1 **Response to the Reviewers:**

2 Reviewer 1:

3 Thank you very much for the detailed review and comments to improve this work. We will4 update the manuscript with the following responses.

5 The responses to questions 1a,c will be added in the model description chapter, section 2.1 as6 additional paragraphs.

7

8

9

1. Does the work apply novel numerical/conceptual techniques?

10 (a) New numerical techniques (e.g. polyhedral mesh) and their advantage

11

The advantages of the numerical method are addressed by Kernkamp (Kernkamp et al., 2010). In this paper, for the first time a real-world case is run with the new model. With a personal computer it is possible to address challenging morphologies combining rectangular and triangular mesh. The advantage of this approach is the reduced number of cell since it is not compulsory for the entire mesh to be triangular, hence less cells and lower computational time.

18

Another advantage is the possibility of directly converting Delft3D grids and settings to DFlow-FM. Delft3D is commonly used models for coastal engineering applications, and DFlow-FM will allow coupling existing Delft3D models with inland models and complex geometry estuaries (van der Wegen and Roelvink, 2012; Guo et al., 2015; Dastgheib et al., 2012; Roelvink, 2006; Lesser et al., 2004).

24

(b) novel algorithms for fast or parallel processing (please note that a computational
focused paper would be more suited for a journal like Computers Geoscience rather
than HESS)

The algorithm is developed by a partner group, and the publication of computational
 developments and new schemes is more suitable to them than us. The publication about the
 grid and solvers is published by Kernkamp (Kernkamp et al., 2010).

4

5 (c) novel transport equations and couplings.

6

7 The transport equations are the same one used in many other models, the advective-diffusion
8 equation, and for the cohesive sediment Krone-Parteniades.

9

10 The novelty here is that the flexible mesh model is directly coupled with water quality, 11 sediment transport and habitat model. In this sense the exchange of model input/output 12 between the hydrodynamics, sediment and ecology model is facilitated. For habitat model 13 having SSC as input improves results of light attenuation therefore better primary production 14 and habitat definition.

- 15
- 16 -----

17 The responses to question 2 will be added in the end of the introduction on page 4. We agree18 that the choice of the specific site was not discussed enough.

19

20

2. What insight can be obtained from choosing this specific field site?

21

22 (a) Is the field site characteristic end member case of transport?

23

The Sacramento San-Joaquin Delta is a typical case of a highly impacted estuary. Being able to numerically simulate and determine sediment transport, budget and turbidity levels in this type of environment open possibilities to better informed political, ecological and management decisions including how to respond to climate change and sea level rise. This type of model is an important management tool that is applicable to other impacted estuaries such as Guanabara Bay (Rio de Janeiro, Brazil) and the Hudson estuary (New York, USA).

1 (b) Is the field site very well suited for calibration/model output comparison due to high quality2 available flow/topographic data?

3

As you already posed, the field site is very well suited for calibration. The Delta and Bay has a big survey network. There is a big data base of freely available data on river stage, discharge and suspended sediment concentration among other parameters, maintained by USGS. The continuous sediment in suspension measurement stations are periodically calibrated by water collection in situ, filtered and weighted in the laboratory.

9

10 The Bay-Delta system has high resolution bathymetry (10m) for all the channels and bays.

11

(c) Is the field site of specific importance, especially with respect to the ecological focus of the paper's title.

14

15 The focus of the paper is to improve the connection between physical and ecological 16 numerical models. In this sense the ecological importance of the Delta is discussed below.

Starting from the bottom of the food web, the Delta is the most important area for primary production in the San Francisco Estuary. The Delta is one order of magnitude more productive than the rest of the estuary (Jassby et al., 2002;Kimmerer, 2004).

20 The Delta is an area for spawning, breeding and feeding for many endemic species of fishes

and invertebrates, including some endangered species like delta smelt (Brown et al., 2013),
Chinook salmon, spring run salmon and steelhead.

Several projects for marshes restoration in the Delta are planned and the success of theseprojects depends on sediment availability (Brown, 2003).

25

What general conclusions can be drawn from the analysis besides comparing (calibrated) simulation with field observations?

28

Here we present the first real case application of the flexible mesh model DFlow-FM. The good agreement with data and reproducibility of the main events gives us confidence in the model. The mesh flexibility makes a perfect model for estuaries and complex geometry areas. 1 DFlow-FM was developed with the possibility of a direct coupling with ecological and water

quality models speeding the knowledge interchange between the two areas (physical andbiological).

4

- 5 Sediment is a key-factor to estuaries water quality and ecology. The already calibrated model
- 6 generates high quality input for the ecological models and is ready for forecast scenarios.
- 7
- 8
- 9 Brown, L., Bennett, W., Wagner, R. W., Morgan-King, T., Knowles, N., Feyrer, F.,

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- 36 Water Resources Research, 41, W09428, 10.1029/2004wr003753, 2005.
- 37

2 Answering the more specific comments:

3

4 P2, Line 24- P3, Line 16 What do the anthropogenic impacts mentioned here have to do
5 with the presented results? Please clarify how this paragraph contributes to the
6 understanding of context of the presented analysis.

7

8 The idea behind these 2 paragraphs is to put the work in a broader context, in terms of world 9 sediment budget (P2 line 24) and the importance of numerical models to forecast scenarios 10 because the conditions and forcing are constantly changing.

11

P2, Line 6: "A robust sediment model ..." Do the authors mean sediment transport?
Please clarify. Also 'chain of models' as the authors describe several.

14

15 Corrected to: "A robust sediment transport model is the first step towards a chain of 16 models...".

17 Regarding the explanation about the chain of models, we will modify P2 Line6-7 to:

18 "The chain of model means that each model provides input to the next model. For example 19 the sediment transport model output is the input to contaminants, phytoplankton and habitat 20 numerical models."

21

22 P2, Line 20: "...are subject..." should be "are subjected"

23

24 The sentence was corrected in the manuscript.

25

P4, Line 3-7: Please explain shortly what the "2DH process based model" physically is?
I suppose it integrates height averaged Saint-Vernant type equations, together with
some transport equations for sediment as mentioned three pages later. However a

(short) physical explanation should be given when the modeling framework is 1 2 introduced in the text. For example adding a phrase like "...solves the 2d height integrated shallow water equations coupled with advective diffusive transport ...'' would 3 help a lot to understand the physics behind the 2DH process based model. 4 5 We will include the sentence in P4, Line 3-7. 6 7 "The 2DH model solves the 2D vertical integrated shallow water equations coupled with 8 advective-diffusive transport." 9 P4 Line 6 "sediment budgets [..] in time (days)" seems to contradict line 8 "... yearly 10 11 sediment budget ...". Please clarify. 12 In P4 Line 6 we were aiming to explain the several time scales, as we present in the results the 13 14 yearly budget and seasonal analysis the sentence will be rephrased as: 15 " This process-based model will be able to quantify high resolution sediment budgets and SSC, both in time (~ monthly/yearly) and space (~10s-100s of m)." 16 17 18 P6 Line 8: Pumping keeps salinity constant. How does this justify the height integrated modeling approach? The (possible) justification for this, (limited saltwater-freshwater 19 interaction in the Delta) is given only 2 pages later. The authors should justify their 20 21 statements at the point where these are made rather than assuming that the reader has 22 already advanced several paragraphs in the text. 23 24 In this case we assumed that it is clearer to remove the sentence "allowing the 2-DH modeling approach." from P6 line 8. 25 And modify P7 line 10 as: 26 "We assume that the main flow dynamics in the Delta are 2D with no vertical stratification. 27

28 This assumption is supported by the lack of salt-fresh water interactions in the Delta due to

the pumping operations. We also assume that temperature differences do not govern flow
 characteristics."

3

4 **P10** Line 10-15: Please explain abbreviations at first usage.

5

6 All the abbreviations in these lines are the stations names. They will be written as:

7 " The river water flow hourly input data are from the following stations, at Sacramento River 8 at Freeport (FPT), San Joaquin River near Vernalis (VNS) and Yolo Bypass (YOLO) were 9 obtained from California Data Exchange Center website (cdec.water.ca.gov/) (Fig 2). The sediment input data, for both input stations FPT and VNS, and calibration stations S 10 11 Mokelumne R(SMR), N Mokelumne R (NMR), Rio Vista (RVB), Mokelumne (MOK), Little 12 Potato Slough (LPS), Middle River (MDM), Stockton (STK), Mallard Island (MAL) (Fig 2), was obtained by personal communication from USGS Sacramento; this data is part of a 13 14 monitoring program (http://sfbay.wr.usgs.gov). Since 1998, USGS has continuous measuring stations for sediment concentration which is derived from optical backscatter sensors (OBS) 15 16 measurements every 15 minutes, and nearly monthly calibrated with bottle samples (Wright 17 and Schoellhamer, 2005)."

18

19 Equation 3 and 4 are confusing. The index over which the sum runs is "i" and the 20 summation is from "i-1" to "N". I suppose it should be "i=1" under the sum.

21

22 You are right, there was a typo error. I was already corrected.

23

Fig. 2 The labels and the location points of the calibration stations are too small. Please
increase the font size.

