Please find our responses in italic typeface below the editor's comments in regular typeface.

Our sincere apologies for uploading the wrong manuscript, we have ensured that this upload conforms to the changes reflected in the previous and this minor revision. We have included a highlighted version after this letter indicating the three major changes made in this revision (the two changes listed in the last revision in section 4.3 and the change to the Figure 2 caption).

With respect to your response: I do not see the stronger diurnal variations at the end of the melt season 1 that you are referring to here "note the greater diurnal variation in melt season 1 during the end-of-season low flow, Figure 4".

Do you mean that they are greater than in the first half of the melt season? I do not see that. The fluctuations at the end of the season are about 100m or less, while in the first half they can also be more than 100m.

We apologize for this confusion. We meant to indicate that the entire melt season 1 was endof-season low flow and shows considerable diurnal variation- mostly compounded classification error. We agree that this is written in a confusing manner, and have simply dropped the parenthetical phrase, as the paragraph is still comprehensible without it.

Please also mention the fact that there are some misclassified areas in the last image of Figure 2 either in the text or in the caption of the figure.

As requested, we have noted that there are misclassified areas in the final panel of figure 2 in the caption. We write, "The final classification as shown the bottom panel has some obvious errors, including speckling and some misclassifications of both water and non-water."

- 1 **Title:** Technical Note: Semi-automated effective width extraction from time-lapse RGB
- 2 imagery of a remote, braided Greenlandic river.
- 3 **Authors**: Colin J. Gleason<sup>1</sup>, Laurence C. Smith<sup>1</sup>, David C. Finnegan <sup>2</sup>, Adam L. LeWinter <sup>2</sup>,
- 4 Lincoln H Pitcher <sup>1</sup>, Vena W. Chu<sup>1</sup>

## 5 **Author affiliation**:

23

- 6 1) Department of Geography, University of California-Los Angeles. 1255 Bunche Hall, 405
- 7 Hilgard Avenue, Los Angeles, California 90095-1524
- 8 2) U.S. Army Cold Regions Research & Engineering Laboratory, Hanover, NH 03755
- 9 **Corresponding Author**: C.J. Gleason, cjgleaso@ucla.edu
- **Abstract:** River systems in remote environments are often challenging to monitor and 10 11 understand where traditional gauging apparatus are difficult to install or where safety concerns prohibit field measurements. In such cases, remote sensing, especially terrestrial 12 13 time lapse imaging platforms, offer a means to better understand these fluvial systems. One such environment is found at the proglacial Isortoq River in southwest Greenland, a river 14 with a constantly shifting floodplain and remote Arctic location that make gauging and *in* 15 situ measurements all but impossible. In order to derive relevant hydraulic parameters for 16 17 this river, two RGB cameras were installed in July of 2011, and these cameras collected over 10,000 half hourly time-lapse images of the river by September of 2012. Existing 18 19 approaches for extracting hydraulic parameters from RGB imagery require manual or 20 supervised classification of images into water and non-water areas, a task that was 21 impractical for the volume of data in this study. As such, automated image filters were developed that removed images with environmental obstacles (e.g. shadows, sun glint, 22

snow) from the processing stream. Further image filtering was accomplished via a novel

- automated histogram similarity filtering process. This similarity filtering allowed
- 2 successful (mean accuracy 79.6%) supervised classification of filtered images from training
- data collected from just 10% of those images. Effective width, a hydraulic parameter highly
- 4 correlated with discharge in braided rivers, was extracted from these classified images,
- 5 producing a hydrograph proxy for the Isortog River between 2011 and 2012. This
- 6 hydrograph proxy shows agreement with historic flooding observed in other parts of
- 7 Greenland in July 2012 and offers promise that the imaging platform and processing
- 8 methodology presented here will be useful for future monitoring studies of remote rivers.

