{sw,t+1}) \quad (17)$$

387 with EFC being the expected future costs. Given initial states and once the GA has chosen
388 the end states, the immediate cost minimization problem becomes linear and hence solvable
389 with LP (see ~~Figure 3~~*Figure 3*). ~~The IBM CPLEX linear programming solver is used to solve~~

390 | ~~the linear programs~~. The expected future costs are found by cubic interpolation of the discrete
391 neighboring future cost grid points in each dimension of the matrix. The GA approaches the
392 global optimum until a fitness limit criteria is met. The total costs are stored, and the
393 algorithm continues to the next state. To reduce the computation time, the outer loop through
394 the groundwater states is parallelized.

395 The performance of the GA-SDP model is compared to a fully deterministic ~~Dynamic~~
396 ~~Program (DP)DP~~, which finds the optimal solution given perfect knowledge about future
397 inflows and groundwater recharge. The DP model uses the same algorithm as the SDP model
398 and 1-dimensional state transition matrices with $p=1$ between the deterministic monthly
399 runoff data. For low storage capacity and long time scales, the effect of the end storage
400 volume becomes negligible. Similar to the SDP model, the DP model was looped and run
401 until the end of period condition does not affect the present management.

402 **3 Results**

403 Without any regulation or consideration of the expected future costs arising from over-
404 exploitation of the groundwater aquifer, the water users will continue maximizing immediate
405 profits (producers) or utility (consumers). Because there are only electricity costs for
406 groundwater, the users will continue pumping groundwater until the marginal groundwater
407 cost exceeds the curtailment costs. At 1 CNY/kWh the marginal cost of lifting groundwater
408 | 200 m (typical depth of wells observed in the study area) can be found with Eq. (13)-(14) to
409 be 0.8 CNY/m³ and thereby less than the lowest curtailment cost at 2.3 CNY/m³. It requires
410 an electricity price higher than 2.8 CNY/kWh before the lowest-value user stops pumping
411 from 200 m below surface.

412 The backward recursive SDP algorithm was run with a looped annual dataset until
413 equilibrium water values, i.e. no inter-annual changes, were obtained. The water values
414 increase fastest during the first years, and after approximately 100 years the annual increases
415 become small. Due to the large storage capacity of the groundwater aquifer, equilibrium is
416 however not achieved until after 150-180 years. These marginal water values represent the
417 true values of storing a unit volume of water for later use, and vary with reservoir storage
418 levels, runoff flow class and time of the year. A sample of the resulting equilibrium water
419 value tables are presented in Figure 4. This figure shows the temporal variations of water
420 values as a function of one state variable, keeping the other state variable at a fixed value. The
421 state variables are fixed at empty, half full and full storage respectively. During the rainy

422 season from June to August, high precipitation rates reduce water scarcity, resulting in lower
423 ~~the~~ surface water values. Because the groundwater storage capacity is much larger, increased
424 recharge can easily be stored for later use, and groundwater values are therefore not affected.
425 Addition of stream-aquifer interactions to the model is expected to affect this behavior, but
426 since the flow in rivers/canals in the case study area is small most of the year, and since most
427 areas are far from a river, it is a reasonable assumption to ignore these dynamics. -The water
428 values after 1980 are clearly higher than in the period before 1980 due to increased water
429 scarcity caused by a reduction in the regional precipitation. In contrast, the groundwater value
430 tables are uniform, with variation only with groundwater storage. The detailed water value
431 tables are included as supplementary information.

432 We simulate management using the equilibrium water value tables as ~~decision rules~~pricing
433 policy and force the system with 51 years of simulated historical runoff. Time series of the
434 simulated groundwater storage levels can be seen in Figure 5 for different initial storage
435 scenarios for the dynamic groundwater aquifer. The groundwater aquifer approaches an
436 equilibrium storage level around 260 km³ (95% full). If the storage in the aquifer is below this
437 level, the average recharge will exceed average pumping until the equilibrium storage is
438 reached. If the storage level is above equilibrium, average pumping will exceed average
439 recharge ~~and over time until~~ equilibrium ~~storage~~ is reached. In Figure 6, the surface water and
440 groundwater storages are shown for a situation with equilibrium groundwater storage. In most
441 years, the surface water storage falls below 1 km³, leaving space in the reservoir for the rainy
442 season. The potential high scarcity costs of facing a dry scenario with an almost empty
443 reservoir is avoided by pumping more groundwater. These additional pumping costs seem to
444 be exceeded by the benefits of minimizing spills in the rainy season.

445 ~~The surface water reservoir storage level varies over time, and in contrast to the findings by~~
446 ~~(Davidsen et al., 2015) the storage capacity now becomes close to zero almost every year.~~
447 ~~This can be explained by the increased groundwater availability in the model, which allows~~
448 ~~increased groundwater allocation in multi-annual dry periods. To To~~ demonstrate the
449 business-as-usual solution, the simulation model is run for a 20 year period with the present
450 water demands and curtailment costs and with a discount rate set to infinity (= zero future
451 costs). The resulting groundwater table is continuously decreasing as shown in Figure 5
452 5.

453 In the simulated management runs, water will be allocated to the users up to a point where
454 reductions in immediate cost are compensated by increases in expected future costs. The
455 user's price, which can be applied in an opportunity-marginal cost pricing (MCP) scheme, is
456 the marginal value of the last unit of water allocated to the users. The user's price is the sum
457 of the actual pumping cost (electricity used) and the additional opportunity-marginal cost
458 given by the equilibrium water value tables. In Figure 7Figure-6, the user's prices for
459 groundwater and surface water are shown for the 51 year simulation at and below the long
460 term sustainable groundwater storage level. When the groundwater storage level is close to
461 equilibrium, the user's prices of groundwater and surface water are equal during periods with
462 water scarcity. In wet months with reduced water scarcity, the model switches to surface
463 water allocation only, and the groundwater user's price is undefined (gaps in the time series in
464 Figure 7Figure-6). If the groundwater storage level is below equilibrium, the groundwater
465 user's price will be higher causing an increase in water curtailments and increasing storage
466 level as shown in Figure 5Figure-5. Under these circumstances the surface water user's price
467 increases up to a point where the two prices meet. With an initial aquifer storage at one third
468 of the aquifer capacity (100 km³), the groundwater value is 3 CNY/m³ (see Figure 7). As the
469 aquifer slowly recovers, the groundwater price decreases gradually.

470 At the equilibrium groundwater storage level, the user's prices for groundwater is stable
471 around 2.15 CNY/m³ as shown in Figure 7Figure-6. This indicates frequent-curtailment of
472 wheat agriculture in the downstream Hebei Province, which has a willingness to pay of 2.312
473 CNY/m³ (see Table 1). The allocation pattern to this user is shown in Figure 8Figure-7: the
474 model switches between high curtailment and high allocations, depending on water
475 availability and storage in the reservoirs. Groundwater allocations fluctuate between
476 satisfying 0% and 80% of the demand. Inclusion of the steady state Thiem drawdown cones in
477 the optimization model increases the marginal groundwater pumping cost with increased
478 pumping rates. Groundwater allocations are distributed more evenly over the months, which
479 results in less local drawdown. The total curtailments remain constant, while 1% of the total
480 water abstraction is shifted from groundwater to surface water, if the stationary Thiem
481 drawdown is included. Inclusion of well drawdown significantly changed the simulated
482 management but resulted in only slightly increased computation time.

