# HESS 2015-325

# "Parameter Regionalization of a Monthly Water Balance Model for the Conterminous United States"

# Major Manuscript Changes

Pg. 2, Line 1-2 (Abstract): removed "and model uncertainty information"

# 1. Introduction

Pg. 4, lines 81-89, 98: Re-formatted text to include some recent work recommended by referee 2

Pg. 5, lines 110-111: Removed the sentence "These methods ignore parameter interaction and often assume that model algorithms have linear responses to different parameters (Cuo et al., 2011)"

# 2.1 Monthly Water Balance Model

Pg. 6, lines 155-160: Removed definitions and descriptions of parameters and their ranges.

Pg. 7, lines 177-178: Substituted reference for url. Added reference (USEPA and USGS, 2010) to references list.

Pg. 7, lines 171: Added "Processes influenced by" to Figure 1 Caption

# 2.2 Fourier Amplitude Sensitivity Test

Pg. 8, lines 200, 217, 218; Pg. 9 lines 221, 230: Changed "first order partial variance (FOPV)" to a more meaningful term.

Pg. 8, line 204: deleted the statement "requiring much less information and parameter sets than other global methods

Pg. 8, lines 205-206: Added the sentence for seasonal adjustment factors.

Pg. 8, line 211: Deleted Fig.3 reference (Fig. 3 is replaced by a different figure).

Pg. 9, lines 222-224: Deleted caption for old Fig. 3

Pg. 9, lines 225-228: Simplified sentence based on Ref. 1 recommendation.

Pg. 9, lines 230-235: Clarified information on the number of FAST runs.

# **3** Parameter regionalization procedure

Pg. 10, lines 252-254: Added Fig. 3 (Schematic Flowchart)

# 3.1 Parameter sensitivities

Pg. 10, lines 265: Deleted "The patterns of" to start sentence

Pg. 11, lines 299-302: Added details for FAST results

Pg. 12, line 305: Changed FOPV to parameter sensitivities for caption of Figure 5.

## **3.2** Calibration regions

Pg. 12, lines 312-315: Modified this sentence to add clarification on the HRU grouping procedure

Pg. 12, lines 318-333: Substantially modified this paragraph to better explain the derivation and purpose of the two HRU classifications based on sensitivity analysis results

Pg. 13, lines 335-336: Changed reference for this sentence to Tang et al., 2007 from Pianosi et al., 2015 (removed Pianosi et al., 2015 from reference list)

Pg. 13, lines 342-333, 345-347: Modified text to add some clarification to additional classification steps.

Pg. 13, lines 349-351: Moved this sentence from earlier in the paragraph the end

Pg. 13, Lines 352-356: Modified figure caption

# **3.4.2** Grouped streamgage calibration

Pg. 16, lines 422-24: Added sentence defining USGS Reference streamgages

Pg. 16-17, lines 434-441: Added equation (Eq. 1) for the objective function used for grouped calibration.

Pg. 17, line 444: changed "variable" to "streamflow"

Pg. 17, lines 451-453: Added sentences clarifying how 25% error bound for SWe worked

Pg. 17, lines 456-457: added comment re-call this papers definition of a sensitive parameter

Pg. 17, line 460: deleted "on a monthly basis"

# 4.2.2 Nash-Sutcliffe efficiency

Pg. 21, lines 571-573: modified caption for figure 14.

## 5.1 Regionalized Parameters

Pg. 22, lines 582-586: Added some more details to help explain the results of Fig. 13

## 6 Conclusions

Pg. 24, line 638: removed "and model uncertainty information"

## 7 References

Added the following references:

Pg. 25, lines 688-690:

Bárdossy, A., Huang, Y., and Wagener, T.: Simultaneous calibration of hydrological models in geographical space, Hydrol. Earth Syst. Sci. Discuss., 12, 1123-11268, doi:10.5194/hessd-12-11223-2015, 2015.

Pg. 31, lines 806-808:

Qamar, M.U., Ganora, D., and Claps, P.: Monthly Runoff Regime Regionalization Through Dissimilarity-Based Methods, Water Resour. Manag., 29, 4735-4751, doi:10.1007/s11269-015-1087-7, 2015.

Pg. 32, lines 844-847:

USEPA (United States Environmental Protection Agency) and USGS (United States Geological Survey): NHDPlus User Guide, available at ftp://ftp.horizonsystems.com/NHDPlus/documentation/NHDPLUS\_UserGuide.pdf (last access 12 Nov 2014), 2010.

Removed the following references:

Pg. 31, 804-805:

Pianosi, F., Sarrazain, F., and Wagener, T.: A Matlab toolbox for Global Sensitvity Analysis, Environ. Modell. Softw., 70, 80-85, doi:10.1016/j.envsoft.2015.04.009, 2015.

## Figures

Fig. 3, Pg. 37 – removed previous fig. 3 (FAST parameter waves), replaced with schematic flowchart of Section 3

Fig. 5, Pg. 40 – modified y-axis to "Param Sens." From "FOPV"

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Interactive Comment

# Interactive comment on "Parameter regionalization of a monthly water balance model for the conterminous United States" by A. R. Bock et al.

# A. R. Bock et al.

abock@usgs.gov

Received and published: 4 January 2016

Reviewer Minor Points: -p. 10024, lines 1-2 "to transfer ... model uncertainty information". What type of uncertainty information is transferred and how? This is mentioned here and in the conclusion but it is not clearly discussed throughout the paper.

AB: Mean Monthly errors for each calibration region (visualized in Figure 8 and 10a) can be estimated and added back to simulated streamflow estimates at ungaged locations as a source of model uncertainty. This is not explicitly discussed or applied in detail in this paper, so authors may need to add more detail at one location of text or remove.





-p. 10026: lines 28: "these methods ignore parameter interaction, and often assume that model algorithms have linear responses to different parameters". I think this sentence is misleading and I would suggest to delete it. Parameter interactions can be evaluated in local SA by computing second-order derivatives (see for example Norton, 2015). Also, when estimating local sensitivities the linearity assumption finds its rationale in the Taylor series expansion and hence it is quite reasonable.

AB: Authors agree to removing sentence.

-p. 10028, line 25 to the end of page: this list of parameter names and meaning does not add much to the information provided in the Table, I would probably avoid it.

AB: Authors agree, parameter definitions and functions are also well-explained in the cited McCable/Wolock papers listed in the reference section. Sentence at p. 10028, lines 23-24 "Table 1 lists...", and climate adjustments sentence beginning with "The Ppt\_adj and..." (p. 10029, lines 2-3) should be moved to the concluding sentence of the opening paragraph of Section 2.1. The remaining sentences can be removed.

-p. 10030, line 8: the term FOPV is not particularly self-explaining to readers not familiar with GSA. I would explain what it is ("contribution to output variance from ...")

AB: Authors agree. The sentence "FAST is a variance-based global sensitivity algorithm that estimates the first-order partial variance (FOPV)..." can be re-worded to "FAST is a variance-based global sensitivity algorithm that estimates the contribution to output variance..." . "Output Variance' should replace FOPV in text, including the Y-axis labels for Figure 5 will be re-named to "Output Variance" with a single axis label.

