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# On the validity of effective formulations for transport through heterogeneous porous media

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HESSD

12, 12281–12310, 2015

## Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Abstract

Geological heterogeneity enhances spreading of solutes, and causes transport to be anomalous (i.e., non-Fickian), with much less mixing than suggested by dispersion. This implies that modeling transport requires adopting either stochastic approaches that model heterogeneity explicitly or effective transport formulations that acknowledge the effects of heterogeneity. A number of such formulations have been developed and tested as upscaled representations of enhanced spreading. However, their ability to represent mixing has not been formally tested, which is required for proper reproduction of chemical reactions and which motivates our work. We propose that, for an effective transport formulation to be considered a valid representation of transport through Heterogeneous Porous Media (HPM), it should honor mean advection, mixing and spreading. It should also be flexible enough to be applicable to real problems. We test the capacity of the Multi-Rate Mass Transfer (MRMT) to reproduce mixing observed in HPM, as represented by the classical multi-Gaussian log-permeability field with a Gaussian correlation pattern. Non-dispersive mixing comes from heterogeneity structures in the concentration fields that are not captured by macrodispersion. These fine structures limit mixing initially, but eventually enhance it. Numerical results show that, relative to HPM, MRMT models display a much stronger memory of initial conditions on mixing than on dispersion because of the sensitivity of the mixing state to the actual values of concentration. Because MRMT does not reconstitute the local concentration structures, it induces smaller non-dispersive mixing than HPM. However long-lived trapping in the immobile zones may sustain the deviation from dispersive mixing over much longer times. While spreading can be well captured by MRMT models, non-dispersive mixing cannot.

## HESSD

12, 12281–12310, 2015

### Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

[Title Page](#)

[Abstract](#)

[Introduction](#)

[Conclusions](#)

[References](#)

[Tables](#)

[Figures](#)



[Back](#)

[Close](#)

[Full Screen / Esc](#)

[Printer-friendly Version](#)

[Interactive Discussion](#)



## 1 Introduction

Transport is anomalous in heterogeneous porous media. Anomalous transport observations include tailing in concentration breakthrough curves and plumes, or the strong increase in the rate of spreading of plumes. Several frameworks have been developed to generalize the Advection Dispersion Equation (ADE) and overcome its limitations (Frippiat and Holeyman, 2008). All these alternative frameworks share the goal to model complex permeability, velocity and concentration patterns in unified parsimonious effective equations. The limited number of parameters makes them efficient for the limited quantity of data usually available. In fact, they can be parameterized from breakthrough curves. They comply with the broad residence time distributions and non-local transport processes observed in reality (Gjetvaj et al., 2015; Le Borgne and Gouze, 2008; Willmann et al., 2008). They represent the consequences of complex concentration patterns, of simultaneous concentration trapping and fast progress on residence times while averaging out all the fine concentration structures in the up-scaling process. These anomalous transport frameworks have proven to be highly effective for residence times, transport time distribution and effective spreading both phenomenologically and practically (Berkowitz et al., 2006; Neuman and Tartakovsky, 2009). However, their ability to reproduce mixing, which is required for properly reproducing chemical reactions, has not been tested.

We argue that an effective transport formulation should honor not only the mean advection, and spreading observed in Heterogeneous Porous Media (HPM), but also the evolution of mixing. This should not be understood as limiting anomalous transport frameworks but at extending them to handle broader ranges of physical and chemical processes, and at further promoting the approach of effective equations that upscale out the fine scale structures to retain only their main consequences in terms of transport, reactivity and reactive transport couplings. Here, we investigate the relevance of Multi-Rate Mass Transfer (MRMT) framework to model not only spreading but also mixing. MRMT is taken as a typical anomalous transport framework. Its advantage lies in

# HESSD

12, 12281–12310, 2015

## Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

[Title Page](#)

[Abstract](#)

[Introduction](#)

[Conclusions](#)

[References](#)

[Tables](#)

[Figures](#)



[Back](#)

[Close](#)

[Full Screen / Esc](#)

[Printer-friendly Version](#)

[Interactive Discussion](#)



providing local concentrations, which can be straightforwardly used to evaluate concentration variance, mixing and mixing induced reactivity (Babey et al., 2014; Carrera et al., 1998; de Dreuzy et al., 2013; Haggerty and Gorelick, 1995), as well as the apparent reduction in the rate of kinetic reactions (Dentz et al., 2011). The question is whether its validity as a representation of transport through heterogeneous porous media (HPM) can be extended to reproduce the effects of the evolution of mixing rates resulting from the stretching and folding associated to complex velocity structures (de Anna et al., 2014b; Jimenez-Martinez et al., 2015; Le Borgne et al., 2015).

This comparison is especially appropriate as anomalous transport processes are currently extended to simulate reactive transport processes (Cirpka and Valocchi, 2007; Clement, 2001; de Barros et al., 2012; Donado et al., 2009; Hochstetler et al., 2013; Luo and Cirpka, 2011; Luo et al., 2008; Orgogozo et al., 2013; Schneider et al., 2013). They deal with chemical reactivity either in a stochastic manner, representing reactivity with molecular analogies, or in classical approaches by means of concentrations (Bolster et al., 2010; Cirpka et al., 2012; Ding et al., 2013; Hayek et al., 2012; Knutson et al., 2007; Zhang et al., 2013). Extensions are both required for applications purposes and attractive for capturing the consequences of anomalous transport to potential “anomalous” and enhanced reactivity (Battiato et al., 2009; Sadhukhan et al., 2014; Scheibe et al., 2015; Tartakovsky et al., 2009).

Some assessment of MRMT to model reactivity in HPM has been made in former works (Willmann et al., 2010). Equivalent reactivity has been evaluated at some well-defined travel distances on MRMT calibrated on residence time distributions. Here we follow a different perspective by analyzing the temporal development of spreading and mixing. We extend the integrated assessment of mixing-induced reactivity at given travel distances to its temporal development.