# 1. Introduction

9

10

11

12

13

14

15

16

17

18

19

20

21

22

Proglacial streams and rivers along land-terminating edges of the Greenland Ice Sheet are among the world's most difficult fluvial systems to study in the field, owing to their remoteness, harsh climate, and braided morphology. Discharge variations in large proglacial rivers are of particular scientific interest, as these systems typically derive water from the interior ablations surface Greenland Ice Sheet and are thus useful for inferring runoff mass losses from the ice sheet (Rennermalm et al., 2013; Smith et al. 2014). However, their high sediment loads, unstable banks, and dynamic braided channels present challenges to traditional *in situ* river gauging techniques, and long term hydrographs for these rivers are rare. While not unique to Greenland, these challenges are particularly evident there, with more than 100 large (> 1 km width) large braided rivers exiting the ice sheet with no observations of discharge whatsoever.

Where *in situ* methods are impractical, remotely sensed imagery offers an increasingly viable option for obtaining scientifically useful estimates of river discharge in

- remote or otherwise inaccessible areas (Smith et al., 1997, Ashmore and Sauks, 2006,
- 2 Durand et al., 2010, Gleason and Smith, 2014). Braided rivers in particular typically display
- 3 a power-law relationship between floodplain inundation area (which can be remotely
- 4 sensed) and discharge, which has been exploited using satellites, aerial imagery, and
- 5 terrestrial time-lapse photography (Smith 1995; 1996, Chandler et al., 2002; Ashmore and
- 6 Sauks, 2006; Egozi and Ashmore 2008; Smith and Pavelsky, 2008; Bertoldi *et al.*, 2009;
- Hundey and Ashmore, 2009; Bertoldi et al., 2010; Bird et al., 2010; Ashmore et al., 2011;
- 8 Welber *et al.*, 2012; Williams *et al.*, 2013; Young *et al.*, 2015).

10

11

12

13

14

15

16

17

18

19

20

21

22

23

Regardless of the technology used, each remotely sensed image must first be classified into areas of water and non-water, a task for which numerous methodologies exist. In satellite remote sensing, NIR wavelength image bands can reliably detect open water surfaces. However, satellite imagery often lacks the required spatial and temporal resolution to adequately capture hydrologic phenomena, especially for smaller rivers. This has led to the use of non-metric, true color (RGB) digital camera imagery to capture water surfaces as an inexpensive and image-on-demand alternative to satellite and airborne platforms, especially for braided rivers. To calculate hydraulic parameters (e.g. effective width, braiding index, sinuosity, or bed slope elevation), these studies have commonly classified water surfaces within images either manually or by supervised classification (Egozi and Ashmore 2008; Bertoldi et al., 2009; Hundey and Ashmore, 2009; Ashmore et al., 2011; Welber et al., 2012). Another parameter estimation approach relies on water surface delineation from automatically generated DEMs constructed from stereo-imagery and other data sources (Chandler et al., 2002; Ashmore and Sauks, 2006; Bird et al., 2010; Bertoldi et al., 2010). Additionally, Young et al (2015) recently demonstrated the

- 1 effectiveness of calculating water stage change at a station from terrestrial
- 2 photogrammetry, which they combined with assumptions of channel geometry and
- 3 roughness to calculate river discharge via Manning's equation. This approach is highly
- 4 effective, but limited to situations where bathymetry is known or channel geometry may be
- 5 simply described. Finally, structure-from-motion, a technique that leverages multiple
- 6 vantage points of the same scene to reconstruct topography, has also been successfully
- 7 leveraged to calculate floodplain geometry and water surface elevation, but is again
- 8 impractical for long term monitoring with large data volumes (e.g. Fonstad et al., 2013,
- 9 Javernick *et al.*, 2014).

11

12

13

14

15

16

17

18

19

20

21

22

While each of these studies successfully calculated hydrologic parameters from remotely sensed images, their manual or time-intensive approaches are impractical for large data volumes. This is especially an issue for long term hydrologic monitoring sorely needed in many remote rivers, as using the image platform and processing developed by Ashmore and Sauks(2006) and Welber *et al.* (2012), for instance, could easily generate tens of thousands of images per year. Automated DEM generation methods would seem a ready alternative, yet these require numerous fixed targets of known position to persist from image to image, which are seldom found or are difficult to install on dynamic braided river systems owing to their constantly shifting morphology. If such image platforms are to be viable for long term monitoring studies, a systematic procedure for automatic image quality selection and classification, preferably for RGB image data, is needed.