-p. 10030, line 12: "much less information and parameter sets". What do you mean by "information"? Unclear. As for "parameter sets, it is possibly less ambiguous if you call them parameter samples or even directly model evaluations

AB: This second half of the sentence is pretty ambiguous. The application of FAST discussed in the paper used a larger number of parameter sets than the minimally

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sufficient number suggested by R, so the sentence should probably be removed.

-p. 10032, lines 12-13: "The patterns of ...". Sentence needs rewording

AB: Agreed. Change to: "Tsnow, Train, and Meltcoef all share similar patterns of areas with higher sensitivity across the CONUS."

-p. 10032, lines 12-13: "The patterns of ...". Sentence needs rewording

AB: Agreed. Change to: "Tsnow, Train, and Meltcoef all share similar patterns of areas with higher sensitivity across the CONUS."

-p. 10033, line 25: "While this idea...". What idea? The one described in the previous sentence? But then is it really in contrast with the one illustrated on lines 27-28? Please clarify.

AB: This is a good point on semantics. The emphasis should be on grouping proximate areas based on similar model behavior, rather than physiographic characteristics. The authors suggest this sentence be changed to: "This idea is rooted in the hypothesis that geographically proximate HRUs share similar forcings and conditions, and thus will behave similarly. This application uses similarity in SA results as a basis for organization, rather than similarity in physiographic characteristics."

-p. 10034, lines 11-13: citation of Pianosi et al., 2015 does not seem to be appropriate here. That paper introduces a toolbox for Sensitivity Analysis but it does not discuss the issue of setting the threshold for sensitive and non-sensitive parameters. The threshold issue is (partially) discussed in Tang et al. (2007). Pianosi et al. have another paper under review which is more focused on the threshold issue, however it has not been published yet. The authors might cite that paper when it will be published (title is "Global Sensitivity Analysis of environmental models: Convergence and validation", journal is Environmental Modelling and Software).

AB: We have referenced Tang et al (2007) already in the paper, so we will replace Painosi et al 2015 with Tang et al (2007). Thanks for the heads up on the upcoming

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paper.

-p. 10037, line 17: Please give a very brief definition of a reference streamgage.

AB: Reference quality stramgages are judged to be largely free of human alterations. From Kiang et al. (2013), these sites were "categorized as either reference quality or nonreference quality by calculating a hydrologic disturbance index (presence of dams, change in reservoir storage, number of canals, road density, proximity to major pollutant discharge site, estimates of water withdrawals, and fragmentation of undeveloped land), reviewing historical digital maps and imagery for evidence of hydrologic alteration and human activity, and reviewing comÂňments in USGS annual water data reports for information on regulation or diversions.". We will paraphrase this to: "Reference streamgages are USGS streamgages deemed to be largely free of anthropogenic impacts and flow modificaitons, and can subsequently be used for estimation of natural flow statistics (Falcone, 2010; Kiang et al., 2013)."

-p. 10038, line 9: "simulated streamflow" should be "simulated variable" (since one of the four is SWE and not runoff)

AB: Authors agree to the suggested change.

-p. 10038, line 20: Recall here that a parameter is deemed insensitive if sensitivity index is below 5%

AB: Authors agree to the suggested addition to text

-p. 10038, lines 21-23: "on a mean monthly based". Unclear. Possibly it might just be dropped, since it was already said that monthly variables are used to compute the Zscores.

AB: Authors agree "mean monthly" should be removed from 21 and 23 since other terms were used in the objective function.

-caption of Fig. 1: "model parameters used in..." Maybe better: "processes influenced

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by the model parameters used in..."

AB: Authors agree to the suggested change

-Figure 3: maybe not needed. Anyway, if maintained, vertical axis should show units of measurements. Also, it would probably be better to show Drofac and Rfactor in a separate panel.

AB: Between the two suggestions (Remove or make two panesl), authors would prefer remove graphic (though Andy B. really likes this graphic).

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Interactive Comment

# Interactive comment on "Parameter regionalization of a monthly water balance model for the conterminous United States" by A. R. Bock et al.

# A. R. Bock et al.

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Received and published: 5 January 2016

Author's Responses to referee #1:

**Reviewer Main Points:** 

-[1] The parameter regionalisation procedure could be explained more effectively. In the first place, it would be good to have a schematic of the procedure to clearly see what is the role, inputs and outputs of each step (sensitivity analysis, classification of regions, individual calibration, grouped calibration, etc.).

AB: A schematic is a good idea. We will create an example for inclusion into the paper,



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as well as locations within the text where references to the schematic would be helpful.

-The structure of Sec 2-3-4 could be revised to better separate out the methodology from the illustration of results. For instance, I find a bit odd that sensitivity analysis results are presented in Sec. 3.1, before describing how they will be used in the proposed methodology. Another example is the first paragraph of Sec. 4.1, which explains why the individual streamgage calibration is needed, it would fit better in a "methods" section rather than the "results" section.

AB: We built the main methodology presented in the paper based on the results from the sensitivity analysis. Because we wanted to stress the independence of the sensitivity analysis from the calibration and regionalization procedure, we pushed the sensitivity analysis results to 3.1. The way sensitivity estimates are used for regionalization (described on page 10034,line 7 onward) needs to be explained more clearly, especially since this is the most novel aspect of the proposed methodology. Specifically:

-What is the connection between the first and second classification? They are independent from each other and then intersected to obtain the actual classification? Please clarify

-Description of the second classification (lines 10-11) is also unclear. What are the "unique combinations of parameter sensitivities"? How are they defined? What is their meaning?

AB: Correct, the two classifications are independent. The first classification (p. 10034, lines 6-9) derives regions based on hydrologic response units (HRUs) with unique combinations of magnitudes of the five parameter sensitivities (highest to lowest). The intent is to identify geographic regions of similar model response or behavior based on the numerical orders of the sensitivities.

AB: The second classification (p.10034, lines 10-11) identifies regions based on HRUs with unique combinations of parameters with FAST-based parameter sensitivities that

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exceed 5% (of the cumulative 100%). The intent with the second classification is to identify geographic regions with similar important parameters identified by the sensitivity analysis). The resulting polygons of these two classifications are merged to create the final region classification.

-Be more specific on how the two classification approaches work. Sentence on lines 7-8 of page 10034 is too generic, does it mean that the parameter ranking is the same in each region?

AB: Yes, all hydrologic response units (HRUs) identified within each region in classification 1 (p. 10034, lines 6-9) have the same ranking of the 5 model parameters from highest sensitivity to lowest sensitivity. Additionally, all HRUs within each region identified in classification 2 (p. 10034, lines 10-11) had the identical subset of parameters which exceeded 5% from FAST.

-From lines 17-20, I understand that the sensitivity-based classification is further refined using a more 'conventional' approach that looks at proximity and topographic divides. How does this refinement step works? Does it introduce significant changes in the classification? This is important to know in order to understand the value of the proposed sensitivity-based classification versus proximity or topography-based classification.

AB: This more conventional approach was necessary because of the lack of streamgages available for calibration in some of the calibration regions. The density of the stream gage network can be very sparse for some geographic regions of the U.S., especially in arid/semi-arid areas (see Kiang et al., 2013 in the discussion paper reference list). Following the "unsupervised" merging of the two classifications, the authors conducted a "supervised" classification of regions with less than 3 gages, where these regions where merged with geographically proximate regions with adequate streamgage representation that also shared the most similar parameter sensitivity results.