26

27 Figure updated.

1	Fig. 3: It is hardly visible that the blue line is dashed. The authors may increase the dash			
2	spacing or simply plot a blue line.			
3				
4	Figure updated.			
5				
6	Fig.4: Same problem with red dashed line as in Fig. 3. Increasing the dash spacing and			
7	plotting the dashed line on top of the solid line may also improve the visibility where			
8	both lines overlap considerably.			
9				
10	Figure updated.			
11				
12	Fig.8: The 3D flow effects mentioned in the caption are not discussed in the main text.			
13	Please update the manuscript accordingly.			
14				
15	I will add a line in the manuscript as follow:			
16	"Seawards from MAL, figure 8 shows preliminary sediment flux for the bay in dashed line,			
17	because we don't have confidence on them. The model here presented is 2DH, seawards from			
18	MAL stratification takes place in the water column due to salinity intrusion, meaning that 3D			
19	effects become important and they are not capture by a 2DH model. "			
20				
21	Figure A1: The figure is too small			
22	Figure updated.			
23	The manuscript will be updated accordingly. We hope to have properly addressed your			
24	comments.			
25				

1 Reviewer 2:

- 2 Thank you very much for the detailed review and comments to improve this work. We will
- 3 update the manuscript with the following responses.
- 4 The responses to questions 1 will be added in the model description chapter, section 2.1 as
- 5 additional paragraphs.
- 6

7 1) Calibration/validation of the numerical model. I understand that the technical details on the numerical method and validation/comparison with other schemes have
9 been given in Kernkamp et al. (2010), but it would be helpful to comment on alternative methods for this specific site.

11

Are there other numerical models done for the Delta, e.g. just the hydrodynamic part? For example, Delft3D in the depth-averaged mode seems to be a close competitor which handles both hydrodynamic and morphodynamic transport. A few comments would be nice, such as:

a) how many grid cells are involved in the calculation of the model setup presented in
 this work, versus how many grid cells would have to be used for the total triangular
 mesh?

19

b) How much time would it take for other methods to reach the same level of accuracy with the same computation power (only one desktop)?

22

23

Structured grid models as Delft3D and ROMS (Regional Oceanic Modeling System) have been widely used and accepted in estuarine hydrodynamics and morphodynamics modeling including San Francisco Estuary (Ganju and Schoellhamer, 2009;Ganju et al., 2009;van der Wegen et al., 2011). In all theses cases of San Francisco estuary modeling efforts, the Delta was schematized as 2 long channels since the grid is not flexible to have a 2D modeling of the rivers, channels and flooded island of the system.

30

There are three widely known unstructured mesh models the TELEMAC-MASCARET (Hervouet, 2007), the UnTRIM (Casulli and Walters, 2000;Bever and MacWilliams, 2013) and D-Flow FM (Kernkamp et al., 2010). The two first models are purely triangle based and are not coupled with sediment transport and/or water quality and ecology model.

DFlow-FM allows for a combination of curvilinear grid and triangles, in the present model DFlow-FM grid has 63.844 cells, from which about 80% are rectangles. In the case of only triangle grid all the rectangles would be 2 triangles summing up to 114.920 grid cells. Not

triangle grid all the rectangles would be 2 triangles summing up to 114.920 grid cells. Notcounting the triangular grid orthogonality and circumscription issues, in the case of entirely

39 triangular grid the running time for a 1 year simulation would increase from ~72 clock hours

40 to \sim 192 hours.

- 1
- 2 -----

3 The responses to questions 2 will be added partly in the discussion description chapter,

- 4 section 4.2 as additional paragraphs, and partly in the conclusion (chapter 5).
- R: (2) Insights of the specific field site. As the authors stated, the Delta provides
 critical habitats for ecological purposes. I can see how this model and its predictions
 become handy for a chain of models.
- 8
- 9 R: Could the author use the model results to inform more into the monitoring and data collecting practice?
- 11
- R: For example, what data set should be collected to help the model make better predictions? And where could be a good location for building a new monitoring
- 14 station?
- Addressing issues like these, or even in a discussion of future work, will expand the scope of this work and gives a higher scientific impact.
- 17
- 18

19 A: The Delta is already very well measured in terms of observation stations. However, this 20 work show that the substantial sediment is exported trough the pumping stations at the 21 Southern Delta where no data in SSC is available.

- The connections between Sacramento and San-Joaquin River (DCC and GLS) are currently is
 being surveyed.
- 24

The present model opens the possibility for forecast and operational modeling. Forecasting the time frame of high levels of SSC (turbidity) allows planning of measurements campaigns for ecologists, as well as the possibility of tracking potentially contaminated sediment and be able to make a contingency plan. The forecast could also be a tool to guide management decisions concerning seasonal barriers as well as pumping operations.

30

The 2D model is a step towards the 3D model. The 2D model allows faster runs facilitating sediment parameter definition and understanding of sediment dynamics in the Delta. However, for the fully Delta-Bay coupling it is necessary fully 3D model, which is the next step of this work.

35

36 -----

37 The responses to questions 3 will be added in conclusion chapter (chapter 5) as additional

38 paragraphs.

- 1 R: (3) Information on suitable systems for model application.
- 2

R: The Delta is a relatively well-surveyed region. What should be taken into account
for potential users of this model, when applying to a different region with less
available data?

- 6
- 7

A: As a well surveyed area the Delta offers the chance of testing how much data it is necessary to define sediment budget. This work shows with simple sediment settings as one fraction at the input boundary and simple distribution of bed sediment availability, it is possible to reproduce seasonal variations as well as define yearly sediment budget with more than 90% of accuracy. It shows also that it is extremely important to have discharge and SSC measurements at least in the input boundaries and close to the system output in order to be able to calibrate the model settings applied for hydrodynamics and suspended sediment.

15

16

17 R: Also, it seems that the configuration of the channel network does play a role 18 determining the SSC level (the effects of closing/opening DCC). Would the authors 19 give more insights into how changing the configuration of the network affects the 20 distribution of fluxes?

21

A: The channel network influence in sediment budget and deposition is the topic of the workin development right now.

24

As for from this work results, we note that the Sacramento San-Joaquin River connecting channels DCC and GLS are important bridges to export sediment from Sacramento to Eastern Delta. On the other hand the smaller channels of the network play a minor role in the Delta sediment budget, since the discharges in these channels are considerably smaller than in the rivers.

- 30
- 31
- 51
- 32
- 33
- 34 Below are some minor comments:
- R: Page 1514, Eq. (2): "M" is also used later as the subscript for "Model".
- 36 A: We changed the from M to m in the other equations
- 37
- R: Page 1515, Line 3: Winterwerp (2006) is missing in the reference list.
- 39

1 A: It will be included 2 3 (Winterwerp et al., 2006) 4 Winterwerp, J. C., Manning, A. J., Martens, C., de Mulder, T., and Vanlede, J.: A heuristic 5 formula for turbulence-induced flocculation of cohesive sediment, Estuarine, Coastal and 6 Shelf Science, 68, 195-207, 10.1016/j.ecss.2006.02.003, 2006. 7 8 9 R: Page 1515, Line 5: sentence is not complete, "the second term in equation (Eq. 1): is close to zero?" 10 11 12 A: Thanks for the remark, already added. It is indeed close to zero. 13 14 R: Page 1515, Line 20-25: will changing the 5m threshold for mud/sand affect the 15 results? 16 17 A: The 5 meter threshold is not fixed, what we observed is a big change when considering 18 available mud in the entire channel as presented in session 3.4. 19 We did some test varying the 5m threshold. From 3 to 10 meters the final results are quite 20 similar to the one showed. Considering no mud availability also does not disturb the final 21 results for more than 10 days. However, considering mud availability in the channels deeper 22 than 10 meters starts to disturb the SSC levels. 23 24 I will include this analysis in the session 3.4 and as discussion as well. 25 26 R: Page 1519, Line 15-20: uRMSe given by Eq.(3) cannot give negative values. 27 Seems that it needs to be multiplied by the sign of the difference between modeled and observed SD (see Eq. (7) in Bever and MacWilliams, 2013). 28 29 30 A: Yes, you are right, I forget the signal difference. Proper changes will be done in figure 6 31 and analysis. 32 33 R: Page 1519, Line 23: "Willmott, 1981" is missing in the reference list. 34 35 A: It will be included 36 (Willmott, 1981) 37 Willmott, C. J.: ON THE VALIDATION OF MODELS, Physical Geography, 2, 184-194, 38 10.1080/02723646.1981.10642213, 1981.

1			
2 3	R: Page 1524, Eq.(7): Why is the term "[A'C'][U]" eliminated?		
4			
5 6 7	A: We use the same methodology as suggested by (Morgan-King and Schoellhamer, 2013). The combination of the advective, dispersive and Stokes Drift already account for more than 97% of the flux.		
8			
9			
10 11 12	R: Page 1542, Fig. 5: The values of the parameters are vague by presenting them as multiples of the standard run (e.g., w*0.5, E*100). I personally prefer the form in Fig. 6 with absolute values.		
13			
14	A: Figure 5 x axis will be updated according with the axis presented in figure 6.		
15			
16			
17	R: Page 1544, Fig. 7: the black rectangle is missing in the left panel.		
18			
19	A: Figure 7 updated.		
20			
21			
22			
23	References:		
24 25 26	Bever, A. J., and MacWilliams, M. L.: Simulating sediment transport processes in San Pablo Bay using coupled hydrodynamic, wave, and sediment transport models, Marine Geology, 345, 235-253, <u>http://dx.doi.org/10.1016/j.margeo.2013.06.012</u> , 2013.		
27 28 29	Casulli, V., and Walters, R. A.: An unstructured grid, three-dimensional model based on the shallow water equations, International Journal for Numerical Methods in Fluids, 32, 331-348, 10.1002/(sici)1097-0363(20000215)32:3<331::aid-fld941>3.0.co;2-c, 2000.		
30 31 32	Ganju, N. K., and Schoellhamer, D. H.: Calibration of an estuarine sediment transport model to sediment fluxes as an intermediate step for simulation of geomorphic evolution, Continental Shelf Research, 29, 148-158, 10.1016/j.csr.2007.09.005, 2009.		
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36 37	Hervouet, JM.: in: Hydrodynamics of Free Surface Flows, John Wiley & Sons, Ltd, 1-360, 2007.		

