We would like to thank Dr Serinaldi and the anonymous reviewer for their review of the 1 2 paper "Non-stationarity in annual maxima rainfall across Australia-implications for Intensity-3 Frequency-Duration (IFD) relationships". We have considered the Reviewers' comments and 4 provided detailed descriptions of how each comment will be addressed in the revised 5 manuscript below:

6 Response to Reviewer 1 (Dr Serinaldi)

7 Specific comments

8 Dr Serinaldi's specific comments center around three main themes. The first is the use of 9 change point analysis to test for non-stationarity in a data series, the second is the application 10 of the CUSUM test (and identifying multiple change points) and the third is the use of the term "regime shift" in preference to "non-stationarity" when discussing variability in the 11 12 annual maxima rainfall timeseries. Each of these issues were further built on in the technical 13 remarks provided by Dr Serinaldi and therefore are addressed in detail below.

14 **Technical remarks**

15	1. Please, consider to check the significance of (bias-corrected) serial correlation (if the
16	was not done) because it can affect the results of change point analyses (see e.g.
17	Springldi and Kilshy (2015a) and references therein for a discussion on Mann-

Serinaldi and Kilsby (2015a) and references therein for a dis 18

Kendall and Pettitt, which however holds true also for e.g. CUSUM and similar). As 19 shown above, apparent regime shifts can be artifacts resulting from hidden

21 Response: In response to the Dr Serinaldi's suggestion, the Durbin-Watson (DW) statistic

22 was used to test for autocorrelation (serial correlation) in the annual maxima timeseries

23 (Durbin and Watson (1950, 1951)). The Durbin-Watson statistic tests the null hypothesis that

24 the residuals from an ordinary least-squares regression are not autocorrelated against the

25 alternative that the residuals follow an AR1 process. The Durbin-Watson statistic ranges in

26 value from 0 to 4. A value near 2 indicates non-autocorrelation; a value toward 0 indicates

²⁰ persistence.

positive autocorrelation; a value toward 4 indicates negative autocorrelation. Typically, tabulated bounds are used to test the hypothesis of zero autocorrelation against the alternative of positive first-order autocorrelation. For the sample size in our case (~100) and a linear trend model with intercept the $d_{lower} = 1.522$ and $d_{upper} = 1.562$ for 1% significance.

All DW statistic values were found to be greater than the 1.562 (the upper bound for 1%
significance) providing no evidence to reject the null hypothesis (see figure 1 showing the
distribution of all DW statistic values for the 1-day annual maxima timeseries at each site).

A discussion of the DW test results are included in the revised paper (Section 2.2) to demonstrate that the annual maxima data does not suffer from serial correlation and therefore the statistical tests used in the change point analysis is appropriate.



37



39 stations (red line indicates $d_{upper}=1.562$ for 1% significance)

40	2.	P3453L20-25: In my opinion, such lines reflect some confusion on this topic. Trends
41		or change points in finite time series do not imply nonstationarity. Nonstationarity
42		cannot be in principle significant or not significant, because it is an assumption made
43		on the underlying process that can be introduced only if we know the underlying
44		nonstationary dynamics (physical equations, well-defined changes with a clear cause
45		such as flow regime changes due to dams operation, etc.). Please consider to reword

49 Response: Dr Serinaldi's review has highlighted some important points regarding the use of 50 the term "stationarity" and if/where it is applicable in our study. In particular he questioned 51 whether the term "regime shifts" was more fitting in describing our findings of change points 52 in the annual maxima rainfall timeseries that are possibly attributable to climate shifts. Dr Serinaldi stated that (following Koutsoyiannis and Montanari, 2014; Serinaldi and Kilsby, 53 54 2015) stationarity is a concept referring to models rather than to timeseries. In our case the 55 model is the IFD curve. Thus the text that describes the assumptions in the IFD development 56 of stationarity in the underlying processes (i.e. the statistical properties of the rainfall do not 57 change over time and that the chance of an extreme event occurring is the same at any point in time (past or future)) is relevant, however we agree that the text discussing change points 58 59 in the rainfall data may be misleading where the term non-stationarity has also been used. 60 However it is interesting to note that many studies have also used the word stationarity and non-stationarity when describing similar timeseries (e.g. Ishak et al. 2013, Westra and Sisson 61 62 2011, Wagesho et al 2013, Wilby 1998, etc), therefore there appears to be widespread disagreement on the use of this term. Despite this, on further review of the journal papers 63 provided by the Reviewer we agree that in our case the term "regime shift" is more suitable. 64 65 Given the above we have revised the text in the paper to reflect this. In particular, we have use the term non-stationarity only when referring to the IFD development, however the 66 67 sections of the paper that are focused on identifying change points in the rainfall timeseries have been edited and the term "regime shift" has been used in preference. Further, as per the 68 69 reviewer's suggestion (in his specific comments) we also changed the title of the paper to 70 "Regime shifts in annual maxima rainfall across Australia - implications for Intensity-

- 71 Frequency-Duration (IFD) relationships". The associated text also includes the references
- 72 provided by the reviewer.
- 73 74 75

3. P3457L25: Please consider to reword, e.g. "LP3 was not rejected at x% significance level for all series (or n series out of N)".

Response: The sentence in question was reworded as suggested to read "Here the null hypothesis is that the data fits the Log-Pearson III distribution (the alternate is that the data does not follow the Log Pearson III distribution). All p-values were greater than 0.05 (average p-value was 0.75), for all series (30min to 72hr durations at Brisbane, Sydney and Melbourne), therefore we accept the null hypothesis at the 5% significance level.

81 4. P3458L12-15: I do not know AR&R, but it is not clear to me why return periods
82 defined on annual maxima should be adjusted for PDS. Usually we do the opposite
83 when we start from PDS and we need the actual AMAX return periods (under suitable
84 conditions such as Poisson arrival dynamics, etc.). Please clarify.

85

86 Response: The methodology adopted in this paper to calculate return periods of annual 87 maxima specifically follows that outlined in Australian Rainfall and Runoff (1987), 88 Engineers Australia's guide to estimating and utilising IFD information. Published IFD currently used by industry in Australia are based on this method. The updated IFD (which are 89 90 NOT currently used in operation) are based on a revised statistical methods (for example, a 91 Generalised Extreme Value (GEV) frequency distribution was fitted to the annual maxima 92 rather than Log Pearson III and extension of sub-daily rainfall statistics to daily read stations 93 is conducted with Bayesian Generalised Least Squares Regression rather than PCA). The 94 purpose of adopting the AR&R 1987 method was to assess the implications of varying data 95 lengths and climatic variability on the resulting IFD (which have been historically used and are currently still in use) and to highlight the issue of underlying variability in the annual 96

97	maxima that should be appropriately considered and addressed in the current (and future)
98	revision of the IFD estimates.
99 100 101	5. <i>P3459L9: As mentioned above, step changes and nonstationarity are very different concepts and surely not synonyms.</i>
102	Response: This has been revised and clarified as per discussion above (Comment 2).
103 104 105 106 107 108 109	6. P3459L17-24: Leaving aside the use of the term nonstationarity, CUSUM identifies automatically the change point location and does not split the time series in two halves. If the Authors mean that the test proceeds based on subsequent dyadic partitions, this is right, but for such short time series it is actually quite difficult (and not meaningful) to go beyond 2-4 changes. Please note that many other refined techniques are available for segmentation of course, a question rises about the (physical) meaning of such refined segmentations

111 Response: This has been clarified in the text. The data is not split in equal halves for the

112 CUSUM test, it is split into two portions, which may or may not be equal. However, unless a

113 moving window is used (say 20 years as we did for the Mann- Whitney, multiple regime

114 shifts could still be missed using this method.

115 7. P3459L25-P3460L10: Following the previous remark, my interpretation of P3473Fig5 is a bit different. The almost uniform spread of changes across the 116 decades denotes that such changes occur quite randomly, and sincerely I cannot see a 117 118 tendency to cluster in the east coast. We may see something in panel (b), but the 119 spatial distribution of the stations is not uniform and we cannot exclude that such 120 stations are spatially correlated, as they are subject to similar climate forcings (thus 121 reducing the evidence for changes). Note that spatial correlation is another factor that can strongly affect the outcome of such a type of tests (see e.g. Douglas et al. 122 123 (2000), Yue et al. (2003), Guerreiro et al. (2014), among others) 124

125 Response: We agree with the Reviewer that there is almost a uniform spread of changes

126 across the decades based on panel (a). We state in the paper that ", the large-scale climate

127 phenomena impact various regions of Australia at different times of the year and to varying

128 degrees, therefore it is not surprising that the timing of shifts in the annual maxima timeseries

129 varies spatially and temporally." However has been further clarified in the revised paper.

