#### The Response to Comments on Manuscript

# "Evolution of the human-water relationships in Heihe River basin in the past 2,000 years"

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Dear Reviewer.

We would like to sincerely thank and acknowledge your efforts for spending your valuable time reviewing our manuscript. We are pleased to resubmit for publication the revised version of hess-2014-560 "Evolution of the human–water relationships in Heihe River basin in the past 2000 years". We appreciate your constructive comments and criticisms. We have addressed your concerns and provided our response (in red) to your comments (in black) as below. Over the past few weeks and the revision process, we have also improved the paper for clarity. We have also provided two version of the manuscript: one with all changes <u>using the "Track Changes" function in Microsoft Word</u>; and the other one is not highlighted but with has the same content.

Thank you and regards,

Dr. Yongping Wei

Corresponding Author

#### **Responses to major comments of Reviewer #1:**

**1. Comment:** The estimation of E is not clear to me. Is E estimated by equation (2) or (3) for the basin or for cultivated oases and natural oases separately? Is w=3.5 for all the historical periods? Should the value of w be different between cultivated oases and natural oases? Should we even change with time depending on the type and intensity of crops?

E is estimated by equation (2) and (3) for the basin. There are two equations because of different emphasis on  $E_0$  and P, respectively. Throughout the paper w=3.5 for all the historical periods.

We totally agree that the value of w will vary among different land use types and could change with time depending on the type and intensity of crops. However, due to the lack of historical documents or data for natural oases (forest and grassland) in this region, it is impossible for us to characterize w for the natural oases in the historical periods. In addition, in equations (2) and (3), when w is larger than 3, the impact of changes in w on E is likely to be small, especially in this arid region where  $E_0/P$  is large, and the available water for evapotranspiration becomes the determining factor (Zhang et al., 2001; Zhang et al., 2004). Therefore, we consider that using w=3.5 for all the historical periods is reasonable. However, we have discussed this issue as a limitation of this manuscript in the Discussions and Conclusions section; see line 41 page 5 to line 6 page 6 and lines 30-36 page 14.

**2. Comment:** Water supply is computed as the summation of local precipitation and irrigation (or groundwater ET). Is irrigation water pumped from groundwater or surface water withdrawal?

Irrigation water was obtained by surface water withdrawal from upstream reaches in historical periods. It has been both pumped from groundwater and diverted from surface water since the establishment of New China in 1949 as the surface water resource was insufficient for the rapid development of agriculture. The development of pumping and drilling technology during this period also facilitated this change. We have improved the description of irrigation in the revised manuscript to clarify this, see lines 25-29 page 6.

**3.** Comment: Is a portion of local precipitation recharged to the groundwater? If not, the groundwater is fully replenished by the precipitation recharge at the upstream (mountain).

Yes, a very small part of local precipitation recharges the groundwater in extremely wet years in the mid and lower stream reaches of our study area. We agree that most of the groundwater is replenished by the precipitation recharge in the upper catchment (mountains).

4. Comment: Line 15 page 1061: change to ", e.g., water"

We agree. We have made this change in our revised manuscript, see line 43 page 1.

**5. Comment:** Lines 1-3 page 1062: There are some recently published papers which are for explanatory and predictive purpose, e.g., "A prototype framework for models of sociohydrology: identification of key feedback loops and parameterisation approach" by Elshafei et al. (2014 HESS)

Thanks for introducing to us this very useful reference to improve the quality of our manuscript. We have read it and referenced it in our revised manuscript, see lines 11-14 page 2.

6. Comment: Lines 24-25 page 1062: some information is repeated at lines 4-9.

We agree. We have deleted the repeated information in our revised manuscript, see lines 35-36 page 2.

7. Comment: Line 15 page 1064: "Budyko and Miller, 1974;" Double check this.

We agree. We have doubled check it and changed it in our revised manuscript, see line 5 page 4 and line 19 page 15.

**8.** Comment: Line 20 page 1064: change to "respectively;" Similar changes are applicable for other locations.

We agree. We have made this suggested change throughout our revised manuscript, see line 10 page 4.

9. Comment: Line 16 page 1066: correct "Fu (1981) fFor details,"

We agree. We have made this suggested change in our revised manuscript, see line 26 page 5.

10. Comment: Line 24 page 1066: change "PET" to "E0" or define PET.

We agree. We have changed "PET" to " $E_0$ " in our revised manuscript.

**11.** Comment: Lines 24 page 1066 - line 1 page 1067: E0 is assumed to be the same between the historical period and the instrumental period. This assumption needs to be justified or the uncertainty on estimated E due to this assumption needs to be discussed.

Thanks for the point raised. We have added some sentences to discuss the uncertainty of estimated E in our revised manuscript, see lines 30-41 page 5 and lines 30-34 page 14.

**12. Comment:** Line 15 page 1067: since "I" has been used for irrigation in Equation (5), you can use "J" to replace "I" in equation (5).

Thanks for the point raised. We have made the suggested change, see lines 8-11 page 6.

**13. Comment:** Line 2 page 1074: "m3/year"? Check the unit in Table 2 too.

We agree. We have changed "m<sup>3</sup>" to "m<sup>3</sup>/year" in our revised manuscript.

**14. Comment:** Lines 10-11 page 1076: The period from 2000-2010 is short. I am not sure whether it has already reached a new equilibrium stage. Natural oasis may continue to increase from Figure 5.

Thanks for the points raised. We have developed in-depth discussion on the equilibrium stage in our manuscript and changed "a new equilibrium stage" to "a new state", see lines 11- 18 page 8 and lines 22-23 page 13.

**15. Comment:** Line 13 page 1077: Are predictions of its possible future dynamics discussed? How to predict future dynamics?

Thanks for the point raised. We did not mean that the future dynamics were predictable at this stage, rather that our findings can inform attempts towards this. We have changed, see lines 1-9 page 15.

**16. Comment:** Lines 15-18 page 1077: I think the claim is over stated. The manuscript can be shortened, but the "transition theory" needs more description and discussion.

Thanks for the point raised. We agree. We have rewritten this section and added more description and discussion on transition theory in our revised manuscript, see lines 11-18 page 8 and lines 1-9 page 15.

#### Additional references:

Zhang, L., Dawes, W., and Walker, G.: Response of mean annual evapotranspiration to vegetation changes at catchment scale. Water resources research, 37, 701-708, 2001. Zhang, L., Hickel, K., Dawes, W., Chiew, F.H., Western, A., and Briggs, P.: A rational function approach for estimating mean annual evapotranspiration. Water resources research 40, W02502, doi:10.1029/2003WR002710, 2004.

#### **Responses to major comments of Reviewer #2:**

**1. Comment**: How the past 2000 years were divided into several different periods is not entirely clear. First, Table 1 provides vague timelines for the different dynasties; it would be much better if the authors provided start and end years to these periods. It would also help the reader understand whether these were successive contiguous periods. Second, it is mentioned in Section 2.3.1 that the authors used "precipitation in each historical period reconstructed by Ren et al. (2010)". Are Ren et al. (2010)'s historical periods the same as the seven dynastic periods chosen in this study? If not, how different are Ren et al.'s divisions of the historical period?

Thanks for this point. We have listed the start and end years of related dynasties in the past 2000 years in our revised manuscript. The reason why we selected seven periods, not seven whole dynasties, was because the data of reconstructed land use and land cover were only available during these periods (Xie, 2013; Xie et al., 2013). Ren et al. (2010) reconstructed a complete precipitation sequence spanning 2000 years with a resolution of 50 years, so the precipitation data for the seven chosen periods in this study were directly extracted from Ren et al. (2010). See lines 29-36 page 3.

**2. Comment**: In Section 2.3.3, three land use types are considered: cultivated oases, natural oases, and unused land. Equation 4 provides how the P (water supply) in the first two land use types was estimated, to be used in equations 2 and 3. However, for the unused land, was precipitation the only water supply considered? If yes, please state it explicitly; if not, please explain how water supply was calculated for unused land.

Yes. Precipitation is the only water considered for the unused land. We have stated this in our revised manuscript, see lines 30-30 page 6.

**3. Comment**: Sticking with Section 2.3.3, in equation 4, the groundwater irrigation I is kept constant at 500 mm throughout the entire historical period. This assumes that the types of crops cultivated in this basin did not change over 2000 years, and does not take into account the evolution in agricultural technology. Moreover, it directly contradicts the statements made in Section 3.6, such as "In the middle of the Qing Dynasty, the Hexi corridor was politically stable and free from wars and innovative farming and engineering methods were introduced, such as better seeds, new crops, and the steel farm implements".

We fully agree with your comment. We have investigated more historical documents on irrigation development in this region. According to Wang' (2003) research on the development history of water conservancy facilities in Heihe River basin, the main crop varieties, water conservancy facilities, irrigation method and farming conditions almost remained constant from the Han dynasty to the early modern period, so the irrigation was set at 500 mm for the whole historical period. However, it was increased from 500 to 650 mm when the cropping pattern evolved from single wheat to wheat and maize after the 1980s (Wang et al., 2005; Shi et al., 2011), and this is discussed as a limitation of this manuscript in Section Discussions and Conclusions, see lines 12-29 page 6, lines 30-38 page 12 and lines 30-37 page 14.