- 1 Morgan-King, T., and Schoellhamer, D.: Suspended-Sediment Flux and Retention in a
- 2 Backwater Tidal Slough Complex near the Landward Boundary of an Estuary, Estuaries and
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 of decadal deposition patterns in San Pablo Bay, California, 1856–1887, Journal of
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- 7 Willmott, C. J.: ON THE VALIDATION OF MODELS, Physical Geography, 2, 184-194,
 8 10.1080/02723646.1981.10642213, 1981.
- 9 Winterwerp, J. C., Manning, A. J., Martens, C., de Mulder, T., and Vanlede, J.: A heuristic
- 10 formula for turbulence-induced flocculation of cohesive sediment, Estuarine, Coastal and
- 11 Shelf Science, 68, 195-207, 10.1016/j.ecss.2006.02.003, 2006.
- 12
- 13
- 14

A 2D Process-Based Model for Suspended Sediment 3 Dynamics: a first Step towards Ecological Modeling

4

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11

12 Abstract

In estuaries <u>Suspended Sediment Concentration (SSC) is one of the most important</u> contributors to turbidity, which influences habitat conditions and ecological functions of the system. Sediment dynamics differ depending on sediment supply and hydrodynamic forcing conditions that vary over space and over time. A robust sediment transport model is the first step towards a chain of models <u>connecting models on contaminants</u>, phytoplankton and habitat.

19 This works aims to determine turbidity levels in the complex-geometry Delta of San 20 Francisco Estuary using a process-based approach (Delft3D Flexible Mesh software). Our 21 approach includes a detailed calibration against measured SSC levels, a sensitivity analysis on 22 model parameters, the<u>and the</u> determination of a yearly sediment budget as well as an 23 assessment of model results in terms of turbidity levels for a single year (Water Year 2011).

Model results show that our process-based approach is a valuable tool in assessing sediment dynamics and their related ecological parameters over a range of spatial and temporal scales and which may act as the base model for a chain of ecological models assessing the impact of climate change and management scenario forecasting. Here we present a modelling approach with limited data producing trustworthy predictions, useful findings for less monitored estuaries. With plenty of available data it was possible to analyze the strictly necessary data
 for the simulations which are water level, discharge and SSC for the input and output
 boundaries.

4

5 1 Introduction

Rivers transport water and sediments to estuaries and oceans. Sediment dynamics will differ
depending on sediment supply and hydrodynamic forcing conditions varying over space and
over time. Many river basins are subjected to slow morphodynamic adaptation due to
(gradually) changing forcing conditions, ranging from sea level rise and climate change to
anthropogenic developments such as reservoir construction in the watershed.

11 The human impact on sediment production dates from 3000 years ago, and has been 12 accelerating over the past 1000 years due to considerable engineering works (Syvitski and 13 Kettner, (2011). Milliman and Syvitski (1992) estimate that the budget of sediment delivered 14 to the coastal zone varies between 9.3 and 58 Gt per year. Estimating the world sediment budget is still a challenge either due to lack of data or detailed model studies in this field 15 (Vörösmarty et al., 2003). Adding to that, there is considerable uncertainty in hydraulic 16 17 forcing conditions and sediment supply dynamics due to variable adaptation timescales over 18 seasons and years (such as varying precipitation and river flow), decades (such as engineering 19 works) and centuries to millennia (sea level rise and climate change).

20 Examples of anthropogenic changes in sediment dynamics in river basins and estuaries are 21 manifold, e.g. San Francisco Bay-Delta (Schoellhamer, 2011), Yangtze Estuaries (Yahg, 22 1998) and Mekong Delta (Manh et al., 2014). These three systems present similar conditions 23 of anthropogenic forced sediment supply. After an increase in sediment supply (due to 24 hydraulic mining and deforestation respectively) each had a steep drop in sediment discharge (30%) due to reservoir building and further estuarine clearance after depletion of available 25 26 sediment in the bed. This implies a) continuous change in sediment dynamics and hence 27 sediment budget in the estuary; b) change in sediment availability leading to change in 28 turbidity levels.

29 Turbidity is a measurement of light attenuation in water and it is a key ecological parameter.
30 Fine sediment is the main contributor to turbidity. Therefore suspended sediment
31 concentration (SSC) can be translated into turbidity applying empirical formulations. Besides

1 SSC, algae, plankton, microbes and other substances may also contribute to turbidity levels 2 (ASTM International, 2002).High turbidity levels limit photosynthesis activity by 3 phytoplankton and microalgae, therefore decreasing associated primary production (Cole et 4 al., 1986). Turbidity levels also define habitat conditions for endemic species (Davidson-5 Arnott et al., 2002). We can cite the Delta Smelt as an example seeking for regions where the 6 turbidity is between 12-18NTU to hide from predators (Bakersville-Bridges, 2004;(Brown et 7 al., 2013).

8 Examples of other ecological impacts related to SSC are for vegetation stabilization (Morris
9 et al., 2002; Whitcraft and Levin, 2007), and salt marsh survival under sea level rise scenarios
10 (Kirwan et al., 2010; Reed, 2002).

11 To assess the aforementioned issues, the goal of this work is to provide a detailed analysis of 12 sediment dynamics concerning a) SSC levels, in the Sacramento-San Joaquin Delta (Delta) 13 area, b) sediment budget and c) translate theses results in turbidity levels, by means of a two 14 dimensions in the horizontal, averaged in the vertical dimension (2DH), process-based, 15 numerical model. The 2DH model solves the 2D vertical integrated shallow water equations 16 coupled with advective-diffusive transport. This process-based model will be able to quantify 17 high resolution sediment budgets and SSC, both in time (~ daysmonthly/yearly) and space 18 (~10s-100s of m). We selected the Delta area as a case study, since the area has been well 19 monitored so that detailed model validation can take place, it hosts endemic species, and 20 allow us to use a 2DH model approach.

21 The Delta and Bay are covered by a large survey network, offering freely available data on river stage, discharge and suspended sediment concentration (SSC) amongst other parameters, 22 23 and maintained by USGS (nwis.waterdata.usgs.gov), by Californian Department of Water Resources (http://cdec.water.ca.gov/) and National Oceanic and Atmospheric Administration 24 (http://tidesandcurrents.noaa.gov/). The continuous SSC measurement stations are 25 periodically calibrated by water collection in situ, filtered and weighted in the laboratory. On 26 top of that, the Bay-Delta system has a high resolution bathymetry available (10m) for all the 27 channels and bays (http://www.d3d-baydelta.org/). 28 29 Regarding ecological value, starting from the bottom of the food web, the Delta is the most important area for primary production in the San Francisco Estuary. The Delta is one order of 30 magnitude more productive than the rest of the estuary (Jassby et al., 2002;Kimmerer, 2004). 31

32 It is an area for spawning, breeding and feeding for many endemic species of fishes and

1 invertebrates, including some endangered species like delta smelt (Brown et al., 2013),

2 Chinook salmon, spring run salmon and steelhead. Additionally, Several projects for marshes

- 3 restoration in the Delta are planned and the success of these projects depends on sediment
- 4 <u>availability (Brown, 2003).</u>

5 The SSC model results are compared to in situ measured SSC data. The calibration process 6 assesses the sensitivity of sediment characteristics such as fall velocity (w), critical shear 7 stress (τ_{cr}) and erosion coefficient (E). The model outputs are the spatial and temporal 8 distribution of SSC (turbidity), yearly sediment budget for different Delta regions, and the 9 sediment export to the bay.

10 For the first time, a detailed, process-based model is developed for San Francisco Bay-Delta, 11 to focus on the complex Delta sediment dynamics (MacWilliams, 2015). From this model it is possible to describe the spatial sediment (turbidity) distribution and deposition patterns that 12 are important indicators to assess habitat conditions. Analyzing seasonal and yearly variations 13 14 in sediment dynamics and translating these into turbidity levels to be used as indicators for 15 ecological modeling (Janauer, 2000), this work fills gap the between the physical aspects 16 (hydrodynamic and sediment modeling) and ecology modeling. Previous work focused on understanding the San Francisco Bay-Delta system through data analysis (Barnard et al., 17 18 2013; Manning and Schoellhamer, 2013; McKee et al., 2006; McKee et al., 2013; Morgan-King and Schoellhamer, 2013;Schoellhamer, 2011;Schoellhamer, 2002;Wright and Schoellhamer, 19 20 2004, 2005), while similar work in other estuaries around the world does not give the direct 21 link to ecology (Manh et al., 2014).