130 The clustering along the east coast can only be clearly seen in panel B. The text has been 131 revised to clarify this.

The spatial correlation of the annual maxima timeseries was investigated as per the reviewer's suggestion. We found that less than 9% of all possible pairings of rainfall data sets display a significant (yet weak) correlation at the 5% level (r > 0.2, significance based on n=100). Only 8 pairings (out of 4465) were correlated at 0.5 or higher. It was also found that stations located more than 500km apart were unlikely to be correlated and that the strength of the correlation reduced as distance increased between the pairs (see Figure 2).





141	8.	Section 3.2: Again, my interpretation of P3474Fig6 and P3475Fig7 is a bit different.
142		If I'm right, box plots for IPO(-) summarize the distribution of 41 AMAX (1913-1920
143		and 1945-1977), while we have 67 AMAX for IPO(+) box plots. For such sample
144		sizes, inferring difference in distribution based on box plots is a bit hard (at least). My
145		suggestion is to use some formal two-sample goodness-of-fit tests such as the two-
146		sample Kolmogorov-Smirnov or similar, thus accounting for sampling uncertainty
147		and different sample sizes. In any case, comparing box plots (overlooking the large
148		uncertainty of the quantile estimates) is not informative and does not provide a
149		quantitative assessment, especially in this case where differences between IPO(-) and
150		<i>IPO(+) regimes are really hard to recognize.</i>

152	Response: As suggested a two sample KS test was applied to the data (to test the
153	significance of the difference between the two IPO distributions). It was found only the
154	results for the Sydney were statistically significant (p-value <0.1). A discussion of the
155	two-sample KS test have been included in the revised paper as suggested. Given the
156	results of the significance test we only further investigate the impact of the IPO step
157	induced regime shifts on the IFD for Sydney (rather than all three stations).

158
9. The same holds for P3475Fig7: if I'm right, this diagram shows the differences Δ (in %) between the point estimates of rainfall return levels obtained by LP3 distributions fitted on 41 and 67 AMAX. It is almost superfluous to highlight how large the

161 uncertainty of such a point estimates can be. I suggest a fairer check based on a

simple bootstrap procedure. For each duration:

163	1.	resample with replacement IPO(-) and IPO(+) time series to obtain two new B-
164		samples;

- 165 2. for each B-sample refit LP3, compute the required LP3 return levels and calculate the difference $\Delta^{(B)}$ as for the observed data;
- 167 3. repeat previous steps B times (e.g. 1000) and store the obtained B differences (for 168 each ARI). These values can be used to build the empirical distribution of the 169 differences $\Delta^{(B)}_{i}$, i = 1, ..., B. This distribution describes the effects of sampling and 170 parameter estimation uncertainties under the hypothesis of existence of two different 171 regimes;

172	4.	Use the $B \Delta^{(B)}_{i}$ values to build confidence intervals (CIs) at a given confidence level
173		(e.g.95%). If these CIs include $\Delta = 0$, then there is not evidence for a significant
174		difference, otherwise we can conclude the opposite.

176 I think this is a better way to provide a quantitative assessment. Of course, conclusions
177 concern the effects of possible regime shifts and not of nonstationarity. Section 3.2
178 should be reworded according to the results of the analyses suggested above.

179 Response: We would like to thank Dr Serinaldi for this suggestion and we have completed 180 the analysis as per the steps outlined above. This has resulted in a more robust test of the 181 effects of the IPO induced regime shifts. As per the previous comment, we only apply this 182 method to the Sydney station (since this was the only station where the KS statistic was 183 significant). Figure 7 has been replaced by the figure below that shows the percentage 184 difference in the rainfall intensity estimate between the IPO positive and negative phases. 185 Positive (negative) values represent an increase (decrease) in rainfall intensity during IPO 186 positive compared to IPO negative. The relationships observed are robust for most durations 187 and return intervals given that the CIs do not include 0 (other than 30mins 5 years, 2 hours 50 188 years, 6 hours 20 years and 24 hour s5 years).





192	10. Section 4: as for Section 3.2, this section should be reworded according to the
193	updated results.
194	

195 Response: Noted, this has been revised.

196 <i>11. Please avoid sentences such as that in P346</i>	4L27-29 and P3465L1-3: even after more
197 <i>accurate analyses, there is not way to make</i>	unquestionable conclusions about
198 nonstationarity if we do not identify a well-d	efined mechanism of evolution which is
almost perfectly predictable (at least, at the	time scales of interest).
200	• ,

201	Response: The sentence "Based on the results of this study, and literature cited within this
202	paper, we emphasise that there undoubtedly is non-stationarity in historical short duration
203	rainfall extremes but the characteristics and causes of this non-stationarity vary from
204	location to location and decade to decade - something which must be considered and
205	accounted for when attempting to estimate IFD design rainfalls and prior to quantifying

206 how those IFD estimates might change in the future." will be revised and expanded to 207 read "This study has highlighted the existence of regime shifts in annual maxima rainfall 208 data in Australia. The driving mechanisms of these regime shifts are likely to vary from 209 location to location and decade to decade. However, these shifts are typical of many 210 natural phenomena and can be described by processes characterized by long range 211 dependence (or regime-switching processes) and captured by hidden Markov models (or 212 similar), resulting in a mixture of distributions that alternate stochastically according to 213 the transition probability from one regime to the next (e.g. Serinaldi and Kilsby, 2015a). 214 While the strategy of defining IFDs for two (or more) different regimes (e.g Serinaldi andKilsby (2015a)) currently only partially solves the problem, as we often do not know 215 216 the beginning or the end of a specific regime (be it rainfall or climate driver), recent work 217 has focused on optimizing designs and planning strategies based on the range of what is 218 plausible rather than a reliance on knowing the current and future climate state (e.g. 219 Mortazavi-Naeini et al., 2015). At the same time, work is also underway on seamless 220 prediction at a range of timescales and if/when this eventuates the results discussed here 221 become even more important/useful. Nevertheless, the immediate usefulness of the 222 insights presented here occurs when first establishing the IFD, as an approach similar to 223 that employed here can be used to determine if the underlying data are biased to a mostly 224 wet or mostly dry regime (or a mix of both) which then provides an indication as to 225 whether the IFD is likely to be an over- or underestimate of the true risk. Importantly, this 226 issue needs to be considered and accounted for when attempting to estimate IFD design 227 rainfalls and prior to quantifying how those IFD estimates might change in both the near 228 and long-term future."

230 **Response to Reviewer 2 (Anonymous reviewer)**

231 Specific comments:

- 232 1. Recent studies (Montanari and Koutsoyiannis, WRR, 2014; and references therein, 233 Koutosoyiannis, JH, 2006) show that modeling approaches which consider non-234 stationarity of real world time series without examining the properties of the 235 stochastic processes, may be inappropriate. The way the authors test and claim for 236 non-stationarity in the extremes is inadequate and quite limited. In fact, it has been shown in some studies (Serinaldi and Kilsby, AWR, 2015) that non-stationary models 237 may increase the uncertainties and that traditional concepts should still be retained 238 239 as bench- marks. Thus, the authors' skepticism about the BoM and ARR's existing 240 approaches may not be justified.
- 241 Response: Based on this review and the review of our paper provided by Dr Serinaldi we
- have improved the paper as follows:
- 243a. We have Included a test for serial correlation (Section 2.2) using the Durbin-Watson244(DW) statistic. The Durbin-Watson statistic tests the null hypothesis that the residuals245from an ordinary least-squares regression are not autocorrelated against the alternative246that the residuals follow an AR1 process. All DW statistic values were found to be247greater than the 1.562 (the upper bound for 1% significance) providing no evidence to248reject the null hypothesis (see figure below showing the distribution of all DW249statistic values for the 1-day annual maxima timeseries at each site).
- 250 b. We have addressed the issue of potential spatial correlation among rainfall sites. We found that less than 9% of all possible pairings of rainfall data sets display a 251 significant (yet weak) correlation at the 5% level (r >0.2, significance based on 252 253 n=100). Only 8 pairings (out of 4465) were correlated at 0.5 or higher. It was also 254 found that stations located more than 500km apart were unlikely to be correlated and 255 that the strength of the correlation reduced as distance increased between the pairs. 256 This is not surprising given annual maximum rainfall events are due to synoptic scale 257 processes. This is include in revised Section 2.2 of the paper. c. We have revised the text in the paper with respect to the use of the term non-258
- stationarity. In particular, we use the term non-stationarity only when referring to the
 IFD development (which is deemed appropriate given the IFD is essentially a model),
 however the sections of the paper that are focused on identifying change points in the
- rainfall timeseries have been edited and the term "regime shift" has been used in

263	preference. Further, we have also change the title of the paper to "Regime-shifts in
264	annual maxima rainfall across Australia - implications for Intensity-Frequency-
265	Duration (IFD) relationships". The associated text also includes the references
266	provided by the reviewer

268 2. Also, the definition of return period itself (and equivalently that of ARI) may change 269 in the non-stationary setting (Salas and Obeysekara, ASCE JHE, 2013). Moreover, 270 although the title mentions "- implications for Intensity-Frequency-Duration (IFD) 271 relationships", this paper only presents a discussion (Section 3.2) which contains 272 rather generic discussion on how non-stationarity may affect such relationships, 273 without carrying out any analysis on how the observed-period IFD relationships 274 actually change because of non-stationarity (such as that done by Cheng and AghaKouchak, Sc. Re- ports, 2014), bringing into question the novelty, utility and 275 276 scientific contribution of this study.