AB: The primary topographic divides utilized were topographically-derived boundaries

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from the NHDPlus or Hydrologic Unit Code boundaries (black lines in Figure 2). These are sub-boundaries of both the model discretization and many USGS water resource management efforts. We felt it was important to maintain these boundaries, especially for the western United States where orographic climate effects are very important to the hydrologic cycle.

-[2] Some of the numerical results are a bit surprising and should be double-checked. In particular, in Fig. 5.a the fact that one parameter has sensitivity of exactly 100 and all others of exactly 0 seems odd.

AB: For regions with homogeneous, sub-tropical type climatic conditions, such as the Southeast, results such as Figures 5a and 5c were consistent across many of the objective functions we had used to measure parameter sensitivity with FAST (including parameter sensitivity measured for NSE at select reference streamgages). In areas such as the Lower Mississippi (Region 8, Figure 2), the amount of snowfall is negligible, so the three parameters that control snowfall and snowpack accumulation have negligible effect on total runoff. If there is minimal occurrence of snow in a region, then snow parameters won't be important, even in a complex model. For further discussion, a colleague has submitted a paper to HESS that strictly examines results of FAST applied to a 35-parameter daily streamflow model across the U.S.

-Also, the result of Figure 13 is very puzzling. As the authors note on page 10041 (line 27 onward), the groupNSE values are expected to be lower than the gageNSE values. Figure 13 instead shows many cases were groupNSE is much larger than gageNSE. I really struggle to believe that NSE can be increased so much and so often when using a model calibrated with a different objective function. The only explanation I can think of is that either the calibration algorithm in the gageNSE failed (for instance by getting stuck in a local minimum or being terminated too early) or that the comparison is not fair (for example that NSE refer to different time periods?). This needs clarification.

AB: What we wanted to emphasize with this plot is not a comparison of two calibration

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methods, but that the grouped calibration strategy that focuses on the sensitive parameters can provide just as much information as traditional NSE-based individual streamgage calibration. There is quite a bit of difference between the two calibration methods compared: the individual gage calibration used the entire period of record for each streamgage, while the grouped calibration uses an odd/even calibration/validation calibration strategy, different objective functions are used (NSE versus multi-term weighted objective function for the grouped calibration), and climate adjustments are derived for each streamgage in the individual calibration, and for the entire region in the grouped calibration.

Author's response to remaining minor remarks

-p. 10030, lines 14-15: please justify why you do not incorporate the adjustment factors in the FAST analysis

AB: We viewed the adjustment factors as more related to the forcing data itself and independent from the model structure.

-p. 10031, lines 6-7 "parameter ranges were based..." Are these the ranges in Table 1 and already commented on p.10028, line 24? If so, just refer to the Table here.

AB: The parameters listed in Table 1 were the bounding ranges. We will make the adjustment to the text.

-p. 10031, lines 7-9: What do you mean by "standard application"? Also, I suppose the R package uses the equation N = 2Nharm max(!) + 1 to determine the minimal number of runs. If so, better cite Cuckier et al (1973), which is where the formula comes from. Also, please mention what is this number in your case, it would help readers to get an idea of how computationally demanding is the proposed approach.

AB: "In standard application" is ambiguous and should be removed. We accept the referee's recommendations and the sentence on lines 8-10 should be re-worded to: "The fast R package can determine the minimal number of runs necessary to estimate

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the output variance all parameters (Cukier et al., 1973). For our application we generated an ensemble of 1000 sets (above the minimal number of 71 suggested by fast) to have the potential capability for further research into comparison of different sensitivity analysis methods."

-p. 10038, lines 1-4: The definition of the multi-term objective function is unclear. Are the four terms summed up? Why considering both mean monthly runoff and annual runoff (I would imagine that they convey the same information, the former being equal to the latter divided by 12)? I think inserting an equation with the mathematical definition of the objective function would help here.

AB: The objective function minimizes the sum of difference between the Z-scores of measured and simulated variables for four terms: Mean Monthly Streamflow (As shown in figures 8 and 10a), Monthly Streamflow (Raw monthly time series), Annual Streamflow (Time Series aggregated to annual time steps), and Mean Monthly SWE with a 25% error bound. The first three terms of the objective function were chosen because they conveyed information that can be used to easily inform other models (such as daily time-step models). Just to note, annual and mean monthly objective functions as we defined them are very different; the former is not equal to the latter/12.

We can include an equation with the appropriate text to help elucidate the different terms of the objective function.

-p. 10038, lines 14-16: Please clarify how the error bounds were taken into account. Did you modify the definition of the Z score for the SWE?

AB: Z-scores were calculated from basin mean monthly measured (SNODAS) and simulated SWE values. Upper and lower bounds of 25% were calculated for the mean monthly SWE Z-score (Zobs \*1.25 for the upper bound; Zobs \*.75 for the lower bound). If the simulated MWBM SWE value was contained within the upper and lower 25% bounds, the absolute difference was designated as 0. If the simulated MWBM SWE Z-score value was above the upper 25% error bound, the absolute Z-score difference

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was calculated between the simulated Z-score value and the value of the upper 25% Z-score bound. If the simulated Z-score was below than the lower 25% error bound, the absolute Z-score difference was calculated between the simulated Z-score value and the value of the lower 25% Z-score bound.

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# Interactive comment on "Parameter regionalization of a monthly water balance model for the conterminous United States" by A. R. Bock et al.

# A. R. Bock et al.

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Received and published: 6 January 2016

Author's Responses to referee #2:

**Reviewer Main Points:** 

[1] Method Process diagram: A conceptual flow diagram is needed to explain the parameter regionalisation procedure more effectively. To visualize the most innovative aspect of the proposed methodology, connection between the first and second classification need more clarification. Figure 4 of the paper titled, "A different light in predicting ungauged basins: regionalization approach based on eastern USA catchments"



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[Shoaib et al., 2013] can be seen as an example.

AB: Authors agree and will add a schematic/flowchart, and modify the appropriate locations in the text to reference the figure. Thanks for the example

[2] My other concern on the paper is that the best MWBM results were not shown relatively as achieved in simulating low –and median flows across the CONUS. Representation of the relative variability of MWBM results in low –median and high flow will enhance the importance of the paper

AB: Analyzing the variability and distribution of the mean monthly model error (Figures 8 and 10) is something we are looking at analyzing and incorporating into further modeling efforts. We felt there are enough interesting results with the monthly model error alone to justify focusing on that aspect of the model for follow-up work and publication, and we decided not to try and push more of that content into this paper.

**Reviewer Minor Points:** 

[3] -p. 10030, line 8: the term FOPV needs more explanation. It is not particularly self-explaining to readers who are not familiar with GSA.

AB: Authors agree to change "FOPV" to output variance. For the sentence on p. 10030 "FAST is a variance-based global sensitivity algorithm that estimates the first-order partial variance (FOPV)..." will be changed to "FAST is a variance-based global sensitivity algorithm that estimates the parameter contribution to output variance...". We will also change the Y-axis labels for Figure 5 to "Output variance".