22 2 Background

23 San Francisco Estuary is the largest estuary on the U.S. West Coast. The estuary comprises 24 San Francisco Bay and the inland Sacramento-San Joaquin Delta (Bay-Delta system), which together cover a total area of 1235 km2 with a mean water depth of 4.6 meter (Jassby et al., 25 26 1993). The system has a complex geometry consisting of interconnected sub-embayments, 27 channels, rivers, intertidal flats, and marshes (Fig 1). The Sacramento-San Joaquin Delta 28 (Delta) is a collection of natural and man-made channel networks and leveed islands, where 29 the Sacramento River and the San-Joaquin River are the main tributaries followed by 30 Mokelumne River (Delta Atlas, 1995). San Francisco Bay has 4 sub-embayments. The most 31 landward is Suisun Bay followed by San Pablo Bay, Central Bay (connecting with the sea through Golden Gate) and, further southward, South Bay. 32

1 Tides propagate from Golden Gate into the Bay and most of the Delta up to Sacramento 2 (FPT) and Vernalis (VNS) when river discharge is low. Suisun Bay experiences mixed diurnal and semidiurnal tide that ranges from about 0.6 m during the weakest neap tides to 1.8 3 4 m during the strongest spring tides. During high river discharge the 2psu isohaline is located 5 in San Pablo bay while during low river discharge it can go landwards of Chipps Island 6 (westernmost reach of the black rectangle, Fig 1). The topography highly influences the wind 7 climate in the Bay-Delta system. Wind velocities are strongest during spring and summer presenting afternoon north-westerly gusts of about 9 ms⁻¹ (Hayes et al., 1984). 8

9 San Francisco estuary collects 40% of the total Californian fresh water discharge. It has a 10 Mediterranean climate, with 70% of rainfall concentrated between October and April (winter) 11 decreasing until the driest month September (summer) (Conomos et al., 1985). The 12 orographic lift of the Pacific moist air linked to the winter storms and the snowmelts in early 13 spring govern this wet (winter) and dry (summer) season variability. This system leads to a 14 local hydrological 'Water Year' (WY) definition from 1st October to 30th September, including 15 a full wet season in one WY.

16 It is important to notice that Sacramento and San Joaquin Rivers, together, account for 90% of 17 the total fresh water discharge to the estuary (Kimmerer, 2004). The daily inflow to the Delta follows the rain and snowmelt seasonality, with average dry summers with discharges of 50-18 150 m³s⁻¹ and wet spring/winter reaching peak discharges of 800-2500 m³s⁻¹. The geographic 19 and seasonal flow concentration leads to several water issues related to agricultural use. 20 habitat maintenance and water export. On a yearly average 300m³s⁻¹ of water is pumped from 21 South Delta to southern California. The pumping rate is designed to keep the 2psu (salinity) 22 23 line landwards of Chipps Island avoiding salinity intrusion in the Delta. - aAllowing the 2DH 24 modeling approach.

The hydrological cycle in the Bay-Delta determines the sediment input to the system, thus biota behavior. McKee (2006) and Ganju and Schoellhamer (2006) observed that a large volume of sediment passes through the Delta and arrives to the Bay in a yearly pulse. They estimated that in 1 day approximately 10% of the total sediment volume could be delivered and in extremely wet years up to 40% of the total sediment volume can be delivered in 7 days. During wet months more than 90% of the total sediment inflow is supplied to the Delta.

The recent Delta history is dominated by anthropogenic impacts. In the 1850's hydraulic mining started after placer mining in rivers became unproductive. Hydraulic mining remobilized a huge amount of sediment upstream of Sacramento. By the end of the nineteenth century the hydraulic mining was outlawed leaving approximately 1.1×10^9 m³ of remobilized sediment, which filled mud flats and marshes up to 1 meter in the Delta and Bay (Wright and Schoellhamer, 2004;Jaffe et al., 2007b). At the same time of the mining prohibition, civil works such as dredging and construction of levees and dams started, reducing the sediment supply to the Delta (Delta Atlas, 1995;Whipple et al., 2012).

7 Typical SSC in the Delta ranges from 10 to 50 mg L-1, except during high river discharge when SSC can exceed 200 mg L-1 reaching values over 1000mgL⁻¹ (McKee et al., 8 9 2006; Wright and Schoellhamer, 2005). A sediment budget reflects the balance between 10 storage, inflow and outflow of sediment in a system. Studies based on sediment inflow and 11 outflow, estimated that about two-third of the sediment entering the system deposits in the de Delta (Schoellhamer et al., 2012; Wright and Schoellhamer, 2005). The remaining third is 12 13 exported to the Bay, and represents on average 50% of the total Bay sediment supply (McKee et al., 2006), the other half comes from smaller watershed around the Bay (McKee et al., 14 15 2013).

16 Several studies have been carried out to determine sediment pathways and to estimate 17 sediment budgets in the Delta area (Schoellhamer et al., 2012;Jaffe et al., 2007a;Gilbert, 18 1917; McKee et al., 2013; McKee et al., 2006; Wright and Schoellhamer, 2005). These studies 19 were based on data analysis and conceptual hindcast models. Although the region has a 20 unique network of surveying stations, there are many channels without measuring stations. This might lead to incomplete system understanding and knowledge deficits for the 21 22 development of water and ecosystem management plans. The monitoring stations are located 23 in discrete points hampering spatial analysis. Also, the impact of future scenarios related to 24 climate change (i.e. sea level rise and changing hydrographs) or different pumping strategies 25 remains uncertain.

26 2.1 Model description

Structured grid models such as Delft3D and ROMS (Regional Oceanic Modeling System)
 have been widely used and accepted in estuarine hydrodynamics and morphodynamics
 modeling including San Francisco Estuary (Ganju and Schoellhamer, 2009;Ganju et al.,
 2009;van der Wegen et al., 2011). In all these cases the Delta was schematized as 2 long

- channels since the grid is not flexible to have a 2D modeling of the rivers, channels and
 flooded island of the system together with the Bay.
- 3 In case of complex geometry unstructured grids or finite volume model is more suitable.
- 4 There are three widely known unstructured grid models the TELEMAC-MASCARET
- 5 (Hervouet, 2007), the UnTRIM (Casulli and Walters, 2000;Bever and MacWilliams, 2013)
- 6 and D3D FM (Kernkamp et al., 2010). The two first models are purely triangle based and are
- 7 not coupled (yet) with sediment transport and/or water quality and ecology model.

8 The numerical model applied in this work is Delft3D Flexible Mesh (D3D FM). D3D FM 9 allows straightforward coupling of its hydrodynamic modules with water quality model, 10 Delft-WAQ (DELWAQ), which gives flexibility to couple with the habitat (ecological) 11 model. D3D FM is a process-based unstructured grid model developed by Deltares (Deltares, 12 2014). It is a package for hydro- and morphodynamic simulation based on a finite volume 13 approach solving shallow-water equations applying a Gaussian solver. The grid can be 14 defined in terms of triangles, (curvilinear) quadrilaterals, pentagons and hexagons, or any 15 combination of these shapes. It is important to note that (orthogonal) quadrilaterals are the 16 most computationally efficient cells. Kernkamp (2010) and the D3D FM manual (Deltares, 17 2014) describe in detail the grid aspects and the numerical solvers.

18 The Bay area and river channels are defined by consecutive curvilinear grids (quadrilateral). Different resolution grid, the river discharging in the Bay, and channel junctions are 19 20 connected by triangles (Fig 2). The average cell size ranges from 1200m x 1200m, in the 21 coastal area, to 450x600m in the Bay area down to 25x25m in the Delta channels. In the Delta, each channel is represented by at least 3 cells in the across-channel direction (Fig 2). 22 23 The grid flexibility allows including the entire Bay-Delta in a single grid containing 63.844 24 cells from which about 80% are rectangles, but still keepings the computer run times at an 25 acceptable level. It takes 6 real days to run 1 year of hydrodynamics simulation and 12 hours to run the sediment module on an 8 cores desktop computer. Not counting the triangular grid 26 27 orthogonality and circumscription issues, in the case of entirely triangular grid the running time for a 1 year simulation would increase from ~72 clock hours to ~192 hours. 28 29 The advantages of the numerical method are addressed by Kernkamp (2010). With a personal

- 30 <u>computer it is possible to address challenging morphologies combining rectangular and</u>
- 31 triangular mesh. The advantage of this approach is the reduced number of cells since it is not

compulsory for the entire mesh to be triangular, leading to less cells and lower computational
 time.

3 We assume that the main flow dynamics in the Delta are 2D no vertical stratification, -since 4 no salt-fresh water interactions occur in the Delta due to the pumping operations and we 5 assume that temperature differences do not govern flow characteristics. D3D FM generates 6 hydrodynamic output for off-line coupling with water quality model DELWAQ (Deltares, 7 2004). Off-line coupling enables faster calibration and sensitivity analysis. DFlow-FM 8 generates time series of the following variables: cell link area; boundary definition; water 9 flow through cell link; pointer file gives information concerning neighbors' cells; cell surface; 10 cell volume; and shear stress file, which is parameterized in DFlow-FM using Manning's n. 11 Given a network of water levels and flow velocities (varying over time) DELWAQ can solve 12 the advection-diffusion-reaction equation for a wide range of substances including fine sediment, the focus of this study. DELWAQ solves sediment source and sink terms by 13 14 applying the Krone-Parteniades formulation for cohesive sediment transport (Krone, 15 1962; Ariathurai and Arulanandan, 1978) (Eq.1, Eq.2).

16
$$D = w_s * c * (1 - \tau_b / \tau_d)$$
 (1)

17
$$E = M * (\tau_b / \tau_e - 1)$$
 (2)

18 Where; *D* Deposition flux of suspended matter (mg m⁻²s⁻¹), w_s settling velocity of suspended 19 matter (ms⁻¹), *c* concentration of suspended matter near the bed (mg m⁻³), τ_b bottom shear 20 stress (Pa) τ_d critical shear stress for deposition (Pa), *E* erosion rate (mg m⁻²s⁻¹), *M* first order 21 erosion rate (mg m⁻²s⁻¹), τ_e critical shear stress for erosion (Pa).

Note: Following Winterwerp (<u>{Winterwerp, 2006 #402</u>}2006) we assume that deposition takes place regardless of the prevailing bed shear stress. τ_d is thus considered much larger than τ_b and the second term in equation (Eq. 1) is close to zero.