483 The average total costs of the 51 years simulation for different scenarios can be seen in Table
484 3. The average reduction in the total costs, associated with the introduction of the SNWTP

485 canal can be used to estimate the expected marginal economic impact of the SNWTP water.
486 The minimum total costs after the SNWTP is put in operation are compared to the scenario
487 without the SNWTP (pre-2008) and divided by the allocated SNWTP water. The resulting
488 marginal value of the SNWTP water delivered from Shijiazhuang to Beijing (2008-2014
489 scenario) is 3.2 CNY/m³, while the SNWTP water from Yangtze River (post-2014 scenario)
490 reduces the total costs with 4.9 CNY/m³. Similarly, a comparison of the total costs for the
491 post-2014 scenarios shows a marginal increase of 0.91 CNY/m³ as a consequence of
492 introducing a minimum in-stream flow constraint.

493 ~~A simple local sensitivity analysis is used to assess the uncertainty of the model. used Monte~~
494 ~~Carlo simulations based on 50 samples to estimate the uncertainty of the model outputs.~~
495 ~~However, the inclusion of an additional state variable has increased the optimization time~~
496 ~~significantly and made such an approach infeasible. Approximately 4000 CPU hours per~~
497 ~~climate period are needed to reach equilibrium in the present model, equivalent to two weeks,~~
498 ~~if the maximum of 12 parallel processors are used in MATLAB R2013a. TheA local~~
499 ~~sensitivity analysis was focused on the local sensitivity related to the water demands and~~
500 ~~water curtailment costs used directly in the objective function (Eq. (1)) and the~~
501 ~~transmissivity used to estimate the local drawdown (Eq. (14)). The uncertain input parameters~~
502 ~~were increased by 10% and the sensitivity evaluated based on the simulation results. The~~
503 ~~resulting total costs can be seen in Table 3. A 10% increase in the curtailment costs is~~
504 ~~returned as a 6.0% increase in the total costs, while a similar increase of the demands~~
505 ~~generates a 2.1% increase in costs. The transmissivity can vary over many orders of~~
506 ~~magnitude because it is a log-normally distributed variable. The sensitivity of $\log(T)$ is high:~~
507 ~~a 1.3% change of $\log(T)$ from the baseline value results in a 1.5% change in the cost. The~~
508 ~~benchmark DP run was run for the post-2014 scenario with Thiem drawdown and minimum~~
509 ~~ecosystem flow constraint. The minimum total costs of this run is $8.46 \cdot 10^9$ CNY/year. This is~~
510 ~~1.3% lower than the equivalent SDP run ($8.56 \cdot 10^9$ CNY/year).~~

511 The minimum total costs were lowered from 10.50 billion CNY/year (Davidsen et al., 2015)
512 to 8.56 billion CNY/year (18% reduction) by allowing the groundwater aquifer to be utilized
513 as a buffer instead of a fixed monthly volume. This difference highlights the problem of
514 defining realistic boundaries to optimization problems and shows that simple hard constraints,
515 here fixed groundwater pumping limits, can highly limit the optimal decision space. With a
516 dynamic groundwater aquifer, the model can mitigate dry periods and stabilize the user's

517 price of surface water as shown in Figure 7. Finally, policies like minimum in-stream
518 ecosystem flow constraints can be satisfied with less impact on the expensive users. The total
519 costs without restrictions on the groundwater pumping have been estimated to 2.98 billion
520 CNY/year (Davidsen et al., 2015). To end the groundwater overdraft in the basin, the present
521 study thus estimates a cost increase of 5.58 billion CNY/year, once the groundwater aquifer is
522 at equilibrium storage. The cost of recharging the aquifer from the present storage level below
523 the equilibrium is significantly higher. In Table 3, the LGW scenario shows that the average
524 cost of sustainable management from an initial storage at 100 km³ (one third full) is 13.32
525 billion CNY/year.

526 From any initial groundwater reservoir storage level, the model brings the groundwater table
527 to an equilibrium storage level at approximately 95% of the aquifer storage capacity. Only
528 small variations in the aquifer storage level are observed after the storage level reaches
529 equilibrium as shown in Figure 6. While addition of the Thiem stationary drawdown has only
530 a small effect on total costs and total allocated water, it is clear from Figure 8 that the
531 additional Thiem drawdown highly impacts the allocation pattern to some of the water users.
532 High groundwater pumping rates result in larger local drawdown and thus in higher pumping
533 costs. This mechanism leads to a more uniform groundwater pumping strategy, which is
534 clearly seen in Figure 8 and results in much more realistic management policy.

535

536

537 **4 Discussion**

538 This study presents a hydroeconomic optimization approach that provides macro-scale
539 economic ~~pricing policy~~decision rules in terms of water values for joint-conjunctive surface
540 water – groundwater management. The method was used to demonstrate how the water
541 resources in the Ziya River Basin should be priced over time, to reach a sustainable situation
542 at minimum cost. We believe that the presented modelling framework has great potential use
543 as a robust decision support tool in real-time water management. However, a number of
544 limitations and simplifications need to be discussed.

545 ~~A great advantage of SDP-based water value method is the capability to obtain optimal~~
546 ~~decision rules for any combination of system states.~~ A first limitation of the approach is the
547 high level of simplification needed. There are two main reasons for the high level of

548 simplification: Limited data availability and the limitations of the SDP method. SDP-
549 approach is, however, the curse of dimensionality as previously mentioned. The curse of
550 dimensionality number of sub-problems to be solved in the backward moving SDP scheme
551 increases exponentially with the number of state variables. In our case we are therefore
552 limiteds the approach to 2-3 inter-linked storage facilities and higher dimensional
553 management problems will not be computationally feasible with SDP today. This limit on the
554 number of surface water reservoirs and groundwater aquifers requires a strongly simplified
555 representation of the real world situation in the optimization model. ~~These requirements can~~
556 ~~be relaxed in t~~The simulation phase ~~that follows~~ing the optimization. ~~While the curse of~~
557 ~~dimensionality applies to the backward moving SDP scheme, the forward moving simulation~~
558 is not limited to the same extent, as since just only one a single sub-problem is solved at each
559 stage. The water values determined by the SDP scheme can thus be used to simulate
560 management using a much more spatially resolved model with a high number of users; this
561 was not demonstrated in this study. The advantage of SDP is that it provides a complete set of
562 decision rules pricing policies that can be applied in adaptive management, provided that the
563 system can be simplified to a computationally feasible level. An alternative approach known
564 as stochastic dual dynamic programming (SDDP, Pereira and Pinto, 1991; Pereira et al.,
565 1998) has shown great potential for multi-reservoir river basin water management problems.
566 Instead of sampling the entire decision space with the same accuracy level, SDDP samples
567 with a variable accuracy not pre-defined in a grid, focusing the highest accuracy around the
568 optimal solution. This variable accuracy makes SDDP less suitable However, because SDDP
569 ~~only samples around the optimal decisions, this method will not be able to provide the~~
570 ~~complete set of shadow prices for all state combinations and is therefore less suitable~~ for
571 adaptive management. Despite the highly simplified system representation, we believe that
572 the modeling framework provides interesting and non-trivial insights, which are extremely
573 valuable for water resources management on the NCP.