Our contribution concerns the comparison of different models much more than the HPM and MRMT model themselves. For the sake of completeness, we recall model equations and simulation methods in Sect. 2 (models and methods) and measures of spreading and mixing in Sect. 3. We use these measures to propose the conditions

## Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

[Title Page](#)

[Abstract](#)

[Introduction](#)

[Conclusions](#)

[References](#)

[Tables](#)

[Figures](#)



[Back](#)

[Close](#)

[Full Screen / Esc](#)

[Printer-friendly Version](#)

[Interactive Discussion](#)



that should be met by effective (upscaled) transport formulations to be considered valid representations of transport through heterogeneous porous media (Sect. 4). We then test whether MRMT formulations meet the proposed conditions (Sect. 5). While this last section depends on the specific choice of the MRMT framework as an equivalent transport model, the comparison methodology is independent of it and can be used to assess transport equations respecting both spreading and mixing.

## 2 Model and methods

We present sequentially the Multi-Rate Mass Transfer (MRMT) and Heterogeneous Porous Media (HPM) models. As they are both well known, we present only the main equations and highlight the critical assumptions of importance in this study.

### 2.1 Multi-Rate Mass Transfer model (MRMT)

Multi-Rate Mass Transfer models express anomalous transport by the interaction between transport in a mobile zone and a series of immobile zones (Haggerty and Gorelick, 1995). Transport in the mobile zone is advective and dispersive with a mean solute velocity  $v$  (water flux divided by mobile porosity,  $\phi$ ) and a dispersion coefficient  $d$ . Each immobile zone  $i$  is parameterized by a characteristic rate  $\alpha_i$  (inverse of a characteristic exchange time) and an immobile porosity  $\phi_i$ . The concentrations  $c$  and  $c_i$  ( $i = 1 \dots N$ ) in the mobile and immobile zones, respectively, are determined by the following set of equations:

$$\phi \frac{\partial c}{\partial t} + \sum_{i=1}^N \phi_i \frac{\partial c_i}{\partial t} = -q \frac{\partial c}{\partial x} + d \frac{\partial^2 c}{\partial x^2}. \quad (1)$$

$$\frac{\partial c_i}{\partial t} = \alpha_i (c - c_i) \quad \text{for } i = 1, \dots, N \quad (2)$$

## Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The ratio of immobile to mobile water volumes is rated by the total capacity ratio  $\beta$ . The term capacity derives from the fact that MRMT formulations were originally devised to represent trapping by sorption in hard-to-reach sorption sites, which were characterized by capacity (including both dissolved and sorbed solute mass) (see, e.g., Haggerty and Gorelick, 1995). We use here an equivalent MRMT formulation for non-sorbing solutes, so as to facilitate comparison with HPM.

Initial and boundary conditions will be described later for both MRMT and HPM models. MRMT models differ by the distributions of characteristic rates  $\alpha_i$  and immobile porosities  $\phi_i$ . Among the available models (Cvetkovic, 2012; Haggerty et al., 2000), we choose a uniform distribution for characteristic times ( $1/\alpha_i$ ) bounded by the two extreme rates  $\alpha_1 = 1/t_1$  and  $\alpha_N = 1/t_N$  ( $t_1 < t_N$ ) and a power-law distribution for  $\phi_i$ :

$$\phi_i \sim \alpha_i^{m-3}. \quad (3)$$

The power-law distribution is consistent with the generally observed breakthrough curves in Heterogeneous Porous Media with long tailings (Gouze et al., 2008; Haggerty et al., 2004; Li et al., 2011; Silva et al., 2009; Willmann et al., 2008). The power-law exponent  $m$  is generally found to be in the interval [1.5; 2.5].  $m = 1.5$  corresponds to fracture/matrix types of exchanges (Haggerty and Gorelick, 1995).

We simulate MRMT models with a standard time- and space-adaptative method that preserves mass (de Dreuzy et al., 2013) and always complies with the CFL conditions (Daus et al., 1985). The advective and the diffusive processes in the mobile zone as well as the exchange with the immobile zones are treated with a sequential non-iterative coupling method. These methods lead to efficient simulations of large spatial domains and extended times with initial refined resolutions. We have successfully compared them with a more classical fixed-time Galerkin finite element methods integrated with the 4th order Runge–Kutta method (ode45 function of Matlab) and found relative differences less than  $10^{-3}$  %. Simulations have been performed over the time required for transport to reach its asymptotic regime.

## 2.2 Heterogeneous Porous Media (HPM)

For reference purposes, we restrict the analysis to heterogeneity of hydraulic conductivity ( $K$ ) as represented by the classical 2-D Gaussian correlated multi-Gaussian log- $K$  fields. These are characterized by their isotropic correlation function:

$$C(r) = \sigma_Y^2 \exp\left(-\left(\frac{r}{\lambda}\right)^2\right) \quad (4)$$

with  $\lambda$  the correlation length, which is used to scale distances, and  $\sigma_Y^2$  the variance of the logarithm of  $Y = \log-K$ . We use simulation results performed in previous studies (de Dreuzy et al., 2012) obtained on 2-D domains of sizes  $L_L$  and  $L_T$  in the direction parallel and orthogonal, respectively, to the mean flux.  $L_L$  is large enough to avoid any finite-size effects (from  $10^2$  to  $10^3$  correlation lengths  $\lambda$ ). Boundary conditions for flow and transport are periodic in the transverse direction to minimize boundary effects.  $L_T$  is of the order of 100 times  $\lambda$  to ensure initially ergodic transport conditions. Under such uniform extended injection conditions, transport in HPM can be considered ergodic and can be fundamentally compared with a 1-D MRMT model. The immobile zones of MRMT can be viewed as representing the low velocity zones of HPM, so that the mobile zone may represent the high velocity channels.

Flow is solved with a finite volume scheme with permeameter-like boundary conditions under a unit head gradient. Transport is simulated using the ADE, with heterogeneous advection and homogeneous diffusion. Therefore, it is characterized by the Peclet number  $Pe$  equal to the mean velocity times the correlation length divided by the diffusion coefficient. Transport is simulated with a random walk Lagrangian method. Numerical methods are exhaustively described in several previous papers (Beaudoin et al., 2006, 2007, 2011).

## Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## 2.3 Injection and boundary conditions

The same type of injection and boundary conditions are used for both models. Flow has a major flow direction imposed in HPM by a head gradient in the longitudinal direction and periodic boundary direction in the transverse direction. For transport, reflecting and absorbing boundary conditions are used respectively upstream and downstream (Beaudoin and de Dreuzy, 2013). Injection is performed downstream to the inlet boundary to minimize boundary effects.

Extended injection conditions are used for the HPM and MRMT models. Concentrations are homogeneous orthogonally to the main flow direction within a square wave of longitudinal and transverse widths  $\Delta L_0$  and  $\Delta T_0$ , respectively. In the HPM case, concentration is a sole function of the coordinate  $x_L$  along the flow direction:

$$c(x, t = 0) = c_0(x_L) \quad (5)$$

with  $c_0$  given by:

$$c_0(x) = \begin{cases} \frac{m_0}{\phi_T \Delta T_0 \Delta L_0} & \text{if } x_0 < x < x_0 + \Delta L_0 \\ 0 & \text{otherwise} \end{cases} \quad (6)$$

$\phi_T$  is the total porosity. To ensure that the same mass  $m_0$  is injected in the HPM and MRMT cases, we adapt the initial state of the MRMT model to:

$$c(x, t = 0) = c_i(x, t = 0) = c_0(x) \quad \text{for } i = 1 \dots N. \quad (7)$$

Spreading becomes independent of the injection length when the longitudinal plume size becomes significantly larger than  $\Delta L_0$ . Mixing depends more critically than spreading on the injection conditions, as the initial concentration value depends on the injection width  $\Delta L_0$  (Eq. 6).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion







In HPM, spreading comes both from diffusive exchanges with low velocity zones and from spatial fluctuations of the velocity field (de Dreuzy et al., 2007; Salandin and Fiorotto, 1998). The asymptotic dispersivity increases both with the correlation length  $\lambda$  and with the log- $K$  variance  $\sigma_Y^2$ :

$$\alpha_{LA}(\text{HPM}) = \lambda g \left( \sigma_Y^2 \right) h \left( \sigma_Y^2, \text{Pe} \right) \quad (11)$$

where  $g$  is either a linear function for small values of  $\sigma_Y^2$  ( $\sigma_Y^2 < 1$ ) and a quadratic function at larger values (de Dreuzy et al., 2007).  $h(\sigma_Y^2, \text{Pe})$  is a correction factor accounting for diffusion (Beaudoin et al., 2010). Local diffusion reduces the effective dispersivity in the high heterogeneity cases by releasing solutes from the low velocity zone and truncating the trapping times induced by slow advection.

Any concentration plume can be approximated by a Gaussian concentration profile  $c_D(x, t)$ , defined by the two first moments,  $m_L^{(1)}(t)$  as mean and  $\sigma_L^2(t)$  as variance. It is the smoothest equivalent profile. Both MRMT and HPM converge asymptotically to this profile. However, it is far away from the full concentration profile  $c(x, t)$  at any time as shown by the comparison of Fig. 1. At early times (left snapshots on Fig. 1), the concentration profile remains heterogeneous especially in the transverse direction with both higher and lower concentrations. Deviation reaches its maximum around the advection time when the Gaussian concentration profile has become much more diluted than the real concentration field (second from the left snapshot of Fig. 1). Concentration inhomogeneities decrease very slowly and remain over very long times even though the range of concentration values decreases (two right-most snapshots of Fig. 1) (de Anna et al., 2014a; Jimenez-Martinez et al., 2015; Le Borgne et al., 2011).

### 3.2 Mixing

The Gaussian profile only gives a crude approximation of the concentration field with a strong deviation on the distribution of concentration values, especially at early times when diffusion has not homogenized the concentration field in the transverse direction

## Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



(Fig. 1). Actual concentrations remain much higher and closer to the initial concentration value than in the Gaussian profile prediction. The Gaussian profile  $c_D(x, t)$  thus sets a lower bound to the effective concentration variability. Therefore, it is most natural to compare the actual distribution of concentration values to that of the Gaussian profile in order to describe the mixing state. Notice that, contrary to spreading, we are not concerned here with the spatial distribution, but only with the values of concentration and their time evolution, which are most simply characterized by the second moment. We quantified the deviation from the Gaussian mixing regime as the ratio of the actual concentration second moment  $M(t)$  to the second moment  $M_D(t)$  of the Gaussian profile concentration  $c_D(x, t)$  integrated over the whole domain  $\Omega$  minus 1 (de Dreuzy et al., 2012):

$$\gamma(t) = \frac{M(t)}{M_D(t)} - 1. \quad (12)$$

with

$$M(t) = \int_{\Omega} c^2 d^d x \quad (13)$$

and the second moment of the reference Gaussian concentration:

$$M_D(t) = \frac{m_0^2}{2\sqrt{\pi}\Delta T_0\sigma_L}. \quad (14)$$

$M_D$  is directly the square of the injected mass  $m_0^2$  divided by an effective area occupied by the plume  $2\sqrt{\pi}\Delta T_0\sigma_L$ . As  $M(t)$  is always larger than  $M_D(t)$ ,  $\gamma$  is always positive.  $\gamma$  is initially and asymptotically very close to zero. It is significantly positive at intermediary times when the concentration distribution is far from the Gaussian profile.  $M(t)$ , which we have introduced here as a measure of global concentration variability, is