25 **2.2** Initial and Boundary Conditions

I

The Bay-Delta is a well measured system; therefore all the input data to the model are in situ data. Initial bathymetry has 10m grid resolution, which is based on an earlier grid (Foxgrover et al., <u>http://sfbay.wr.usgs.gov/sediment/delta/</u>), modified to include new data by Wang and Ateljevich

 $30 \quad (\underline{http://baydeltaoffice.water.ca.gov/modeling/deltamodeling/modelingdata/DEM.cfm}) \qquad and \qquad \\$

1 further refined. The bathymetry is based on different data sources including bathymetric 2 soundings and LiDAR data. The hydrodynamic model includes real wind, which results from the model described by (Ludwig and Sinton, 2000). The wind model interpolates hourly data 3 4 from more than 30 meteorological stations into regular 1km grid cells. Levees and temporal 5 barriers are included in the model considering their deployment time (http://baydeltaoffice.water.ca.gov/sdb/tbp/web_pg/tempbsch.cfm). 6

The hydrodynamic model has been calibrated for the entire Bay-Delta system (see appendix 7 A and <u>http://www.d3d-baydelta.org/</u>). Initial SSC was set at $0mgL^{-1}$ over the entire domain. 8 9 The initial bottom sediment availability defined available mud at places shallower than 5 meters below Mean Sea Level (MSL) including intertidal mud flats, and sand at places deeper 10 11 than 5 meter below MSL, which are primarily channel regions. This implies that the main 12 Delta channels such as, Sacramento, San Joaquin, Mokelumne are defined as sandy with few 13 mud patches. The smaller channels and the flooded islands such as Franks Tract are initialized 14 with a muddy bottom. DELWAQ does not compute morphological changes or bed load 15 transport.

In this study we applied 5 open boundaries. Seaward we set hourly water level time series 16 17 derived from Point Reves station (tidesandcurrents.noaa.gov/). The other four landward 18 boundaries are river discharge boundaries at Sacramento River (Freeport), Yolo Bypass 19 (upstream water divergence from Sacramento River), San Joaquin River and Mokelumne 20 River. Studies show that Sacramento River accounts for 85% of the total sediment inflow to 21 the Delta, while San Joaquin accounts for 13% (Wright and Schoellhamer, 2005), so it is 22 reasonable to apply 2 sediment discharge boundaries at Sacramento and San Joaquin River. 23 All river boundaries present unidirectional flow, excluding tidal influence.

24 The river water flow hourly input data are from the following stations, at-Sacramento River at 25 Freeport (FPT), San Joaquin River near Vernalis (VNS) and Yolo Bypass (YOLO) were obtained from California Data Exchange Center website (cdec.water.ca.gov/) (Fig 2). The 26 27 sediment input data, for both input-(stations FPT and VNS,) and calibration stations S Mokelumne R((SMR), N Mokelumne R (NMR), Rio Vista (RVB), Mokelumne (MOK), 28 29 Little Potato Slough (LPS), Middle River (MDM), Stockton (STK), Mallard Island (MAL) (Fig 2), was obtained by personal communication from USGS Sacramento; this data is part of 30 a monitoring program (http://sfbay.wr.usgs.gov). Since 1998, USGS has continuous 31 32 measuring stations for sediment concentration which is derived from backscatter sensors (OBS) measurements every 15 minutes, and nearly monthly calibrated with bottle samples
 (Wright and Schoellhamer, 2005).

3 The SSC data that is used to compare to model results are derived from optical backscatter 4 sensors (OBS). This type of sensor converts scattered light from the particles in photocurrent, 5 which is proportional to SSC. To define the rating curve it is necessary to sample water, filter 6 and weight the filter. However, in some locations the cloud of points when correlating 7 photocurrent and filtered weight shows a large scatter. Large scatter leads to errors in 8 translating photocurrent to SSC. These errors are due to (amongst others) particle size, 9 desegregation (cohesiveness, flocculation, organic-rich estuarine mud); shape effects; 10 sediment-concentration effect (Kineke and Sternberg, 1992;Downing, 2006;Sutherland et al., 11 2000; Gibbs and Wolanski, 1992; Ludwig and Hanes, 1990). Wright and Schoellhamer (2005) 12 showed that for the Sacramento-San Joaquin Delta these errors can sum up to 39%, when calculationg sediment fluxes throught Rio Vista. 13

In this work we modeled the 2011 water year - 1st October 2010 until 30th September 2011. First, we ran D3D FM for this year to calculate water level, velocities, cell volume and shear stresses. Then, the 1 year hydrodynamic results were imported in DELWAQ which calculated SSC levels.

18 **3** The next five centuries

Our focus is to represent realistic SSC levels capturing the peaks, timing and duration, and to develop a sediment budget to assess sediment trapping in the Sacramento-San Joaquin Delta, (Fig 1, highlighted by the black rectangle). Throughout the following sections the results are analyzed in terms of tide averaged results, meaning that the data and model results are filtered to frequencies lower than 2 days. We applied a Butterworth filter with cut off frequency of $1/30h^{-1}$ as presented in Ganju and Schoellhamer (2006).

25 **3.1 Calibration**

The results shown below are the derived from an extensive calibration process where the different sediment fractions, w, τ_{cr} and E were tested. The first attempt applied multiple fraction settings presented in previous works (van der Wegen et al., 2011;Ganju and Schoellhamer, 2009). However, tests with a single mud fraction proved to be consistent with the data, representative of the sediment budget, and allow a simpler model setting and better understanding of the SSC dynamics. The choice of one mud fraction follows (Beven et al.,
 2011) argument that it is better to keep a model simple in order to understand the processes.
 In addition, with a single fraction it was possible to reproduce more than 90% of the sediment
 budget for the Delta.

5 After defining for 1 fraction runs, the sensitivity analysis and metrics helped to choose the 6 standard run presented below, the fraction selected has w of 0.25mms⁻¹, τ_{cr} of 0.25Pa and E of 7 10^{-4} kgm⁻²s⁻¹. The same applies to the bed sediment availability defined by 1 mud (shoals) and 8 1 sand (channels) fraction.

9 **3.2** Suspended Sediment Dynamics (water year 2011)

The 2011 simulation reproduces the SSC seasonal variation in the main Delta regions such as the North (Sacramento River) represented by Rio Vista station (RVB); the South (San Joaquin River) represented by Stockton (STK); Central-East Delta represented by Mokelumne station (MOK) and Delta output represented by Mallard Island (MAL) (Fig 4).

All stations clearly reproduce SSC peaks during high river flow periods during November to July and lower concentrations during the remainder of the year (apart from MAL during the July-August period). The good representation of the peak timing indicates that the main Delta discharge event is reproduced by the model as well as the periods of Delta clearance. These two periods are critical for ecological models, and a good representation generates robust input to ecological models. The differences found between the model and data are further discussed in appendix B.

21 **3.3 Sensitivity analysis**

22 **3.3.1 Sediment fraction analysis**

We considered one fraction for simplicity and because it reproduces more than 90% of the sediment budget throughout the Delta as well as the seasonal variability of SSC levels. Although more mud fractions considerably increase running time, several tests with multiple fractions were done to explore possibilities for improving the model results.

Including heavier fractions changes the peaks timing and also lowers the SSC curve. Comparing the standard run (w=0.25mm s⁻¹, T=0.5Pa, E= 10^{-4} kg m⁻²s and bottom composition with mud available shallower than 5 meters) to another run considering 15% of heavier fraction (w=1.5mm s⁻¹) and 30% of a lighter fraction (fall velocity of 0.15mm s⁻¹),
showed that the peak magnitudes were underestimated but the first peak timing is closer to the
data and the spurious peak mid May is lower.

4 To be able to find the best parameter setting a sensitivity analysis was done varying the main
5 | parameters in the Krone-Parteniades formulation (<u>Table 1</u>).

6 Regarding sediment flux, these tests show that some stations, such as RVB and MAL, are 7 more sensitive to parameter change than others, such as STK (Fig 5). The model results are 8 most sensitive to the critical shear stress and least sensitive to the erosion coefficient. 9 Analyzing the time series, one concludes that in stations where the fluxes are higher, the 10 change in critical shear stress is less important, since during most of the time the shear stress 11 is already higher than any given critical shear stress.

It is important to note that reaching a perfect fit for one station does not mean reaching it for the others. We are analyzing two metrics, the unbiased Root Mean Square Error (uRMSe, Fig 6) and Skill (Skill, Fig 6) (Bever and MacWilliams, 2013). The uRMSe analyzes the variability of the model relative to the data, in this case 0 is the case when the model and data have equal variability, positive values indicate more model variability and negative values indicate less model variability.

18
$$| uRMSe = \left(\frac{1}{N}\sum_{i=-1}^{N} \left[\left(X_{\mathcal{M}mi} - \overline{X_{m\mathcal{M}}}\right) \left(X_{0i} - \overline{X_{0}}\right) \right]^2 \right)^{0.5}$$
19 (3)

20 Where *N* is the time series size, *X* is the variable to be compared, in this case SSC, and \overline{X} is 21 the time-averaged value. Subscript <u>M-m</u> and O represent modeled and observed values, 22 respectively.

Skill is a single quantitative metric for model performance (Willmott, 1981). When skill
equals 1 the model perfectly reproduces the data. The 2 metrics where evaluated at RVB,
STK and MAL, representing respectively Sacramento River, San Joaquin River and Delta
output.