Response: We agree with the reviewer that the ARI may change in the non-stationary 277 278 setting. In fact, that is the point we are making in our paper; that depending on when the 279 data is sampled from to generate the IFD, it may be biased to either a wet or a dry phase 280 (or surplus or absence of high intensity events) and therefore would have consequences 281 on the resulting return period for individual rainfall depths. Indeed we suggest in our discussion "that a separate set of IFDs could be developed for use in high risk modelling 282 283 for engineers who need to account for the 'worst case' (in a similar manner to climate 284 change allowances). This second set of IFDs could be developed based on the periods of 285 elevated annual maxima alone (for those stations with clearly defined epochs of annual 286 maxima) such that if we were to enter such an epoch, designs based on these estimates would be robust for the duration of such a period." We disagree with the reviewer that we 287 288 do not show how nonstationarity (which we will now term regime shifts) may affect the IFD relationships. This is demonstrated in Section 3.2 "Effect of non-stationarity on IFD 289 290 estimation" where we recalculate the IFD curves for the difference phases of the IPO. 291 Importantly we show that the return period is different for the various rainfall depths and 292 durations depending on the underlying rainfall dataset (i.e. depending on whether it is

- sampled from the IPO positive or IPO negative distributions).
- 294
 3. The authors mention, in their conclusions, "The research presented here demonstrates that information currently available on natural variability..can act as a guide to the base- line..." - this is a fat-fetched conclusion. The present research, however, doesn't provide any guidelines on how this baseline can be defined.
- Response: This has been clarified in the revised paper. Our intention here was to emphasize that, for regions where large-scale climate drivers operate on a multi-year to multi-decadal timescales and are known to influence extreme rainfall events, we can use this information to determine if the climate statistics on which the IFD are based are likely to be biased or missing crucial information.
- 4. Is IPO the same as PDO (Pacific Decadal Oscillations)? If it was known apriori that locations such as Melbourne are not affected by the IPO, why was it chosen for the analysis? Perhaps a more appropriate approach would consider several natural variability modes, as well as forced changes and investigate their individual effects on rainfall extremes.
- Response: The IPO is not the same as the PDO. The IPO is a Pacific Basin wide phenomena rather than just the north Pacific that is represented by the PDO. There are similarities between the two timeseries however and they are significantly correlated. According to Salinger et al 2001 "The IPO may be a Pacific-wide manifestation of the
- 312 PDO, excluding subdecadal time scales, and seems to be part of a continuous spectrum of
- low frequency modulation of ENSO, and so may be partly stochastic".
- It is true that some existing studies suggest that the IPO signal on rainfall tends to be weaker in Melbourne due to competing influences from the Southern Ocean, however we cannot say that Melbourne rainfall/climate is "not affected by the IPO". Some studies suggest IPO significantly effects rainfall characteristic in Melbourne (e.g. Verdon et al 2004, Gallant et al 2012) while others do not, therefore in our study we do not make any assumptions about IPO effects on rainfall maxima in Melbourne and include it in our investigation. Our results suggest there is a relationship where "all events (other than 72

hours) with a 2-year ARI are associated with a higher rainfall intensity estimate in IPO

322 positive for Melbourne, however the reverse is true for the less frequent events."

323 While we agree with the reviewer that there are several modes of natural climate 324 variability that may have an effect the extreme rainfall from year to year, in our study we 325 were specifically interested in climate drivers that are likely to force a regime shift in extreme rainfall (similar to that observed for flood risk). Therefore we were interested in 326 327 drivers that operate on a decadal to multi-decadal timescale (as is the case for the IPO). Other drivers (such as ENSO, Indian Ocean Dipole, Southern Annular Mode) tend to 328 329 influence rainfall in Australia on much shorter timescales. However, if this method was to 330 be applied to regions other than east coast Australia (i.e. where IPO is known not to be 331 the primary driver on decadal to multi-decadal timescales), other potential sources of 332 decadal to multi-decadal variability would need to be identified.

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Response: The Reviewer's point is correct, the GEV distribution does indeed fit the 336 337 annual maxima data well. In fact the updated IFD (which are NOT currently used in 338 operation) are based on a revised statistical methods that includes fitting the GEV 339 distribution to the data in preference to the Log Pearson III. However, the methodology 340 adopted in this paper (including fitting the Log Pearson III) to calculate return periods of annual maxima deliberately follows that outlined in Australian Rainfall and Runoff 341 342 (1987), Engineers Australia's guide to estimating and utilising IFD information. The 343 purpose of adopting the AR&R 1987 method was to assess the implications of varying data lengths and climatic variability on the resulting IFD (which have been historically 344 345 used and are currently still in use) and to highlight the issue of underlying variability in

the annual maxima that should be appropriately considered and addressed in current and 346 future revisions of the IFD estimates. Further, as an additional check the KS goodness of 347 348 fit test was applied to test if the Log Pearson III was a reasonable fit to the data. Here the 349 null hypothesis is that the data fits the Log-Pearson III distribution (the alternate is that the 350 data does not follow the Log Pearson III distribution). All p-values were greater than 0.05 351 (average p-value was 0.75), for all series (30min to 72hr durations at Brisbane, Sydney 352 and Melbourne), therefore we accept the null hypothesis at the 5% significance level. 353 This has been clarified in the revised paper.

Claims such as "we emphasize that there undoubtedly is non-stationarity in historical
 short duration rainfall extremes" might be inappropriate for reasons stated above

Response: We agree and this has been revised and the discussion extended. The followingtext replaces the sentence above:

358 "This study has highlighted the existence of regime shifts in annual maxima rainfall data in Australia. The driving mechanisms of these regime shifts are likely to vary from 359 360 location to location and decade to decade. However, these shifts are typical of many natural phenomena and can be described by processes characterized by long range 361 dependence (or regime-switching processes) and captured by hidden Markov models (or 362 363 similar), resulting in a mixture of distributions that alternate stochastically according to 364 the transition probability from one regime to the next (e.g. Serinaldi and Kilsby, 2015a). 365 While the strategy of defining IFDs for two (or more) different regimes (e.g Serinaldi 366 andKilsby (2015a)) currently only partially solves the problem, as we often do not know 367 the beginning or the end of a specific regime (be it rainfall or climate driver), recent work 368 has focused on optimizing designs and planning strategies based on the range of what is plausible rather than a reliance on knowing the current and future climate state (e.g. 369 370 Mortazavi-Naeini et al., 2015). At the same time, work is also underway on seamless