To that end, this paper proposes a semi-automated processing stream designed to classify and extract hydraulic parameters of interest from large volumes of RGB image data

- collected from a fixed terrestrial platform, and demonstrates its efficacy in a remote
- 2 Greenlandic river. Automated filters are developed that remove obstacles to image
- 3 classification based on easily calculated environmental variables, and an image similarity
- 4 filter is developed that allows supervised classification of many images from minimal
- 5 training data. Here, these filtering and classification techniques are employed to extract
- 6 effective width (W<sub>e</sub>, inundation area divided by reach length), a hydraulic parameter that
- 7 has been shown to be highly correlated with discharge in braided rivers and has been
- 8 successfully extracted from remotely sensed data in proglacial environments (Smith et
- 9 *al.*,1996; Smith, 1997; Ashhmore and Sauks, 2006; Smith and Pavelsky, 2008; Ashmore *et*
- 10 *al.*, 2011). To evaluate the robustness of the extraction, we assess image classification
- accuracy using manually generated ground truth data.

# 2. Data

12

- This study was conducted on the proglacial Isortog River in southwestern
- Greenland. The Isortog, one of the largest braided rivers draining the Greenland ice sheet,
- issues from the Issunguata Sermia glacier terminus with discharge dominated by
- meltwater outflow from the ablating ice surface (Smith et al., 2014). In July 2011, two
- Nikon D200 model RGB cameras (focal lengths of 24 and 50mm) were installed 250m
- above a reach of the Isortog braid plain approximately 3.1 km downstream of the ice edge.
- 19 The camera system was identical to that developed by the Extreme Ice Survey project
- 20 (www.extremeicesurvey.org) for use in severe Arctic conditions. In addition to the
- cameras, a modified battery pack and electronic controller were housed inside a
- 22 weatherproof case with an abrasion-resistant viewing window. The case was mounted on

a survey tripod and powered by a 12V gel battery recharged by solar panel. The cameras

were oriented so as to image sections of the braid plain of approximately 1.5km x 2.0km

and 2.0km x 2.3km, respectively (Figure 1), and captured one image every 30 minutes

when light conditions permitted.

Camera data collection commenced July 22nd, 2011, and over 10,000 images were retrieved from the cameras by September 10th, 2012, covering most of two melt seasons. The camera setup proved robust: the light sensor operated properly, the position of the cameras remained unchanged, and the batteries powering the cameras were still functional after the one year collection period for the wide focus camera. However, a presumed Arctic fox chewed through the cables connecting the battery to the camera for the more narrowly focused platform and halted data collection only two months after installation. Therefore, all analyses presented in this paper refer to the wide focus camera, which remained continuously operable throughout the study period July 22nd 2011 – September 10th 2012.

#### 3. Methods

Classifying the RGB image data into water and non-water areas to extract  $W_e$  presented several technical challenges for the 10,327 images that were collected by the wide focus camera from July 2011 to September 2012. Existing approaches for hydraulic parameter extraction from RGB data require either manual or supervised classification of water within each image and are thus inappropriate for the large data volumes generated in this study. Unsupervised classification techniques provide a straightforward alternative for large time-lapse camera datasets, yet also present additional challenges as the images collected here are extremely diverse and differing soil moisture in the braid plain gives the

- appearance of multiple classes of output. Environmental factors such as time-varying solar
- 2 angles, blowing sand, dense fog, shadowing, snow and rain on the camera lens, and acute
- 3 sun-glint from water surface are especially prevalent in the Isortoq image data. These
- 4 factors were all addressed, and W<sub>e</sub> accurately extracted, by the processing workflow
- 5 described below and presented in Figure 2.

# 3.1 Environmental Filtering

6

7

8

9

10

11

12

13

14

15

16

17

18

19

20

21

22

The first task for extracting W<sub>e</sub> was to filter the large amount of image data into those images that were most easily classified into water and non-water areas by eliminating images containing the environmental obstacles described above. Once images are classified, water area (and therefore W<sub>e</sub>) may be calculated. Several filters were developed to remove these poor quality images. First, images acquired during periods of non-flow (before and after melt season activity) were culled. Next, images with shadowing were culled by calculating the zenith and azimuth angles of the sun relative to the river plain. Through visual inspection of the image time series, zenith angles less than 65 degrees and azimuth angles between 245-290 and between 70-100 degrees were found to produce shadows created by steep valley walls that prevented accurate classification (note valley walls, Figures 1 and 2). Next, images that exhibited excessive sun glinting were removed. Sun glint was defined as when an image exhibited either a ratio of the 95th brightness percentile to the 5th brightness percentile greater than 1.8 or contained more than 1% of pixels with brightness value greater than 215. This filter was necessary, as sun glint was observed both on open water and saturated sand, making distinction between these very different fluvial environments difficult (Figure 2). Successful application of these winter,