27
$$\left| \begin{array}{c} Skill = 1 - \left[\sum_{i=-1}^{N} \left| X_{m\mathcal{M}i} - \overline{X_{0i}} \right|^2 \right] / \left[\sum_{i=1}^{N} \left(\left| X_{m\mathcal{M}i} - \overline{X_{0}} \right| + \left| X_{0i} - \overline{X_{0}} \right| \right)^2 \right] \right]$$

28 (4)

One notices that changing a parameter can lead to better results in one station but worse in other stations (Fig 6). The choice of the standard run analyzed throughout the paper comes from this analysis as well as the budget analysis. Following the argumentation from (Beven et al., 2011) statistical analyses are needed in order to make a better choice of the model settings. We note that both uRMSe and Skill varies up to 50% over the different runs.

6 **3.4 Initial bottom composition**

To study the importance of initial bottom sediment availability we considered 2 cases; one
excluding sediment (concrete bed) and the other by defining available mud at places
shallower than 5 meters below Mean Sea Level (MSL) including intertidal mud flats and sand
at places deeper than 5 meter below MSL being mainly channel regions (standard run).

11 We did some test varying the 5m threshold. From 3 to 10 meters the final results are quite

12 similar. However, considering mud availability in the channels deeper than 10 meters starts to

<u>disturb the SSC levels.</u> Time series of SSC comparing the 2 cases and data show that bottom
 composition has virtually no influence on SCC after the first couple of days. This result also
 applies for different mud fractions availability and opens horizons for modeling less measured

16 estuaries where virtually no bottom sediment data is available.

17 Another test shows that it is better to initialize the model with a 'concrete' bed than with mud 18 available in the entire domain. Initializing the channels with loose mud generates 19 unrealistically high SSC levels through the years, which can take up to 5 years to be 20 reworked.

21 **4 Discussion**

22 In the previous section we presented the model calibration, a normal practice in the modeling 23 process. In this section we discuss the new insights that were derived from the model results. Although these insights are specific to the San Francisco Bay-Delta system, the same 24 25 approach can be applied to other estuaries and deltas. The model shows detailed sediment 26 dynamics and the main paths that sediment is transported in the Delta. Sediment flux 27 calculations define the sediment dynamics while gradients in sediment describe the sediment 28 distribution and deposition pattern in the Delta. We also discuss daily and seasonal variation 29 of turbidity levels.

1 4.1 Spatial sediment distribution

2 Starting the analysis with the general Delta behavior, during dry periods SSC in the entire Delta is low (<20mgL⁻¹) and the Delta water is relatively clear. The current model results 3 confirm and compile data showing that the Sacramento River is the main sediment supplier 4 5 into the Delta (Wright and Schoellhamer, 2004;Schoellhamer et al., 2012). Sacramento River 6 peak flow fills the north and partially fills the central/east Delta with sediment. However, the rest of the Delta keeps quite low levels (~ 20mgL^{-1}) of SSC all year long. Passing Vernalis 7 8 (VNS), San Joaquin River main branch flows to the east, however the SSC peak reaches no 9 much further than STK. The west branch goes toward the water pumping stations where to the 10 sediment is pumped out of the system. This behavior reflects in very low SSC in the South 11 Delta (Old River and Franks Tract) region, which are deposition areas.

12 Three Mile Slough (TMS) and the Delta Cross Channel (DCC) connect the Sacramento River with the central and eastern Delta. Model results show that together they carry 60Kton per 13 14 year of sediment southward. DCC operation defines SSC levels in the eastern/central Delta to a large extent. To show the importance of DCC we run the model twice, one with DCC 15 always open and one always closed. When DCC is open, high SSC Sacramento river water 16 (~150mgL⁻¹) flows towards Mokelumne River and Eastern Delta increasing the overall SSC 17 in the area. When it is closed SSC levels in central and eastern delta are about 30mgL⁻¹ lower 18 than in the previous case ($\underline{Fig 7}$). The effect of opening DCC can be observed in the SSC level 19 20 at the San Joaquin River from MOK station seawards. In the Sacramento River, the opening decreases SSC levels, by about 10mgL⁻¹. It affects the river SSC all the way to Mallard Island 21 22 (Fig 7).

During peak river discharge, Sacramento River sediment reaches Mallard Island in approximately 3 days, Carquinez Straight in 5 days, and the Golden Gate Bridge in approximately 10 days. This timing is proportional to river discharge. However from Mallard Island seawards it is a rough estimate due to the 2D approximation. The San Joaquin River sediment remains largely trapped in the southern Delta. The flooded islands, breached levees like Franks Tract, present a different behavior. During the entire year the SSC levels are below 15mg L⁻¹, the river peak discharge signal does not affect them.

Sediment flux is a useful tool for a quantitative and qualitative analysis of the sediment path
and its derivative gives sedimentation/erosion patterns. It is defined by the product of water
velocity (U), times Cross-sectional area (A) times SSC (C) (Eq. 5).

$$1 \quad F_{sed} = U * A * C \tag{5}$$

The yearly sediment flux through FPT from model results is 1132Kt yr⁻¹ (thousand metric 2 tons per year) against 1096Kt yr⁻¹ from data, following Sacramento River we have RVB with 3 832Kt yr⁻¹ (994Kt yr⁻¹, data), then MAL with 617Kt yr⁻¹ (654 Kt yr⁻¹) (Fig 8). We calculate 4 that 30Kt yr⁻¹ of Sacramento River sediment flows to the eastern Delta through DCC, 30 Kt/yr 5 through TMS and 20Kt yr⁻¹ from Georgina Slough. San Joaquin River carries 490Kt yr⁻¹ 6 (498) through VNS, heading to STK with 205Kt yr⁻¹ (190Kt yr⁻¹). It was estimated that 100Kt 7 yr⁻¹ was exported through pumping. To close the system in central Delta, the flux through JPT 8 is 126Kt yr⁻¹ (no data) and DCH approximately 0 (no data) (Fig 8). 9

Seaward from MAL considerable salt-freshwater stratification takes place in the water column
 These 3D effects are not captured by our 2DH approach so that the model results in this
 region are disregarded. Therefore, Fig 8 shows preliminary sediment flux to the Bay by a
 dashed line.

14 4.2 Sediment budget

From the previous section one can see that more sediment enters (~1600 Kt yr⁻¹) than leaves 15 (~600 Kt yr⁻¹) the Delta. So by mass conservation law, the difference between inflow and 16 outflow deposits in the Delta. Jaffe et al.(2007b) developed a box model based on bathymetry 17 18 data to define sediment budget of the Delta and Bay to define sediment availability for 19 ecology purposes. The model results agree with data estimations that about two third of the 20 sediment input is retained in the Delta (Schoellhamer et al., 2012; Wright and Schoellhamer, 21 2005), and it is consistent throughout the years (Cappiella et al., 1999;Jaffe et al., 22 1998; Wright and Schoellhamer, 2004). Because of the detailed description of the sediment 23 path, it is possible to further understand and describe the sediment budget in Delta subregions (north, central and south), comparing our results to data when available (Morgan 24 25 King, 2012, personal communication).

Besides the overall trend, different parts of the Delta present different trap efficiency. Model
results show that Northern Delta (the least efficient) traps ~ 23%; Central/Eastern Delta traps
32%, Central/Western 65%, and the most efficient is the Southern Delta region trapping 67%
of the sediment input. The highest trapping efficient regions correspond to islands inundated
through levee breaching (Wright and Schoellhamer, 2005).

From the total Sacramento River sediment input 40% stays in the northern Delta and about 40% is exported to Bay area. The remaining 20% deposits in the central/eastern Delta and only 2% travel all the way to South Delta. About 70% of San Joaquin sediment deposits in the southern Delta, 10% go to central Delta, 15% is exported via Clifton Court pumping facilities and 5% is exported to the Bay. This transport is reflected in the bottom composition of the Delta, Sacramento River sediment dominates the Northern and Central Delta and San Joaquin River sediment dominates the Southern Delta bottom composition (Fig 9).

8 It is possible to divide the sediment budget analysis for the wet and the dry season, since the 9 Delta presents different dynamics for each season. Water year 2011 was a wet year, with the 10 wet season lasting from mid-January until the end of May. During the wet period 60% of the 11 yearly sediment input budget entered the Delta through FPT and VNS and 70% of the yearly budget was exported through MAL. In the wet season the high river water discharges and 12 SSC pulses flushes the entire Delta with sediment. In this season high SSC gradients are 13 14 observed in the plume fronts leading to rapid changes in habitat conditions for many species. After the front the high SSC level can last for more than a month, indicating changing in 15 16 habitat conditions

During the dry season the Delta experiences lower river discharges and SSC levels thus the sediment transport is lower as well. In the dry season SSC levels are more uniform not presenting peaks, at this time the water is clear and the advective flux is lower, which is going to be discussed in the next section.

21 **4.3 Sediment flux analysis**

SSC peaks at FPT can be tracked down the estuary. At the RVB station the SSC peak follows the same dynamic as observed at FPT; however, this behavior does not apply for the entire Delta. Schoellhamer and Wright (2005) observed that the river signal is attenuated through the estuary. This attenuation can be understood by analyzing changes in the dominant sediment flux component.

27 Dyer (1974) decomposed the tidally averaged fluxes in three main components: tidal mean, 28 the advective term; tidal fluctuation, the dispersive term; and the Stokes Drift. This 29 decomposition was possible considering that the measured valued is the sum of a tidally mean 30 component [x], and a fluctuating component x', so x = [x] + x' (Eq.6), substituting in Eq 5 31 and simplifying the small contribution terms, three main terms remain (Eq.7). The first term of Eq7 is the advective term, it is the river flow as it is calculated by the mean discharge, area and concentration; the second one is the dispersive term that accounts for tidal pumping, which is the compensation flow for the inward transport of the tidal wave the 2 first terms already account for more than 95% of the flux and the Stokes Drift which is the transport due to a variation in the cross-sectional area.