4. Comment: I think Section 4 of the paper needs to include a paragraph or two on the limitations/assumptions/caveats of the methods used. Historical reconstruction of annual

water fluxes over such a long period will most definitely involve huge uncertainties and assumptions (one example pointed out in my point 3 above). These need to be mentioned and discussed in this section.

Thanks for this point. We agree. Several points on key limitations/ assumptions/ caveats of the methods have been raised above and we have used those, plus a careful consideration of other limitations to develop a more detailed discussion of these in an additional paragraph in our revised manuscript, see lines 30-48 page 14.

**5.** Comment: What is k in Figure 6? I did not find any explanation in the article text.

Thanks for this point. k is the change rate of the factors and it was estimated by dividing the difference between the values at the start and end of the period to the years of the period. We have explained it in lines 25-27 page 8 of the Method Section.

#### Additional references:

Xie, Y.: Dataset of cultivated oasis distribution in the Heihe River Basin during the historical period. Heihe Plan Science Data Center, DOI: 10.3972/heihe.092.2013.db, 2013.

Xie, Y., Wang, X., Wang, G., and Yu, L.: Cultivated land distribution simulation based on grid in middle reaches of Heihe River basin in the historical periods, Advances in Earth Science, 28, 71-78, 2013.

Ren, Z., Lu, Y., and Yang, D.: Drought and flood disasters and rebuilding of precipitation sequence in Heihe River basin in the past 2000 years, Journal of Arid Land Resource and Environment, 24, 91-95, 2010.

Shi, M., Wang, L., and Wang, X.: A study on changes and driving factors of agricultural water supply and demand in Zhangye after Wwater reallocation of the Heihe River, Resources Science, 33, 1489-1497, 2011.

Wang, G., Yang, L., Chen, L., and Jumpei, K.: Impacts of land use changes on groundwater resources in the Heihe River basin, Acta Geogr. Sin., 60, 456–466, 2005.

Wang, Y.: The development history of water conservancy facilities in Heihe River basin, Gansu Nationalities Press, Lanzhou, 2003.

#### 1 Abstract

This paper quantitatively analyzed the evolution of human-water relationships in the Heihe 2 River basin of northern China over the past 2000 years by reconstructing the catchment water 3 balance by partitioning precipitation into evapotranspiration and runoff. The reconstruction 4 results provided the basis for investigating the impacts of human societies on hydrological 5 systems. Based on transition theory and the rates of changes of the population, human water 6 consumption and the area of natural oasis, the evolutionary processes of human-water 7 relationships can be divided into four stages: predevelopment (206 BC - 1368 AD), take-off 8 (1368 - 1949 AD), acceleration (1949 - 2000 AD), and the start of a rebalancing between the 9 human and ecological needs (after post 2000 AD). Our analysis of the The evolutionary 10 process-analysis revealed that there were large differences in the rate and scale of changes 11 and the period over which they occurred, and. The transition of the human-water relationship 12 had no fixed pattern. This understanding of the dynamics of the human-water relationship 13 will assist policy makers to identify management practices that require improvement by 14 understanding how today's problems were created in the past, which may lead to for-more 15 sustainable catchment management in the future. 16

#### 17 **1 Introduction**

The development of land and water resources within catchments over thousands of years has led to spectacular growth in agricultural production along with increased human consumption of water, significant modification of catchment vegetation, and serious degradation of ecosystems, worldwide (Carpenter et al., 2011; Falkenmark and Lannerstad, 2005; R öckstrom et al., 2009; V ör ösmarty et al., 2010). The <u>fF</u>uture of human wellbeing may be seriously compromised if we pass a critical threshold that tips catchment ecosystems into irreversible degradation.

25 Understanding the connections and feedback mechanisms between changes in human activities and hydrological systems in the long term, and uncovering the mechanisms 26 governing the human-water feedback loop, can help us to understand how today's conditions 27 and problems were created in the past, and have important implications for future 28 management (Sivapalan et al., 2012; Liu et al., 2013; Montanari et al., 2013; Savenije et al., 29 2013). However, there is at present, there is limited understanding of the major modes of 30 interactions between the human and hydrologic systems over long time scales, ... although 31 Developing such understanding is the aim of social-hydrology as a new discipline emerged in 32 2012 (Savenije et al., 2013). 33

34 Historical analysis is a key method of socio-hydrology in which hydrological analysis over a long timeframe is a key component. Accurate historical data for hydrology, climate, land use, 35 ecology and geomorphology are often unavailable, but hydrological reconstruction that aims 36 to generate long-term datasets, could provide a basis for the identification, description and 37 parameterization of feedback mechanisms between human activities and water (Thompson et 38 al., 2013). Empirical reconstructions of changes in single hydrological elements at specific 39 locations have been reported, such as including precipitation, streamflow, water salt content 40 and lake levels (Turner et al., 2008; Lowry and Morrill, 2011). Whilst these studies are 41 empirically informative, few of them have been conducted on water balance in basins that are 42 facing significant threats e.g., such as water over-abstraction, sea level rise, or land use change, 43 or in basins and that experience major transitions in different ways (V ör ösmarty et al., 2010). 44

In the social science literature transition is a well-established concept. It is: "a non-linear process of social change in which the structure of a societal system (energy sector, water management and agriculture) transforms" (Rotmans, 2005). Although there is a considerable number of empirical studies focusing on the dynamics of transition, and in particular on the

different stages and processes of transition, they are-have been criticized for empiricism: 1 good at description but weak at explanation (Wimmer, 2006). There have already been 2 several early attempts at exploring the co-evolution of human and water systems. For 3 example, Xiao and Xiao (2004) divided the evolutionary processes of the human-land 4 relationships affected by the water resources in the Ejin region, downstream of the Heihe 5 River basin, into four periods. Geels (2005) studied the trajectories of the co-evolution of 6 7 water technology and society in present-day Netherlands. Kallis (2010) studied the co-evolution of water resource development in ancient Athens. Pataki et al. (2011) provided 8 an outline of the interplay of sociological and ecological processes in urban water 9 management. Unfortunately, most of them adopted "thick descriptive" approaches that have 10 poor explanatory and predictive ability. Recently, Elshafei et al. (2014) developed a prototype 11 framework for models of social-hydrology including identification of some important 12 feedback loops. The framework aims for more explanatory capability although the fully 13 parameterized dynamic coupled model has yet to be applied. 14

The Hexi Corridor, located in western Gansu Province, China, is an important part of the 15 ancient Silk Road established in the Han Dynasty (206 BC-220 AD), and was a trade route 16 between China and western countries that facilitated cultural and economic exchange for 17 approximately 1500 years. It is an arid area supported by oases oasis ecosystems where water 18 dominates the dynamics of human society and natural systems, and therefore the interactions 19 and feedbacks between humans and water are very prominent. The region has a rich written 20 history of over 2000 years. Over-development of land and water resources over thousands of 21 years has significantly modified the catchment vegetation conditions and desertification is a 22 continuing process causing environmental degradation in the region (Xiao and Xiao, 2004, 23 2008). 24

The overarching goal of this paper is to reveal the evolutionary processes of human-water 25 relationships in the Heihe River basin, an important part of the Hexi Corridor over for a 26 period spanning approximately 2000 years, in over which hydrologic, social and 27 environmental systems were connectedinteracted. The specific objectives are to reconstruct 28 the water balance at the basin scale over the past 2,000 years and to determine the 29 30 development stages of the evolutionary processes of the human-water relationships. The analysis is usedIt is expected to gain important understanding of the human-water 31 relationships and provide guidance for the region's sustainable development. 32

### 33 **2 Methods**

### 34 **2.1 Study area**

We selected the Heihe River basin (HRB) in northwest inland China as our case study area. 35 The HRB is located in the central Hexi Corridor; an important part of the ancient Silk Road, 36 and one of the most arid regions in the world, as our case study area. It covers The HRB, 37 covering approximately 130,000 km<sup>2</sup>, and is located at the climatic intersection between the 38 Westerlies and the East Asian summer monsoon (Fig. 1). Many civilizations and cultures 39 have flourishedwere found there, such asincluding the Siba culture, and the Juyan Wooden 40 Slips and the Literature of Heishui city in the HRB (Cheng et al. 2011; Shi, 2007). The rise 41 and fall of civilizations in the HRB is closely linked with water: when there is water there are 42 43 oases and flourishing societies, when there is no water, there is desert and diminished human activities. 44

45 **Figure 1.** Location of the Heihe River basin and locations of data of ice core, tree ring and

1 lake sediment <u>data</u>.