574 Computation time was a a major limitation limitation in this study. Three factors increased the
575 computational load of the optimization model: 1) inclusion of the groundwater state variable
576 resulted in an exponential growth of the number of subproblems; 2) the non-convexity
577 handled by the slower GA-LP formulation caused an increase in the computation time of 10-
578 100 times a single LP; and 3) the SDP algorithm needed to iterate through more than 200
579 years to reach steady-state. A single scenario run required 4,000 CPU hours and was solved in
580 two weeks using 12 cores at the high performance cluster (HPC) at the Technical University

581 of Denmark. This is 50,000 times more CPU hours than a single reservoir SDP model
582 (Davidsen et al., 2015). Since the water value tables can be used offline in the decision
583 making, this long computation time can be accepted.

584 The long computation time made the use of, e.g., Monte Carlo-based uncertainty analysis
585 infeasible. The local sensitivity analysis showed that a 10% increase in the curtailment costs
586 is returned as a 6.0% increase in the total costs, while a similar increase of the demands
587 generates a 2.1% increase in costs. The transmissivity can vary over many orders of
588 magnitude because it is a log-normally distributed variable. The sensitivity of $\log(T)$ is high:
589 a 1.3% change of $\log(T)$ from the baseline value results in a 1.5% change in the cost. At the
590 same time, the simple system representation needed in SDP required assumptions of inflow
591 and storage discretization, aggregation of the surface water reservoirs, generalized estimates
592 of pumping cost and a lumped groundwater model which all contribute to the uncertainty.
593 Further, poor data availability for the case study area required some rough estimates of the
594 natural water availability, single-point demand curves and perfect correlation between rainfall
595 and groundwater recharge. The method-driven assumptions generally limit the decision
596 support to basin-scale, while the simple estimates caused by poor data availability contribute
597 to raising the general uncertainty of the model results. Given the computational challenges
598 and the diverse and significant uncertainties, the model results should be seen as a
599 demonstration of the model capabilities rather than precise cost estimates. Better estimates
600 will require access to a more comprehensive case dataset and involve a complete sensitivity
601 analysis.

602 ~~and the transition from the previous much simpler linear single state SDP model (Davidsen et~~
603 ~~al., 2015) to the presented non-linear SDP model with two state variables proved to require~~
604 ~~around 50,000 times more CPU hours. We used the high performance cluster (HPC) at the~~
605 ~~Technical University of Denmark to solve the SDP, and as the optimization can be run~~
606 ~~offline, an optimization time of 2 weeks on 12 cores can be accepted.~~

607 ~~Replacement of the hard upper groundwater pumping constraint used by (Davidsen et al.,~~
608 ~~2015) with a dynamic groundwater aquifer, lowered the total costs from 11.39 billion~~
609 ~~CNY/year to 8.47 billion CNY/year. This difference highlights the problem of defining~~
610 ~~realistic boundaries to optimization problems and shows that simple hard constraints, here~~
611 ~~fixed groundwater pumping limits, can highly limit the optimal decision space. With inclusion~~
612 ~~of a dynamic groundwater aquifer, the model can use the large groundwater storage capacity~~

613 as a buffer to the system, which significantly stabilizes the user's price of surface water as
614 shown in Figure 6. Finally, policies like minimum ecosystem flow constraints can be satisfied
615 with less impact on the expensive users, which results in reductions in the respective shadow
616 prices.

617 Another addition to the modelling framework was the Thiem stationary drawdown. The long
618 time steps (monthly) make stationarity a realistic assumption. Inclusion of well drawdown
619 significantly changed the simulated management but resulted in only slightly increased
620 computation time. While addition of the Thiem stationary drawdown has only a small effect
621 on total costs and total allocated water, it is clear from Figure 7 that the additional Thiem
622 drawdown highly impacts the allocation pattern to some of the water users. High groundwater
623 pumping rates result in larger local drawdown and thus in higher pumping costs. This
624 mechanism leads to a more uniform groundwater pumping strategy, which is clearly seen in
625 Figure 7 and results in much more realistic management policy.

626 From any initial groundwater reservoir storage level, the sustainable management brings the
627 groundwater table to an equilibrium storage level at approximately 95% of the aquifer storage
628 capacity. Only small variations in the aquifer storage level are observed after the storage level
629 reaches equilibrium. Intuitively, one would expect the equilibrium groundwater storage level
630 to be as close as possible to full capacity, while still ensuring that any incoming groundwater
631 recharge can be stored. Finding the exact equilibrium groundwater storage level would
632 require a very fine storage discretization, which, given the size of the groundwater storage, is
633 computationally infeasible. Therefore the equilibrium groundwater storage level is subject to
634 significant discretization errors. The long time steps (monthly) make the stationarity required
635 for using the Thiem stationary drawdown method a realistic assumption.

636 In the previous study by (Davidsen et al., 2015), total costs, without restrictions on the
637 groundwater pumping, were estimated to 3.09 billion CNY/year. The present study estimates
638 an increase to 8.56 billion CNY/year for a comparable setup but with sustainable groundwater
639 pumping and groundwater storage at equilibrium. This increase of 5.47 billion CNY/year
640 reflects the expected cost of ending the groundwater overdraft in the basin once the
641 groundwater aquifer is at equilibrium storage. The cost of recharging the aquifer from the
642 present storage level below the equilibrium is significantly higher. In Table , the LGW
643 scenario show that the average cost of sustainable management from an initial storage at 100
644 km³ (one third full) is 13.32 billion CNY/year.

645 The difference between total costs with SDP and with DP (perfect foresight) is surprisingly
646 small (1.3%). ~~While this difference indicates a very good performance of SDP, the model~~
647 ~~setup also simulates small economic consequences of wrong decisions. With the SNWTP in~~
648 ~~operation (post 2014) the most expensive user~~ Apart from Beijing, which has ~~(Beijing) will~~
649 ~~always have access to enough the SNWTP water, and~~ the remaining downstream users have
650 unlimited access to groundwater. The large downstream groundwater aquifer serves as a
651 buffer to the system and eliminates the economic consequences of a wrong decision. ~~If too~~
652 ~~much water is allocated to a user in month 1, the same user will simply receive a bit less water~~
653 ~~in a following time step. The model almost empties the reservoir every year as shown in~~
654 Figure 6. ~~And~~ wrong decisions are ~~therefore~~ not punished with curtailment of expensive
655 users as observed by ~~(~~Davidson et al., (2015). The groundwater aquifer reduces the effect of
656 wrong decisions by allowing the model to minimize spills from the reservoir without
657 significant economic impact of facing a dry period with an empty reservoir. A dynamic
658 groundwater aquifer thereby makes the decision support more robust, since it is the timing
659 and not the amount of curtailment being affected ~~but will shift allocations in time and~~
660 ~~between the users with curtailment costs close to the long term equilibrium water price (in~~
661 ~~this study the farmers). The inclusion of a dynamic groundwater aquifer thereby makes the~~
662 ~~model self regulate, as periods with too strict policy will be compensated by periods with a~~
663 ~~more unrestrained policy. The robustness is also supported by the simple local sensitivity~~
664 ~~analysis. The impact of changes in the input parameters on the total costs is small, as it is~~
665 ~~mainly the timing and not the amount of curtailment being affected. Inclusion of the large~~
666 ~~groundwater aquifer reduces the effect of wrong timing, which is reflected in small~~
667 ~~differences between the total costs with and without perfect foresight.~~