6 [F] = [U][A][C] + [[U'C'][A]] + [[U'A'[C]]] (7)

7 The model allows for a detailed temporal and spatial analysis of the three flux components. 8 The temporal analysis are done in 3 steps, the first one considering the whole year and then 9 splitting in the wet and dry season. For the spatial analysis, we defined 4 stations for each 10 river where the first station is dominated by the river flux and the last experience a mix of 11 tidal and river fluxes. The stations were determined following Sacramento River, starting with 12 FPT, followed by RVB down to Mallard Island the delta output and following San Joaquin 13 River from VNS, to STK and MOK. Three Mile Slough (TMS) and San Joaquin Junction 14 (SJJ) represent the Delta smaller channels.

Sacramento River at FPT, the most landward station, experiences no tidal influence so the flux is purely advective. RVB, seaward, experiences tidal fluctuations and the dispersive flux is responsible for 22% of the total flux; however no Stokes drift flux is present (Fig 10). On the other hand, Stokes Drift component accounts for 33% of the total flux in MAL station implying that tides have a bigger influence in this region.

An analogue can be drawn to San Joaquin branch, where VNS and STK experience only advective terms. In MOK and SJJ dispersive (20% and 63% respectively) and Stokes flux start (5 and 11%) to change the total flux (Fig 10). The analyses of the 3 different flux components in smaller Delta channels show that river and tidal signals are equally important. In other words the river peak signal is less important inside smaller channels than in rivers. At TMS, the dispersive flow accounts for 60% of the total flux.

The fluxes analysis shows that there is no change in the Delta net circulation when comparing wet and dry seasons. In other words, there is not a major signal change in the flux signal direction when comparing the seasons. However, there is a change in importance of each flux component.

30 <u>Fig 10</u> shows that dispersive flux and Stokes Drift relative contributions vary seasonally: 31 when river discharge is high the relative contribution of dispersive flux and Stokes Drift is lower than during low flow conditions. This pattern is better observed in stations where the river signal is stronger. At RVB the dispersive flux contribution is about 15% during the wet season and 26% in the dry season, the same applies for MAL and STK. In smaller channels, like TMS and SJJ, The dispersive flux seasonal variation is milder, varying about 10%, from 55% in the wet season and 65% in the dry. In the dry season the change in fluxes contribution, from advective to dispersive and Stokes Drift, leads to a lower net export of sediment from the Delta, even though the concentrations in the Delta about 30mg L⁻¹.

8 4.4 Sediment deposition pattern

9 The flux change from completely advective to dispersive and Stokes drift sheds some light on
10 the Delta deposition areas. The places where the dispersive flux starts to play a role, near
11 RVB and MOK areas, are the same places where net deposition is observed (Fig 11). Other
12 locations where considerable sedimentation takes place are in flooded islands areas, such as
13 Frank Tract and the Clifton Court. The 2D model allows determining such areas (Fig 11).

The San Joaquin River downstream of Stockton experiences high deposition. This finding is confirmed by constant dredging works need to maintain Stockton navigation channel. The river discharge modulates the deposition pattern in the main channels. In the Sacramento River, Rio Vista area (RVB), a rapid deposition takes place just after the peak discharge. Later this deposited sediment is gradually washed away and transported to the mud flats at the channel margins, until the next peak.

At flooded island the sedimentation process is gradually and steady, and is not transported again. The deposition pattern does not change from wet to dry season, except for some small bends in the Sacramento River that goes from eroding (wet) to depositing areas (dry). The deposition pattern provides insight into the best areas for marsh restoration.

24 **4.5 Turbidity**

So far the discussion was presented in terms of SSC levels, budgets and fluxes, while ecological analysis is often based on turbidity levels. SSC and turbidity are correlated by rating curves as log10 (SSC) = a*log10 (Turb) +b, where a and b are local parameters empirically defined for each Delta area. The Northern area a=0.85 and b=0.35; Central/Western area a=0.91 and b=0.29, Central/Eastern a=0.72 and b=0.26; Southern a=1.16 and b=0.27; Eastern a=0.914 and b=0.29 (USGS Sacramento, personal communication
 2014).

3 In this section we present average values for turbidity within a specific Delta region as well 4 as its seasonal and daily variations (Fig 12). Generally, the mean turbidity levels and spatial 5 variations are higher during the wet season than during the dry season. During the wet season, 6 the Southern area presents the highest mean value (50 ntu), and deviation (15ntu), caused by a 7 combination of large sediment supply and low flow velocities. The Northern region is the 8 second most turbid area (45±10ntu), where sediment transported by Sacramento River flows 9 in the channels, increasing the turbidity levels. The Central Eastern region is the least turbid 10 area (5±2ntu) and, as previously shown, it presents the highest trapping efficiency of the 11 entire Delta. In the dry season the mean turbidity daily variation decreases in the whole Delta, 12 excepting the Central/Eastern region. The opening of the DCC during the dry season lets sediment from the Sacramento River entering these areas, increasing the mean turbidity level. 13 14 The spatial distribution of the most turbid areas is the same as in the wet season. The daily 15 deviation is mostly proportional to the turbidity level and to the distance from the sea. In the 16 Southern and Western areas the daily variation is higher during the dry season. It shows that 17 there is a strong tidal signal in these parts of the Delta.

As for from this work results, we note that the Sacramento to San-Joaquin River connecting
channels DCC and GLS are important bridges to export sediment from Sacramento to Eastern
Delta. On the other hand the smaller channels of the network play a minor role in the Delta
sediment budget, since the discharges in these channels are considerably smaller than in the
rivers.

23 4.6 Data input discussion

As a well surveyed area, combining with a complex process-based model, the Delta offers the
 chance of testing how much data it is necessary to develop a robust model to sediment
 dynamics and budget. The model offers the possibility of having high temporal and spatial
 resolution, as well as considers multiple physical processes as bottom friction, sedimentation
 and erosion. The available data allows calibration and validation of model results.

As presented before, with simple settings as 1 mud fraction and simple bed sediment
 availability the model is capable of representing the main sediment dynamics processes, the
 peak timing and duration, as well as sediment budget. The necessary data to a trustworthy

1 modeling and further forecasting is a fine resolution bathymetry to correctly reproduce 2 hydrodynamics, SSC and discharge in the inflow and outflow boundaries as well as in 1-3 2stations in the domain in order to properly calibrate the model. The results from the 4 calibrated model using these few data can be extrapolated for the entire domain, allowing 5 closing the sediment budget for the whole system. 6 The 2D model results output are available in high temporal (~hours) and spatial (~20 meters) 7 resolution, allowing to translate model results in water quality parameters for modeling or for 8 descriptive purposes. In other words, with limited input data we can come to a detailed system

9 description with considerable forecast capacity, expanding the applicability of this work to less
 10 measured estuaries.

11

12 **5** Conclusions

- In this work we make a step towards the understanding and simulating sediment dynamics
 from source to sink in a complex estuary. This work shows that it is possible to reproduce the
 main system sediment dynamics as well as a detailed budget for complex areas such as the
 Delta using a 2D process based numerical model coupled with a water quality model.
- 17 Overall, the model reproduces the SSC peaks and event timing and duration (wet season) as
- 18 well as the low concentration in dry season throughout the Delta, except at Mallard where
- 19 water column is stratified due to salt intrusion. Stratification issues are not solved in a 2D
- 20 model. For this reason we are working on a 3D model in order to include the Bay area,
- 21 <u>leading to a unique model from source to sink.</u>
- The Delta is well covered by observation stations. However, this work shows that the
 substantial sediment is exported trough the pumping stations (100kt yr⁻¹) at the Southern
 Delta where no data in SSC is available. The sediment exporting needs further investigation,
- 25 since it is possible that has being deposited in the channels before the pumps.
- We show that with simple sediment settings as one fraction at the input boundary and simple distribution of bed sediment availability, it is possible to reproduce seasonal variations as well as define yearly sediment budget with more than 90% of accuracy. It shows also that it is extremely important to have discharge and SSC measurements at least in the input boundaries and close to the system output in order to be able to calibrate the model settings applied for

hydrodynamics and suspended sediment. This methodology now can be applied in less
 measured estuaries.

- 3 Sediment is a key-factor to estuaries water quality and ecology. The D3D FM software allows 4 direct coupling to water quality, sediment transport and habitat modeling. Our work provides 5 the basis to a chain of models, which goes from the hydrodynamics, to suspended sediment, to 6 phytoplankton, to fish, clams and marshes. The turbidity and deposition pattern analysis may 7 guide ecologists in future works to define areas of interest and/or venerable areas to be study, 8 as well as guide data collecting efforts. The present model opens the possibility for forecast 9 and operational modeling. Forecasting the time frame of high levels of SSC (turbidity) allows 10 planning of measurements campaigns for ecologists, as well as the possibility of tracking potentially contaminated sediment and be able to make a contingency plan as well as 11 12 temporary barriers and pumping operations. 13 The Sacramento San-Joaquin Delta is a typical case of a highly impacted estuary. Being able
- to numerically simulate and determine sediment transport, budget and turbidity levels in this
 type of environment open possibilities to better informed political, ecological and
 management decisions including how to respond to climate change and sea level rise. This
 type of model is an important management tool that is applicable to other impacted estuaries
- 18 worldwide.