The Qilian Mountains are the principal water source areas of the Heihe River. They-and have 2 an elevation varying between 2000 and 5500 m and a mean annual precipitation varying from 3 250 to 500 mm. The midstream oases area is are a part of the Hexi Corridor with elevations 4 between 1000 and 2000 m and mean annual precipitation ranginges from 100 to 250 mm. The 5 lower reaches are located on the arid Alaxa Plateau where the mean annual precipitation is 6 7 less than 50 mm (Qin et al, 2010). The Heihe River is the second largest-longest inland river in China with a length of 821 km. Starting from the upstream Yingluoxia Hydrological 8 Station (YLX), the Heihe River flows northward into its midstream area., and finally fFlows 9 out of the midstream area, that is are measured at the Zhengyixia Hydrological Station (ZYX), 10 and if finally flows into the terminal lakes in the downstream areas. Its upstream flowFlow 11 from the upper catchment provides water supplies for agricultural production and ecosystem 12 stabilization in the middle and lower reaches of the HRB. 13

The HRB is an important area for grain production in China and is a highly developed 14 irrigation district with an unremittingly agricultural history dating back nearly 2000 years. 15 The iIntensive and non-unsustainable utilization of water resources in the middle reaches of 16 the basin has led to a sharp decrease of in water supply in to the lower reaches during the last 17 50 years (Zhou and Yang, 2006). As a consequence the ecosystems in the lower reaches has 18 have been degraded with by land desertification, more frequent sandstorms, and the drying 19 out of terminal lakes. Therefore the HRB is a compelling case study area for an analysis on of 20 the co-evolution of the human-water systems at the basin scale. 21

#### 22 **2.2 Study period**

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We selected the past 2000 years as our study period. This time scale represents a period in which dramatic changes in climate, land uses, runoff, management policy, population, societal development and catchment ecological conditions have occurred. These are major variables affecting the river basin water cycles of river basins. It is also a time of significant civilisation development in China. T, for which there is a wealth of documentary evidence available (Holmes et al., 2009; Zheng and Wang, 2005).

We reconstructed the co-evolutionary processes of societal development and hydrological
systems based on seven periods
were selected in the past 2000 years according tousing the results of Shi (2010), Wang et al.
(2013) and Xie et al. (2013) (Table 1). This reconstruction was dependent on adequate land
use information, which resulted in some limitations to developing a contiguous hydrological
data set over this long timeframe. We reconstructed the co-evolutionary processes of societal

35 development and hydrological system based on seven Dynastic periods.

<b>Table 1.</b> Seven periods selected in the past 2000 years.				
Dynasty	<u>Start</u>	<u>End</u>	Main production	Selected periods
Han Dynasty	<u>206 BC</u>	<u>220 AD</u>	<u>Agriculture</u>	The beginning of the 1st century AD
Wei-Jin Era	<u>220 AD</u>	<u>420 AD</u>	<u>Animal husbandry</u>	The end of the third century AD
Tang Dynasty	<u>618 AD</u>	<u>907 AD</u>	Agriculture	The mid- 8th century AD
Yuan Dynasty	<u>1271 AD</u>	<u>1368 AD</u>	Animal husbandry	The end of the 13th
Ming Dynasty	<u>1368 AD</u>	<u>1644 AD</u>	Agriculture	century AD The mid- 16th century AD

 Table 1. Seven periods selected in the past 2000 years.

Qing Dynasty	<u>1644 AD</u>	<u>1912 AD</u>	Agriculture	The mid- 18th century
The Republic of China Era	<u>1912 AD</u>	<u>1949 AD</u>	<u>Agriculture</u>	AD The 1940s

#### 2.3 Reconstructing the evolutionary processes of catchment water balance

We used annual water balance partitioning to provide insights into the evolutionary processes of human-water relationships at basin scale., which Such partitioning is widely used as a signature of hydrologic regimes when catchments experience changes in precipitation regimes, temperature and land use change (Budyko and Miller, 1974; Sivapalan et al., 2003), to provide insights into the evolutionary processes of human-water relationships at basin scale. For this study, the water balance equation can be written as:

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 $P + R_{in} = E + R_{out}$ (1)

9 Where where P and E are precipitation and evapotranspiration in the mid- and down- stream areas of HRB, respectively;  $R_{in}$  is the streamflow in from the upstream part of HRB flowing 10 into the midstream area, and  $R_{out}$  is the amount volume of water flowing into the terminal 11 lakes of the downstream areas. In arid regions soil water content is very small and the 12 groundwater levels were stable over historical periods, so changes of soil water content and 13 groundwater are negligible and not included in Eq. (1). 14

Due to <u>a</u> lack of measured data in historical periods the reconstruction of P,  $R_{in}$  and E, and 15 validation of the derived  $R_{out}$  from Eq. (1) are important steps for developing catchment water 16 balance over the long-term timeframe necessary for understanding the co-evolutionary 17 process of human-water relationships. 18

#### 2.3.1 Reconstruction of Reconstructing precipitation (P) in the mid- and down-19

#### stream areas

We estimated precipitation (P) in historical periods based on instrumental data in the most 21 recent period and changes in paleoclimatic conditions. Ren et al. (2010) reconstructed the 22 mean precipitation sequence of the whole HRB in-for the past 2000 years using historical 23 drought and flood sequences, based on the good correlation between drought and flood 24 disasters and precipitation in the 40 years from 1956 to 1995 ( $R^2 = -0.892$ ). We reconstructed 25 26 the distributed precipitation (P) in the mid- and down- stream areas in historical periods as follows. by multiplying tThe mean precipitation from the instrumental record fordata from 27 1956 to 1995, when there were continuous records at ten meteorological stations, was 28 multiplied by the proportion ratio of the precipitation in each historical period reconstructed 29 by Ren et al. (2010) to the precipitation that in the measured period. The instrumental 30 precipitation data in the recent period were obtained from the China Meteorological 31 Administration of Meteorology. 32

#### 2.3.2 <u>Reconstruction of Reconstructing streamflow flowing into the midstream</u> 33

#### area: R<sub>in</sub>

Dendrochronologically-based hydrological reconstructions have been widely used to extend 35 existing instrumental streamflow records, as streamflow variations correlate well with tree 36 ring-width series (Woodhouse et al., 2006; Saito et al., 2008). There are many studies on the

streamflow reconstruction <u>studies in Qilian Mountains</u> based on tree ring analyses <u>for the</u>
 <u>Qilian Mountains</u>. The longest streamflow record in this region is about 1400 years, <u>obtained</u>
 <u>developed</u> by Yang et al. (2012), then 1300 years by Kang et al. (2002) and 1000 years by
 Qin et al. (2010). None of these streamflow reconstructions <u>completely</u> spanned <u>the</u> 2000
 years <u>or moreof interest</u>.

In order to reconstruct the historical streamflow in the upstream area of HRB  $(R_{in})$  in over the 6 past 2000 years, we firstly analyzed the consistency of the historical streamflow 7 reconstructions by of Yang et al. (2012), Kang et al. (2002) and Qin et al. (2010), and 8 9 selected the two more-most reasonable two-reconstructions based on the humidity changes of climate in this region as reflected by other proxy indices, e.g. lake sediments and ice cores; 10 then among the two selected streamflow reconstructions, the The shorter one was used to 11 12 extend the historical series and the longer one was used to validate the extension where these two reconstructions did not overlapin the gap period between the shorter and the longer. We 13 then extended the selected reconstructed streamflow up to 2000 years by using the 14 reconstructed precipitation based on theand a established relationship established between the 15 selected streamflow reconstruction and the existing precipitation reconstructions in the 16 upstream area. As a All the streamflow reconstructions focused on the mountainous region of 17 the mainstream of the Heihe River (Figure  $1_{\frac{1}{2}}$ ). Therefore, in order to obtain the streamflow 18 flowing into the midstream area  $(R_{in})$ , we multiplied the streamflow at YLX by a 19 proportion the ratio of the total streamflow of from the upstream area of the HRB to the one 20 streamflow at YLX, based on the instrumental data in the recent 50 year period. In addition to 21 the meteorological data mentioned in 2.3.1, tThe instrumental streamflow in the recent period 22 was obtained from the Hydrographic Service of Gansu provinceProvince. 23

#### 2.3.3 Estimating *E* based on the reconstructed ion of land use

*E* in Eq. (1) was calculated by <u>using</u> the top-down method of the Budyko hypothesis. We
used the equations developed by Fu (1981) (For details, see Fu, 1981 and Zhang et al., 2004)
which are expressed as:

28 
$$\frac{E}{P} = 1 + \frac{E_0}{P} - [1 + (\frac{E_0}{P})^w]^{1/w}$$
 (2)

24

29 
$$\frac{E}{E_0} = 1 + \frac{P}{E_0} - [1 + (\frac{P}{E_0})^w]^{1/w}$$
 (3)

Where where  $E_0$  is potential evapotranspiration.  $E_0$  on a daily timescale was estimated on a 30 31 Penman-Monteith Equation which was has been acknowledged as the best method for this 32 region (Zhao and Ji, 2010). It is known that many factors influence  $E_0$  and it is difficult to 33 clearly determine the changes of  $E_0$  in historical periods without instrumented data. However, 34 35 it is recognized that air temperature is one of most importantkey factors that influences influencing  $E_0$  the *PET*. As the oscillating range of the temperature change over the study 36 period was not more than 2 °C in the past 2000 yearsthis region (Zheng et al., 2010), and as 37 Zhang et al. (2014) found that the  $E_0$  increased by only 1.16 mm per month for a temperature 38 increase of 2 °C in the Gulang River Basin (next to the Heihe River basin), it was assumed 39 that the  $E_0$  in the HRB was constant over the study period and the same as the historical 40 period was considered the same as in modern instrumented times. The term w is a catchment 41 scale model parameter determining the evaporation E ratio (E/P) for a given  $E_0/P$ . 42 Theoretically, w should vary between land use types and could change with time depending 43 on the type and intensity of crops. Unfortunately, due to the lack of historical documents or 44 data for the natural oases (forest and grassland) in this region, it is impossible to characterize 45