- 371 prediction at a range of timescales and if/when this eventuates the results discussed here become even more important/useful. Nevertheless, the immediate usefulness of the 372 373 insights presented here occurs when first establishing the IFD, as an approach similar to 374 that employed here can be used to determine if the underlying data are biased to a mostly 375 wet or mostly dry regime (or a mix of both) which then provides an indication as to 376 whether the IFD is likely to be an over- or underestimate of the true risk. Importantly, this issue needs to be considered and accounted for when attempting to estimate IFD design 377 rainfalls and prior to quantifying how those IFD estimates might change in both the near 378 379 and long-term future."
 - *Literature review pertains mostly to studies on Australian datasets, whereas much work on similar ideas are also carried out elsewhere.*
 - 383 Response: The literature review has been extended to include the references mentioned by
 - 384 the Reviewer as well as the following international papers:
 - Douglas EM, Vogel RM, Kroll CN. (2000) Trends in floods and low flows in the
 United States: impact of spatial correlation. Journal of Hydrology 240: 90–105.
 - Guerreiro S.B., Kilsby C.G., Serinaldi F. (2014) Analysis of time variation of rainfall in transnational basins in Iberia: abrupt changes or trends? Int J Climatol 34(1):114– 133
 Koutsoviannis D., Montanari A. (2014) Negligent killing of scientific concepts: the
 - Koutsoyiannis D., Montanari A. (2014) Negligent killing of scientific concepts: the sta- tionarity case Hydrol Sci J http://dx.doi.org/10.1080/02626667.2014.959959
 - Salas J.D., Obeysekera J. (2014) Revisiting the concepts of return period and risk for nonstationary hydrologic extreme events. J Hydrol Eng, 19(3):554–68.
 - Serinaldi F., Kilsby C.G. (2015a), Stationarity is undead: Uncertainty dominates the distribution of extremes. Advances in Water Resources, 77, 17-36
 - Serinaldi F., Kilsby C.G. (2015b). Stochastic Environmental Research and Risk
 Assessment
 - Yue S, Pilon P, Phinney B. (2003) Canadian streamflow trend detection: impacts of serial and cross-correlation. Hydrological Sciences Journal 48(1): 51–64.
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403	<u>Regime shifts</u> in Annual Maxima Rainfall across Australia– Implications	
404	for Intensity-Frequency-Duration (IFD) relationships	Danielle Verdon-Kidd 11/9/2015 3:29 PM Deleted: Non-stationarity
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407	Danielle C. Verdon-Kidd ¹ , Anthony S. Kiem ¹	
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409	¹ Environmental and Climate Change Research Group, School of Environmental and Life	
410	Sciences, University of Newcastle, AUSTRALIA	
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423 Abstract

424 Rainfall Intensity-Frequency-Duration (IFD) relationships are commonly required for the 425 design and planning of water supply and management systems around the world. Currently 426 IFD information is based on the 'stationary climate assumption' - that weather at any point in 427 time will vary randomly and that the underlying climate statistics (including both averages 428 and extremes) will remain constant irrespective of the period of record. However, the validity 429 of this assumption has been questioned over the last 15 years, particularly in Australia, 430 following an improved understanding of the significant impact of climate variability and 431 change occurring on interannual to multidecadal timescales. This paper provides evidence of 432 regime shifts in annual maxima rainfall timeseries using 96 daily rainfall stations and 66 sub-433 daily rainfall stations across Australia. Further, the effect of these regime shifts on the 434 resulting IFD estimates are explored for three long-term sub-daily rainfall records (Brisbane, 435 Sydney and Melbourne) utilising insights into multidecadal climate variability. It is 436 demonstrated that IFD relationships may under- or over-estimate the design rainfall 437 depending on the length and time period spanned by the rainfall data used to develop the IFD 438 information. It is recommended that regime shifts in annual maxima rainfall be explicitly 439 considered and appropriately treated in the ongoing revisions of Engineers Australia's guide 440 to estimating and utilising IFD information, 'Australian Rainfall and Runoff', and that clear 441 guidance needs to be provided on how to deal with the issue of regime shifts in extreme 442 events (irrespective of whether this is due to natural or anthropogenic climate change). The 443 findings of our study also have important implications for other regions of the world that 444 exhibit considerable hydroclimatic variability and where IFD information is based on 445 relatively short data sets.

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452 **1. Introduction**

Information on rainfall event intensity, frequency and duration (IFD, or IDF as it is known in some countries) plays a critical role in the design of dams, bridges, stormwater drainage systems and floodplain management. Dependent upon the application, information is required for event-durations ranging from hours to several days. The development of IFD relationships were first proposed by Bernard (1932) and since then different versions of this relationship have been developed and applied worldwide (e.g. Bara et al. 2009, Chen 1983, Hershfield 1961, IHP-VII 2008, Nhat et al. 2006, Raiford et al. 2007).

460 Historically, in Australia, IFD design rainfall curves were developed by the Australian 461 Bureau of Meteorology (BoM) for durations ranging from 5 minutes to 72 hours and Average 462 Return Intervals (ARI) of 1 year to 100 years (however, recently additional durations and 463 ARIs have also been developed). Up until very recently IFD information available to (and 464 used by) engineers and hydrologists were developed 25 years ago, as part of Engineers 465 Australia publication Australian Rainfall and Runoff (AR&R) in 1987. New IFD information was released early in 2013 after a major revision of IFD information carried out by Engineers 466 467 Australia. Importantly, the revised IFD information is based on a longer and more extensive 468 rainfall data set (http://www.bom.gov.au/water/designRainfalls/ifd/). However, the BoM and 469 Engineers Australia still recommend to use the AR&R 1987 information for existing flood 470 studies and the probabilistic rational method and to conduct sensitivity testing with the 471 IFD revised 2013 AR&R parameters including the new design rainfalls 472 (http://www.bom.gov.au/water/designRainfalls/ifd/index.shtml).

473 At the time of writing, the revised IFD information does not take into account the impact of 474 climate change on IFD estimates. This is part of ongoing research commissioned through 475 Engineers Australia. It is also not yet clear how or if the role of natural climate variability is 476 going to be considered. Of concern is the fact that currently, estimates of IFD are based on 477 the assumption that "climatic trend, if it exists in a region, has negligible effect on the design 478 intensities" (Pilgrim 1987). This is known as the 'stationary climate assumption' n (i.e. 479 thestatistical properties of the rainfall do not change over time) and implies that the chance of 480 an extreme event occurring is the same at any point in time (past or future). However, the 481 validity of this assumption has been questioned over the last 15 years following 482 demonstration of the significant impact of climate variability occurring on interannual to 483 multidecadal timescales in Australia. For example, research has shown that annual maximum 484 flood risk estimates in Australia vary depending on climate state (e.g. Ishak et al. 2013, Kiem et al. 2003, Leonard et al. 2008). Importantly these studies demonstrate that founding flood 485 risk estimates on an unsuitable time period has the potential to significantly underestimate (or 486 487 overestimate) the true risks. This may apply to design rainfall also given that current IFD 488 estimates are based on varying lengths of data spanning different time periods (the latest IFD 489 estimates are based on all daily-read stations with 30 or more years of record and all continuously-recording stations with more than 8 years of record). 490

491 Khaliq et al. (2006) explained that the traditional idea of probability of exceedance and return 492 period are no longer valid under non-stationarity. Recently, Jakob et al (2011a) found that 493 rainfall quantile estimates derived for Sydney Observatory Hill for the period 1976 to 2005 494 show significant decreases across durations from 6 minutes to 72 hours. Jakob et al (2011b) 495 subsequently extended the sub-daily rainfall data analysis to 31 sites located in southeast 496 Australia, assessing variations in frequency and magnitude of intense rainfall events across 497 durations from 6 minutes to 72 hours. This study identified two different trends in the data 498 sets, a decreasing trend in frequency of events at durations of 1-hour and longer for sites in 499 the north of the study region, while sites in the south cluster displayed an increase in 500 frequency of events, particularly for sub-hourly durations. Importantly Jakob (2011a, 2001b) 501 concluded that, for at least some regions of Australia, trends found in historical records has

Danielle Verdon-Kidd 11/9/2015 3:48 PM Deleted: investigated the assumption of stationarity in rainfall frequency analysis and

Danielle Verdon-Kidd 11/9/2015 3:33 PN Deleted: non-stationarity Luke Kidd 17/9/2015 11:40 PM Deleted: regime shifts 506 the potential to significantly affect design rainfall estimates. Westra and Sisson (2011) also 507 investigated evidence of trends in extreme precipitation at sub-daily and daily timescales 508 (1965-2005) using a spatial extreme value model. They identified a statistically significant increasing trend in precipitation extremes for the sub-daily data set, however at the daily 509 510 timescale no change in annual maximum rainfall could be detected with the exception of 511 southwest Western Australia (Westra and Sisson 2011). Further, Yilmaz and Perera (2014) 512 conducted change point analysis for extreme rainfall data for storm durations ranging from 6 513 minutes to 72 hours in Melbourne, and found evidence of regime shifts, concluding the year 514 1966 as a statistically significant change point. Yilmaz et al (2014) then investigated changes in_extreme rainfall through trend analysis, non-stationarity tests and non-stationary GPD 515 516 models (NSGPD) for Melbourne. They found statistically significant extreme rainfall trends 517 for storm durations of 30 minutes, 3 hours and 48 hours, however for above storm durations 518 there was no evidence of a regime shift (which they termed 'non-stationarity') according to 519 statistical non-stationarity tests and non-stationary GPD (Yilmaz et al (2014).