**Effective transport equations**

J.-R. de Dreuzy and J. Carrera

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



a widely used measure, as its time derivative, dissipation rate, determines the physical constrains of chemical reactivity (de Simoni et al., 2005; Le Borgne et al., 2010). The dissipation is also closely related to the dilution index, which is another measure of mixing (Kitanidis, 1994; Rolle et al., 2009). It should be finally noticed that  $\gamma$  and  $M_D$  fully characterize the mixing state given by  $M$ :

$$M = M_D(1 + \gamma) \quad (15)$$

In HPM models, resistance to dispersive mixing  $\gamma$  is enhanced by heterogeneity and reduced by larger diffusion rates (smaller Peclet number) (de Dreuzy et al., 2012).  $\gamma$  sharply increases at initial times to a maximum value  $\gamma_{\max}$ , at a time  $t_{\gamma\max}$  close to the advection time, and slowly decreases back to 0 (Fig. 2). The time range over which  $\gamma$  is significantly non zero can be characterized by  $r_{t_\gamma}$ , which is the ratio of the upper and lower times at which  $\gamma$  is equal to a quarter of its maximal value  $\gamma_{\max}$ . While the amplitude of  $\gamma$  depends on the variability of the velocities and on the rate of advection to diffusion, the shape of the function  $\gamma$  remains unchanged by the  $K$  field heterogeneity ( $\sigma_V^2$ ), the ratio of advection to diffusion (Pe), and the width of the initial conditions ( $\Delta L_0$ ). The time range  $r_{t_\gamma}$  over which  $\gamma$  is non-negligible also remains constant (Fig. 2). Therefore,  $t_{\gamma\max}$  can be used for scaling time, so that  $\gamma$  can be written as:

$$\gamma(t) = \gamma_{\max} f\left(\frac{t}{t_{\gamma\max}}\right) \quad (16)$$

where  $f$  is the characteristic scaling function (Fig. 2, insert). A similar constant shape behavior has been noted for viscous fingering in heterogeneous velocity fields (Jha et al., 2011a, b).

## Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## 4 Conditions for effective formalisms of transport through HPM

We propose four conditions for any effective transport formulation to be considered as a valid representation of transport through heterogeneous media. In essence, an effective transport equation should yield the same mean advection, spreading and mixing as the HPM and be sufficiently flexible to represent real problems. Evaluation of these conditions can be done as follows:

(1) Mean advection simply requires mean water velocity (i.e., mean plume velocity for non-reactive solutes) to equal  $v = q/\phi_T$ . This condition can be met by all published up-scaled transport equations, by imposing some simple constraints on their parameters.

In MRMT, it is sufficient to impose  $\phi_T = \phi + \sum \phi_i$ .

(2) Spreading is characterized by dispersivity, which measures the rate of growth of plume size (Eq. 10). In cases where asymptotic dispersion is reached, this condition implies that dispersivity of the effective equation should tend to the asymptotic dispersivity of the HPM. Otherwise, dispersion (or directly, spread, as measured by  $\sigma_L$ ) can be compared for a spatial scale comparable to the problem dimension (e.g., size of the aquifer, or distance covered by the plume).

In addition, the time required to reach the above dispersion value should also be honored by the effective formulation to ensure that the rate of growth of the plume is reproduced. In our case, where asymptotic dispersion is reached, we propose to define this criterion in terms of  $r_\alpha$ , mean distance covered by the plume at the time  $t_{\alpha_{LA}/2}$  where dispersivity reaches half of its asymptotic value normalized by the asymptotic dispersivity  $\alpha_{LA}$ :

$$r_\alpha = \frac{vt_{\alpha_{LA}/2}}{\alpha_{LA}} \quad (17)$$

where  $t_{\alpha_{LA}/2}$  is implicitly defined by

$$\alpha(t_{\alpha_{LA}/2}) = \frac{\alpha_{LA}}{2}. \quad (18)$$

$r_\alpha$  can also be interpreted as the ratio of advective and dispersive scales like in the definition of the Peclet number.

(3) Mixing is required for properly reproducing fast reactions (slow reactions should be properly reproduced if the resident time distribution is honored, which is assured if mean advection and dispersion are reproduced). As discussed above, mixing is essentially dispersive and well characterized by  $M_D$  (Eq. 14) for late times. Therefore, assuming dispersion to be well reproduced, an effective transport formulation only needs to reproduce the deviation from dispersive mixing, characterized by  $\gamma$  (Eq. 12). In a first stage, the comparison can be restricted to the amplitude of the deviation  $\gamma_{\max}$  and the time range over which it extends  $t_{\gamma}$ . In a more advanced stage, the characteristic shape of the  $\gamma$  function,  $f$ , can be used for comparison.

To compare the timings of spreading and mixing, we define the additional criterion  $r_{\text{MT}}$  as the ratio of the characteristic spreading time  $t_{\alpha_{\text{LA}}/2}$  to the characteristic mixing time  $t_{\gamma_{\max}}$

$$r_{\text{MT}} = \frac{t_{\gamma_{\max}}}{t_{\alpha_{\text{LA}}/2}}. \quad (19)$$

$r_{\text{MT}}$  compares the timing of the development of the resistance to mixing and of spreading and rates the lag between the timing of mixing and spreading.

(4) Flexibility. Most of the work on effective transport is of a theoretical nature, but the ultimate goal should be application to real problems. This implies that a valid transport formulation should be able to accommodate different types of boundary conditions and flow regimes (i.e., transient flow) and dimensions. Most importantly, it should accommodate characterization. Dispersion usually includes the effects of heterogeneity and uncertainty. Whereas the latter is reduced by aquifer characterization, the former is not. Specifically, hydrologists use geology, hydraulics, geophysics, hydrochemistry and isotopes to figure out, among other things, the patterns of spatial variability of hydraulic conductivity. The resulting models display variability not only in the mean log- $K$  but also

## HESSD

12, 12281–12310, 2015

### Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



on their correlation distance and variance. An effective transport formulation should be able to honor this variability.

## 5 Results and discussion

We consider well established that MRMT, and other non-local in time formulations, can reproduce mean advection and spreading, as discussed in the introduction. Mean advection in the MRMT approach is equivalent to that of the HPM provided that flux and total porosity are equivalent. And the distribution of residence times in immobile zones can be adapted so that the asymptotic dispersivity of the MRMT model be equal to that of the HPM model in Eq. (11). It is always possible as dispersivity is an increasing function of the residence times. This imposes a condition on the temporal range of  $t_1, \dots, t_N$  or equivalently on their mean residence time  $\langle \tau_{\text{MRMT}} \rangle$ . As trapping in the immobile zones is the main dispersive mechanism, the mean residence time is logically adapted to calibrate the asymptotic dispersivity. With the total flow imposed to be set by the HPM, the characteristic spatial scale is the typical plume position at  $\langle \tau_{\text{MRMT}} \rangle$ . As the characteristic spatial and temporal scales are interrelated to ensure consistent asymptotic behaviors, comparison of results can be performed on dimensionless terms and should ensure consistent preasymptotic regimes. In fact, MRMT are calibrated on tracer tests and breakthrough information, but this does not ensure a good reproduction of mixing (Luo and Cirpka, 2011). Therefore, we restrict our comparison to mixing criteria and sensitivity to initial conditions.