8 For cultivated oases: 
$$P_{crop,i} = P_i + J$$

9 For natural oases: 
$$P_{veg,i} = P_i + G_{veg}$$

Where where  $P_{crop,i}$  and  $P_{veg,i}$  are the precipitation equivalent in the period i for crop and 10 natural oases respectively;  $P_i$  is <u>the</u> actual precipitation in period i; <u>J</u> is irrigation; <u>and</u>  $G_{veg}$  is 11 the water consumed from consumed groundwater by the natural oases. In this arid region, 12 there is no agriculture without irrigation. According to Xiao and Xiao (2008) flood irrigation 13 has beenwas the main irrigation method in northern China from since the Han dynasty-to the 14 early modern period,. The main crop varieties, water conservancy facilities, irrigation 15 methods and farming conditions have remained almost unchanged from the Han dynasty to 16 the early modern period according to Wang's (2003) research on the development history of 17 farm irrigation in the Heihe River basin which, to our understanding, is the most 18 comprehensive study on historical agricultural irrigation in this region. The wheat was the 19 major crop and single wheat was the main cropping pattern in this region, therefore the 20 annual irrigation volumeand the value of *I* in historical period was set at 500 mm for the 21 historical periods. Since the 1980s, annual irrigation applications in this region have 22 increased from 500 to 650 mm as the cropping pattern has evolved from single wheat to 23 double cropping with wheat and maize (Wang et al., 2005; Shi et al., 2011). 24

Irrigation water was obtained by surface water withdrawal from upstream reaches in
historical periods; however, it has been both pumped from groundwater and diverted from
surface water since 1949 as the surface water resource was insufficient for rapid development
of agriculture. The development of pumping and drilling technology during this period also
facilitated this change. *G<sub>veg</sub>* was set at 225 mm per year for natural oases according tobased
on Wang et al. (2005). For unused land, where the groundwater level was deep, the only
water supply was precipitation.

32 So t The total *ET* of the basin is given by:

$$E_{total} = \sum_{l=1}^{3} E_l \times S_l \quad (4)$$

34 Where where  $E_{total}$  is the total evapotranspiration of the basin; *l* is the number of the land use 35 types: cultivated oases, natural oases, and unused land.  $E_l$  is the evapotranspiration from land 36 use type *l*, and  $S_l$  is the area of land use type *l*.

The maps of cultivated oases in historical periods were downloaded from the Heihe Plan 37 Science Data Center: www.heihedata.org/heihe (Xie, 2013). As a-the historical 38 reconstructions of the natural oases in this region was were not available found in the 39 literature, we used the land use scenario in 1975 as the final land use pattern in order to 40 reconstruct the distribution of natural oases because it is known that the expansion of the 41 farmland was at the expense of the desert after 1975. We made reconstructed them based on 42 the following two assumptions about the reclamation of cultivated oases based on the 43 previous results (Li, 1998; Wu, 2000; Xie et al. 2009): (1) people selected the regions with 44 45 natural oases (grassland and forest) rather than desert for reclamation in the historical periods because the former has better water and soil conditions in these arid regions, and (2) once the 46 47 reclaimed farmlands were abandoned and without the vegetation cover, they were

<sup>1</sup> w for the natural oases in historical periods. In addition, in equations (2) and (3), when w is 2 larger than 3, the impact of changes in w on E is small, especially in arid regions where  $E_0/P$ 3 is large. In such situations, the available water for evapotranspiration becomes the 4 determining factor (Zhang et al., 2001; Zhang et al., 2004). Therefore, It is a catchment 5 parameter and the value of w for HRB was set at to 3.5 according tofollowing Yang et al 6 (2007). The sources of water supply for E include precipitation, groundwater, and irrigation 7 water in this region, so P in equations (2) and (3) can be replaced as follows:

subsequently desertified because of wind-driven sand and burial by dunes<del>dune coverage</del>. The 1 hundreds of ruins of towns along the Silk Road in the vast deserts of northwest China are clear evidence of this change of oases systems. Therefore Then, it is known that the expansion of the farmland from the desert in this region started after 1975, we considered the total area of oases from the first period to the period of 1975 as the largest area of the oases in historical periods, which included cultivated oases and natural oases. In each period the area of cultivated oases had reconstructed data (Xie, 2013), then the area of natural oases was obtained by deducting the area of cultivated oases in the particular-this period and the area of cultivated oases abandoned in the past periods, from the total oases area, and the remainder was considered unused land.

Based on the land use reconstructions, precipitation reconstructions and estimated  $\underline{E}_0 PET$ , the 11 E in Eq. (1) was obtained according to using Eqs. (2) and (3). The data sources used for 12 calculations of E and reconstruction of land use included: land use data obtained for three 13 periods by remote sensing (1975 Landsat MSS, 2000 and 2010 satellite TM and ETM+ data), 14 the historical atlas of China (Tan, 1996), and meteorological data from the China 15 Meteorological Administration of Meteorology, including daily mean, maximum and 16 minimum air temperatures, wind speed, and relative humidity. 17

#### 18 2.3.4 Validating the derived $R_{out}$ with the reconstructed evolution of the terminal

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The input volumes of water to terminal lakes  $R_{out}$  were derived from the reconstructed 20 precipitation, E based on the reconstructed land use, and reconstructed streamflow  $R_{in}$  using 21 Eq. (1). We validated the derived  $R_{out}$  with the lake evolution reconstructed by previous 22 research on the lithology, geochemistry and mineralogy of lacustrine sediment depth profile 23 sequences. 24

As sediment profiles of lakes in arid zones sensitively reflect changes in climate changes and 25 human activities, they are regarded as excellent resources for palaeoclimate research (Jin et 26 al., 2004). Lacustrine sediment sequences have been widely used for deducing the mass 27 balance between the inflow water volume and evaporation from terminal lakes, climate 28 change and human activity (Jin et al., 2004; Jin et al., 2005). For example, grain size 29 30 distributions of lacustrine sediments directly reflect water dynamics, and soluble salt content reflects the chemical characteristics of lake water, which is affected by climate and inflow 31 water (Jin et al., 2004). 32

Due to the unavailability of systematic and consistent studies on lake evolution in the HRB, 33 we validated the derived  $R_{out}$  values based on the changes of input volumes of water to the 34 35 terminal lakes in downstream areas as they reflect changes of the hydrologic cycle involving precipitation, land use, evaporation and runoff in the upper and middle reaches. Rout directly 36 influences the processes of expansion and shrinkage of surface area, sediment deposition and 37 salinization of the terminal lakes. When the input volume of water to the terminal lakes is 38 relatively abundant, lake area extends, lake water level rise, lake water has smaller salt 39 concentrations, and the sediment deposition environment is relatively stable, and vice versa. 40

The data and information sources used for reconstruction of the evolution of the terminal 41 lakes include all collected research achievements studies on the palaeoenviromental evolution 42 in the downstream area of the HRB from Lakes Sogo Nur, Gaxun Nur and Juyanze. The 43 evolution of the terminal lakes in the Heihe River experienced three periods: Juyanze from 44 Warring States Period to Yuan dynasty, Juyanze-Gaxun Nur from Yuan dynasty to Ming 45 dynasty and Gaxun Nur-Sogo Nur from Ming dynasty to 1961 AD (Chen, 1996). The data 46

include granularity, soluble salt, sedimentary pigment, organic carbon content, and
 groundwater level (Jin et al., 2004; Jin et al., 2005; Qu et al., 2000; Zhang et al., 1998).

## 3 2.4 Determining the development stages of evolutionary processes of human-water

### 4 relationships

River basins are co-evolvinged social-ecological systems in which water management 5 decisions affect environmental outcomes that are subject to societalsociological conditions. 6 We interpreted and determined the key states of the evolutionary processes of the 7 human-water relationships in the HRB based on the transition theory of social science. 8 Transition theory is one of the most relevant approaches to understand the evolution of 9 societal systemsevolution and support the management of sustainable developmentsocietal 10 adaptation to sustainability (T abara and Ilhan, 2008). In general terms a transition can be 11 understood as the process of change of a system from one stage of a dynamic state 12 equilibrium to another. According to Rotmans (2005) a set of typological phases can be 13 identified in a transition: (1) predevelopment, (2) take-off, (3) acceleration, and (4) 14 15 stabilisation. In the predevelopment stage, a change occurs marginally or imperceptibly, while after take-off a rapid process of societal change occurs until another state is reached, in which 16 the speed of change decreases again (T abara and Ilhan, 2008). Transitions can fail at any 17 stage. 18

A transition can be measured and assessed by indicators which that could be variables with 19 actual physical meanings, or their surrogates. In this study we used human water consumption 20 and natural oases oasis area as the indicators to understand the evolutionary processes of the 21 human-water relationships in the HRB over the past 2000 years. Human water consumption, 22 the difference between evapotranspiration and precipitation in cultivated land areas, reflects 23 the consequence of human societal development on water cycles. The area of natural oases 24 area reflects water supporting the environment. We used direction and rate of change (k)the 25 change trend and rate of these two indicators over time to divide the human-water relationship 26 into different development stages. Both the natural oases area and human water consumption 27 were obtained using the methods above. 28