520 A limitation of the analysis presented by Westra and Sisson (2011) and Jakob et al (2011a, 521 2011b) is that they tested for linear trends in the rainfall data series based on the hypothesis 522 that extreme rainfall events would have either decreased, increased or exhibited no trend over 523 the time period being investigated. However these are not the only attributes of trend 524 detection, since annual rainfall maxima may also cycle through interannual to multidecadal 525 periods (note that Westra and Sisson (2011) also investigated possible links between extreme rainfall and annual fluctuations in the El Niño/Southern Oscillation (ENSO)). Therefore, 526 527 depending on what time period the annual rainfall maxima data are derived from (in reference to any long term cycles or epochs) the observed trends may be misleading or even not 528 529 apparent (leading to the misconception that regimes shifts are non-significant or not an important consideration). Recently Yilmaz et al (2014) investigated the potential impact of 530

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Danielle Verdon-Kidd 11/9/2015 3:43 PM Deleted: non-stationarity Danielle Verdon-Kidd 11/9/2015 3:43 PM Deleted: is 537 the Interdecadal Pacific Oscillation (IPO) on extreme rainfall and resulting IFD for a case 538 study in Melbourne. They concluded that, the IPO negative phase can be the driver of higher 539 rainfall intensities for long durations and high return periods. However, the trends in extreme 540 rainfall data and differences in rainfall intensities for short storm durations and return periods could not be explained with the IPO influence. Given that Melbourne is located in south-east 541 542 Australia, where the influence of the IPO is temporally variable due to other climate drivers 543 operating (acting to enhance or suppress impacts, see Kiem and Verdon-Kidd 2010; 2009), 544 the research by Yilmaz et al (2014) provides promise for developing relationships between 545 extreme rainfall and the IPO for regions where the IPO may have a more consistent influence (due to fewer competing climate modes), such as north-eastern Australia, 546 Therefore this paper aims to establish if there is evidence of regime shifts in the annual 547 548 maxima rainfall timeseries (1-hour to 7-days) across Australia by testing for shifts (regardless 549 of direction or timing) in the long term sub-daily and daily data. Further, the implications on 550 IFD estimation are explored, along with the potential influence of the IPO on extreme rainfall and resulting IFD, Recommendations are then provided as to how these insights may be 551

552 incorporated in future revisions of AR&R.

553 **2. Data and methods**

554 2.1 Data

555 2.1.1 Rainfall data

556 Sub-daily and daily rainfall data for Australia were obtained from the BoM. Sub-daily data 557 records from continuously recording (i.e. pluviograph) rainfall stations in Australia tend to be 558 relatively short, hindering the ability to conduct trend and attribution studies. In this study 559 pluviograph rainfall stations were chosen with data spanning at least 40 years and at least 560 90% complete, resulting in 66 stations (see Figure 1a). In order to address the concerns raised Luke Kidd 14/9/2015 8:05 PM Deleted: S

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ARI should not be extrapolated from more than twice the record length), three long-term data
sets, highlighted in Figure 1a, were chosen for further analysis that contained data from at
least 1913 onwards (Brisbane Aero, Sydney (Observatory Hill) and Melbourne Regional
Office).

574 Daily rainfall stations with data spanning the period 1900 to 2009 were selected in order to 575 capture as much temporal variability as possible (see Figure 1b). These stations were filtered 576 according to the amount of data missing in order to identify the highest quality stations 577 recording rainfall during this period, resulting in 96 being considered suitable for further 578 analysis. Due to variability in the quality and quantity of rainfall data in each State of 579 Australia, the following selection criteria were applied:

- New South Wales, Queensland and Victoria selected stations are at least 97%
 complete;
- Tasmania- selected stations are at least 90% complete; and
- South Australia, Northern Territory and Western Australia selected stations are at
 least 85% complete.
- 585 ****Figure 1 about here****

586 2.1.2 Climate index data

The climate of Australia has experienced a number of <u>regime shifts in climate during its</u> history, resulting in sustained periods of above average rainfall and storminess and abnormally cool temperatures, followed by the reverse conditions (i.e. droughts and elevated bushfire risk) (e.g. Erskine and Warner 1988, Franks and Kuczera 2002, Kiem et al. 2003, Kiem and Franks 2004, Verdon et al. 2004). These shifts have tended to occur every 20 to 30

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Danielle Verdon-Kidd 11/9/2015 3:44 PM Deleted: ' Danielle Verdon-Kidd 11/9/2015 3:44 PM Deleted: ' 596 years and are associated with changes in the Interdecadal Pacific Oscillation (IPO, Power et 597 al. 1999). The IPO represents variable epochs of warming (i.e. positive phase) and cooling 598 (i.e. negative phase) in both hemispheres of the Pacific Ocean (Folland et al. 2002). 599 Importantly, the IPO has been shown to influence the magnitude and frequency of flood and 600 drought cycles across eastern Australia (Kiem et al. 2003, Kiem and Franks 2004). In New 601 Zealand, the IPO is also associated with similar shifts in flood frequency (McKerchar and 602 Henderson 2003). It has been noted that, following the abrupt shift in the IPO in the mid 603 1970s, the period, amplitude, spatial structure and temporal evolution of ENSO markedly 604 changed (Wang and An, 2001). Historically, during negative phases of the IPO there tends to 605 be more La Niña (wet) events and fewer El Niño (dry) events (Kiem et al. 2003, Verdon and 606 Franks 2006), resulting in an overall 'wet' epoch for eastern Australia and New Zealand . 607 While during the positive phase of the IPO there tends to be a higher frequency of El Niño 608 events and fewer La Niña events (Kiem et al. 2003, Verdon and Franks 2006), resulting in an 609 overall 'dry' epoch. In this study negative phases of the IPO were defined as 1913-1920 and 610 1945-1977, while positive phases included 1921-1944 and 1978 to 2010.

611 2.2 Statistical tests

612 A 20 year moving window was used to test for low frequency variability in the annual 613 maxima timeseries (1-hour, 1-day and 7-day). A Mann-Whitney U test was then used to 614 determine the statistical significance of possible regime shifts by testing if the first 10 years 615 of data was significantly different from the second 10 years, within the 20 year window (the 616 null-hypothesis in this case was that the data was independently distributed). If the difference 617 in medians was found to be statistically significant (i.e. p-value < 0.05) and there was a 618 change in sign of the median values (e.g. switch from negative to positive), a climate shift 619 was postulated to have occurred during the 10th year of the window. The Mann-Whitney U 620 test is a robust test that does not place implicit assumptions on the underlying distribution of

Danielle Verdon-Kidd 11/9/2015 3:49 PM Deleted: phase 622 the data (i.e. it is a distribution free test), which is particularly appropriate here due to the 623 small number of years used in each window (Kundzewicz and Robson 2004). Note that a 624 number of different size windows were also tested, however this did not change the results or 625 conclusions.

626 A second test was also applied to identify step changes in the 1-day and 7-day annual maxima 627 time series known as the distribution free CUSUM with resampling (note that the test was not 628 applied to the shorter sub-daily data as longer data sets are recommended for this method). 629 CUSUM tests whether the means in two parts of a record are different (for an unknown time 630 of change). The second test was applied as it does not require the use of a moving window 631 (which is a limitation of the Mann-Whitney U test described above). However the CUSUM 632 test sequentially splits the timeseries into two potions (which are not necessarily equal). 633 which may be a problem if more than one cycle/shift is present in the timeseries.

634 The existence of serial correlation (or autocorrelation) in a time series will affect the 635 ability of tests (such as the Mann-Whitney U and CUSUM) to assess the site significance of a trend (e.g. Yu et al. 2003, Serinaldi and Kilsby 2015b). The presence 636 637 of cross-correlation among sites in a network will also influence the ability of the test 638 to evaluate the field significance of trends over the network (e.g. Yu et al. 2003, Douglas et al. 2000, Guerreiro et al. 2013). Therefore, prior to applying the change point 639 640 analysis as described above, the Durbin-Watson (DW) statistic was used to test for 641 autocorrelation in the annual maxima timeseries (Durbin and Watson (1950, 1951)). In this 642 case the null hypothesis is that the residuals from an ordinary least-squares regression are not 643 autocorrelated against the alternative that the residuals follow an AR1 process. All DW 644 statistic values were found to be greater than the 1.562 (the upper bound for 1% significance 645 and a sample size of ~100) providing no evidence to reject the null hypothesis. Therefore,

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648 any regime shifts detected using the change point methods above are not likely to be artefacts

649 <u>resulting from hidden persistence.</u>

650 The potential issue of cross- correlation was also investigated. It was found that less than 9% 651 of all possible pairings of rainfall data sets display a significant (yet weak) correlation at the 652 5% level (r >0.2, significance based on n=100). Only eight pairings (out of 4465) were 653 correlated at 0.5 or higher. It was also found that stations located more than 500km apart 654 were unlikely to be correlated and that the strength of the correlation reduced as distance 655 increased between the pairs. This is not surprising given annual maximum rainfall events are 656 due to synoptic scale processes. Therefore observations relating to spatial consistency of 657 regime shifts are unlikely to be due to spatial correlation between sites.