- shadow, and sun glint filters culled 9,487 images from the image time series, leaving 840
- 2 images free of environmental obstacles that still captured every day of the two melt
- 3 seasons.

### 3.2 Similarity Filtering

Even with these stringent filters, unsupervised classification was still unable to delineate water surfaces with satisfactory accuracy, and the number of images remaining was still too large for supervised classification to be feasible. As such, a semi-supervised classification approach was developed. To perform this classification, another image filtering was needed to find images that were similar enough to one another to share training data from a small sample of images in a supervised classification. The presence of dense fog, blowing sand, or cloudiness changes the brightness values of the imagery, so even images collected with identical solar geometry can be difficult to classify in an unsupervised manner. A similarity filter was developed that selected images that not only had similar solar geometry, but also had the same brightness and illumination and were all free of environmental obstacles not covered by the first filtering.

This similarity filtering was accomplished by calculating and comparing the histograms of each of the red, green, and blue bands for each image. Histograms of brightness values that fell into 100 bins evenly spaced from 0 to 255 (reflectance values) were calculated for each band of each image. Using the same bins for each image ensured that cross comparison of images would not be affected by stretching of the image data. Once these histograms were generated, the root mean square error (RMSE) between histogram counts per bin was computed in a band-by-band pairwise permutation, giving a per-image and per-band indication of the similarity of every image to each other image.

The pairwise permutation tests all possible image pairs for similarity. That is, for any given image, the histogram bin counts in each of its RGB bands is compared against bin counts of every other image and the RMSE (across all bins) of each comparison is recorded. Then, the process is repeated for every other image in the set, which yields (n²-n)/2 RMSE values per image, where n is the number of images. These band-by-band RMSE values were then averaged to arrive at an overall measure of image similarity: here termed an image's similarity index. This metric was used to identify the 20% of the images that were most similar to each other, resulting in 168 images that were collected at similar sun angles without any environmental obstacles. Importantly, the similarity filter also produced images that contained four basic elements: dark (non-sun lit, turbid) water, bright (sun lit or non-turbid) water, dark (wet) sand, and bright (dry) sand (see Fig. 1c), thus producing images easily classified from lumped training data- a process described next.

## 3.3 Georectification and classification

Once the final filtering of images was complete, images were cropped to exclude the wide upstream floodplain and georectified into ground coordinates using a 4<sup>th</sup> degree polynomial transformation implemented in ENVI v4.8 (Figure 2). Eighty ground control points were manually extracted from a 2 m panchromatic World View 2 image acquired on September 23<sup>rd</sup>, 2011 (paired with a camera image collected 10 minutes later) and used to define the basis for the transformation. This warping polynomial was subsequently applied to all filtered images. After georectification, each image pixel had dimensions of 1m by 1m, an appropriate resolution for camera data collected at this scale. These georectified pixels allowed calculation of water surface area, and thus W<sub>e</sub>, from the classified images.

To classify images into water and non-water areas for  $W_e$  extraction, training data representing four classes (dark water, bright water, dark sand, and bright sand) were manually collected from a random 10% sample (16 images) of the similarity filtered images. The RGB statistics generated from these training polygons were applied to all images passing the similarity filtering and used to train a maximum likelihood supervised classification method performed in ENVI v4.8 for each image. This process requires that each image has nearly identical RGB composition in order to be successful, which was guaranteed by the similarity filtering.