29 **3 Results** 

## 30 **3.1 Reconstructed precipitation** (*P*) in mid- and down- stream reaches

The proportions ratios of the precipitation for seven selected historical periods to the current 31 period forof the whole HRB, derived using drought and flood sequence information and data, 32 to that in the most recent period in the seven selected historical dynastic periods over the past 33 2000 years, were 0.7, 0.95, 1, 0.9, 1, 0.98 and 0.96, respectively. The precipitation in mid-34 and down- stream areas in the for each historical periods was then obtained by multiplying the 35 mean instrument-measured precipitation from data 1966 to 1995 at ten meteorological 36 stations by these proportions-is shown in (Fig. 2). The precipitation in historical periods 37 decreased from the midstream to downstream reaches, and it was least in the Han Dynasty, 38 and was similar to the present level in the Tang and Ming Dynasties. 39

40 Figure 2. The reconstructed precipitation in historical periods in mid- and down- stream41 areas of the HRB.

#### 1 3.2 Reconstructed streamflow flowing into midstream reaches $(R_{in})$

The streamflow reconstruction of Qin et al. (2010) was used to extend the measured 2 streamflow reconstruction, and the streamflow reconstruction of Yang et al. (2012) was used 3 to validate it. It was found that the streamflow record reconstructions obtained by Yang et al. 4 (2012), Qin et al. (2010), and Kang et al. (2002) for the period AD 1000-2000 are generally 5 consistent,; however, discrepancies occurred around the years of 1290, 1530, 1690, 1840 and 6 1910. The streamflow reconstructions by Yang et al. (2012) and Qin et al. (2010) were more 7 consistent with the changes in regional humidity suggested by paleoclimate results for this 8 region. The paleoclimatic series were derived from Based on the results from the 9 paleoclimate established in this region with tree rings from living trees or archaeological 10 woods in the Qilian Mountains and the Tibetan Plateau (Yang et al., 2014; Sheppard et al., 11 2004; Shao et al., 2010), lake-sediments in Qinghai Lake (Shen et al., 2001) and ice cores in 12 Dunde (Liu et al., 1998), the. Therefore these two reconstructions by Yang et al. (2012) and 13 Qin et al. (2010) were more consistent with the changes of regional humidity and were 14 considered to be the more reasonable. 15

It is known that the annual streamflow at YLX and mean precipitation in the Qilian 16 Mountains region changed consistently over in the past 50 years (Xiao and Xiao, 2008). It 17 was found that precipitation reconstructions of Yang et al. (2014) (Fig. 3a) and streamflow 18 19 reconstructions of Qin et al. (2010) (Fig. 3b) changed consistently in the last 1000 years. We derived the <u>a</u> linear relationship between them as follows:  $R_{\text{Qin et al. (2010)}} = 0.2771 * P_{\text{Yang et al.}}$ 20 (2014)+80.632. We used this relationship to extend the streamflow reconstruction from back to 21 the period 0 to 1000 AD at YLX (Fig. 3c). The extended streamflow reconstruction is 22 consistent with the streamflow reconstruction of Yang et al. (2012) for the period from 575 23 AD to 1000 AD. Over the whole study period, From historical periods to now the 24 25 reconstructed streamflows into midstream areas  $(R_{in})$  were-varied between about 2.6 and 4.0 billion m<sup>3</sup> per year. It-Streamflow peaked in recent years due to abundant precipitation 26 together with glacier and snow melt in the upstream areas due to rises in temperature. 27

Figure 3. (a) Yang et al.'s (2014) annual precipitation reconstruction for the Qilian Mountains over the last 2000 years, with 50-y smoothing in Qilian Mountains region over the last 2000 years. (b) Qin et al.'s (2010) annual streamflow reconstruction spanning the last millennium with 50-y smoothing at YLX. (c) The Our extension of the streamflow from 0 to 1000 AD by comparing and analyzingbased on the Yang et al. (2014) precipitation reconstruction.

#### 34 **3.3 Reconstructed historical land use and land cover**

The reconstructions of land use in for the seven historical periods and three land use maps for 35 1975, 2000 and 2010 in modern New China (since 1949) obtained by image interpretation are 36 37 shown in Fig. 4. The <u>areas of cultivated oases areas</u> changed significantly in over historical periods. It had awas large size in the Han Dynasty, and then gradually decreased in area until 38 the Yuan Dynasty. From the Ming Dynasty it increased gradually, and finally reached a peak 39 in the period of New-modern China. The cultivated areas were mainly distributed in the 40 downstream area of the basin in the first period, and then moved toward the upstream area, 41 and finally ending up focused on the middle reaches. This might suggests could reflect that 42 43 land reclamation was directly affected by the available water resources.

**Figure 4.** Land use reconstructions in historical periods and land use through image interpretation in recent periods in mid- and down- stream areas of HRB. (It should be noted

that the grassland, forest and water or wet land were combined <u>into-with</u> the natural oases,
and the farmland and built-up land were combined <u>into-with</u> cultivated oases in the land use
in 1975, 2000 and 2010.)

#### 4 **3.4** Validation of derived $R_{out}$ with the reconstructed evolution of the terminal lakes

5 The <u>average annual</u> volume of water that entered the terminal lakes  $(R_{out})$  in the historical 6 periods is shown in Table 2. The <u>estimatesdata</u> were obtained <u>usingfrom</u> Eq. (1) <u>together</u> 7 <u>withbased on</u> the reconstructed precipitation, streamflow and *E* related to land use. The 8 <u>reconstruction of the evolution of the lake</u>, <u>evolution reconstruction</u> based on lithological, 9 geochemical and mineralogical data from the lacustrine sediment profile sequences in 10 terminal lakes, together with some interpretation, is also described in Table 2.

There are relatively good relationships between the input volumes of water to the terminal 11 lakes  $(R_{out})$  and <u>the</u> evolution of the terminal lakes in historical periods. The input of water to 12 terminal lakes was not only determined by the streamflow from upstream, but also affected 13 14 by land use in the mid- and down- stream areas of the basin. When the streamflow from the upstream area was high and the cultivation activity in the middle stream was not intense, the 15 input of water to terminal lakes was high, such as in-during the Tang and Ming Dynasties, 16 and vice versa. This was reflected by the pigmentation and organic carbon content of the 17 sediments of the terminal lakes (Qu et al., 2000; Zhang et al., 1998). After the turn of this the 18 <u>21st</u> century  $R_{out}$  became negative which meant that there was a deficit in groundwater 19 recharge because of over-extraction of water for irrigation to meet the need of food (Wei, 20 21 2013).

Periods	$R_{out}$ /10 <sup>8</sup> m <sup>3</sup> /year	Evolution of terminal lakes	
Han Dynasty	7.5	The lake was shrinking (Qu et al., 2000), and fine magnetic minerals peaked in the sediment profile (Qu et al., 2000; Zhang et al., 1998). This might be affected by low $R_{out}$ and intense reclamation in the downstream areas around the terminal lake.	
Wei-Jin Era	9.2	The lake was still shrinking (Qu et al., 2000), and the primary productivity of the lake was low, such as Osc <u>illatoria flavin</u> , Myx and CD-chlorophyll derivative (Qu et al., 2000; Zhang et al., 1998). This may be because of low $R_{out}$ and weakening reclamation due to war and other factors.	
Tang Dynasty	18.1	There were stable water dynamics, a large lake area and deep water reflected by the sediments with higher contents of silt and clay, and relatively low contents of coarse grains content (Jin et al., 2005; Jin et al., 2004). This was consistent with indicated a large $R_{out}$ during this period.	
Yuan Dynasty	14.9	Same as the Tang dynasty.	
Ming	18.9	The salinity of lake water decreased and the lake extended expanded further (Zhang et al., 1998). This was consistent	

Table 2. The input volumes of water to the terminal lakes  $(R_{out})$  and evolution of the terminal lakes in historical periods.

Dynasty		<u>with indicated by a large <math>R_{out}</math>.</u>
Qing Dynasty	11.8	Same as the Ming dynasty.
New-Modern China in 1949	15.4	The lake <u>s kept-maintained</u> a relatively large area (Zhang et al., 1998; Xiao et al., 2004). This was <u>consistent with</u> indicated by a large $R_{out}$ .
1975	2.0	Terminal lake Gaxun nur dried up, <u>and</u> Sogo nur came and wentbecame ephemeral (Xiao et al., 2004). This is was because of intense <u>exploitation of water for</u> <u>agriculturereclamation</u> in the midstream area, <u>which led to and</u> the streamflow decreased and was unstable.
2000	-2.8	The lakes dried out, and the groundwater <u>depth_levels</u> decreased (Xiao et al., 2004). This <u>is-was</u> because of intense <u>usage of water for agriculturereclamation</u> in the midstream area and overexploitation of the groundwater in the basin.
2010	-0.5	Lake restoration <u>started</u> .