658 2.1 IFD Calculation

659 The standard process for obtaining IFD information for a location is to refer to the six master 660 charts of rainfall intensity for various durations and ARIs covering all of Australia in Volume 2 of AR&R 2001. Alternatively, IFD curves can be obtained for any location in Australia via 661 662 the BoM website (both the AR&R 1987 and revised IFDs are available). This information is 663 based on regionalised estimates of IFDs that are spatially and temporally consistent. 664 However, this approach cannot be adopted when using the instrumental rainfall data required 665 for the analysis presented in this study. As such, the IFD information generated for this 666 project follows the methodology on which the IFDs were based for AR&R 1999 which utilises point source data with no regionalisation. It should be noted that it is not the purpose 667 of this paper to compare different methods of generating IFDs, rather, one method has been 668 669 adopted in order to provide a comparative assessment of the impact of non stationarity on 670 IFD estimation. The AR&R 1999 procedure used to generate IFDs from raw rainfall data (i.e. 671 point based estimates) is summarised as follows:

673	• A	log-Pearson III distribution was fitted to the annual maxima timeseries using the	
674	m	ethod of moments (for annual maxima series of 30 minutes to 72 hours duration).	
675	T	is is the standard distribution that has historically been adopted for generating IFDs	
676	in	Australia; however other distributions have recently been tested as part of the	
677	re	vision of AR&R. To test if this distribution is suitable for the region being studied,	
678	th	e goodness of fit for the log-Pearson III was tested using a Kolmogorov Smirnov	
679	(K	S) test. Here the null hypothesis is that the data fits the Log-Pearson III distribution	
680	<u>(t</u>	ne alternate is that the data does not follow the Log Pearson III distribution). All p-	
681	<u>va</u>	lues were greater than 0.05 (average p-value was 0.75), for all series (30min to 72hr	
682	<u>dı</u>	rations at Brisbane, Sydney and Melbourne), therefore we accept the null	
683	<u>h</u> y	pothesis at the 5% significance level.	Danielle Verdon-Kidd 11/0/2015 4:00 PM
684	• T	ne coefficient of skewness was determined for each duration (30 minutes to 72	Deleted: It was found that distribution gave a statistically significant fit to the annual maxima
685	ho	purs);	umeseries data
686	• T]	ne coefficient of skewness was then used to obtain a frequency factor, K_{Y} , for use	
687	W	th Log-Pearson III Distribution. $K_{\rm Y}$ was obtained from Table 2.2 (positive skew	
687 688	w	ith Log-Pearson III Distribution. K_Y was obtained from Table 2.2 (positive skew efficients) and Table 2.3 (negative skew coefficients) in AR&R 1999 Book 4;	
687 688 689	w cc • Ra	ith Log-Pearson III Distribution. K_Y was obtained from Table 2.2 (positive skew efficients) and Table 2.3 (negative skew coefficients) in AR&R 1999 Book 4; infall intensities for a range of ARI were calculated using the following formula:	
687 688 689 690	w cc • R	ith Log-Pearson III Distribution. K_Y was obtained from Table 2.2 (positive skew refficients) and Table 2.3 (negative skew coefficients) in AR&R 1999 Book 4; uinfall intensities for a range of ARI were calculated using the following formula: • $\log RI_Y = M + K_Y S$ (1)	
687 688 689 690 691	w cc • R W	ith Log-Pearson III Distribution. K_Y was obtained from Table 2.2 (positive skew refficients) and Table 2.3 (negative skew coefficients) in AR&R 1999 Book 4; ainfall intensities for a range of ARI were calculated using the following formula: • $\log RI_Y = M + K_Y S$ (1) here: $RI_Y = rainfall$ intensity having an ARI of 1 in Y	
 687 688 689 690 691 692 	w cc • R W	ith Log-Pearson III Distribution. K_Y was obtained from Table 2.2 (positive skew refficients) and Table 2.3 (negative skew coefficients) in AR&R 1999 Book 4; ainfall intensities for a range of ARI were calculated using the following formula: • $\log RI_Y = M + K_Y S$ (1) here: $RI_Y = rainfall$ intensity having an ARI of 1 in Y M = mean of the logarithms of the annual maxima rainfalls	
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 687 688 689 690 691 692 693 694 	w cc • Ri W	ith Log-Pearson III Distribution. K_Y was obtained from Table 2.2 (positive skew refficients) and Table 2.3 (negative skew coefficients) in AR&R 1999 Book 4; ainfall intensities for a range of ARI were calculated using the following formula: • $\log RI_Y = M + K_YS$ (1) here: $RI_Y = rainfall$ intensity having an ARI of 1 in Y M = mean of the logarithms of the annual maxima rainfalls S= Standard deviation of the logarithms of the annual maxima rainfalls $K_Y =$ frequency factor for the required ARI of 1 in Y	
 687 688 689 690 691 692 693 694 695 	w cc • Ri W	ith Log-Pearson III Distribution. K_Y was obtained from Table 2.2 (positive skew befficients) and Table 2.3 (negative skew coefficients) in AR&R 1999 Book 4; ainfall intensities for a range of ARI were calculated using the following formula: • log RI _Y = M + K _Y S (1) here: RI _Y = rainfall intensity having an ARI of 1 in Y M = mean of the logarithms of the annual maxima rainfalls S= Standard deviation of the logarithms of the annual maxima rainfalls K _Y = frequency factor for the required ARI of 1 in Y ARIs of 2 years to 10 years were adjusted to partial-duration series estimates. In	

700	than 10 years, no corrected factor is required): 2 year ARI - 1.13, 5 year ARI -
701	1.04, 10 year ARI – 1.0.

702	It should be noted that this approach is likely to result in different estimates of IFDs than
703	those obtained from the standard maps provided by AR&R 1999 or the revised IFD
704	estimates released in 2013. Here we are using point based rainfall data, whereas AR&R
705	1999 have derived regionalised estimates based on multiple rainfall stations with varying
706	lengths of data, varying resolution (daily and pluviograph) and varying quality of records.
707	It is recognised that analysis of rainfall data from single stations is often unreliable, is not
708	temporally or spatially consistent and should generally not be used for design purposes.
709	However, the use of point based rainfall data satisfies the specific aims of this study
710	(which is a comparative analysis) and is therefore considered appropriate.

711 3. Results

712 Test for <u>regime shifts</u> in the annual maxima rainfall timeseries 3.1 713 Significant step changes identified in the extreme rainfall timeseries are shown in Figure 2. 714 Of the 66 sub-daily rainfall stations tested, 40 (61%) displayed at least one step change in the 715 1-hour annual maxima timeseries (Figure 2a), with some stations exhibiting multiple shifts. 716 Of the 96 daily rainfall stations tested, 86 displayed at least one step change in the 1-day 717 annual maxima timeseries (Figure 2b), while 92 exhibited at least one shift in the 7-day 718 annual maxima timeseries (Figure 2c), and some stations exhibited multiple shifts. Figure 2 719 collectively shows that observed step changes (or regime shifts) in annual maxima rainfall are 720 not confined to any one particular region of Australia, with most stations analysed exhibiting 721 at least one statistically significant shift.

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725 As shown in Figure 3, the CUSUM test yielded fewer stations with statistically significant 726 step change in the annual maxima timeseries (only 18 stations out of 96) and many of these 727 were clustered along the coastal fringe of eastern Australia (note that, although the total number of stations displaying significant change points was the same for both the 1-day and 728 729 7-day annual maxima, in some cases the location of the stations differed between the two). 730 However, as stated previously a limitation to this method is that only one significant change 731 can be detected using the CUSUM test (given that the data is sequentially split into two portions during testing). This can be a problem if more than one step change or cycle in the 732 733 data is present (see example timeseries in Figure 4). Therefore, while the number of stations 734 displaying a step change is reduced using the alternative method, the results do in fact support 735 the theory that regime shift(s) in the annual timeseries are, present for some stations at 736 different durations.