#### 4. Results and discussion

## 4.1 Image Filtering

The environmental and similarity filters developed in this study substantially reduced the number of images available for  $W_e$  extraction from image collection to classification. The automated environmental filtering removed 9,487 images with sun glint, shadowing, or winter conditions, leaving 840 images for further operations. The similarity filtering further reduced the image pool to 168 images that were ultimately passed to classification and  $W_e$  extraction. This is obviously a large percentage of images removed, but this stringent filtering left only very high quality images that were easily classified using the semi-supervised approach. However, this high degree of culling still left images with daily (or better) temporal resolution available for  $W_e$  extraction. If hourly or better resolution images are needed, then the similarity filtering would need to be performed on iterative batches of images, as there are other groups of images similar to one another that are not similar to all images as a whole that are removed by the similarity filter. Each of these groups could also be classified using their own lumped training data and output

1 classes determined by their composition. This would extend the temporal coverage of the

2 record, but since the similarity filter we propose yielded near daily coverage of the river we

felt this simplest case to be sufficient for the river in this study and did not identify further

groups of similar images.

Water turbidity could have effected this successful filtering. As sediment load and river velocities change, water can appear darker or brighter depending on river turbidity, thus affecting our choice of two water classes ('dark' and 'bright'). In the Isortoq, the monitoring section is very close to the glacial terminus (~3.1km), and as such the sediment load is fairly constant, the river well mixed, and sediment relatively unsorted, so 'bright' water corresponds to sunlight water, rather than less turbid water. Given these conditions, the two classes do cover nearly all the turbidity values observed in the Isortoq River after image similarity filtering. In rivers with more variable turbidity or places where the bed is visible at low flows, more water/non-water classes and different filters might be needed to adequately cover the range of observed sediment loads.

# 4.2 Accuracy assessment

The semi-supervised classification described here proved an effective and unbiased classification method. Figure 3 shows the overall accuracy, user's accuracy for water, and user's accuracy for non-water as a function of  $W_e$  from a random sample of 56 images (33% of filtered images). Accuracy was assessed using approximately 500 semi-random, manually derived assessment points for each class (water and non-water) per image. Of particular interest were both the overall accuracy (total number of correctly classified assessment points divided by total number of assessment points,  $\sim$ 500), and the user's accuracy for water and non-water (percentage of image pixels classified correctly as

- assessed by the training data). These metrics provide an assessment of classification
- 2 performance from the standpoint of each classified image: the paradigm that speaks
- 3 directly to the fidelity of extracted W<sub>e</sub>. Accuracy assessment indicates that overall accuracy
- 4 is acceptable (mean accuracy for the assessment sample is 79.6%), and neither overall
- accuracy (r = -0.11) nor water user's accuracy (r = 0.35) show strong correlation with W<sub>e</sub>.
- 6 This lack of correlation indicates that the classification of water is not affected by the extent
- of water inundation in the scene. There is a strong correlation (r=-0.79) between the user's
- 8 accuracy of non-water pixels and W<sub>e</sub>, but this negative correlation is a reflection of the
- 9 difficulty of classifying the small number of non-water pixels remaining in scenes where
- the braid plain was nearly completely flooded. The reason for this successful classification
- was the similarity of filtered images, which was guaranteed by the similarity index
- procedure described above. After classification, We was calculated as the area of classified
- water within a 1,000m reach located where the image data provided complete bank to
- bank coverage, indicated by the magenta polygons (dashed) in Figure 2.
- 15 *4.3 Extracted* W<sub>e</sub> hydrograph
- The W<sub>e</sub> hydrograph shown in Figure 4 is a proxy for discharge variations in the
- 17 Isortog River from 2011-2012. Gaps in the date record indicate that there were no images
- that passed filtering on those dates, even though images were acquired half hourly. In the
- first melt season, there were 30 days with missing data over a total melt season length of
- 49 days, but this includes a 15 day gap in late August where there is no data due to
- 21 inclement weather. The second melt season has better temporal coverage, with only 31 of
- 22 104 days missing. This miss rate of about one-third would occur with or without similarity
- 23 filtering, as the majority of these data gaps are due to rain, fog, and snow events that