-p. 10030, lines 14-15: Please clarify why you have avoid incorporating the seasonal adjustment factors in the FAST analysis

AB: We viewed the adjustment factors as more related to the forcing data itself and independent from the model structure.

-p. 10037, lines 24-25: The multi-term objective function is unclear. Inserting an equa-

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tion with the mathematical definition of the objective function would help here. Though NSE, logNSE are mentioned in the manuscript, it is not clear what the authors try to represent the term multi-term objective function. What about SQRT NSE?

AB: The objective function for the final grouped calibration minimizes the sum of difference between the Z-scores of measured and simulated variables for four terms: Mean Monthly Streamflow (As shown in figures 8 and 10a), Monthly Streamflow (Raw monthly time series), Annual Streamflow (Time Series aggregated to annual time steps), Mean Monthly SWE with a 25% error bound. The first three terms of the objective function were chosen because they conveyed information that can be used to easily inform other models (such as daily time-step models).

We will include an equation into the text.

-p. 10058 and 10066, Figure 6 and Figure 14 are not that clear. It is understandable after reading the text, but it could be much improved

AB: For figure 10058 We can add details to the figure caption to indicate these are the final calibration regions representing the merge/intersection of the two classifications discussed in Section 3.2. i.e.: "Final 110 Monthly Water Balance Model calibration regions derived across the CONUS differentiated by color. Streamgages in each calibration group were calibrated in a group-wise fashion to produce a single optimized parameter set for the entire region."

For Figure 14. We will more detail to the figure caption, i.e., "Median Nash Sutcliffe Efficiency (NSE) of all streamgages used for calibration within each calibration region".

-p.10056, Figure 4 could be improved by showing the relative scale of sensitivity. The figure can be more quantifiable, to make the methods more applicable.

AB: We considered this, but we really wanted the symbology in this figure to emphasize the gradation of the parameter sensitivities across the US, along with the "hot spots" for individual parameters. We felt the scaling the sensitivity relative to each parameter

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was the best way to illustrate this.

-A recently published papers on Monthly Runoff Regime Regionalization through Dissimilarity –based Methods[Qamar et al., 2015] and Simultaneous calibration of hydrological models in geographical space [Bárdossy et al., 2015] can be seen as added reference.

AB: Thanks for the additional information. We will read through the recommended works and consider where they might fit in the background and/or discussion.

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# Parameter regionalization of a monthly water balance model for the conterminous United States

3	A.R. Bock <sup>1</sup> , L.E. Hay <sup>2</sup> , G.J. McCabe <sup>2</sup> , S.L. Markstrom <sup>2</sup> , and R.D. Atkinson <sup>3</sup>
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### 25 Abstract

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26	A parameter regionalization scheme to transfer parameter values and model uncertainty	
27	information from gaged to ungaged areas for a monthly water balance model (MWBM) was	_
28	developed and tested for the conterminous United States (CONUS). The Fourier Amplitude	
29	Sensitivity Test, a global-sensitivity algorithm, was implemented on a MWBM to generate	
30	parameter sensitivities on a set of 109,951 hydrologic response units (HRUs) across the CONUS.	
31	The HRUs were grouped into 110 calibration regions based on similar parameter sensitivities.	
32	Subsequently, measured runoff from 1,575 streamgages within the calibration regions were used	
33	to calibrate the MWBM parameters to produce parameter sets for each calibration region.	
34	Measured and simulated runoff at the 1,575 streamgages showed good correspondence for the	
35	majority of the CONUS, with a median computed Nash-Sutcliffe Efficiency coefficient of 0.76	
36	over all streamgages. These methods maximize the use of available runoff information, resulting	
37	in a calibrated CONUS-wide application of the MWBM suitable for providing estimates of water	
38	availability at the HRU resolution for both gaged and ungaged areas of the CONUS.	
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**Comment [Bock1]:** Response to Ref. 1 Minor Points: -p. 10024, lines 1-2 "to transfer ... model uncertainty information". What type of uncertainty information is transferred and how? This is mentioned here and in the conclusion but it is not clearly discussed throughout the paper.

### 48 1 Introduction

The WaterSMART program (http://water.usgs.gov/watercensus/WaterSMART.html) was started 49 by the United States (U.S.) Department of the Interior in February 2010. Under WaterSMART, 50 the National Water Census (NWC) was proposed as one of the U.S. Geological Survey's (USGS) 51 key research directions with a focus on developing new hydrologic tools and assessments. One 52 of the major components of the NWC is to provide estimates of water availability at a sub-53 watershed resolution nationally (http://water.usgs.gov/watercensus/streamflow.html) with the 54 goal of determining: (1) if the Nation has enough freshwater to meet both human and ecological 55 56 needs and (2) if this water will be available to meet future needs. Streamflow measurements do not provide direct observations of water availability at every location of interest; approximately 57 72 percent (%) of land within the conterminous U.S. is gaged, with approximately 13% of these 58 gaged areas being unaffected by anthropogenic effects (Kiang et al., 2013). This creates the 59 challenge of determining the best method to transfer information from gaged catchments to data-60 poor areas where results cannot be calibrated or evaluated with measured streamflow (Vogel, 61 2006). This transfer of model parameter information from gaged to ungaged catchments is 62 known as hydrologic regionalization (Bloschl and Sivapalan, 1995). 63 Many hydrologic regionalization methods have focused on developing measures of similarity 64 between gaged and ungaged catchments using spatial proximity and physical characteristics. 65

66 These methods are highly dependent on the complexity of the terrain and scale at which the

67 relations are derived. Spatial proximity is considered the primary explanatory variable for

hydrologic similarity (Sawicz et al., 2011) because of the first-order effects of climatic and

69 topographic controls on hydrologic response. Close proximity, however, does not always result

in hydrologic similarity (Vandewiele and Elias, 1995; Smakhtin, 2001; Ali et al., 2012).

71 Physical characteristics have been used as exploratory variables to develop a better

<sup>72</sup> understanding of the relation between model parameters that represent model function, and

73 physical properties of the catchment (Merz and Bloschl, 2004). The relation between model

74 parameters and the relevant physical characteristics, expressed for example as a form of

<sup>75</sup> multivariate regression, can be transferred to ungaged catchments (Merz and Bloschl, 2004).