#### 1 **3.5 Reconstructed catchment water balance in the past 2000 years**

2 We reconstructed the catchment water cycles at in the HRB in-for the past 2000 years from the precipitation reconstruction (P) in mid- and down- stream areas, the streamflow 3 reconstruction  $(R_{in})$ , land use reconstruction, evapotranspiration reconstruction (E) and the 4 derived streamflow reconstruction into terminal lakes ( $R_{out}$ ). Comparison of Through 5 validation with the reconstructed lake condition using the sediment record and Rout (Table 6 2)evolution reconstruction, shows that the reconstructed water cycles reasonably reflects 7 variations inreflected the reality of water balance partitioning at-in the HRB in-over the past 8 2000 years. 9

Fig. 5 shows the evolution of the catchment water balance elements in the HRB in the past 10 2000 years. Human water consumption changed clearly, especially after the founding of 11 modern China in 1949, when streamflows from upstream areas were approximately 12 unchanged. The main cause of the water balance changesfactor for this was rapid expansion 13 of the cultivated areas around oases, reflecting the increasing population, which was a the 14 primary driver-for this. The cultivated oases-oasis areas shrank from the Han to the Yuan 15 Dynasty but thereafterand have expanded until now, the natural oases oasis areas were 16 continually shrinking shrank until 2000, and the areas of desertified land have increased as 17 cultivated land was abandoned due to war, disasters or other causes. The volumes of 18 streamflow water flowing into terminal lakes remained about 1 billion m<sup>3</sup> per year, even 19 more in historical periods, but it decreased sharply after 1975, and even became negative. 20 The negative values (which would be zero in reality) probably indicate that which meant the 21 groundwater was being overexploited so that there was a negative mass balance, which is 22 consistent with falling water tables in recent times. After a water reallocation scheme was 23 24 implemented in 2000, the ecological and environment deterioration was halted and the lakes were restored. At the same time, the cultivated oases oasis areas, population and human water 25 consumption increased. This was at the expense of groundwater in midstream areas combined 26 withand the benefits of a wet period of about ten years. 27

Figure 5. Changes in elements of theof reconstructed catchment water balance elements
 inover the past 2000 years in the mid- and downstream area.

## **3.6 Determination of the development stages of evolutionary processes of human-water**

#### relationships

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The hHuman water consumption and natural oases-oasis\_areas changed with at\_different rates
(k) in different periods (Fig. 5). Based on their change rates of change with time we divided
the evolutionary processes of the human-water relationships in the HRB in-over the past 2000
years into four phases (Fig. 6): (1) predevelopment (206 BC\_- 1368 AD), (2) take-off (1368 1949 AD), (3) acceleration (1949 - 2000 AD), and (4) the start of rebalancingStarting to
rebalance between the human and water relationships (after 2000 AD).

11 **Figure 6.** The development stages of evolutionary processes of human-water relationships.

The predevelopment phase started after the Han Dynasty. In the Han Dynasty an 12 unprecedented expansion of manmade cultivated areas based on oases occurred associated 13 with. This happened because of defence needs, immigration and settling of farms, which 14 changed the production mode from nomadic herding into settled farming (Cheng et al., 2011). 15 It also corresponded with the warm and humid climate in the early Western Han Dynasty 16 (Ren et al., 2010; Xie et al., 2009). However, in the late eastern Han Dynasty, agricultural 17 declined due population production levels to loss and damage to water 18 19 conservation<del>conservancy</del> facilities after long-term warfare. During-From the Southern and Northern Dynasties (420-581) to the Yuan Dynasty (1271-1368), the people led nomadic 20 lifestyles and the Hexi corridor was in the a state of frequent wars and dynastic changes, and 21 22 the The HRB landuse was primarily pastoral-land as most agricultural oases were abandoned (Li, 1998; Xie et al., 2013). In this predevelopment stage of about 1500 years, the population 23 did-was stablenot increase, and cultivated areareclamation was small and focused on 24 downstream areas. As a result, so humans had little few impacts on the water system and the 25 area of natural oases area did not change significantly. 26

27 Since the Ming Dynasty, in which when agricultural civilization revived, the evolutionary processes of the human-water relationships in the HRB entered the take-off stage. During this 28 phase oases oasis reclamation activities were promoted and moved upstream to the midstream 29 30 area (Wu, 2000). In the middle of the Qing Dynasty, the Hexi corridor was politically stable and free of wars and the basic requirements for agricultural development were provided by 31 the government, includinginnovative farming and engineering methods were introduced, such 32 as better seeds, new cropscattle, and the steel farming implements. This led to expansion of 33 the cultivated land and agricultural development. Therefore the population increased quickly 34 (Shi, 2010). At the same time irrigation technology hardly changed, and with the area of 35 cultivated land expanding, water resource utilization became increasingly intense (Wang, 36 2003). It was also during this period that water-disputes about water arose (Cheng et al., 2011; 37 Shen and He, 2004). This phase was relatively short, lasting about 580 years. During this 38 phase, human water consumption increased at a rate of 1.09 million m<sup>3</sup> per year on average, 39 and the area of natural oases-area decreased at an average rate of 1.38 km<sup>2</sup> per year. The 40 hHuman intervention of in the water system was gradually increasing. 41

After <u>New-modern</u> China was founded <u>in 1949</u>, <u>the socialsocietal</u> development in the HRB <u>stepped-moved\_into</u> the acceleration stage. During this stage the population, <u>the area of</u> cultivated land and human water consumption increased sharply, especially after the <del>world-wide</del> green revolution <u>in-of</u> the 1960s, and China's reform and opening-up in 1978. In

addition, food self-sufficiency has dominated Chinese agricultural and water resources 1 development policy. Many wells, reservoirs and channels were built during this stage. This 2 stage was the shortest, only 50 years long, but the human water consumption increased at an 3 alarming rate of 35.1 million m<sup>3</sup> per year, and the <u>area of natural oases area decreased at an</u> 4 average rate of 58 km<sup>2</sup> per year. The influence of human activities on water resources reached 5 its peak and the environment was seriously degraded as natural wetlands, rivers, and lakes 6 7 dwindled rapidly (Xiao et al., 2004).

In order to prevent continuing environmental degradation a series of actions and measures 8 9 were carried outimplemented, such as the Natural Forest Protection Project after 2001, and another large project of that turning the cultivated land into forests or grasslands from 2002 to 10 2004, and. In addition, Zhangye city, in the midstream area, was selected as the first 11 construction experimental construction site by the Water Saving and Conservation Society 12 (WSCS) of China in 2002. This was supported by a water reallocation scheme in 2000 by in 13 which the midstream area should discharge 950 million m<sup>3</sup> of water in normal years (as 14 measured at the ZYX station) to downstream areas in normal years (when the upstream YLX 15 discharges 1,580 million m<sup>3</sup> of water). The Central Government's No.1 Water Document in 16 2011, which limits total water diversion, promotes water use efficiency and reduces water 17 pollution, reinforces the changessignaled a big step in the relationship between humans and 18 water emerging since the early 2000s. All of these actions have resulted in some 19 improvements to downstream ecosystems, includingsuch as halting the ecological and 20 environment deterioration and restoring the lakes. The area of natural oases-area increased at 21 an average rate of 28 km<sup>2</sup> per year from 2000 to 2010. A new stateequilibrium stage between 22 the humans and water emerged since after 2000. 23

#### **4** Discussions and conclusions 24

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This paper represents an attempt to reveal the evolutionary processes of the human-water relationships in the HRB over the past 2000 years. We quantitatively analyzed the dynamics of coupled human and hydrological systems as well as the associated climatic and ecological changes in the past more than 2000 years within the HRB by reconstructing the catchment 28 water balance. Based on transition theory we divided the evolutionevolutionary processes of the human-water relationships into four stages, which are including predevelopment (206 BC - 1368 AD), take-off (1368 - 1949 AD), acceleration (1949 - 2000 AD), and rebalancing 32 (after 2000 AD).

This study for the first time provided provides new understandings of how societal drivers 33 34 and societal responses over time interact and feedback with catchment water cycles over a timescale of 2000 years. The pace of This the evolutionary process was not at a uniform 35 pacevaried. The predevelopment stage lasted forexperienced 1500 years, and the take-off 36 37 periodtake-off was shorter at only 580 years, and aAfter that, in a period of only 50 years, only 60 years' the acceleration period occurred when the population increased up from 38 0.5 to 1.9 million, and the area of cultivated oases areas expanded by 3649  $\text{km}^2$ , which was 39 about two times double that in at the beginning of the acceleration this stage, and h. Human 40 water consumption increased by 1.9 billion m<sup>3</sup> per year, resulting in a doubling of water use 41 over, which was more than two times of that in the beginning of this the stage. This resulted 42 in volumes of water from midstream areas being discharged flowing into terminal lakes 43 decreasing from more than 1 billion  $m^3/year$  to 0. This situation became the trigger for a 44 sustainability transition in the HRB in 2000 when a water reallocation scheme was 45 46 relationships in the basin enteredstarted a new stage: rebalancing. This understanding of the 47 dynamics of transitions will assist policy makers to identify management practices that 48

require improvement by understanding how today's conditions and problems were created in
 the past. It could also help integrate management of land and water use to allow for more
 sustainable catchment management to combatagainst desertification in this region.