737 ****Figure 3 about here****

738 ****Figure 4 about here****

The temporal consistency of step changes in the annual maxima timeseries was further
investigated (Figure 5a) and it was found that <u>the timing of observed shifts</u> were not
necessarily consistent across Australia. However, for some regions (e.g. the east coast of
Australia) periods such as the 1940s (Figure 5b) and to a lesser degree 1970's (Figure 5c)
display a higher degree of spatial consistency,

744 ****Figure 5 about here****

745 Instability and storminess can result during periods when a number of climate driving 746 mechanisms interact (e.g. El Niño/Southern Oscillation, Indian Ocean Dipole and the 747 Southern Annular Mode) to influence the occurrence of regional weather systems such as east Deleted: non-stationarity

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Luke Kidd 14/9/2015 7:48 PM **Deleted:**, particular along the east coast (Figure 5b). coast lows and cut off lows (Pook et al. 2006, Verdon-Kidd and Kiem 2009). However, the large-scale climate phenomena impact various regions of Australia at different times of the year and to varying degrees, therefore it is not surprising that the timing of shifts in the annual maxima timeseries varies spatially and temporally. This highlights the limitations of trying to assess and attribute variability in annual maxima rainfall based on a single climate driver (e.g. ENSO) or attempting to address the issue of <u>climate trends</u> for the whole of Australia using one simple approach or model.

763 3.2 Effect of non-stationarity on IFD estimation

764 Section 3.1 provided evidence of non-stationarity in the annual maxima timeseries for a range 765 of durations. This non-stationarity may ultimately influence the IFD estimation depending on 766 the length of data, and the time period it comes from, and therefore the underlying climatic state (or combination of states). Current IFD estimates for Australia (both the 1987 and 2013 767 768 versions) are based on data as short as 30 years for the daily-read stations and 8 years for the 769 sub-daily data. Therefore IFD estimates based on relatively short-term data sets may under-770 or over-estimate rainfall intensities, depending on where the data series fits within the long 771 term context (i.e. before or after a shift in annual maxima).

772 For many east coast stations a shift in 1-day annual maxima (along with the 7-day) occurred 773 around the 1940s - 1950s and again in the 1970s. This timing also corresponds to well-known periods of change in the IPO (see Section 2.1.2 for a description of the IPO and its 774 775 influences). Therefore, to further explore the issue of regime shifts, breakpoints in the IPO 776 were used to stratify the annual maxima rainfall timeseries into IPO positive and negative 777 epochs for the three long sub-daily data sets described in Section 2.1.1 (i.e. Brisbane, Sydney 778 and Melbourne, see Figure 1a for location). The reason for selection of these stations was 779 twofold. Firstly, for all three stations, a shift in the annual maxima timeseries (for 1-day and Danielle Verdon-Kidd 11/9/2015 3:54 PM Deleted: non-stationarity

Danielle Verdon-Kidd 11/9/2015 3:54 PM Deleted: non-stationarity 782 7-day) was observed during the 1940s and again in the 1970s, and secondly the stations 783 contain long records of pluviograph data (the shortest being from 1913 onwards). Figure 6a 784 shows the modulating effect of the IPO on total annual rainfall for the three east coast 785 stations. Annual maxima at the three east coast stations during the two IPO epochs are also 786 shown in Figure 6 (b-d) for event durations of 30 minutes to 72 hours (durations that are 787 critical for flood design applications). A two-sample Kolmogorov-Smirnov (KS) test was 788 applied to determine if the observed differences between the IPO positive and negative 789 rainfall distributions are statistically significant. Here the null hypothesis is that the two 790 samples are drawn from the same distribution.

791 ****Figure 6 about here****

792 It is evident from Figure 6a that the effect of the IPO on annual rainfall totals (as measured by 793 the largest difference between the two rainfall distributions associated with each climate 794 phase and the results of the KS test) is greatest for Sydney. Although there does appear to be 795 some impact in Brisbane, the result was not statistically significant according to the KS test. 796 Melbourne does not appear to be greatly influenced by the IPO in terms of annual rainfall 797 variability. This is due to the fact that the southern regions of Australia are affected by other 798 climate modes than those arising from the Pacific (i.e. the Southern Annular Mode and the 799 Indian Ocean Dipole (e.g. Kiem and Verdon-Kidd 2010, Gallant et al, 2012)). Regions such 800 as Brisbane and Sydney tend to be dominated by Pacific Ocean influences (e.g. Verdon et al. 801 2004). Figure 6b shows annual maxima rainfall tends to be higher during IPO negative on 802 average for durations 6 hours and longer at Brisbane (though not statistically significant 803 according to the KS test), while Figure 6c confirms the same to be true for Sydney for durations 2 hours and longer (statistically significant at 95%). However, for Sydney, the 804 805 outliers (represented by circles) tend to be larger during IPO positive, indicating that the less frequent events might be more intense during this phase, 806



Luke Kidd 17/9/2015 10:32 PM Deleted: using a Wilcoxon Rank Sum test Luke Kidd 17/9/2015 10:32 PM

Deleted: Interestingly, it was found that IPO positive periods have resulted in higher annual maxima for durations 30 minutes and 2 hours at Brisbane. This result is consistent with Yilmaz et al (2014) who also found that the IPO could be responsible for higher rainfall intensities for long durations and high return periods.

819 Irrespective of the fact that the annual rainfall totals for Melbourne do not show any 820 significant difference between the two phases of the IPO, there does appear to be a consistent 821 relationship between IPO and the sub-daily and daily statistics (Figure 6d), whereby the 822 median of the IPO positive distribution is higher across all durations, however IPO negative is associated with less frequent but more extreme events (although results were not 823 824 statistically significant based on the KS test). For events 24 hours and longer, the IPO 825 negative distribution also shows a much higher degree of variability than smaller durations, 826 with the 'box and whiskers' extending beyond the IPO positive counterpart for these longer 827 durations. This suggests that while IPO might not be as dominant in southeastern Australia as 828 it is further to the north it still has some influence that needs to be better understood.

829 Based on the analysis presented in Figure 6 and the results of the KS test, the Sydney record 830 was chosen to further investigate the effects of regime shifts on IFD estimation. IFD 831 information was generated for the Sydney record using rainfall data from the two IPO phases 832 and the methodology outlined in Section 2.1 for durations 6 minutes through to 72 hours and 833 ARI 2 years to 200 years. In order to test the robustness of the point estimates of rainfall 834 return levels and estimate the uncertainty in their calculation, a simple bootstrap procedure 835 was carried out. Firstly the IPO positive and IPO negative rainfall timeseries were resampled 836 with replacement to obtain two new B-samples. Then for each B-sample the log-Pearson III 837 distribution was fitted and the rainfall intensities calculated for the various return intervals. 838 The difference between the rainfall intensities (of the two B-samples) was then calculated. 839 This procedure was repeated 100 times to build the empirical distribution of the differences 840 (which represents the effects of sampling and parameter estimation uncertainties under the 841 hypothesis of the existence of two different regimes).

Figure 7 shows the difference in rainfall intensity between IPO positive and IPO negative
estimates, along with the 95% confidence intervals (CIs) derived using the procedure above.

Luke Kidd 17/9/2015 10:34 PM Deleted: 30 Luke Kidd 17/9/2015 10:34 PM Deleted: 100 847 Figure 7 demonstrates clear differences in the resulting rainfall intensities for Sydney 848 estimated for each duration and ARI using the two regimes (i.e. rainfall data from either IPO 849 negative or IPO positive). The difference in rainfall intensity estimated is as great as 65% in some cases, In all cases, the magnitude of the difference in rainfall intensity estimated using 850 851 the different data regimes is greater for less frequent events (e.g. 50-year, 100-year, 200-year 852 ARIs), highlighting that uncertainty is greatest with less frequent events. The rainfall 853 intensity is greater in IPO positive for the very short duration events (6 minutes) at all return 854 intervals and for 30min duration events for return intervals of 10 years or more. Similarly, for 855 the 24 and 72 hour duration events rainfall intensity in the positive IPO phase is higher for return intervals of 5 years or more. For 2 hour and 6 hour events, the negative phase results in 856 857 higher intensity events for more frequent return levels (20 years or less) but lower intensities 858 for less frequent events (50 years or more)

859 4. Discussion and conclusions

860 An analysis of regime shifts in the annual maxima timeseries (1-hour, 1-day and 7-day) has 861 been carried out using a set of high quality rainfall stations across Australia. It was found that the annual maxima timeseries does indeed exhibit statistically significant step changes/shifts, 862 for the majority of stations for various durations. Further it was demonstrated using three 863 long term sub-daily rainfall stations along the east coast that this impacts upon the resulting 864 865 IFD estimation. The potential for Pacific Influences (i.e. the IPO) to influence the resulting 866 IFD estimation was explored in order to demonstrate this issue. The authors acknowledge that 867 the IPO is unlikely to be the only driver of variability in the annual maxima timeseries across 868 Australia, and it is recommended that future research should aim to identify other potential 869 drivers of this variability.