668 The derived equilibrium groundwater value tables in Figure 4 ~~Figure-4~~ (and the supplementary
669 detailed water value tables); show that that the groundwater values vary with groundwater
670 storage alone and are independent of time of the year, the inflow and recharge scenario and
671 the storage in the surface water reservoir. This finding is important for future work, as a
672 substitution of the groundwater values with a simpler cost function could greatly reduce the
673 number of states and thereby the computation time. The equilibrium groundwater price, i.e.
674 the groundwater values around the long term equilibrium groundwater storage, can possibly
675 be estimated from the total renewable water and the water demands ahead of the optimization,
676 but further work is required to test this. Further work should also address the effect of
677 discounting of the future costs on the equilibrium water value tables and the long term steady

678 state groundwater table. In the present model setup, the large groundwater aquifer storage
679 capacity forces the backward moving SDP algorithm to run through 200-250 model years,
680 until the water values converge to the long term equilibrium. Another great improvement,
681 ~~given the availability of the required data if data allow~~, would be to replace the constant water
682 demands with elastic demand curves in the highly flexible GA-LP setup.

683 A significant impact of including groundwater as a dynamic aquifer is the more stable user's
684 prices shown in ~~Figure 7~~Figure-6. The user's price of groundwater consists of two parts: the
685 immediate groundwater pumping costs (electricity costs) and the expected future costs
686 represented by the groundwater value for the last allocated unit of water. As the model is run
687 to equilibrium, the user's prices converge towards the long term equilibrium at approximately
688 2.2 CNY/m³. ~~This long term equilibrium is not affected by the actual electricity price, as~~
689 ~~increasing electricity prices will be offset by a similar reduction in the opportunity costs.~~
690 ~~The~~A constant electricity price can ~~therefore~~ be used as a policy tool to internalize the user's
691 prices of groundwater shown in ~~Figure 7~~Figure-6. Stable water user's prices will ease the
692 implementation of e.g. ~~an opportunity cost pricing (OCP) MCP~~ scheme, which is one of the
693 available policy options to enforce long-term sustainability of groundwater management.

694

695 **5 Conclusions**

696 This study ~~presented~~~~describes development and application of how~~ a hydroeconomic
697 ~~optimization~~ approach ~~to optimally manage conjunctive use of groundwater and surface~~
698 ~~water. The model determines the water allocation, reservoir operation and groundwater~~
699 ~~pumping that minimizes the long-term sum of head and rate dependent groundwater pumping~~
700 ~~costs and water curtailment costs. The model is used to quantify potential savings of joint~~
701 ~~water management of the Ziya River Basin in Northern China, but the model can be applied~~
702 ~~to other basins as well. Estimates of natural runoff, groundwater recharge, water demands and~~
703 ~~marginal user curtailment costs are cast into a SDP-based optimization framework. Regional~~
704 ~~and Thiem stationary drawdown is used to estimate rate and head dependent marginal~~
705 ~~groundwater pumping costs. The resulting optimization subproblems become nonlinear and~~
706 ~~non-convex and are solved with a hybrid GA-LP setup. A central outcome from the SDP~~
707 ~~framework is tables of shadow prices of surface and groundwater for any combination of~~
708 ~~time, inflow class and reservoir storage. These tables represent a complete set of pricing~~
709 ~~policies for any combination of system states and can be used to guide real-time water~~

710 management. Despite a significant computational demand to extract the water value tables,
711 the method provides a suitable approach for basin-scale decision support for conjunctive
712 groundwater and surface water management.

713 The model provides useful insight to basin-scale scarcity-driven tradeoffs. The model outputs
714 time series of optimal reservoir storage, groundwater pumping, water allocation and the
715 marginal economic value of the water resources at each time step. The model is used to derive
716 a pricing policy to bring the overexploited groundwater aquifer back to a long-term
717 sustainable state. The economic efficient recovery policy is found by trading off the
718 immediate costs of water scarcity with the long term additional costs of a large groundwater
719 head. From an initial storage at one third of the aquifer capacity, the average costs of ending
720 groundwater overdraft are estimated to be 13.32 billion CNY/year. The long-term cost-
721 effective reservoir policy is to slowly recover the groundwater aquifer to a level close to the
722 surface by gradually lowering the groundwater value from an initial level of 3 CNY/m³. Once
723 at this sustainable state, the groundwater values are almost constant at 2.15 CNY/m³ which
724 suggests that wheat agriculture should generally be curtailed under periods with water
725 scarcity. The dynamic groundwater aquifer serves as a buffer to the system and is used to
726 bridge the water resources to multiple years. The average annual total costs are reduced with
727 18% to 8.56 billion CNY compared to a simpler formulation with fixed monthly pumping
728 limits. The stable user's prices are suitable to guide a policy scheme based on water prices
729 and the method has great potential as basin-scale decision support tool in the context of the
730 China No. 1 Policy Document. can be used to derive a pricing policy to bring an
731 overexploited groundwater aquifer back to a long-term sustainable state. The model quantifies
732 potential savings of joint water management of a complex river basin in China. Surface water
733 and groundwater management was optimized in a SDP framework based on a coupled GA-LP
734 setup. The derived equilibrium water value tables represent the shadow prices of surface and
735 groundwater for any combination of time, inflow class and reservoir storage. The
736 groundwater values at equilibrium were found to be almost constant at 2.2 CNY/m³,
737 independent of the time of the year, the surface water storage and the inflow class. Non-
738 convexity caused by the groundwater reservoir could be accommodated with the use of a GA
739 and was further extended to include stationary Thiem local drawdown cones. Inclusion of a
740 dynamic groundwater aquifer greatly reduced the total costs of water scarcity, compared to a
741 setup with fixed monthly pumping limits. The sustainable management will recharge the
742 aquifer until the equilibrium storage level is reached. From an initial storage at one third of

743 ~~the aquifer capacity, the average costs of ending groundwater overdraft are estimated to be~~
744 ~~13.32 billion CNY/year. After equilibrium is reached, the average costs are estimated to be~~
745 ~~5.47 billion CNY/year. The aquifer serves as buffer and allows for overexploitation in dry~~
746 ~~years and this mechanism stabilizes the user's water prices. These stable user's prices are~~
747 ~~suitable for use in an OCP scheme. While the representation of the management problem~~
748 ~~must be kept simple in the optimization model, the OCP prices can be used to drive a much~~
749 ~~more detailed simulation model, which includes a detailed physical representation of the~~
750 ~~system.~~

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942

943 | Table 1: Annual water demands and curtailment costs for the users in the Ziya River Basin.
 944 | *Based on the dataset from Davidsen et al. (2015).*