This paper, through reconstruction, incorporated metrics of human-water interaction into 4 fundamental understanding of complex human-water systems. The quantitative historical 5 analysis not only improved our understanding of past human-water relationships but also 6 facilitated improved predictions of its possible future dynamics. It has added a valuable case 7 study for comparative socio-hydrologic studies across different human-water systems around 8 the world. This paper has suggested some guidelines toward an analytical approach to water 9 related societal transitions that should be, on one hand, strongly attached to social science 10 theory, and on the other hand, firmly based on formal hydrological modeling. It can be seen 11 from the four stages of evolutionary processes of human-water relationships in the HRB that 12 transitions have no fixed pattern. The stabilization, a typological phase in the standard 13 transition theory, did not appear. In addition, there were large differences in the rates and 14 scales of changes and the period of time over which they occurred. This happened because 15 during a process of change, humans are able to adapt to, learn from and anticipate new 16 situations (Chen. 2005). 17

An important part of the paper was reconstructing the This paper reconstructed catchment 18 water balance. This relied onby using a range of data sources, including paleo-climates and 19 paleo-environments reflected by dendrochronology, ice cores, lake sediments and historical 20 drought and flood sequences, a historical atlas of China, remote sensing images and 21 instrumented streamflow and climate data. The resulting reconstructed water balance was 22 consistent with the dynamics of the terminal lakes, which explained the evolutionary-process 23 of human-water relationship in the HRB in over the past 2000 years with relatively good 24 agreement. The reconstruction provided the a basis for generating baseline data against which 25 to evaluate recent changes, for investigating the impact of human societies on hydrological 26 27 systems in historical contexts and for generating datasets for improving models of hydrological systems over timescales that exceed the length of the instrumented record 28 (Savenije et al., 2013). 29

-There are some important limitations on the methods and with the data collection and 30 analysis. Several assumptions and uncertainties in the 2000 year hydrological reconstruction 31 exist due to lack of data. Values for  $E_0$  in the historical periods were assumed to be the same 32 as in recent periods, which is reasonable given the variation in average temperature was less 33 than 2  $^{\circ}$ C. Values for w may vary among different land use types and could change with time 34 depending on the type and intensity of crops in historical periods, however due to the lack of 35 data, the value of w for HRB was set at 3.5. For the same reason, irrigation was set at 500 mm 36 per year for the whole historical period. There was also some inconsistencyHowever, there 37 were some discrepancies among the reconstruction methods. They might come from: 1) 38 Inconsistency between the data extracted from the different proxy materials, for example, the 39 streamflow reconstructions by Yang et al. (2012), Qin et al. (2010), and Kang et al. (2002) 40 using tree rings were not completely consistent. There were; 2) limitations of due to the 41 available data's representativeness of locations, for example the data from tree rings only 42 focused in the upstream area of the mainstream areas of the Heihe River, and the samples of 43 lake sediment mainly focused in on the terminal lake Sogo Nur. Problems of ; and 3) 44 non-representativeness of data in various time periods and varyingdifferent resolutions of 45 data also occurred. For, for example, the land use maps only covered several periods, the tree 46 ring dating can be specific to the annual scale, and the information from ice cores and lake 47 sediment profiles was at the century scale. In future, we should improve the consistency, 48 length and quality of historical datasets by advancing data analysis techniques. 49

The transitions seen from the four stages of evolutionary processes of human-water 1 relationships in the HRB did not follow the standard theoretical processes. Stabilization, a 2 typological phase in the standard transition process, did not appear. In addition, there were 3 large differences in the rates and scales of changes and the period of time over which they 4 occurred. Some further theoretical research is needed to explain the transition pattern, but this 5 result evidences that transitions have no fixed pattern. This paper provides a path toward an 6 analytical approach to water related societal transitions that should be, on one hand, strongly 7 attached to social science theory, and on the other hand, firmly based on formal hydrological 8 modeling. 9

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#### 18 **References**

- 19 Budyko, M.I.: Climate and Life. Academic, San Diego, CA, pp. 508, 1974.
- Carpenter, S. R., Cole, J. J., Pace, M. L., Batt, R., Brock, W. A., Cline, T., Coloso, J.,
  Hodgson, J. R., Kitchell, J. F., Seekell, D. A., Smith, L., and Weidel, B.: Early warnings of
  regime shifts: a whole-ecosystem experiment, Science, 332, 1079–1083, 2011.
- 23 Chen, L.: Land desertification and its control strategies in the low reaches of the Heihe River,
- 24 J. Nat. Resour., 1, 35–43, 1996.
- Cheng, H., Huang, Y., and Zhao, L.: The human activity in Hexi Corridor during historical time, 10 available at: http://www.paper.edu.cn (last access: 5 April 2014), 2011.
- Elshafei, Y., Sivapalan, M., Tonts, M., and Hipsey, M.: A prototype framework for models of
   socio-hydrology: identification of key feedback loops and parameterisation approach,
   Hydrol. Earth Syst. Sci., 18, 2141-2166, 2014.
- Falkenmark, M. and Lannerstad, M.: Consumptive water use to feed humanity curing a
  blind spot, Hydrol. Earth Syst. Sci., 9, 15–28, doi:10.5194/hess-9-15-2005, 2005.
- Fu, B.: On the calculation of the evaporation from land surface, Sci. Atmos. Sin., 5, 23–31,
  1981.
- Geels, F. W.: Technological transitions as evolutionary reconfiguration processes: a
   multi-level perspective and a case-study, Res. Policy, 31, 1257–1274, 2002.
- Holmes, J. A., Cook, E. R., and Yang, B.: Climate change over the past 2000 years in
  Western China, Quatern. Int., 194, 91–107, 2009.
- Jin, H., Xiao, H., Sun, L., Zhang, H., Sun, Z., and Li, X.: Vicissitude of Sogo Nur and
  environmental-climatic change during last 1500 years, Sci. China Ser. D, 47, 61–70, 2004.
- Jin, H., Xiao, H., Zhang, H., and Sun, Z.: Evolution and climate changes of the Juyan Lake
  revealed from grain size and geochemistry element since 1500aBP, J. Glaciol. Geocryol.,
  27, 233–240, 2005.
- 43 Kallis, G.: Coevolution in water resource development the vicious cycle of water supply and

- Kang, X., Cheng, G., Kang, E., and Zhang, Q.: Mountainous runo\_ reconstruction of Heihe
  River during past 1000 years using tree-ring, Sci. China Ser. D, 32, 49–53, 2002.
- Li, B.: An investigation and study on the desertification of the ancient oases from Han to
  Tang dynasties in the Hexi corridor, Acta Geogr. Sin., 53, 106–115, 1998.

Liu, K., Yao, Z., and Thompson, L. G.: A pollen record of Holocene climatic changes from the Dunde ice cap, Qinghai-Tibetan Plateau, Geology, 26, 135–138, 1998.

Liu, Y., Tian, F., Hu, H., and Sivapalan, M.: Socio-hydrologic perspectives of the
co-evolution of humans and water in the Tarim River basin, Western China: the Taiji–Tire
model, Hydrol. Earth Syst. Sci., 18, 1289–1303, doi:10.5194/hess-18-1289-2014, 2014.

- Lowry, D. P. and Morrill, C.: Changes in the Global Hydrological Cycle: Lessons from
   Modeling Lake Levels at the Last Glacial Maximum, American Geophysical Union, Fall
   Meeting, San Francisco, California, USA, 2011.
- Montanari, A., Young, G., Savenije, H., Hughes, D., Wagener, T., Ren, L., Koutsoyiannis, D.,
  Cudennec, C., Toth, E., and Grimaldi, S.: "Panta Rhei Everything Flows": change in
  hydrology and society the IAHS Scientific Decade 2013–2022, Hydrolog. Sci. J., 58,
  1256–1275, 2013.
- Pataki, E., Carreiro, M., Cherrier, J., Grulke, E., Jennings, V., Pincetl, S., Pouyat, V.,
  Whitlow, H., and Zipperer, C.: Coupling biogeochemical cycles in urban environments:
  ecosystem services, green solutions, and misconceptions, Front. Ecol. Environ., 1, 27–36,
  doi:10.1890/090220, 2011.
- Qin, C., Yang, B., Burchardt, I., Hu, X., and Kang, X.: Intensified pluvial conditions during
  the twentieth century in the inland Heihe River Basin in arid northwestern China over the
  past millennium, Global Planet. Change, 72, 192–200, 2010.
- Qu, W., Wu, R., Wang, S., and Zhang, Z.: Sedimentary pigment and its environmental
  signification of East Juanyanhai in Inner Mongolia since the past 2600 years, Acta
  Sendiment. Sin., 48, 13–17, 2000.
- Ren, Z., Lu, Y., and Yang, D.: Drought and flood disasters and rebuilding of precipitation
  sequence in Heihe River basin in the past 2000 years, J. Arid Land Resour. Environ., 24,
  91–95, 2010.
- Röckstrom, J., Karlberg, L., Wani, S. P., Barron, J., Hatibu, N., Oweis, T., Bruggeman, A.,
  Farahanie, J., and Zhu, Q.: Managing water in rainfed agriculture the need for a paradigm
  shift, Agr. Water Manage., 97, 543–550, 2010.
- Rotmans, J.: Societal Innovation: Between Dream and Reality Lies Complexity, DRIFT
   Research Working Paper, Erasmus Research Institute of Management (ERIM), Rotterdam,
   the Netherlands, doi:10.2139/ssrn.878564, 2005.
- Saito, K., Hirai, M., and Yonekura-Sakakibara, K.: Decoding genes with coexpression
  networks and metabolomics –"majority report by precogs", Trends Plant Sci., 13, 36–43,
  2008.
- Savenije, H. H. G., Hoekstra, A. Y., and van der Zaag, P.: Evolving water science in the
  Anthropocene, Hydrol. Earth Syst. Sci., 18, 319–332, doi:10.5194/hess-18-319-2014,
  2014.
- Shao, X., Xu, Y., Yin, Z.-Y., Liang, E., Zhu, H., and Wang, S.: Climatic implications of a
  3585-year tree-ring width chronology from the northeastern Qinghai-Tibetan Plateau,