### 5.1 Comparison of mixing in HPM and MRMT

In Heterogeneous Porous Media (HPM), the temporal extension of the deviation to the dispersive mixing regime  $r_{t_\gamma}$  does not depend significantly on the permeability heterogeneity as also expressed by the constancy of the shape of  $\gamma$  (Fig. 2). We thus compare the shape of  $\gamma$  obtained for the HPM with  $\sigma_\gamma^2 = 9$  ( $f$  function of Eq. 16) to shapes of  $\gamma$  ob-

## Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion















## Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



High-End Computing, edited by: Joubert, G. et al., volume 33 of NIC series, Verlag des Forschungszentrums Jülich, 399–406, 2006.

Beaudoin, A., de Dreuzy, J. R., and Erhel, J.: An efficient parallel tracker for advection–diffusion simulations in heterogeneous porous media, in: *Europar. Lecture Notes in Computer Science*, edited by: Kermarrec, A.-M., Bougé, L., and Priol, T., Springer-Verlag, Berlin, Heidelberg, 705–714, 2007.

Beaudoin, A., de Dreuzy, J.-R., and Erhel, J.: Numerical Monte Carlo analysis of the influence of pore-scale dispersion on macrodispersion in 2-D heterogeneous porous media, *Water Resour. Res.*, 46, W12537, doi:10.1029/2010WR009576, 2010.

Beaudoin, A., Huberson, S., and Rivoalen, E.: A particle method for solving Richards' equation, *CR Mecanique*, 339, 257–261, 2011.

Berkowitz, B., Cortis, A., Dentz, M., and Scher, H.: Modeling non-Fickian transport in geological formations as a continuous time random walk, *Rev. Geophys.*, 44, RG2003, doi:10.1029/2005RG000178, 2006.

Bolster, D., Benson, D. A., Le Borgne, T., and Dentz, M.: Anomalous mixing and reaction induced by superdiffusive nonlocal transport, *Phys. Rev. E*, 82, 021119, doi:10.1103/PhysRevE.82.021119, 2010.

Carrera, J., Sánchez-Vila, X., Benet, I., Medina, A., Galarza, G., and Guinera, J.: On matrix diffusion: formulations, solution methods and qualitative effects, *Hydrogeol. J.*, 6, 178–190, doi:10.1007/s100400050143, 1998.

Cirpka, O. A. and Valocchi, A. J.: Two-dimensional concentration distribution for mixing-controlled bioreactive transport in steady state, *Adv. Water Resour.*, 30, 1668–1679, 2007.

Cirpka, O. A., Rolle, M., Chiogna, G., de Barros, F. P. J., and Nowak, W.: Stochastic evaluation of mixing-controlled steady-state plume lengths in two-dimensional heterogeneous domains, *J. Contam. Hydrol.*, 138, 22–39, 2012.

Clement, T. P.: Generalized solution to multispecies transport equations coupled with a first-order reaction network, *Water Resour. Res.*, 37, 157–163, 2001.

Cvetkovic, V.: A general memory function for modeling mass transfer in groundwater transport, *Water Resour. Res.*, 48, W04528, doi:10.1029/2011WR011657, 2012.

Dagan, G.: Transport in heterogeneous porous formations: spatial moments, ergodicity, and effective dispersion, *Water Resour. Res.*, 26, 1281–1290, 1990.

Daus, A. D., Frind, E. O., and Sudicky, E. A.: Comparative error analysis in finite-element formulations of the advection-dispersion equation, *Adv. Water Resour.*, 8, 86–95, 1985.

**Effective transport equations**J.-R. de Dreuzy and  
J. Carrera[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- de Anna, P., Dentz, M., Tartakovsky, A., and Le Borgne, T.: The filamentary structure of mixing fronts and its control on reaction kinetics in porousmedia flows, *Geophys. Res. Lett.*, 41, 4586–4593, 2014a.
- De Anna, P., Jimenez-Martinez, J., Tabuteau, H., Turuban, R., Le Borgne, T., Derrien, M., and Méheust, Y.: Mixing and reaction kinetics in porous media: an experimental pore scale quantification, *Environ. Sci. Technol.*, 48, 508–16, 2014b.
- de Barros, F. P. J., Dentz, M., Koch, J., and Nowak, W.: Flow topology and scalar mixing in spatially heterogeneous flow fields, *Geophys. Res. Lett.*, 39, L08404, doi:10.1029/2012GL051302, 2012.
- de Dreuzy, J.-R., Beaudoin, A., and Erhel, J.: Asymptotic dispersion in 2D heterogeneous porous media determined by parallel numerical simulations, *Water Resour. Res.*, 43, W10439, doi:10.1029/2006WR005394, 2007.
- de Dreuzy, J.-R., Carrera, J., Dentz, M., and Le Borgne, T.: Time evolution of mixing in heterogeneous porous media, *Water Resour. Res.*, 48, W06511, doi:10.1029/2011WR011360, 2012.
- de Dreuzy, J.-R., Rapaport, A., Babey, T., and Harmand, J.: Influence of porosity structures on mixing-induced reactivity at chemical equilibrium in mobile/immobile Multi-Rate Mass Transfer (MRMT) and Multiple INteracting Continua (MINC) models, *Water Resour. Res.*, 49, 8511–8530, doi:10.1002/2013WR013808, 2013.
- de Simoni, M., Carrera, J., Sanchez-Vila, X., and Guadagnini, A.: A procedure for the solution of multicomponent reactive transport problems, *Water Resour. Res.*, 41, 1–16, doi:10.1029/2005WR004056, 2005.
- Dentz, M., Gouze, P., and Carrera, J.: Effective non-local reaction kinetics for transport in physically and chemically heterogeneous media, *J. Contam. Hydrol.*, 120–121, 222–236, 2011.
- Ding, D., Benson, D. A., Paster, A., and Bolster, D.: Modeling bimolecular reactions and transport in porous media via particle tracking, *Adv. Water Resour.*, 53, 56–65, 2013.
- Donado, L. D., Sanchez-Vila, X., Dentz, M., Carrera, J., and Bolster, D.: Multicomponent reactive transport in multicontinuum media, *Water Resour. Res.*, 45, W11402, doi:10.1029/2008WR006823, 2009.
- Fripiat, C. C. and Holeyman, A. E.: A comparative review of upscaling methods for solute transport in heterogeneous porous media, *J. Hydrol.*, 362, 150–176, 2008.
- Gelhar, L. W.: *Stochastic Subsurface Hydrology*, Prentice Hall, Engelwood Cliffs, New Jersey, 1993.