19 Appendix A: Hydrodynamic Calibration

The hydrodynamic calibration was carried out for 3 month high river flow conditions (December 16, 1999 until March 16, 2000) and a 3 month period of low river flow conditions (July 16, 2001 until October 16, 2001). All presented data is with respect to NAVD88 (vertical datum), UTM 10 (horizontal datum) and GMT (time reference).

Hourly measured water levels at Point Reyes (tidesandcurrents.noaa.gov/) were used as seaward boundary condition. Landward boundary conditions for the Sacramento River were obtained from daily measured river flow data at Freeport (FPT) and for the San Joaquin River near Vernalis (VNS) (cdec.water.ca.gov/). The inflow from the Yolo Bypass was approximated by (Derived after curve fitting data relationship between Qyolo and Qrsac.)

- 29 Measured data for the Bay area were derived from tidesandcurrents.noaa.gov/, for part of the
- 30 Delta from the California Data Exchange Centre cdec.water.ca.gov/ and for station with
- 31 numbers from direct contact with the Department of Water Resources (DWR).

1 Calibration was carried out by systematically varying the value of the Manning's coefficient 2 for different sub-areas of the Bay-Delta system. The calibration data analysis includes (local and time varying) influence of air pressure and wind in the definition of the boundary 3 4 condition as well as in the calibration data inside the modeling domain. These may account 5 for (part of) the error between measurements and modeling results. Also, the NAVD88 reference is not known for all measurement stations, although tidal water fluctuations may be 6 7 modeled properly. To circumvent these distortions a better method to assess the model 8 performance is to focus on water level amplitude and phasing of the different tidal 9 constituents. Boundary conditions, calibration data and model results are thus decomposed by 10 Fourier transformation into tidal components which are then compared. The following table 11 gives the results of this analysis for 34 tidal constituents at Golden Gate (GGT) for high river 12 flow conditions. By far, the main tidal constituents at (GGT) are O1, K1, N2, M2 and S2, 13 with M2 being the largest. The model represents their values quite well. The difference in 14 amplitude is 1.3 % for M2, up to 14% for O1, but the phasing shows a maximum of only 3% 15 (01)).

The Fig A1 gives calibration results for the high and low river flow. The largest (extreme) deviations are explained by the fact that the measured water levels did not have a known reference to NAVD88 (http://www.d3d-baydelta.org/).

19

20 Appendix B: SSC Calibration

All stations clearly reproduce SSC peaks during high river flow periods and lower concentrations during the remainder of the year (apart from MAL during the July-August period). The good representation of the peak timing means that the main Delta event is reproduced by the model as well as the periods of Delta clearance. These two periods are critical for ecological models, and a good representation generates robust input to ecological models. A closer look at Fig 4 reveals differences between model results and data. Following these differences is discussed station by station in this appendix.

One observes that at RVB, SSC levels are directly proportional to Sacramento River discharge (Fig B3), and that the model properly represents the water discharge peak intensity and duration. However, in the model, the first peak remobilizes sediment faster than observed

1 in the data. Analyzing the raw data, it is possible to observe a trend of SSC increase which the 2 model overestimates. A probable explanation lies on the initial sediment composition of the bed. Defining the bottom sediment composition does not account for consolidation processes; 3 4 so the first peak comes after the dry season when the mud in the banks has consolidated. In 5 the simulation case, when river discharge increases, it remobilizes non-consolidated bottom/bank sediment causing an earlier peak than in the data; similar behaviour is observed 6 7 in STK in December. Sediment trapped in subaquatic vegetation and marshes could be 8 another explanation for the slower increase of the first peak as the model discharges for both 9 stations agree with data (Fig 4).

10 Another difference between the data and the model results in RVB is the peak in May (second rectangle, Fig B1), which is not observed in the data. SSC level at RVB station is directly 11 12 proportional to water discharge in FPT (Fig B3, RVB). The May peak is observed in FPT and so should have been transported towards RVB just as the two preceding peaks. However, 13 14 the data set does not reproduce this peak. One of the possible explanations is errors in data 15 meassurements, since it comes after a major event and the equipment might be damaged. 16 Other explanations could be a different composition of the suspended sediment properties 17 and/or floculation.

18 The model underestimates the first and second SSC peaks at MOK. However, the data SSC 19 signal is not consistent with the local water discharge signal. First, we checked that modeled 20 water discharge is reproducing the local conditions, where data is available from mid-February onwards. The last peak in Fig (mid-March) shows that water discharge, in situ and 21 22 modeled SSC have the same rage of variation. Therefore the SSC levels are proportional to 23 the local water discharge. Backwards in time, the January SSC data peak is much higher than 24 the water discharge and the SSC level calculated in the model. The same happens in mid-25 February when no water discharge peak is observed but there is a peak in the SSC data. Again the peaks in SSC could be derived from an error in the measurements or local, diffuse input of 26 27 sediment such as from local farm waste water or biological activity remobilizing the substrate.

The model represents well the wet season SSC peaks at MAL; however, during the three drier periods of the year the model underestimates SSC levels (Fig_B2). From the scatter plot water discharge versus SSC (Fig_B3Error! Reference source not found.), it is possible to explain the weaker performance of the model during low river flow at MAL. These graphs represent river water discharge in FPT lagged by 2 days to SSC in RVB and MAL. Several time lags were tested, as MAL does not present a reasonable correlation with any of the time lags; it is presented here with the same time lag as the one for RVB. RVB station reflects a positive correlation between river discharge and SSC derived from in situ data and model results. The correlation coefficient (R), statistically shows how two variables are correlated, in RVB R=0.58.

6 In MAL station R=0.26, showing that there is almost no correlation between river discharge 7 and SSC levels. The low correlation is due to high SSC level in low water discharge periods, 8 when the model underestimates SSC levels. Under low river discharges conditions, salt water 9 intrudes into Suisun Bay leading to considerable stratification between fresh and salt water 10 and shifting of the ETM landward (http://sfbay.wr.usgs.gov/access/wqdata/) (Brennan et al., 11 2002). In order to better model SSC levels at these conditions a 3D model would be needed to 12 reflect conditions at MAL adequately. With this results we are still able to calculate sediment export, since most of the sediment export occurs in the wet period (McKee et al., 2006), 13 14 when the model reproduces SSC levels.

15

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1 Table 1: Parameters set of sensitivity analysis.

Parameters		Minimum	Maximum
standard		w = 0.25; Tau = 0.25	25; E=1*10-4
Fall velocity	w (mm s ⁻¹)	0.15	0.38
Critical shear stress	τ_{cr} (Pa)	0.125	0.5
Erosion Coefficient	E (kg m ² s ⁻¹)	2.5*10 ⁻⁵	1*10 ⁻²





3 indicate measurement stations.





Fig 2: Numerical mesh, the red dots indicate the calibration stations. (http://san-francisco-baydelta-model.unesco-ihe.org/). Zoom in the computational grid, A) San Pablo Bay connecting
to Petaluma and Napa Rivers, B) Delta channels and Franks Tract.



4 San Joaquin River at VNS and Mokelumne River at Woodbridge respectively.



the show discharge. Data are dashed red lines and model results are solid blue lines. Note th
in the discharge plots of RVB and STK the data line is behind the model line.





4 Fig 5: Sensitivity analysis of sediment flux, for 3 stations, RVB in green squares (Sacramento

- 5 River), STK in red triangles (San Joaquin River) and MAL in blue circles (Delta output). The
- 6 colour lines indicate the data values.









Fig 7: Anomaly of a SSC (mg L⁻¹) snapshot between a run with open and one with closed
DCC, this pattern is representative in time as well. The right panel is a zoom in between the
DCC and MOK (black rectangle). Red shades represent regions where the SSC level is higher

5 in the open than the close scenario, the blue shades where it was lower.

6





В

- 1 Fig 8: Water Discharge (A) and Sediment Flux (B) Path models. The arrows represent the
- 2 water (A) and sediment (B) fluxes through the cross sections. Area of the arrow is
- 3 proportional to the flux. Red/blue indicates the sediment (water) fluxes from data/model.
- 4 Inside each polygon is the sediment budget for the area. The Bay portion is dashed because
- 5 the model is 2D and 3D processes occur in that region.





2 Fig 9: Sediment bottom composition after one year, starting from concrete bed. In Red shades

3 region dominated by Sacramento and in white by San Joaquin, the black line highlights this

4 separation.



Fig 10: Flux calculation for several stations within the Delta. Figs A, C, E, and G show the
flux change following Sacramento branch and B, C, F and H following San Joaquin branch.
The total flux is represented in magenta (in FPT and VNS the total is the same as the

- 1 advective), in blue the advective flux, in red the dispersive flux and in green Stokes drift.
- 2 (Positive is seaward).



2 Fig 11: DELWAQ deposition volume translated in bottom sediment deposition in [mm].



Fig 12: Turbidity histogram for each Delta area, the left hand side bars indicates wet season
and the ones in the right dry season. The light gray bar indicate the mean turbidity over the
area, the darker bar the spatial deviation and the lines the daily deviation. Each horizontal line
represents 10 ntu.



3 Fig A1: Hydrodynamic calibration example.



2 Fig B1: Comparison between SSC levels in RVB station in situ data (dashed red) and model

3 result (solid blue) and FPT station (dotted green).





- 2 Fig B2: Water discharge (model) and SSC level (data and model) in MOK station.





2 Fig B3: Scatter plot Discharge versus SSC. Showing on the left (MAL) for MAL station

- 3 and on the right hand (RVB) side RVB station. The red dots represent the Data and the
- 4 blue model results.