	Upstream	Downstream	
Water demands (10⁶ m³/month/year)			
<i>Industries</i>	539	543	^a
<i>Domestic</i>	223	864	^b
<i>Maize</i>	569	1,522	^c
<i>Wheat</i>	-	6,089	^c
<i>Beijing</i>	-	1,000	^d
<i>Ecosystems</i>	-	100	^e
<i>Total</i>	1,331	10,119	
Curtailment costs (CNY/m³)			
<i>Industries</i>	5.3	5.3	^f
<i>Domestic</i>	3.2	3.2	^f
<i>Maize</i>	1.8	2.8	^g
<i>Wheat</i>	-	2.1	^g
<i>Beijing</i>	-	5.5	^h

945 | ^a*Demands scaled with area, (Berkoff, 2003; Moiwo et al., 2010; World Bank, 2001)*

946 | ^b*Based on daily water demand (National Bureau of Statistics of China, 2011) scaled with the 2007 population from*
 947 | *Landsat (Bright et al., 2008)*

948 | ^c*Based on the land cover (USGS, 2013) and irrigation practices collected in the field. The wheat irrigation demand*
 949 | *is evenly distributed in March, April, May and June. Maize is irrigated in July.*

950 | ^d*Based on plan by The People's Government of Hebei Province (2012), (Ivanova, 2011)*

951 | ^e*Estimated deficit in the Baiyangdian Lake (Honge, 2006)*

952 | ^f*Estimate by World Bank (2001)*

953 | ^g*Based on the water use efficiency (Deng et al., 2006) and producers' prices (USDA Foreign Agricultural Service,*
 954 | *2012)*

955 | ^h*Estimate by Berkoff (2003)*

956 |

957 |

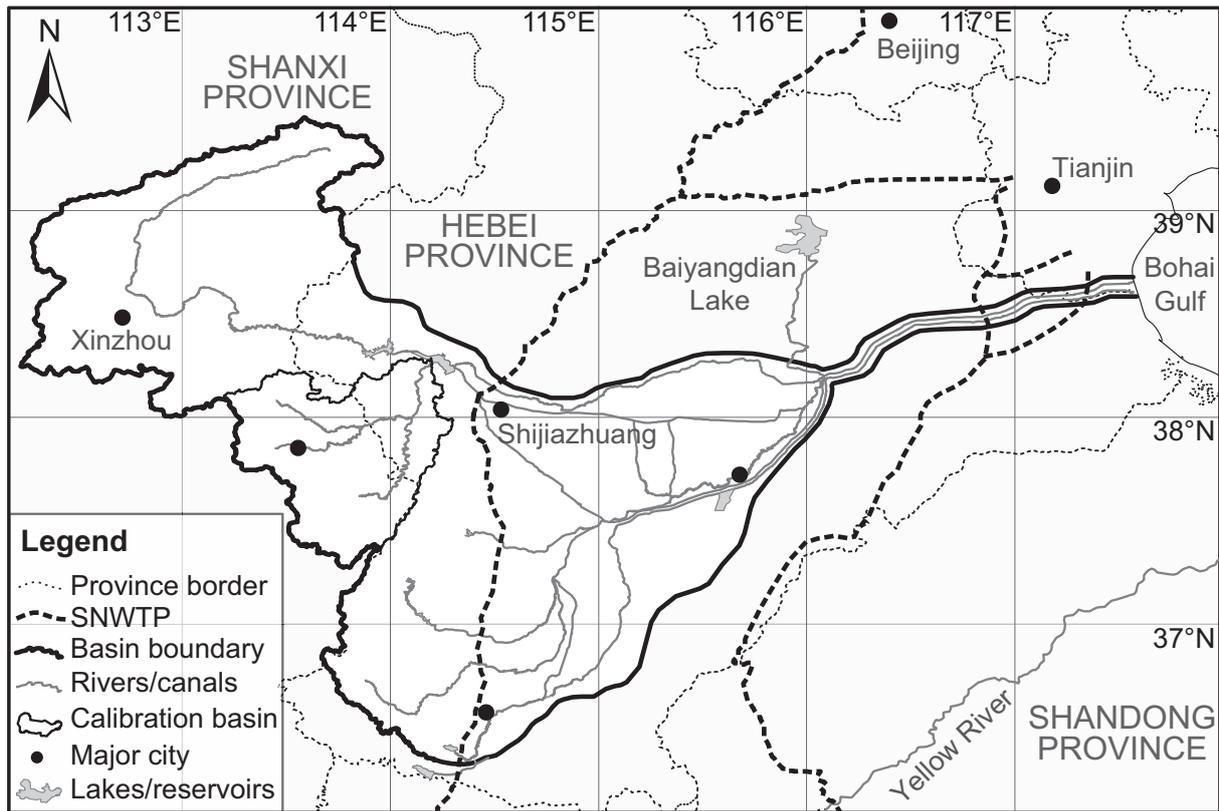
958 Table 2: Nomenclature.

959	<u>F_t^*</u>	<u>optimal value function in stage t (2005 Chinese Yuan, CNY)</u>
960	<u>$V_{gw,t}$</u>	<u>stored volume in the groundwater aquifer, decision variable (m^3)</u>
961	<u>$V_{sw,t}$</u>	<u>stored volume in the surface water reservoir, decision variable (m^3)</u>
962	<u>$V_{max,sw}$</u>	<u>upper storage capacity, surface water reservoir (m^3)</u>
963	<u>$V_{max,gw}$</u>	<u>upper storage capacity, groundwater aquifer (m^3)</u>
964	<u>$Q_{sw,t}$</u>	<u>river runoff upstream reservoirs, stochastic variable ($m^3/month$)</u>
965	<u>$Q_{gw,t}$</u>	<u>groundwater recharge, assumed to be perfectly correlated with $Q_{sw,t}$ ($m^3/month$)</u>
966	<u>m</u>	<u>indicates the M water users</u>
967	<u>gw</u>	<u>groundwater</u>
968	<u>sw</u>	<u>surface water</u>
969	<u>ct</u>	<u>water curtailments</u>
970	<u>x</u>	<u>allocated volume, decision variable ($m^3/month$)</u>
971	<u>c</u>	<u>marginal costs (CNY/m^3). The costs are all constants, except for c_{gw} which is</u>
972		<u>correlated to the specific pump energy. See Eq. (11)-(16)</u>
973	<u>r_t</u>	<u>reservoir releases through hydropower turbines, decision variable ($m^3/month$)</u>
974	<u>R</u>	<u>upper surface water reservoir turbine capacity ($m^3/month$)</u>
975	<u>s_{sw}</u>	<u>reservoir releases exceeding R, decision variable ($m^3/month$)</u>
976	<u>b_{hp}</u>	<u>marginal hydropower benefits (CNY/m^3)</u>
977	<u>k</u>	<u>indexes the K inflow classes in stage t</u>
978	<u>l</u>	<u>indexes the L inflow classes in $t+1$</u>
979	<u>p_{kl}</u>	<u>transition probability from k to l</u>
980	<u>dm_m</u>	<u>water demand for user m ($m^3/month$)</u>
981	<u>u</u>	<u>indexes the U upstream users</u>
982	<u>d</u>	<u>indexes the D downstream users</u>
983	<u>s_{gw}</u>	<u>spills from aquifer when $V_{gw,t} + Q_{sw,t} - x_{gw,t} > V_{max,sw}$ ($m^3/month$)</u>
984	<u>X_{gw}</u>	<u>maximum monthly groundwater pumping in the upstream basin ($m^3/month$)</u>
985	<u>$q_{E,t}$</u>	<u>unused surface water available to ecosystems, decision variable ($m^3/month$)</u>
986	<u>Q_E</u>	<u>minimum in-stream ecosystem flow constraint ($m^3/month$)</u>
987	<u>Bei</u>	<u>Beijing user</u>
988	<u>Q_{SNWTP}</u>	<u>maximum capacity of the SNWTP canal ($m^3/month$)</u>