- 1 preclude classification by any means: this is an issue for any high latitude camera-based
- 2 study. Despite these gaps, the data record still provides good temporal coverage and allows
- analysis of the W<sub>e</sub> hydrograph.
- 4 Historic melting of the Greenland ice sheet occurred in July of 2012 (Hall et al.,
- 5 2013; Tedesco et al., 2013), coinciding with destruction of the Watson River bridge in the
- 6 town of Kangerlussuaq (Smith et al., 2014), located approximately 15km south of the
- 7 Isortoq River. This flood event is clearly evident in Figure 4, as are the rising and falling
- 8 limbs of the hydrograph leading up to this event. Figure 4 also reveals that the relative
- 9 magnitude of W<sub>e</sub> during this melt event was an order of magnitude greater than W<sub>e</sub> in low
- 10 flow stages. This shows that the Isortog River behaves like other braided rivers with non-
- 11 cohesive bed material, as its width adjusts rapidly to changing discharge. In addition, the
- peak W<sub>e</sub> observed here corresponds to almost complete floodplain occupation by the river,
- highlighting the difficulty of installing traditional gauging equipment at this site. These
- variations in W<sub>e</sub> as the melt season progresses are detected even though diurnal variations
- in W<sub>e</sub> can be quite large: melting of the Greenland ice sheet has a strong diurnal forcing
- reflected in Figure 4. Time of day effects are minimized via the similarity filtering (which
- leaves images with similar solar geometry), but measurable changes in W<sub>e</sub> are evident
- despite this insolation matching and are compounded by classification errors. However, the
- 19 filtering and classification procedures here ultimately yield W<sub>e</sub> values that effectively
- capture both diurnal and day-to-day variation in the Isortog River. For the full melt season
- captured in 2012, the W<sub>e</sub> hydrograph has good temporal coverage and diurnal variations
- are small enough so that the larger trends in melting are clearly evident and align with
- expected melt activity in that year.

#### 5. Conclusions

This paper has demonstrated the efficacy of a fixed position RGB time-lapse camera platform for hydraulic parameter extraction for a large proglacial braided river in a remote area of Greenland. The operational camera delivered over 10,000 half hourly images in just over one year of collection, and demonstrated remarkable climactic resilience in the Greenlandic winter. The other camera, however, was lost to a wildlife attack, pointing to the need for stronger housing for all camera components. Such a platform is useful for extraction of multiple hydraulic parameters, including effective width  $(W_e)$ , a proxy for discharge variations. To fully realize this monitoring potential, the  $W_e$  variations extracted for each image could be calibrated with a rating curve built from intermittent field data.

The above accuracy assessments indicate that the semi-supervised classification method produced accurate and unbiased results. An accurately delineated water surface is necessary to preserve the fidelity of extracted hydraulic parameters. The processing techniques described in this paper fall short of completely automated processing, yet this paper does present an analysis protocol that achieves a consistent standard of classification from images that are automatically selected for ease of classification. Furthermore, the similarity filtering presented herein allows for supervised classification of numerous images from minimal training data, enabling long term hydrologic records to be maintained without onerous manual classification of imagery or photogrammetrically challenging DEM extraction.

**Acknowledgements:** This research was supported by the NASA Remote Sensing Theory initiative (grant NNX12AB41G), NASA Cryosphere Program (grant NNX11AQ38G) managed

- by Dr. Thomas Wagner, and NASA Earth and Space Sciences Fellowship NNX12AN32H.
- 2 Field logistical support was provided by CH2M Hill Polar Field Services, the Kangerlussuaq
- 3 International Science Station (KISS), and Air Greenland.

# 4 6. Works cited

- 5 Ashmore, P., and Sauks, E.: Prediction of discharge from water surface width in a braided
- 6 river with implications for at-a-station hydraulic geometry, Water Resources Research,
- 7 42, 10.1029/2005wr003993, 2006.
- 8 Ashmore, P., Bertoldi, W., and Gardner, J. T.: Active width of gravel-bed braided rivers,
- 9 Earth Surface Processes and Landforms, 36, 10.1002/esp.2182, 2011.
- Bertoldi, W., Zanoni, L., and Tubino, M.: Planform dynamics of braided streams, Earth
- 11 Surface Processes and Landforms, 34, 547-557, 10.1002/esp.1755, 2009.
- Bertoldi, W., Zanoni, L., and Tubino, M.: Assessment of morphological changes induced by
- flow and flood pulses in a gravel bed braided river: The Tagliamento River (Italy),
- Geomorphology, 114, 348-360, 10.1016/j.geomorph.2009.07.017, 2010.
- Bird, S., Hogan, D., and Schwab, J.: Photogrammetric monitoring of small streams under a
- riparian forest canopy, Earth Surface Processes and Landforms, 35, 952-970, 2010.
- 17 Chandler, J., Ashmore, P., Paola, C., Gooch, M., and Varkaris, F.: Monitoring river-channel
- change using terrestrial oblique digital imagery and automated digital photogrammetry,
- Annals of the Association of American Geographers, 92, 631-644, 10.1111/1467-
- 20 8306.00308, 2002.