76 Model parameter definitions are by nature ambiguous and often difficult to correlate to a small

77	number of meaningful variables such as physical and climatic characteristics (Zhang et al.,	
78	2008); some studies have found no significant correlation between catchment attributes and	
79	model parameters (Seibert, 1999; Peel et al., 2000), whereas others found that high correlation	
80	does not guarantee parameters that result in reliable model simulations of measured data (Sefton	
81	and Howarth, 1998; Kokkonen et al., 2003; Oudin et al., 2010). Physical and hydrologic	
82	characteristics are also used to derive measures of similarity (or dissimilarity) from multi-	
83	dimensional attribute space, which can be used to identify donor catchments (Qamar et al.,	
84	2015), or classify catchments into discrete regions or clusters (Oudin et al, 2008, 2010; Samuel et	
85	al., 2011). Physical characteristics also are used to classify catchments into discrete regions or	
86	clusters based on similarity in multi dimensional attribute space (Oudin et al, 2008, 2010;	
87	Samuel et al., 2011). While these methods have indicated some success in simulating behavior of	
88	specific hydrologic components, such as base flow (Santhi et al., 2008) or monthly flow regimes	
89	(Qamar et al., 2015), other efforts utilizing discrete clusters performed poorly in explaining	
07		
90	variability of measured streamflow (McManamay et al., 2011).	
90	variability of measured streamflow (McManamay et al., 2011).	
90 91	variability of measured streamflow (McManamay et al., 2011). Two important components of the transfer of parameters to ungaged catchments are the	
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90 91 92 93 94 95 96 97 98 99 100	variability of measured streamflow (McManamay et al., 2011). Two important components of the transfer of parameters to ungaged catchments are the identification of (1) influential (and non-influential) parameters, and (2) geographic extents and scales at which parameters exert control on model function. Reducing the number of parameters is important for calibration efficiency by reducing the structural bias of the model and the uncertainty of results where they cannot be verified or confirmed (Van Griensven et al., 2006). A high number of calibrated, poorly constrained parameters can often mask data or structural errors, which can go undetected and reduce the skill of the model in replicating results outside of calibration conditions (Kirchner, 2006; Bloschl et al., 2013; <u>Bárdossy et al., 2015</u> ). This increases the potential for equifinality of parameter sets and higher model uncertainty that can be propagated to model results (Troch et al., 2003).	

- Sensitivity analysis (SA) has advanced the understanding of parameter influence on model
   behavior and structural uncertainty. SA measures the response of model output to variability in
   model input and/or model parameter values. SA partitions the total variability in the model
   response to each individual model parameter (Reusser et al., 2011) and results in a more-defined
- 105 set of parameters and parameter ranges. Identification of sensitive parameters and their ranges is

**Comment [Bock2]:** Added suggested recent reference from Ref. 2. "-A recently published papers on Monthly Runoff Regime Regionalization through Dissimilarity -based Methods[Qamar et al., 2015] and Simultaneous calibration of hydrological models in geographical space [Bárdossy et al., 2015] can be seen as added reference."

**Comment [Bock3]:** Added this recent reference at suggestion of Ref. 2: "-A recently published papers on Monthly Runoff Regime Regionalization through Dissimilarity -based Methods[Qamar et al., 2015] and Simultaneous calibration of hydrological models in geographical space [Bárdossy et al., 2015] can be seen as added reference."

important for hydrologic model applications as key model parameters can vary spatially across 106 physiographic regions, and also temporally (Tang et al., 2007; Guse et al., 2013). 107 Until recently, the high computational demands of SA have limited most implementations of 108 hydrologic model SA to local sensitivity algorithms that evaluate a single parameter at a time 109 (Tang et al., 2007). These methods ignore parameter interaction, and often assume that model 110 algorithms have linear responses to different parameters (Cuo et al., 2011). Global SA uses 111 random or systematic sampling designs of the entire parameter space to quantify variation in 112 model output (Van Griensven et al. 2006, Reusser et al. 2011). Some of these methods can 113 account for parameter interaction and quantify sensitivity in non-linear systems. Global SA 114 methods are computationally intensive (Cuo et al., 2011), but ever increasing computational 115 efficiency has allowed for the development and application of a large number of global SA 116 117 algorithms. Previous work has suggested that isolating the key parameters that control model performance 118 can be used to infer dominant physical processes in the catchment, as well as which components 119

- 120 of the model dominate hydrologic response (Van Griensven et al. 2006, Tang et al., 2007,
- Reusser et al., 2011). To date, there has been little analysis of the use of SA for deriving
- 122 measures of hydrologic similarity across catchments that can be applied towards hydrologic
- 123 regionalization of model parameters. The spatially-distributed application of SA could be used
- 124 to provide additional information for the delineation of homogeneous regions for parameter
- 125 transfer based on similarity of model results from the SA. This strategy allows for the use of the
- 126 existing model information and configuration to develop a calibration and regionalization
- 127 framework without significantly changing the model structure or implementation
- 128 In this study, we present a hydrologic regionalization methodology for the CONUS that derived
- 129 regions of hydrologic similarity based on the response of a Monthly Water Balance Model
- 130 (MWBM) to parameter SA. Groups of streamgages within each region are calibrated together to
- 131 define a single parameter set for each region. By extending model calibration to a large number
- 132 of sites grouped by similarity through a quantified measure of model behavior, a more specific
- 133 and constrained parameter space that fits each region can be identified.

**Comment [Bock4]:** Response to Ref 1: "these methods ignore parameter interaction, and often assume that model algorithms have linear responses to different parameters". I think this sentence is misleading and I would suggest to delete it."

### 134 2 Methods

### 135 **2.1 Monthly Water Balance Model**

The seasonal adjustment

162

136 The MWBM (Fig. 1) is a modular accounting system that provides monthly estimates of components of the hydrologic cycle by using concepts of water supply and demand (Wolock and 137 McCabe 1999; McCabe and Markstrom, 2007). Monthly temperature (T) is used to compute 138 potential evapotranspiration (PET) and to partition monthly precipitation (P) into rain and snow 139 (Fig. 1). Precipitation that occurs as snow is accumulated in a snow pack (snow storage as snow 140 water equivalent, or SWE); rainfall is used to compute direct runoff (R<sub>direct</sub>) or overland flow, 141 142 actual evapotranspiration (AET), soil-moisture storage recharge, and surplus water, which 143 eventually becomes runoff (R) (Fig. 1). When rainfall for a month is less than PET, AET is equal to the sum of rainfall, snowmelt, and the amount of moisture that can be removed from the soil. 144 145 The fraction of soil-moisture storage that can be removed as AET decreases linearly with decreasing soil-moisture storage; that is, water becomes more difficult to remove from the soil as 146 the soil becomes drier and less moisture is available for AET. When rainfall (and snowmelt) 147 exceeds PET in a given month, AET is equal to PET; water in excess of PET replenishes soil-148 moisture storage. When soil-moisture storage reaches capacity during a given month, the excess 149 150 water becomes surplus and a fraction of the surplus (R<sub>surplus</sub>) becomes R, while the remainder of 151 the surplus is temporarily held in storage. The MWBM has been previously used to examine 152 variability in runoff over the CONUS (Wolock and McCabe, 1999; Hay and McCabe 2002; 153 McCabe and Wolock, 2011a) and the global extent (McCabe and Wolock, 2011b). Table 1 lists the MWBM parameters, with definitions and parameter ranges for calibration. 154 The parameter ranges were determined in previous work (Wolock and McCabe, 1999; Hay and 155 McCabe, 2002). The Drofac parameter specifies the fraction of monthly P that becomes direct 156 runoff. The Rfactor parameter specifies how much surplus in a month becomes runoff. The T<sub>rain</sub> 157 158 parameter specifies the temperature threshold above which all precipitation is rain. The T<sub>snow</sub> parameter specifies the temperature threshold below which all precipitation is snow. The 159 *Meltcoef* parameter specifies the proportion of snowpack that becomes runoff. The *Ppt\_adj* and 160

161 *Tav\_adj* parameters specify seasonal adjustments for precipitation and temperature, respectively.