990 Table 3: Average minimum total costs (TC) and hydropower benefits (HP) over the 51 year
 991 planning period, ~~hydropower benefits and shadow prices~~ for different scenario runs. SNWTP
 992 scenarios: P_{pre 2008} = before the SNWTP canal, 2008 - 2014 = SNWTP partly finished
 993 (emergency plan canal from Shijiazhuang to Beijing), P_{post 2014} = SNWTP finished (water
 994 canal from Yangtze River to Beijing.); Scenarios: -LGW is are results from a run with initial
 995 groundwater storage at 100 km³ below equilibrium (100 km³; all other scenarios are initiated
 996 at equilibrium groundwater storage); - (dm) is are the results with 10 % higher water demands;
 997 -(ct) is are with 10% higher curtailment costs ; and (T) is with 10% higher transmissivity;
 998 TD = is Thiem steady state drawdown included; -E = is minimum ecosystem flow constraint;
 999 “ + ” is active and “ - “ is inactive. (to Baiyangdian Lake); TC = minimum total costs over the
 1000 planning period (51 years tested), b_{hp} = marginal hydropower benefits, DP = dynamic
 1001 programming (perfect foresight), SP = shadow price.

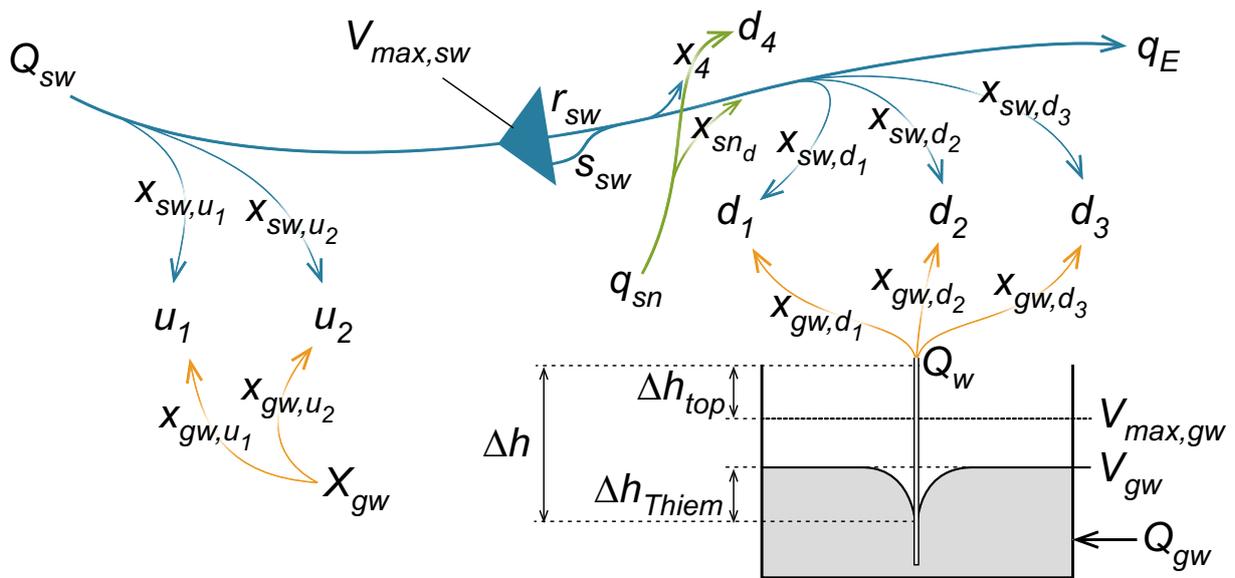
SNWTP scenario	Scenario settings			TC	b_{hp} HP
	Special run	TD	E	SDP 10 ⁹ CNY/y	SDP 10 ⁶ CNY/y
pre-2008	=	+	+	14.87	103.6
2008-2014	=	+	+	11.69	103.5
post-2014	=	=	=	8.43	103.5
post-2014	=	+	=	8.47	103.6
post-2014	=	+	+	8.56	104.3
post-2014	LGW	+	+	13.32	99.2
post-2014	T	+	+	8.69	103.5
post-2014	dm	+	+	8.74	103.3
post-2014	ct	+	+	9.08	103.1

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 1006 Figure 1: The Ziya River Basin. Watershed and rivers automatically delineated from a digital
 1007 elevation map (USGS, 2004) and manually verified and corrected with Google Earth (Google
 1008 Inc., 2013). The SNWTP route-s (Central and Eastern) were sketched in Google Earth and
 1009 verified with field observations. Provincial boundaries from (NGCC, 2009).

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1012 Figure 2: Conceptual sketch of the Ziya River Basin management problem with water users
 1013 located upstream (u) and downstream (d) the surface water reservoir. Allocation decision
 1014 variables are indicated for surface water (blue), SNWTP water (green) and groundwater
 1015 (orange). A conceptual sketch of the downstream dynamic aquifer is included and show how
 1016 the total lift (Δh) is composed of the top layer + the regional groundwater lowering + the local
 1017 Thiem steady state groundwater drawdown.

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Load data

for all stages

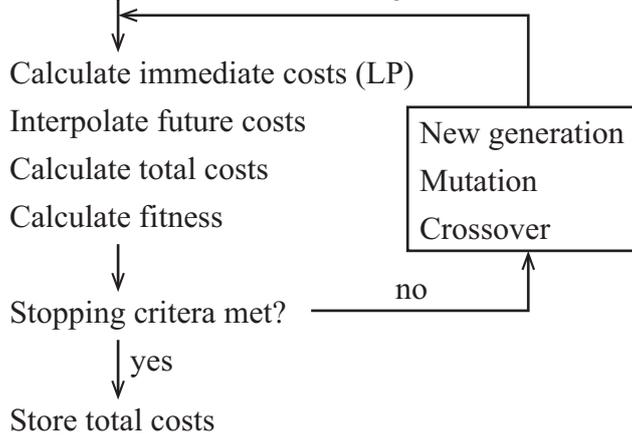
for all runoff flow classes

for all surface water states

for all groundwater states

Upper and lower bounds for GA

Generate initial population ($V_{gw,t+1}$, $V_{sw,t+1}$)