- Durand, M., Rodriguez, E., Alsdorf, D.E., and Trigg, M. Estimating river depth from remote
- 2 sensing swath interferometry measurements of river height, slope, and width. IEEE
- 3 Journal of Selected Topics in Applied Earth Observations and Remote Sensing, 3, 20-31,
- 4 10.1109/jstars.2009.2033453, 2010.
- 5 Egozi, R., and Ashmore, P.: Defining and measuring braiding intensity, Earth Surface
- 6 Processes and Landforms, 33, 2121-2138, 10.1002/esp.1658, 2008.
- Fonstad, M.A., Dietrich, J.T., Courville, B.C., Jensen, J.L., and Carbonneau, P.: Topographic
- 8 structure from motion: a new development in photogrammetric measurement. Earth
- 9 Surface Processes and Landforms, 38, 421-430. 10.1002/esp.3366. 2013.
- Gilvear, D. J., Davids, C., and Tyler, A. N.: The use of remotely sensed data to detect channel
- hydromorphology; River Tummel, Scotland, River Research and Applications, 20, 795-
- 12 811, 10.1002/rra.792, 2004.
- Gleason, C.J., and Smith, L.C. Toward global mapping of river discharge using satellite
- images and at-many-stations hydraulic geometry. Proceedings of the National Academy of
- 15 Sciences, 11, 4788-4791, 10.1073/pnas.1317606111, 2014.
- Hall, D.K., Comiso, J.C., DiGirolamo, N.E., Shuman, C.A., Box, J.E., and Koenig, L.S. Variability
- in the surface temperature and melt extent of the Greenland ice sheet. Geophysical
- 18 Research Letters, 40, 2114-2120, 10.1002/grl.50240, 2013.
- 19 Hundey, E. J., and Ashmore, P. E.: Length scale of braided river morphology, Water
- 20 Resources Research, 45, 10.1029/2008wr007521, 2009.

- 1 Javernick, L., J. Brasington, and Caruso, B.:, Modelling the topography of shallow braided
- 2 rivers using Structure-from-Motion photogrammetry, Geomorphology, 213, 166-182,
- 3 10.1016/j.geomorph.2014.01.006. 2014.
- 4 Rennermalm, A.K., Smith, L.C., Chu, V.W., Box, J.E., Forster, R.R., Van den Broeke, M.R., Van
- As, D., and Moustafa, S.E. Evidence of meltwater retention within the Greenland ice sheet.
- 6 The Cryosphere, 7, 1433–1445, 10.5194/tc-7-1433-2013, 2013.
- 7 Smith, L. C., Isacks, B. L., Bloom, A. L., and Murray, A. B.: Estimation of discharge from three
- 8 braided rivers using synthetic aperture radar satellite imagery: Potential application to
- 9 ungaged basins, Water Resources Research, 32, 2021-2034, 10.1029/96wr00752, 1996.
- Smith, L. C.: Satellite remote sensing of river inundation area, stage, and discharge: A
- review, Hydrological Processes, 11, 1427-1439, 10.1002/(sici)1099-
- 12 1085(199708)11:10<1427::aid-hyp473>3.3.co;2-j, 1997.
- Smith, L. C., and Pavelsky, T. M.: Estimation of river discharge, propagation speed, and
- hydraulic geometry from space: Lena River, Siberia, Water Resources Research, 44,
- 15 10.1029/2007wr006133, 2008.
- Smith, L.C., Chu, V.W., Yang, K., Gleason, C.J., Pitcher, Lincoln H, Rennermalm, A.K.,
- Leglieter, C.J., Behar, A.E., Overstreet, B.T., Moustafa, S.E., Tedesco, M., Forster, R.R.,
- LeWinter, A.L., Finnegan, D.C., Sheng, Y., and Balog, J. Efficient meltwater drainage through
- supraglacial streams and rivers on the southwest Greenland ice sheet. Proceedings of the
- National Academy of Sciences, in press, 2014.