Comment [Bock5]: Response to Ref. 1: "-p. 10028, line 25 to the end of page: this list of parameter names and meaning does not add much to the information provided in the Table, I would probably avoid it."

163	The <i>Ppt_adj</i> and <i>Tav_adj</i> parameters were included to account for errors in the precipitation and
164	temperature data used in this analysis. Sources of systematic and non-systematic errors of
165	climate forcing data are well documented from the precipitation gage-derived sources (Groisman
166	and Legates, 1994; Adam and Lettenmaier, 2003). Interpolation of these systematic errors from
167	point-scale to gridded domains may propagate these biases, especially in complex terrain (Clark
168	and Slater, 2006; Oyler et al, 2015). The use of adjustment factors allows uncertainty associated
169	with forcing data and model parameter values to be treated separately (Vrught et al., 2008).
170	Figure 1. Conceptual diagram of the Monthly Water Balance Model (McCabe and Markstrom
171	2007). Processes influenced by mModel parameters used in Fourier Amplitude Sensitivity
172	Test (FAST) are identified by green arrow and numbered (Table 1).
173	Table 1. Monthly Water Balance Model parameters and ranges.
174	The MWBM was applied to the CONUS with 109,951 hydrologic response units (HRUs) from
175	the Geospatial Fabric (Viger and Bock, 2014), a national database of hydrologic features for
176	national hydrologic modeling applications (Fig. 2). This HRU derivation is based on an
177	aggregation of the NHDPlus dataset (USEPA and USGS, 2010http://www.horizon-
178	systems.com/nhdplus/), an integrated suite of geospatial data that incorporates features from the
179	National Hydrography Dataset (http://nhd.usgs.gov/), the National Elevation Dataset
180	(http://ned.usgs.gov/), and the Watershed Boundary Dataset (http://nhd.usgs.gov/wbd.html). The
181	sizes of the HRUs range from less than 1 square kilometer (km <sup>2</sup> ) up to 67,991 km <sup>2</sup> , with an
182	average size of 74 km <sup>2</sup> .
183	Inputs to the MWBM by HRU are: (1) monthly P (millimeters), monthly mean T (degrees
184	Celsius), (2) latitude of the site (decimal degrees), (3) soil moisture storage capacity
185	(millimeters), and (4) monthly coefficients for the computation of PET (dimensionless).
186	Monthly P and mean T were derived from the daily time step, 1/8° gridded meteorological data
187	for the period of record from January 1949 through December 2011 (Maurer et al., 2002).
188	Monthly P and T data were aggregated for each HRU using the USGS Geo Data Portal

- 189 (http://cida.usgs.gov/climate/gdp/) (Blodgett et al., 2011). Latitude was computed from the
- 190 centroid of each HRU. Soil moisture storage capacity was calculated using a 1 km<sup>2</sup> grid derived

Comment [Bock6]: Changed based on Ref1 suggstion : "-caption of Fig. 1: "model parameters used in..." Maybe better: "processes influenced by the model parameters used in..."

**Comment [Bock7]:** Wanted to add full citation of NHDPlus base dataset instead of a url

from the Soils Data for the Conterminous United States (STATSGO) (Wolock, 1997). The
monthly PET coefficients were calculated by calibrating the Hamon PET values to Farnsworth et
al. (1982) mean monthly free-water surface evapotranspiration. McCabe et al. (2015) describes
these PET coefficient calculations in detail.

Figure 2. Hydrologic Response Units of the Geospatial Fabric, differentiated by color, overlain
by NHDPlus region boundaries (R01-R18).

### 197 2.2 Fourier Amplitude Sensitivity Test

A parameter SA for the CONUS was conducted for the MWBM using the Fourier Amplitude 198 199 Sensitivity Test (FAST) to identify areas of hydrologic similarity. FAST is a variance-based global sensitivity algorithm that estimates the first order partial variance (FOPV) contribution to 200 201 of model output variance (or objective functions) explained by each parameter (Cukier et al. 1973, 1975; Saltelli et al. 2000). Advantages of using FAST over other SA methods are that 202 203 FAST can calculate sensitivities in non-linear systems, and is extremely computationally efficient, requiring much less information and parameter sets than other global methods. The 204 seasonal adjustment factors were not incorporated into the FAST analysis.; We viewed the 205 seasonal adjustment factors as related more to the forcing data, and for this application only 206 207 parameters associated with model structure were included (first five parameters in Table 1). FAST transforms a model's multi-dimensional parameter space into a single dimension of 208 209 mutually independent sine waves with varying frequencies for each parameter, while using the parameter ranges to define each wave's amplitude (Cuker et al., 1973, 1975; Reusser et al., 2011) 210 (Fig. 3). This methodology creates an ensemble of parameter sets numbering from 1 to N, each 211 of which is unique and non-correlated with the other sets. Parameter sets are derived using the 212 corresponding y-values along each parameter's sine wave given a value on the x-axis. The 213 model is executed for all parameter sets using identical climatic and geographic inputs for each 214 simulation. The resulting series of model outputs are Fourier-transformed to a power spectrum 215 of frequencies for each parameter. Parameter sensitivity is calculated as the sum of the powers 216 offor the output variance for each parameter (FOPV), divided by the sum of the powers of all 217 parameters (Total Variance). FOPV-The parameter sensitivities for all parameters are scaled so 218 that the FOPV sensitivities for all parameters sum to 1. Thus, pParameters that explain a large 219

**Comment [Bock8]:** Renamed FOPV at suggestion of Ref 1.: "-p. 10030, line 8: the term FOPV is not particularly self-explaining to readers not familiar with GSA. I would explain what it is ("contribution to output variance from ...")" and Ref. 2: "[3] -p. 10030, line 8: the term FOPV needs more explanation. It is not particularly self-explaining to readers who are not familiar with GSA."

**Comment [Bock9]:** Deleted as per suggestion from Ref. 1: "-p. 10030, line 12: "much less information and parameter sets". What do you mean by "information"? Unclear. As for "parameter sets, it is possibly less ambiguous if you call them parameter samples or even directly model evaluations"

**Comment [Bock10]:** Clarified our justification for why we did not use seasonal adjustments in fast; Ref 1: "-p. 10030, lines 14-15: -p. 10030, lines 14-15: please justify why you do not incorporate the adjustment factors in the FAST analysis"

Comment [Bock11]: Ref 2: "-p. 10030, lines 14-15: Please clarify why you have avoid incorporating the seasonal adjustment factors in the FAST analysis"

**Comment [Bock12]:** Removed Figure 3 (FAST waves figure) per Ref. 1 suggestion: "-Figure 3: maybe not needed. Anyway, if maintained, vertical axis should show units of measurements. Also, it would probably be better to show Drofac and Rfactor in a separate panel."