**Effective transport equations**J.-R. de Dreuzy and  
J. Carrera[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Gjetvaj, F., Russian, A., Gouze, P., and Dentz, M.: Dual control of flow field heterogeneity and immobile porosity on non-Fickian transport in Berea sandstone, *Water Resour. Res.*, doi:10.1002/2015WR017645, 2015.

Gouze, P., Melean, Y., Le Borgne, T., Dentz, M., and Carrera, J.: Non-Fickian dispersion in porous media explained by heterogeneous microscale matrix diffusion, *Water Resour. Res.*, 44, W11416, doi:10.1029/2007WR006690, 2008.

Haggerty, R. and Gorelick, S. M.: Multiple-rate mass transfer for modeling diffusion and surface reactions in media with pore-scale heterogeneity, *Water Resour. Res.*, 31, 2383–2400, 1995.

Haggerty, R., McKenna, S. A., and Meigs, L. C.: On the late-time behavior of tracer test breakthrough curves, *Water Resour. Res.*, 36, 3467–3479, 2000.

Haggerty, R., Harvey, C. F., von Schwerin, C. F., and Meigs, L. C.: What controls the apparent timescale of solute mass transfer in aquifers and soils? A comparison of experimental results, *Water Resour. Res.*, 40, W01510, doi:10.1029/2002WR001716, 2004.

Hayek, M., Kosakowski, G., Jakob, A., and Churakov, S. V.: A class of analytical solutions for multidimensional multispecies diffusive transport coupled with precipitation-dissolution reactions and porosity changes, *Water Resour. Res.*, 48, W03525, doi:10.1029/2011WR011663, 2012.

Hochstetler, D. L., Rolle, M., Chiogna, G., Haberer, C. M., Grathwohl, P., and Kitanidis, P. K.: Effects of compound-specific transverse mixing on steady-state reactive plumes: insights from pore-scale simulations and Darcy-scale experiments, *Adv. Water Resour.*, 54, 1–10, 2013.

Jha, B., Cueto-Felgueroso, L., and Juanes, R.: Fluid mixing from viscous fingering, *Phys. Rev. Lett.*, 106, 194502, doi:10.1103/PhysRevLett.106.194502, 2011a.

Jha, B., Cueto-Felgueroso, L., and Juanes, R.: Quantifying mixing in viscously unstable porous media flows, *Phys. Rev. E*, 84, 066312, doi:10.1103/PhysRevE.84.066312, 2011b.

Jimenez-Martinez, J., Anna, P. D., Tabuteau, H., Turuban, R., Borgne, T. L., and Méheust, Y.: Pore-scale mechanisms for the enhancement of mixing in unsaturated porous media and implications for chemical reactions, *Geophys. Res. Lett.*, 42, 5316–5324, 2015.

Kitanidis, P. K.: The concept of the dilution index, *Water Resour. Res.*, 30, 2011–2026, doi:10.1029/94WR00762, 1994.

Knutson, C., Valocchi, A., and Werth, C.: Comparison of continuum and pore-scale models of nutrient biodegradation under transverse mixing conditions, *Adv. Water Resour.*, 30, 1421–1431, 2007.

**Effective transport equations**J.-R. de Dreuzy and  
J. Carrera[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- Le Borgne, T. and Gouze, P.: Non-Fickian dispersion in porous media: 2. Model validation from measurements at different scales, *Water Resour. Res.*, 44, W06427, doi:10.1029/2007WR006279, 2008.
- Le Borgne, T., Dentz, M., Bolster, D., Carrera, J., De Dreuzy, J. R., and Davy, P.: Non-Fickian mixing: temporal evolution of the scalar dissipation rate in heterogeneous porous media, *Adv. Water Resour.*, 3, 1468–1475, 2010.
- Le Borgne, T., Dentz, M., Davy, P., Bolster, D., Carrera, J., De Dreuzy, J. R., and Bour, O.: Persistence of incomplete mixing: a key to anomalous transport, *Phys. Rev. E*, 84, 015301, doi:10.1103/PhysRevE.84.015301, 2011.
- Le Borgne, T., Dentz, M., and Villermaux, E.: The lamellar description of mixing in porous media, *J. Fluid Mech.*, 770, 458–498, 2015.
- Li, L., Zhou, H., and Jaime Gomez-Hernandez, J.: Transport upscaling using multi-rate mass transfer in three-dimensional highly heterogeneous porous media, *Adv. Water Resour.*, 34, 478–489, 2011.
- Luo, J. and Cirpka, O. A.: How well do mean breakthrough curves predict mixing-controlled reactive transport?, *Water Resour. Res.*, 47, W02520, doi:10.1029/2010WR009461, 2011.
- Luo, J., Dentz, M., Carrera, J., and Kitanidis, P.: Effective reaction parameters for mixing controlled reactions in heterogeneous media, *Water Resour. Res.*, 44, W02416, doi:10.1029/2006WR005658, 2008.
- Neuman, S. P. and Tartakovsky, D. M.: Perspective on theories of non-Fickian transport in heterogeneous media, *Adv. Water Resour.*, 32, 670–680, 2009.
- Orgogozo, L., Golfier, F., Bues, M. A., Quintard, M., and Kone, T.: A dual-porosity theory for solute transport in biofilm-coated porous media, *Adv. Water Resour.*, 62, 266–279, 2013.
- Rolle, M., Eberhardt, C., Chiogna, G., Cirpka, O. A., and Grathwohl, P.: Enhancement of dilution and transverse reactive mixing in porous media: experiments and model-based interpretation, *J. Contam. Hydrol.*, 110, 130–142, 2009.
- Sadhukhan, S., Gouze, P., and Dutta, T.: A simulation study of reactive flow in 2-D involving dissolution and precipitation in sedimentary rocks, *J. Hydrol.*, 519, 2101–2110, 2014.
- Salandin, P. and Fiorotto, V.: Solute transport in highly heterogeneous aquifers, *Water Resour. Res.*, 34, 949–961, 1998.
- Scheibe, T. D., Schuchardt, K., Agarwal, K., Chase, J., Yang, X., Palmer, B. J., and Redden, G.: Hybrid multiscale simulation of a mixing-controlled reaction, *Adv. Water Resour.*, 83, 228–239, 2015.