- Tedesco, M., Fettweis, X., Mote, T., Wahr, J., Alexander, P., Box, J.E., and Wouters, B.
- 2 Evidence and analysis of 2012 Greenland records from spaceborne observations, a
- regional climate model, and reanalysis data. The Cryosphere, 7, 615-630, 10.5194/tc-7-
- 4 615-2013, 2013.
- Welber, M., Bertoldi, W., and Tubino, M.: The response of braided planform configuration
- 6 to flow variations, bed reworking and vegetation: the case of the Tagliamento River, Italy,
- 7 Earth Surface Processes and Landforms, 37, 572-582, 10.1002/esp.3196, 2012.
- 8 Williams, R. D., J. Brasington, M. Hicks, R. Measures, C. D. Rennie, and Vericat, D.: Hydraulic
- 9 validation of two-dimensional simulations of braided river flow with spatially continuous
- adcp data, Water Resources Research, 49, 5183-5205, 10.1002/wrcr.20391, 2013.
- Young, D. S., J. K. Hart, and Martinez, K.: Image analysis techniques to estimate river
- discharge using time-lapse cameras in remote locations, Computers & Geosciences, 76, 1-
- 13 10,/10.1016/j.cageo.2014.11.008. 2015

- 15 **Figure 1.** Figure 1 shows example images taken on July 17, 2012 of the Isortog River by the
- two camera systems as well as the cameras themselves (foreground and background, panel
- a). The Issunguata Sermia Glacier is seen in the background, and nearly all water in this
- river is derived from its melting terminus. Only the wide focus camera (c) has a continuous
- data record from 2011-2012, as a presumed Arctic fox severed the wiring on the narrow
- 20 focus camera (b). The yellow polygon in the wide focus image shows the target reach for
- 21 W<sub>e</sub> extraction, covering an area of approximately 1,000 by 2,000m.

- Figure 2. The processing steps required to extract W<sub>e</sub> from raw images are shown here.
- 2 Every step until the final classification is completely automated, allowing for a vast
- 3 reduction in processing time. Winter images were selected by a manual inspection of first
- 4 and last observed open water flow. Shadowing was defined as when solar zenith angles
- 5 were less than 65 degrees or solar azimuth between 245-290 or 70-100 degrees, and sun
- 6 glint was defined as a ratio of pixel brightness and as a total pixel value threshold. The final
- 7 classification as shown the bottom panel has some obvious errors, including speckling and
- 8 some misclassifications of both water and non-water. As Figure 4 shows, these filters did
- 9 not significantly affect the temporality of the data and about two-thirds of all days during
- the two melt season study duration are represented.
- Figure 3. Accuracy assessment as a function of W<sub>e</sub> from a 33% sample of post filtered
- images is presented here, with overall accuracy (a), water user's accuracy (b), and non-
- water user's accuracy (c) all showing acceptable performance. Overall accuracy and water
- user's accuracy are not strongly correlated with W<sub>e</sub>, suggesting that the amount of water in
- the scene does not strongly influence the calculation of water area. Non-water accuracy,
- 16 however, is strongly affected by the amount of water in the scene as the Isortoq River
- occupies nearly the entire valley at high flow, making classification of a few scattered non-
- 18 water pixels challenging.
- 19 **Figure 4.** Successful image classification allowed for extraction of W<sub>e</sub> across two melt
- seasons from the wide angle camera and gives a proxy for discharge in the braided Isortog
- 21 River. 22 statistical outliers, representing poorly classified images, were removed before
- 22 generating this figure. These W<sub>e</sub> time series clearly show historic flooding in Greenland in

- 1 July of 2012, as well as the abrupt start of the 2012 melt season, and suggest that the
- 2 camera platform and semi-automated classification techniques advanced here are
- 3 sufficient for monitoring of this remote river.