Quaternary Sci. Rev., 29, 2111–2122, 2010. 1 Shen, J., Zhang, E., and Xia, W.: Records from lake sediments of the Qinghai Lake to mirror 2 climatic environmental changes of the past about 1000 years, Quaternary Sciences, 21, 3 508-513, 2001. 4 Shen, M. and He, L.: Comparison of new and old equal water regime in Heihe River basin, 5 Yellow River, 26, 27–29, 2004. 6 Sheppard, P., Tarasov, P., Graumlich, L., Heussner, K.-U., Wagner, M., Österle, H., and 7 Thompson, L.: Annual precipitation since 515 BC reconstructed from living and fossil 8 juniper growth of northeastern Qinghai Province, China, Clim. Dynam., 23, 869-881, 9 2004. 10 11 Shi, J.: Integrating humanistic and scientific studies to reveal the changes of Khara Khoto, Studies in humanity and environment of Khara Khoto, in: Proceedings of international 12 symposium on the humanity and environment of Khara Khoto region, Renmin University 13 of China Press, Beijing, 1-4, 2007. 14 Shi, L.: Study on the spatio-temporal process of oasisization and desertification in the period 15 of the Ming, Qing dynasty and Republic of China in the middle reaches, MD thesis, 16 17 Lanzhou University, Lanzhou, 2010. Shi, M., Wang, L., and Wang, X.: A study on changes and driving factors of agricultural 18 water supply and demand in Zhangye after water reallocation of the Heihe River, 19 Resources Science, 33, 1489-1497, 2011. 20 Sivapalan, M., Blöschl, G., Zhang, L., and Vertessy, R.: Downward approach to hydrological 21 prediction, Hydrol. Process., 17, 2101–2111, 2003. 22 Sivapalan, M., Savenije, H. H., and Blöschl, G.: Socio-hydrology: a new science of people 23 and water, Hydrol. Process., 26, 1270–1276, 2012. 24 T abara, J. D. and Ilhan, A.: Culture as trigger for sustainability transition in the water domain: 25 the case of the Spanish water policy and the Ebro river basin, Reg. Environ. Change, 8, 26 59-71, 2008. 27 Tan, Q.: The Historical Atlas of China, China Cartographic Publishing House, Beijing, 1996. 28 29 Thompson, S. E., Sivapalan, M., Harman, C. J., Srinivasan, V., Hipsey, M. R., Reed, P., Montanari, A., and Blöschl, G.: Developing predictive insight into changing water systems: 30 useinspired hydrologic science for the Anthropocene, Hydrol. Earth Syst. Sci., 17, 31 5013-5039, doi:10.5194/hess-17-5013-2013, 2013. 32 Vörösmarty, C. J., McIntyre, P. B., Gessner, M. O., Dudgeon, D., Prusevich, A., Green, P., 33 Glidden, S., Bunn, S. E., Sullivan, C. A., Reidy Liermann, C., and Davies, P. M.: Global 34 threats to human water security and river biodiversity, Nature, 467, 555–563, 2010. 35 36 Wang, G., Yang, L., Chen, L., and Jumpei, K.: Impacts of land use changes on groundwater resources in the Heihe River basin, Acta Geogr. Sin., 60, 456–466, 2005. 37 Wang, G., Xie, Y., Wang, X., Yu, L., and Shi, Z.: Data reconstruction of Heihe River basin 38 cultivated land area prior to the Ming dynasty, Resources Science, 35, 362–369, 2013. 39 Wang, Y.: The development history of water conservancy facilities in Heihe River basin, 40 Gansu Nationalities Press, Lanzhou, 2003. 41 42 Wei, H.: Groundwater age and sustainability based on process simulation in Zhangye basin, PhD thesis, Cold and Arid Regions Environment and Engineering Research Institute, 43 Chinese Academy of Sciences, Lanzhou, 2013. 44 Wimmer, A.: Models, methodologies, and metaphors on the move, in: Understanding Change 45

Models, Methodologies, and Metaphors, edited by: Wimmer, A. and Kössler, R., Palgrave
MacMillan, London, England, 2006.
Woodhouse, A., Gray, T., and Meko, M.: Updated streamflow reconstructions for the Upper
Colorado River Basin, Water Resour. Res., 40, W05415, doi:10.1029/2005WR004455, 2006.
Wu, X.: Historical variance of the ecological environment in the inland river area along the
Hexi corridor, J. Lanzhou Univ. (Social Sciences), 28, 46–49, 2000.
Xiao, S. and Xiao, H.: Farming-grazing vicissitude and man–land relation evolution of Ejin Banner in historical period, J. Desert Res., 24, 449–451, 2004.
Xiao, S. and Xiao, H.: Advances in the study of the water regime process and driving mechanism in the Heihe River basin, Adv. Earth Sci., 23, 748–755, 2008.
Xiao, S., Xiao, H., Zhou, M., Si, J., and Zhang, X.: Water level change of the west Juyan
Lake in the past 100 years recorded in the tree ring of the shrubs in the lake banks, J. Glaciol. Geocryol., 26, 557–562, 2004.
Xie, Y.: Dataset of cultivated oasis distribution in the Heihe River Basin during the historical
period, Heihe Plan Science Data Center, Lanzhou, China, doi:10.3972/heihe.092.2013.db, 2013.
Xie, Y., Chen, F., and Qi, J.: Past desertification processes of Minqin Oasis in arid China, Int.
J. Sust. Dev. World, 16, 260–269, 2009.
Xie, Y., Wang, X., Wang, G., and Yu, L.: Cultivated land distribution simulation based on grid in middle reaches of Heihe River basin in the historical periods, Adv. Earth Sci., 28, 71–78, 2013.
Yang, B., Qin, C., Shi, F., and Sonechkin, D. M.: Tree ring-based annual streamflow reconstruction for the Heihe River in arid northwestern China from AD 575 and its implications for water resource management, Holocene, 22, 773–784, 2012.
Yang, B., Qin, C., Wang, J., He, M., Melvin, T. M., Osborn, T. J., and Bri_a, K. R.: A
3,500-year tree-ring record of annual precipitation on the northeastern Tibetan Plateau, P.
Natl. Acad. Sci. USA, 111, 2903–2908, 2014.
Yang, D., Sun, F., Liu, Z., Cong, Z., Ni, G., and Lei, Z.: Analyzing spatial and temporal variability of annual water-energy balance in nonhumid regions of China using the Budyko
hypothesis, Water Resour. Res., 43, W04426, doi:10.1029/2006WR005224, 2007. Zhang, L., Dawes, W., and Walker, G.: Response of mean annual evapotranspiration to
vegetation changes at catchment scale, Water resources research, 37, 701-708, 2001.
Zhang, L., Hickel, K., Dawes, W., Chiew, F. H., Western, A., and Briggs, P.: A rational
function approach for estimating mean annual evapotranspiration, Water Resour. Res., 40,
W02502, doi:10.1029/2003WR002710, 2004.
Zhang, L., Pang, B., Xu, Z., and He, R.: Impacts of climate change and LULC on
hydrological processes in the Gulang River basin, Sounth-to-North Water Transfers and Water Science & Technology, 12, 42-46, 2014.
Zhang, Z., Wu, R., Wang, S., Xia, W., Wu, Y., and Qu, W.: Environmental changes recorded
by lake sediments from East Juyanhai Lake in Inner Mongolia during the last 2600 years,
J.Lake Sci., 10, 44–51, 1998.
Zhao, J., Xu, Z., and Zuo, D.: Spatiotemporal variation of potential evapotranspiration in the Heihe River basin, Journal of Beijing Normal University (Natural Science), 49, 164-169,
2013.

1	Zhao, L. and Ji, X.: Quantification of transpiration and evaporation over agricultural field
2	using the FAO-56 Dual Crop Coe_cient Approach – A case study of the maize field in an
3	oasis in the middlestream of the Heihe River Basin in northwest China, Scient. Agr. Sin.,
4	43, 4016–4026, 2010.
5	Zheng, J. and Wang, S.: Assessment on climate change in China for the last 2000 years, Acta
6	Geogr. Sin., 60, 21–31, 2005.
7	Zheng, J., Shao, X., Hao, Z., and Ge, Q.: An overview of research on climate change in China
8	during the past 2000 years, Geogr. Res., 29, 1561–1570, 2010.
9	Zhou, L. and Yang, G.: Ecological economic problems and development patterns of the Arid
.0	Inland River Basin in Northwest China, Ambio, 35, 316–318, 2006.







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