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Deleted: The results presented in Figure 7 also highlight the fact that it cannot be assumed that one particular climate state (e.g. IPO negative) will always result in a higher rainfall intensity estimate for all durations and all ARI for all regions in Australia

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Deleted: For example rainfall intensity for the 30 minute events was estimated to be higher in IPO positive for both Brisbane and Sydney, yet the reverse is true for events of duration 2 hours and 6 hours. Similarly, all events (other than 72 hours) with a 2-year ARI are associated with a higher rainfall intensity estimate in IPO positive for Melbourne, however the reverse is true for the less frequent events. This is consistent with the observations made from Figure 6d that the median annual maxima of the IPO positive distribution is higher (for Melbourne) across the range of durations, however the IPO negative periods were associated with less common but much more extreme events. Deleted: non-stationarity Danielle Verdon-Kidd 11/9/2015 3:55 PM Deleted: of lle Verdon-Kidd 11/9/2015 3:55 PM Deleted: non-stationarity in Danielle Verdon-Kidd 11/9/2015 3:55 PM Deleted: exist

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904 These findings highlight the fact that in some instances the IFD estimates currently being 905 used are likely to be either under- or over-estimated at any one time depending on the length 906 of data, and climatic state, from which they were derived. This is a particular concern given 907 that current regionalised IFD information is based on data of varying length (as short as 8 908 year in the case of sub-daily data) spanning different time periods. An over estimation of 909 rainfall intensity for a given duration could impact on construction costs, while the risks of 910 underestimating rainfall intensities could result in failure of design criteria. That is, the risk is 911 dependent on the application and length of time over which the risk is assessed.

912 Further revisions of AR&R are currently underway to include an assessment of the potential 913 impacts of climate change on IFD estimates. However, there are many uncertainties 914 associated with climate change projections, particularly when extracting information on 915 timescales shorter than a season and particularly for hydrological extremes (e.g. Blöschl and 916 Montanari 2010, Kiem and Verdon-Kidd 2011, Koutsoyiannis et al. 2008, 2009, Montanari et 917 al. 2010, Randall et al. 2007, Stainforth et al. 2007, Stephens et al. 2012, Verdon-Kidd and 918 Kiem 2010). Therefore, assessing future changes in extreme events that occur over short 919 durations (e.g. minutes to days) is inherently difficult. Furthermore, climate projections are 920 presented in terms of a percent change from a particular baseline. However, the baseline is 921 often inconsistent and ill-defined leading to very different estimates of risk depending on the 922 time over which the baseline is calculated (as has been demonstrated in this paper). 923 Importantly, for regions where large-scale climate drivers operate on a multi-year to multi-924 decadal timescales and are known to influence extreme rainfall events, we can use this 925 information to determine if the climate statistics on which the IFD are based are likely to be 926 biased or missing crucial information,

927 It is recommended that <u>regime shifts</u> in annual maxima rainfall be considered and 928 appropriately treated in any further updates of AR&R. One way to do this may be to only Luke Kidd 17/9/2015 11:17 PM Formatted: Font:(Default) Times New Roman

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Deleted: The research presented here demonstrates that information currently available on natural variability is extremely useful in providing a better assessment of the true risk profile and can act as a guide to the baseline on which climate change projections may be applied

Danielle Verdon-Kidd 11/9/2015 3:57 PM Deleted: non-stationarity 936 utilise data sets of similar length ensuring that they span a sufficient number of years in order 937 to capture data from epochs of both high or low annual maxima (to remove bias towards one 938 climatic phase or another). However, it is acknowledged that this would potentially result in 939 discarding a large amount of data. alternatively, a separate set of IFDs could be developed for 940 use in high risk modelling for engineers who need to account for the 'worst case' (in a similar 941 manner to climate change allowances). This second set of IFD could be developed based on 942 the periods of elevated annual maxima alone (for those stations with clearly defined epochs 943 of annual maxima) such that if we were to enter such an epoch, designs based on these 944 estimates would be robust for the duration of such a period. Salas and Obeysekera (2014) 945 provide similar recommendations to deal with changing exceedence probabilities over time . 946 This would have to be assessed and calculated on a region by region basis given that 947 Australia is a country associated with high spatial and temporal rainfall variability caused by 948 numerous large-scale climate drivers and regional weather phenomena. Finally, any revised 949 estimates of annual maxima should be compared in terms of uncertainty bounds (e.g. 950 following Koutsoyiannis (2006)). Uncertainty analysis, which takes into account both the 951 data availability and variability within the observation period would provide relevant 952 information to practitioners about the reliability of IFD estimates.

953 This study has highlighted the existence of regime shifts in annual maxima rainfall data in 954 Australia. The driving mechanisms of these regime shifts are likely to vary from location to 955 location and decade to decade. However, these shifts are typical of many natural phenomena 956 and can be described by processes characterized by long range dependence (or regime-957 switching processes) and captured by hidden Markov models (or similar), resulting in a 958 mixture of distributions that alternate stochastically according to the transition probability 959 from one regime to the next (e.g. Serinaldi and Kilsby, 2015a). While the strategy of defining 960 IFDs for two (or more) different regimes (e.g Serinaldi andKilsby (2015a)) currently only

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Deleted: Based on the results of this study, and literature cited within this paper, we emphasise that there undoubtedly is are non-stationaritystep changes in historical short duration rainfall extremes but the characteristics and causes of this non-stationarityvariability vary from location to location and decade to decade – something which must be considered and accounted for when attempting to estimate IFD design rainfalls and prior to quantifying how those IFD estimates might change in the future

973	partially solves the problem, as we often do not know the beginning or the end of a specific
974	regime (be it rainfall or climate driver), recent work has focused on optimizing designs and
975	planning strategies based on the range of what is plausible rather than a reliance on knowing
976	the current and future climate state (e.g. Mortazavi-Naeini et al., 2015). At the same time,
977	work is also underway on seamless prediction at a range of timescales and if/when this
978	eventuates the results discussed here become even more important/useful. Nevertheless, the
979	immediate usefulness of the insights presented here occurs when first establishing the IFD, as
980	an approach similar to that employed here can be used to determine if the underlying data are
981	biased to a mostly wet or mostly dry regime (or a mix of both) which then provides an
982	indication as to whether the IFD is likely to be an over- or underestimate of the true risk.
983	Importantly, this issue needs to be considered and accounted for when attempting to estimate
984	IFD design rainfalls and prior to quantifying how those IFD estimates might change in both
985	the near and long-term future,

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While the analysis presented here has been conducted using rainfall data from Australia alone, the recommendations provided are likely to be applicable for other regions of the world where IFD information is based on short term records and particularly for locations with a highly variable climate.

991 5. Acknowledgements

The authors wish to acknowledge the Australian BoM for supplying the rainfall data used in this study and the UK Meteorological Office for kindly making the IPO data available. We would also like to thank Mr Andrew Magee for assisting with statistical analysis of the rainfall data and the two reviewers of the paper who provided feedback that greatly improved the paper.

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1135 Note the three long term sub-daily stations used in the IFD analysis are also labelled.





1137 hour, b) 1-day, c) 7-day annual maximum rainfall (using the Mann-Whitney U test)







and b) 7-day annual maximum rainfall (using the CUSUM test with resampling)



- 1142 Figure 4 Example of inadequate identification of non-stationarity using CUSUM test (red line
- 1143 highlights three distinct epochs of high/low rainfall, while green line demonstrates effect of
- 1144 splitting the data into two sections for CUSUM test)





Figure 5 a) number of stations each decade displaying evidence of a step change in 1-day annual max, b) Stations (in red) with at least one statistically significant step change in the 1day annual max during 1940-1950 (using the Mann-Whitney U test), c) Stations (in red) with at least one statistically significant step change in the 1-day annual max during 1970-1980 (using the Mann-Whitney U test)



1151 Figure 6 Relationship between IPO and a) total annual rainfall, and annual maximum rainfall





1155 Figure 7 Difference in rainfall intensity for each duration and ARI. Positive (negative) values

1156 represent an increase (decrease) in rainfall intensity during IPO positive compared to IPO

1157 negative