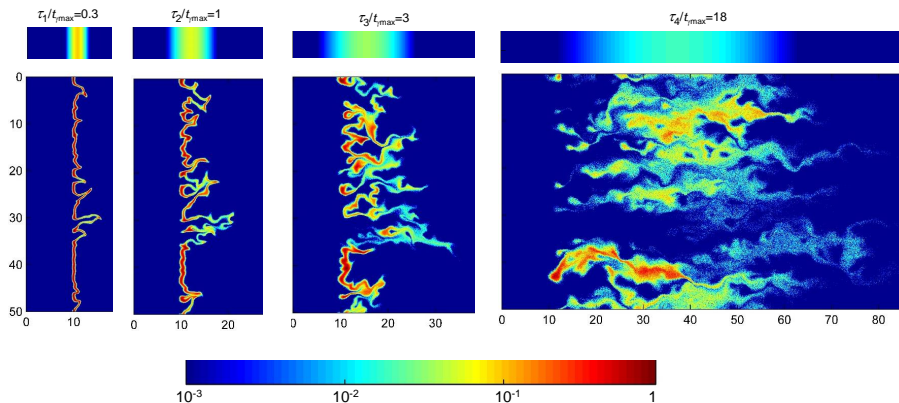


**Effective transport equations**J.-R. de Dreuzy and  
J. Carrera[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- Schneider, B., Paster, A., and Bolster, D.: A numerical investigation of mixing and spreading across an angled discontinuity, *Adv. Water Resour.*, 62, 280–291, 2013.
- Silva, O., Carrera, J., Dentz, M., Kumar, S., Alcolea, A., and Willmann, M.: A general real-time formulation for multi-rate mass transfer problems, *Hydrol. Earth Syst. Sci.*, 13, 1399–1411, doi:10.5194/hess-13-1399-2009, 2009.
- 5 Tartakovsky, A. M., Tartakovsky, G. D., and Scheibe, T. D.: Effects of incomplete mixing on multicomponent reactive transport, *Adv. Water Resour.*, 32, 1674–1679, 2009.
- Willmann, M., Carrera, J., and Sanchez-Vila, X.: Transport upscaling in heterogeneous aquifers: what physical parameters control memory functions?, *Water Resour. Res.*, 44, W12437, doi:10.1029/2007WR006531, 2008.
- 10 Willmann, M., Carrera, J., Sanchez-Vila, X., Silva, O., and Dentz, M.: Coupling of mass transfer and reactive transport for nonlinear reactions in heterogeneous media, *Water Resour. Res.*, 46, W07512, doi:10.1029/2009WR007739, 2010.
- Zhang, Y., Papelis, C., Sun, P., and Yu, Z.: Evaluation and linking of effective parameters in particle-based models and continuum models for mixing-limited bimolecular reactions, *Water Resour. Res.*, 49, 4845–4865, 2013.
- 15



## Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

**Figure 1.** Concentration fields normalized by their maximal value  $c(x,t)/\max(c(x,t))$  and their related Gaussian profile concentrations  $c_D(x,t)/\max(c(x,t))$  in the bar over them at the four evolving times indicated on Fig. 2. In this case, the time at which the non-dispersive mixing reaches its maximum  $t_{y_{\max}}$  is of the same order of the advection time.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

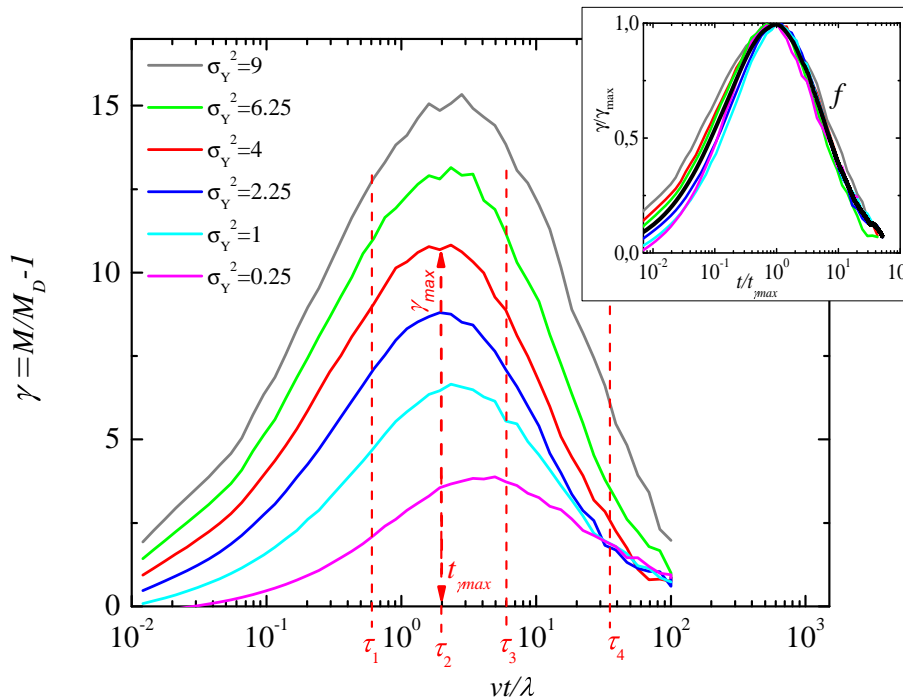
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Full Screen / Esc