220	amount of variability in the model output have higher values of large (i.e. closer to 1) parameter	
221	sensitivityFOPV values.	Comment [Bock13]: More modificaitons related to comment Bock7
222	A portion (0 to 500 parameter sets) of the parameter sampling scheme for the Monthly Water	
223	Balance Model in the Fourier Amplitude Sensitivity Test (FAST). A total of 1000 parameter	
224	sets were generated for implementation in FAST.	
225	FAST was implemented with the MWBM using the 'fast' library in the statistical software R	
226	(Reusser, 2012; R Core Team, 2013). To help constrain the Pparameter ranges used by FAST	
227	for generating wave amplitudes of parameter ensembles across the CONUS, parameter ranges	
228	were based on table 1 information from previous MWBM calibrations at selected streamgages	
229	(Hay and McCabe, 2002). In standard application, Tthe 'fast' R package pre-determines the	<b>Comment [Bock14]:</b> Modified based on
230	minimal number of runs necessary to estimate the sensitivities FOPV for the given number of all	"parameter ranges were based" Are these the
231	parameters (Cukier et al., 1973). For our application we generated an ensemble of 1000	and already commented on p.10028, line 24? If so, just refer to the Table here."
232	parameter sets (as compared to the minimally suggested number of 71 estimated by 'fast') to	
233	have the capability to compare results of different sensitivity analysis methods. The	<b>Comment [Bock15]:</b> Changes at the
234	computational efficiency of the MWBM allowed the parameter sets to be simulated quickly	What do you mean by "standard application"?
235	through parallel processing.	R package uses the equation N = 2Nharm max(!) + 1 to determine the minimal number
		of runs. If so, better cite Cuckier et al (1973), which is where the formula comes from.
236	Many applications of SA in hydrologic modeling have evaluated parameter sensitivity for	Also, please mention what is this number in your case, it would help readers to get an
237	measured streamflow using performance-based measures such as bias, root mean squared error	idea of how computationally demanding is the proposed approach.
238	(RMSE), and the Nash-Sutcliffe Efficiency (NSE) (Nash and Sutcliffe, 1970; Moriasi et al.,	
239	2007). In this study, parameter sensitivity is examined using two hydroclimatic indices that	
240	account for the magnitude and variability of both climatic input and model output: the (1) Runoff	
241	Ratio (RR), a ratio of simulated runoff to precipitation, and (2) Runoff Variability (RV) index,	
242	the standard deviation of simulated runoff to the standard deviation of precipitation	
243	(Sankarasubramanian and Vogel, 2003).	
244	<b>3</b> Parameter regionalization procedure	
245	The MWBM parameter sensitivities from the FAST analysis using an ensemble of 1000 MWBM	
246	parameter sets were evaluated across the CONUS. The spatial patterns and magnitudes of	
247	parameter sensitivities then were used to organize the 109,951 HRUs across the CONUS into	

hydrologically similar regions for parameter regionalization through MWBM calibration. 248

Potential streamgages were identified for use in two automated calibration procedures. The 249

calibration procedures were used to produce an 'optimal' set of MWBM parameters for each 250

calibration region. The following sections describe the parameter regionalization procedure in 251

detail (Fig. 3).-252

Figure 3. Schematic flowchart of the parameter sensivity analysis and regionalization method 253 described in this paper (Section 3). 254

#### **3.1 Parameter sensitivities** 255

The relative sensitivities derived from the FAST analysis using the RR and RV indices at each of 256 the 109,951 HRUs across the CONUS were scaled so that the five MWBM parameter 257 258 sensitivities derived for each HRU summed to 100 (Fig. 4). RR (Fig. 4a) is most sensitive to the parameter *Drofac* in regions where MWBM runoff is not dominated by snowmelt and orographic 259 260 precipitation, such as arid and sub-tropical areas of the CONUS. MWBM parameters that control snowpack accumulation and melt (Meltcoef, Tsnow, and Train) are more important to the 261 RR in the extensive mountain ranges in the Western CONUS, and northerly latitudes around the 262 263 Great Lakes and in the Eastern CONUS. The RR indicates the highest sensitivity to the *Rfactor* parameter in mountainous areas of the CONUS and areas of the West Coast, and moderate to 264 high sensitivity in areas where the sensitivity of RR to *Drofac* is low. The patterns of *Tsnow*, 265 Train, and Meltcoef all share similar patterns across the CONUS. The spatial variability of the 266 sensitivity of RR to Meltcoef indicates different physical mechanisms controlling Metlcoef 267 parameter influence on RR in different areas of the CONUS. In the Western CONUS, the 268 sensitivity of RR to Meltcoef is greatest in mountainous areas that accumulate and hold 269 snowpack through the late spring, such as the Rocky Mountains, Cascade, and Sierra Nevada 270 mountain ranges. In the Eastern and Midwestern CONUS, the sensitivity of RR to Meltcoef is 271 greatest for HRUs with more northerly latitudes. 272 Figure 4. Relative sensitivity of the (a) Rainfall Ratio (RR) and (b) Runoff Variability (RV) 273

274

Comment [Bock16]: Added flowchart schematic at request of both Refs. Ref1: -[1] The parameter regionalisation procedure could be explained more effectively. In the first place, it would be good to have a schematic of the procedure to clearly see what is the role, inputs and outputs of each step (sensitivity analysis, classification of regions, individual calibration, grouped calibration, etc.)."

Comment [Bock17]: Ref 2: [1] Method Process diagram: A conceptual flow diagram is needed to explain the parameter regionalisation procedure more effectively. To visualize the most innovative aspect of the proposed methodology connection between the first and second classification need more clarification. Figure 4 of the paper titled, "A different light in predicting ungauged basins: regionalization approach based on eastern USA catchments" [Shoaib et al., 2013] can be seen as an example.

Comment [Bock18]: Changed based on redundancy notice by Ref1 : "-p. 10032, lines 12-13: "The patterns of ...". Sentence needs rewording"

indices to Monthly Water Balance Model parameters.

The spatial patterns of sensitivities of RV to the five MWBM parameters (Fig. 4b) show both 275 similarities and deviations from the patterns shown in the RR maps. For the central part of the 276 CONUS, the relative sensitivity for the parameter *Drofac* is high for both indices, and low for the 277 parameter Rfactor for both indices. Meltcoef, Tsnow, and Train share the same relations between 278 higher sensitivity and higher elevation (primarily in the western part of the CONUS), and higher 279 sensitivity and more northerly latitude (primarily in the eastern half of the CONUS) for both 280 281 indices. However, *Drofac* and *Rfactor* show distinctly different patterns of relative sensitivities 282 for the eastern part of the CONUS for RV as compared to RR. The other three parameters follow the same general spatial patterns for RV as compared to RR, but with greater fine-scale 283 284 spatial variation and patchiness. The differences between the spatial distributions of the 285 sensitivities between the two indices highlight that applying SA to different model outputs can generate different levels of sensitivities for each parameter. In addition, the choice of objective 286 287 function or model output for which to measure parameter sensitivity is important, as parameter sensitivities will differ depending on whether a user evaluating measures of magnitude, the 288 289 variability of distribution, or timing (Krause et al., 2005; Kapangaziwiri et al, 2012). 290 Figure 5 illustrates the variability of parameter sensitivities between NHDPlus regions 08 (Lower Mississippi) and 14 (Upper Colorado) (see Fig. 2) for the RV and RR indices. The Lower 291 Mississippi and Upper Colorado NHDPlus regions have a similar number of HRUs (4,449 and 292 3,879, respectively) and cover a similar area (26,285 and 29,357 km<sup>2</sup>, respectively). The Lower 293 Mississippi region has homogenous topography, with humid, subtropical climate, while the 294 Upper Colorado region has highly variable topography, and thus highly variable climatic 295 controls on hydrologic processes. For the Lower Mississippi region only one parameter 296 dominates modeled RV variance (Rfactor, Fig. 5a) and modeled RR variance (Drofac, Fig. 5c). 297 In contrast, for the Upper Colorado River region several parameters influence RV variability 298 (Drofac, Rfactor and Meltcoef, Fig. 5b) and RR variability (Drofac and Meltcoef, Fig. 5d). In 299 the Lower Mississippi Region, the amount of snowfall is negligible, so the three parameters that 300 control snowfall and snowpack accumulation in the MWBM have negligible effect on simulated 301 total runoff. The comparison of the parameter sensitivities for these two regions illustrates how 302 303 variable parameter sensitivities are for different regions (i.e. different climatic and physiographic 304 regions)

**Comment [Bock19]:** Added in response to Ref 1. Comment: "-[2] Some of the numerical results are a bit surprising and should be double-checked. In particular, in Fig. 5.a the fact that one parameter has sensitivity of exactly 100 and all others of exactly 0 seems odd."