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next

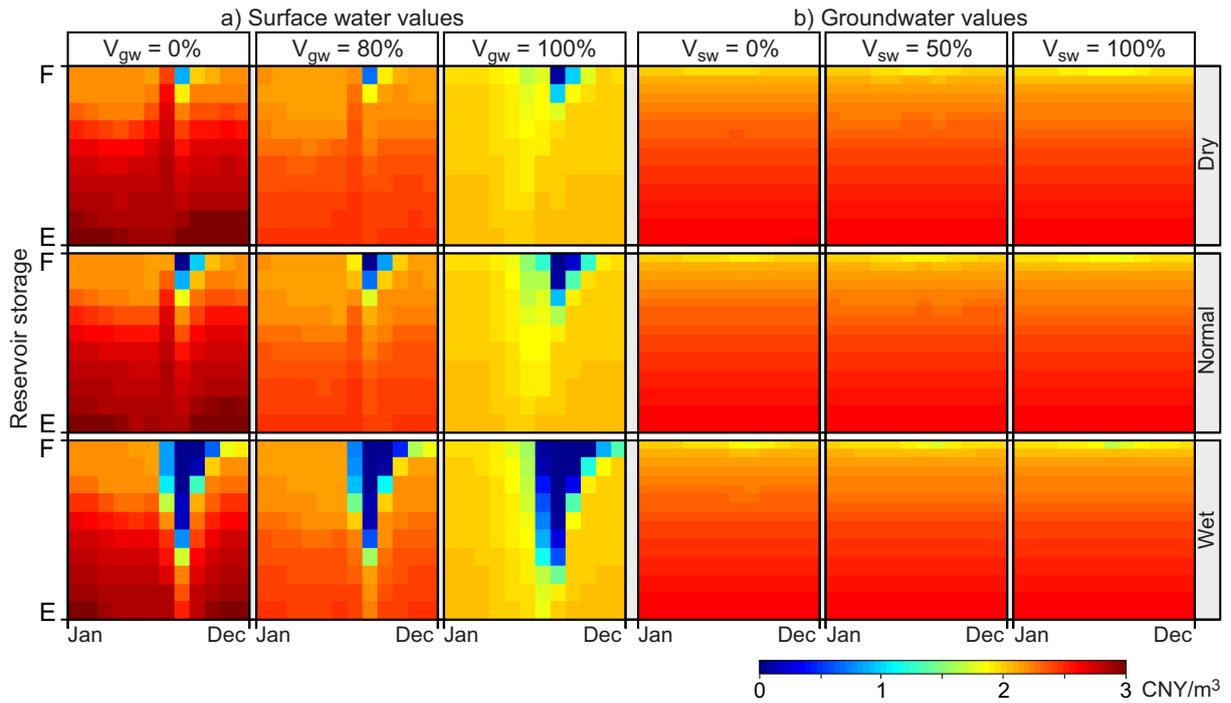
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Figure 3: SDP optimization algorithm design.

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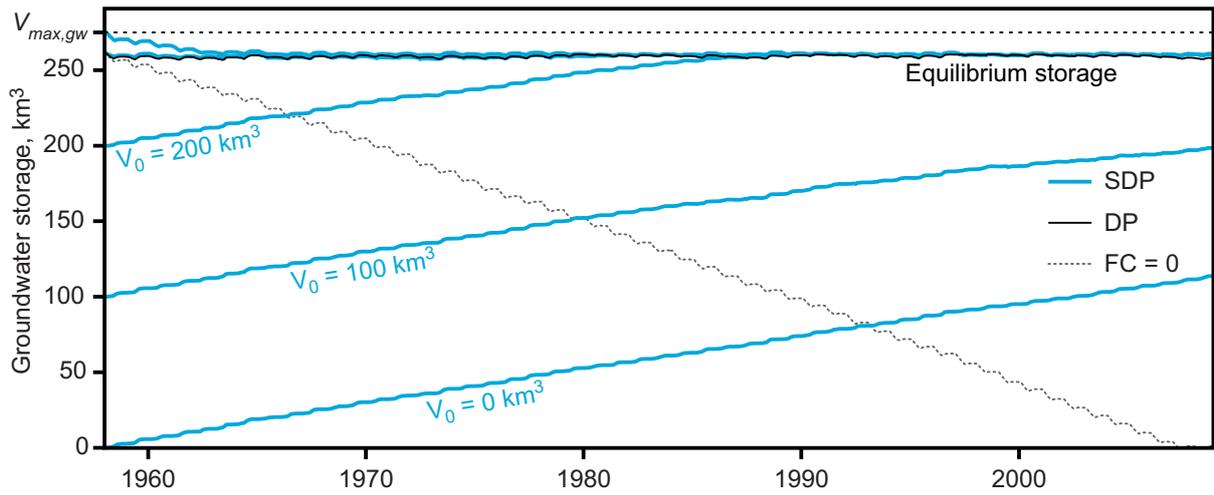
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1027 Figure 4: Temporal changes of the water values (CNY/m³) for the climate period before 1980.
 1028 The marginal water value is the true value of storing a unit volume of water for later use, and
 1029 varies with reservoir storage levels, runoff flow class and time of the year. a) Surface water
 1030 values at fixed [0%, 50%, 100%] groundwater aquifer storage. b) Groundwater values over
 1031 time at fixed [0%, 50%, 100%] surface reservoir. The reservoir storage is shown from E
 1032 (empty) to F (full).

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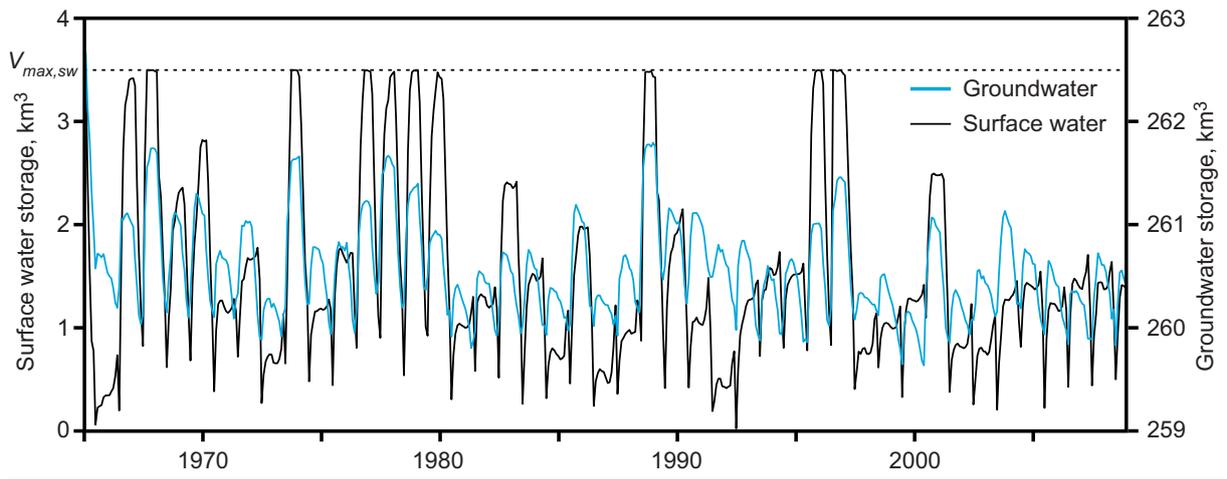


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1037 Figure 5: Simulated groundwater aquifer storage levels for 51 years of historical runoff with
1038 different initial groundwater tables (0, 100, 200, 258 and 275 km³). The perfect foresight DP
1039 and management without consideration of the future (FC = 0) are also shown.