Printer-friendly Version

Interactive Discussion



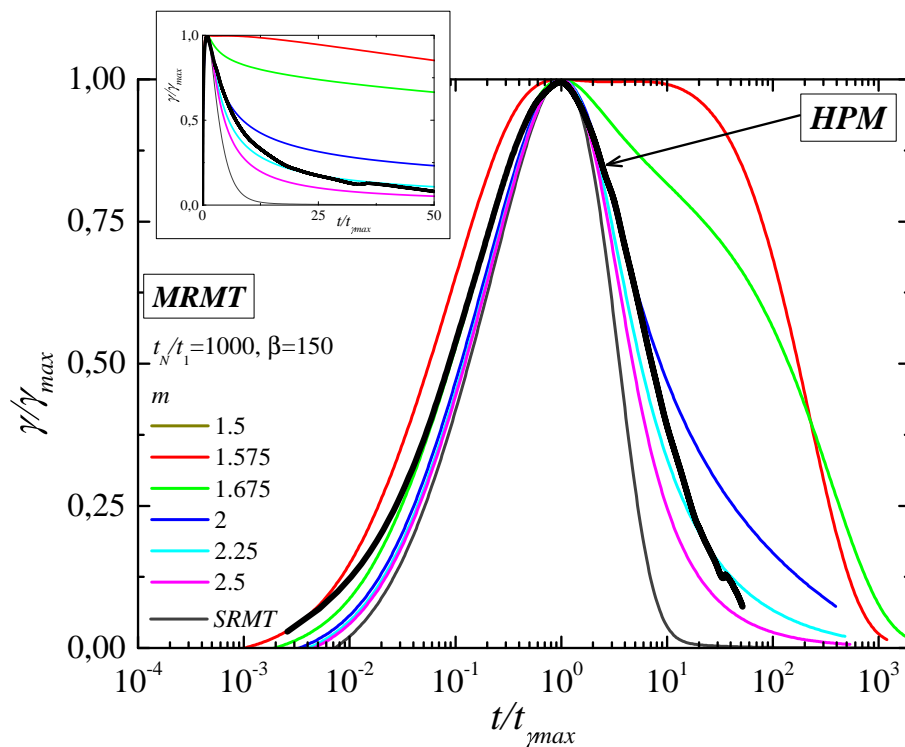


**Figure 2.** Time evolution of the deviation from dispersive mixing  $\gamma(t)$  defined by Eq. (12) in HPM for evolving log- $K$  variances,  $\sigma_Y^2$ , under a small width injection window ( $\Delta L_0/\alpha_{LA} = 0.075$ ), flux weighted injection conditions and  $Pe = 100$  (adapted from de Dreuzey et al., 2012). The similarity of function shapes is highlighted in the insert by the scaling function  $f$  of Eq. (16) where the thick black line is the average of the displayed functions. Note that the time of maximum deviation,  $t_{\gamma_{max}}$ , is hardly affected by  $\sigma_Y^2$  and falls around the characteristic advection time  $\lambda/\nu$ . The four dashed lines indicate the times displayed in Fig. 1.

Title Page	
Abstract	Introduction
Conclusions	References
Tables	Figures
◀	▶
◀	▶
Back	Close
Full Screen / Esc	
Printer-friendly Version	
Interactive Discussion	



## Effective transport equations

J.-R. de Dreuzy and  
J. Carrera

**Figure 3.** Comparison of  $\gamma$  shapes ( $f$  scaling functions defined in Eq. 16) for HPM and MRMT simulations with slopes of the power-law distributions of MRMT transfer rates,  $m$ , between 1.575 and 2.5. HPM is represented by the broad black curve obtained from the insert of Fig. 2. The insert displays the same curves in arithmetic scale.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

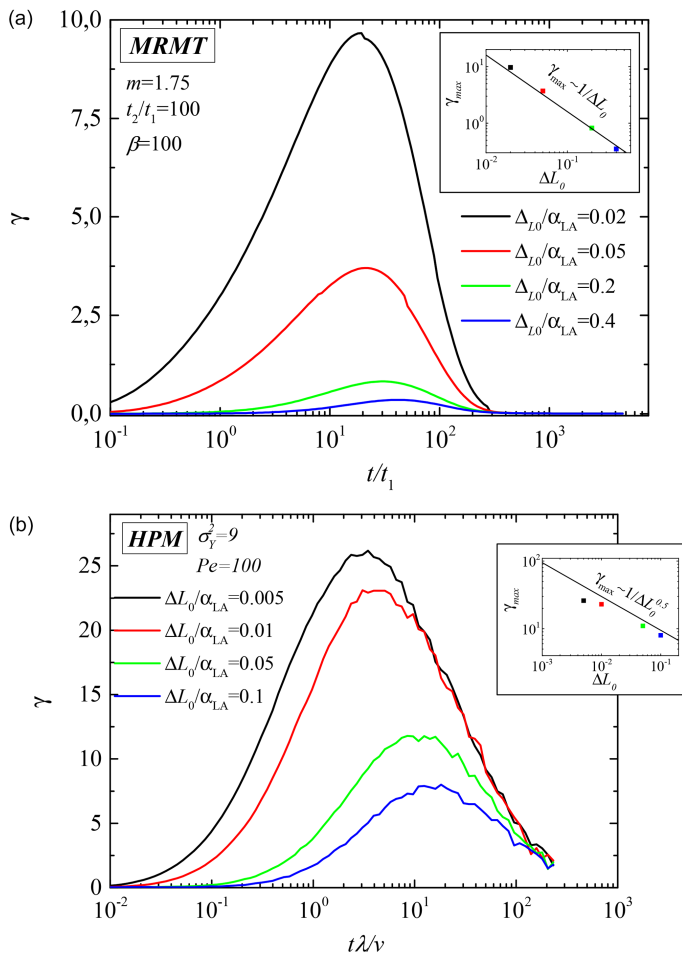
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Printer-friendly Version

Interactive Discussion





**Figure 4.** Dependence of the deviation from dispersive function  $\gamma$  on the injection width  $\Delta L_0$  for **(a)** MRMT and **(b)** HPM models. Insert shows the dependence of  $\gamma_{max}$  on  $\Delta L_0$ .