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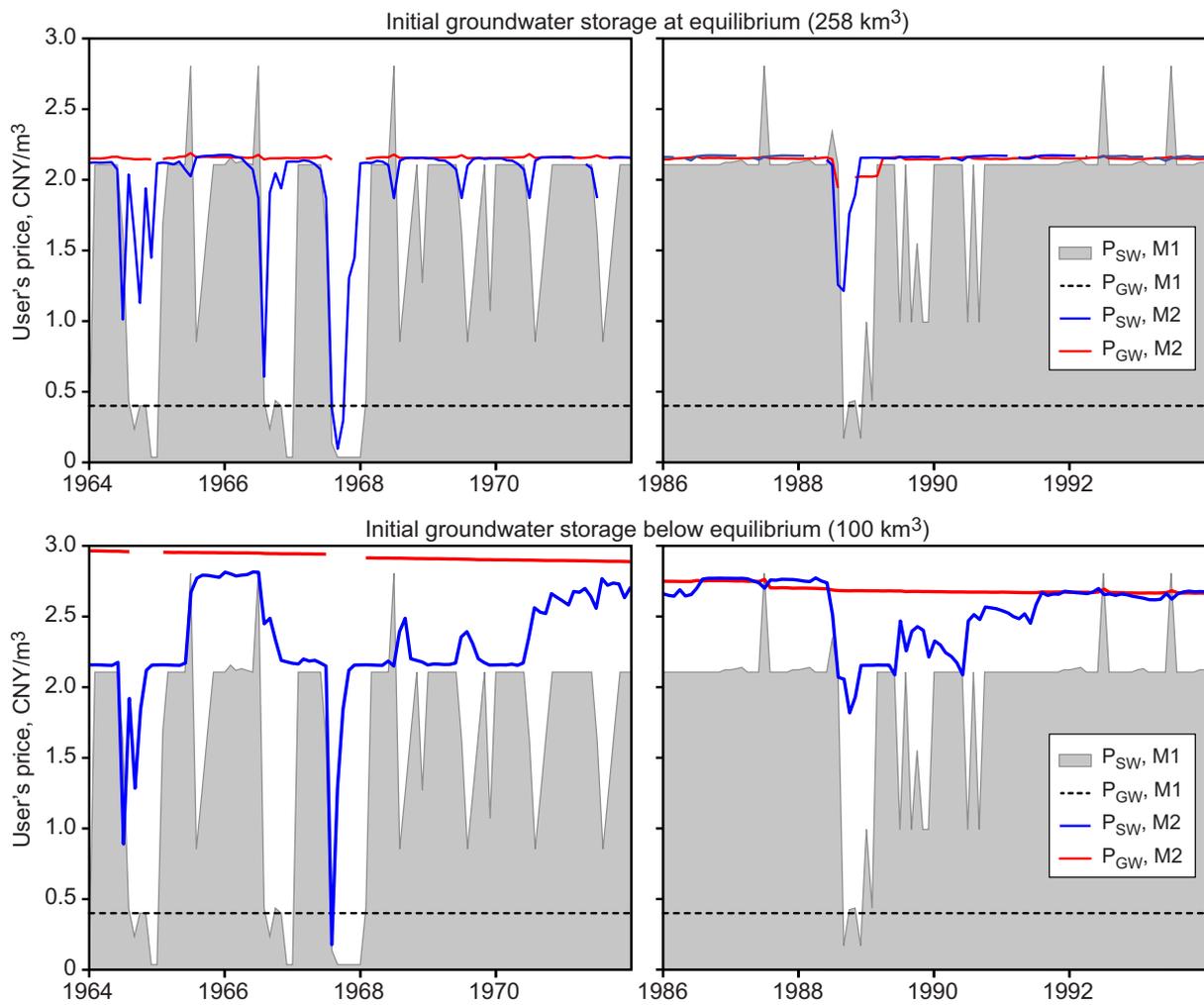
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1043 Figure 6: Simulated storage levels in the surface water reservoir and the groundwater aquifer
 1044 at equilibrium groundwater storage.

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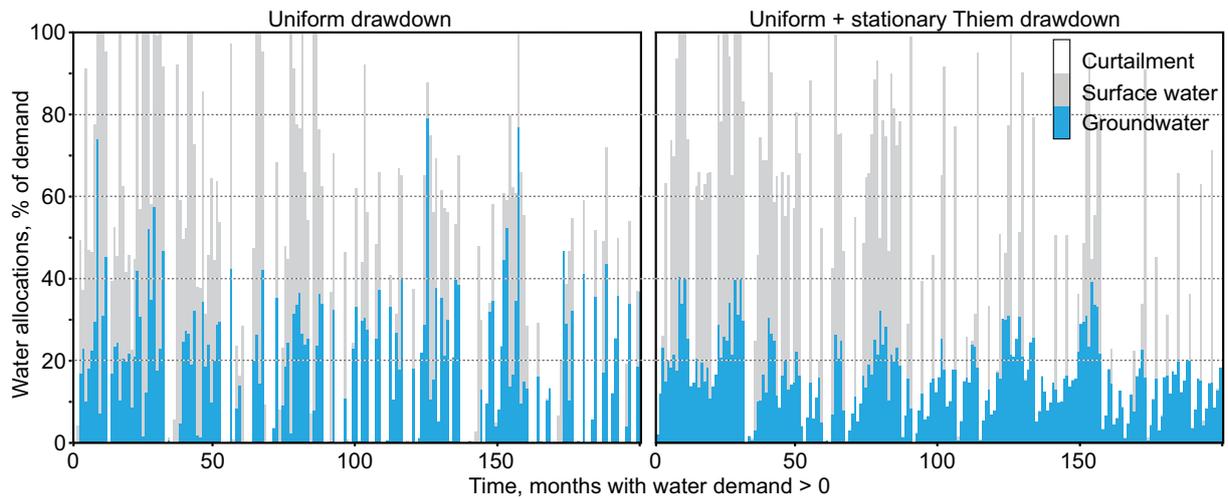
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Figure 76: User's price for groundwater and surface water through for a 51 year simulation based on simulated historical runoff for two initial groundwater storages. P = user's price. M1 = results from for with a single combined surface water reservoir and with a constant groundwater costs (Davidsen et al., 2015). M2 = results from the presented model framework with a combined surface water reservoir and with an additional a dynamic groundwater aquifer. The user's price for groundwater in M2 is the immediate pumping costs added the opportunity marginal costs from the water value tables.



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 1056 | Figure 87: Composition of allocations and curtailments to wheat agriculture in the Hebei
 1057 | Province for the months March, April, May and June through 51 years simulation from an
 1058 | initial groundwater storage at equilibrium (258km³). The results are shown for a simple
 1059 | drawdown model with uniform regional lowering of the groundwater table, and a more
 1060 | realistic drawdown model, which includes the stationary Thiem local drawdown cones.