305	First order partial variance (FOPV)Figure 5. Parameter sensitivities of Runoff Variability (RV;	Comment
306	a-b) and Runoff Ratio (RR; c-d) indices for Monthly Water Balance Model parameters in the	comment/res
307	Lower Mississippi (R08) and Upper Colorado (R14).	
308	3.2. Calibration regions	
309	The spatial patterns and magnitudes of parameter sensitivities across the CONUS were used as a	
310	basis for organizing HRUs into hydrologically similar regions for parameter regionalization	
311	through MWBM calibration. While <u>T</u> this idea is rooted in the hypothesis that geographically	
312	proximate HRUs share similar forcings and conditions, and thus will behave similarly.	
313	application the uses similarity inof SA results as a basis for organization, rather than similarity in	
314	physiographic characteristics provides a quantification of similarity based on similar model	
315	responses to a wide ensemble of model conditions. The derived regions are subsequently used to	Comment [
316	simplify model calibration across the CONUS and provide a basis for the transfer and application	this idea". the previous
317	of parameters to ungaged areas.	sentence? B the one illust clarify."
318	The parameter sensitivities derived using from the RR were used to organize the HRUs into two	
319	independently-derived calibration regions; the first derived by identifying HRUs with unique	
320	combinations of the order of parameter sensitivities to the RR (highest parameter sensitivities to	
321	lowest, i.e. 1-Drofac (78%), 2-Rfactor (16%), 3-Meltcoef (4%), 4-Tsnow (1%), 5-Train (1%)),	
322	and the second classification based upon identifying HRUs with unique sets of parameters whose	
323	sensitivities exceeded a specified threshold of parameter sensitivity (i.e. only Drofac and Rfactor	
324	using a 5% threshold in the first classification example). Using the parameter sensitivities for	
325	each HRU, two different classifications of HRUs were derived. The purpose of the first	Comment [ substantially
326	classification was to delineate regions of similar model response or behavior based on the order	Comment [
327	of importance of the MWBM parameters to the RR for each HRU. Thise first classification	connection to classification
328	identified 16 distinct regions of HRUS across the CONUS based on the order of the FOPV for	from each of the actual cl
329	theparameter sensitivities of the five parameters (derived using the RR index). Sizes of these	-Description
330	regions ranged from 94 km <sup>2</sup> to almost 2 million km <sup>2</sup> . The second classification delineated	"unique com sensitivities"
331	regions with an identical set of the most important parameters to the RR based on parameters	their meaning?
332	whose sensitivities exceeded a 5% threshold. This step e second classification identified 12	Be more s approaches
333	regions of HRUs with unique combinations of parameter sensitivities with FOPV exceeding	that the para in each regio

Comment [Bock20]: Changed FOPV to parameter sensitivities based on comment/response Bock7

omment [Bock21]: Modified this sentence n response to ref. 1: "-p. 10033, line 25: "While his idea...". What idea? The one described in he previous entence? But then is it really in contrast with he one illustrated on lines 27-28? Please larify."

Comment [Bock22]: This section is substantially modified based on the following suggestions/comments by the referees: Comment [Bock23]: Ref1: -What is the

nection between the first and second
ssification? They are independent
n each other and then intersected to obtain
actual classification? Please
ify
scription of the second classification (lines
11) is also unclear. What are the
ique combinations of parameter
sitivities"? How are they defined? What is
ir
aning?
e more specific on how the two classification
proaches work. Sentence on lines
of page 10034 is too generic, does it mean
t the parameter ranking is the same
ach region?

334	5%. There has been progress in providing quantitative thresholds for the identification of	
335	sensitive and non-sensitive parameters for hydrologic modelers (Pianosi et al., 2015 Tang et al.,	
336	2007), but no definitive consensus yet exists. Therefore a 5% threshold was used based on visual	_
337	delineation of major physiographic features such as mountain ranges across the CONUS. The	
338	sizes of this second group of regions ranged from 94 $\text{km}^2$ to more than 15 million $\text{km}^2$ . Maps of	
339	the two groupings of HRUS were intersected to create a total of 49 regions across the CONUS.	
340	NHDPlus region and sub-region-boundaries, proximity, and significant topographic divides were	
341	used to further divide the groups into 159 geographically unique calibration regions across the	
342	CONUS. The lack of streamgages available in some regions, especially areas with arid and	
343	semi-arid climates, necessitated merging regions together. Calibration regions that contained	_
344	less than 3 streamgages from the 8,410 gages present in the Geospatial Fabric (see section 3.3)	
345	were combined with the proximate and most similar group based on the which shared the most	
346	similar parameter sensitivities (both order and magnitude), of parameter sensitivities resulting in	
347	110 calibration regions across the CONUS (Fig. 6). <u>Additionally, W</u> within each region the	
348	FAST results for both the RR and RV indices were used to determine which parameters to	
349	calibrate. Parameters with a median parameter sensitivity of 5% for the RR and RV among the	
350	region's HRUs were selected for group calibration. Parameters not shown as sensitive were kept	
351	at the default value for the group.	
352	Figure 6. Final 110 Monthly Water Balance Model calibration regions differentiated by colors.	
353	A subset of streamgages within each calibration region were calibrated in a group-wise	
354	fashion to produce a single optimized parameter set for the entire region (Fig. 3). Monthly	_
355	Water Balance Model calibra	
356	tion regions differentiated by color.	
357	3.3 Initial streamgage selection	
358	The initial set of streamgages used for testing in the MWBM calibration procedures was selected	
359	from 8,410 streamgages identified in the Geospatial Fabric (Fig. 7). The Geospatial Fabric	

360 includes reference and non-reference streamgages from the Geospatial Attributes of Gages for

361 Evaluating Streamflow dataset (GAGES-II, Falcone et al., 2010). Of the 8,410 streamgages in

**Comment [Bock24]:** Replaced this reference at the recommendation of ref. 1: -p. 10034, lines 11-13: citation of Pianosi et al., 2015 does not seem to be appropriate here. That paper introduces a toolbox for Sensitivity Analysis but it does not discuss the issue of setting the threshold for sensitive and non-sensitive parameters. The threshold issue is (partially) discussed in Tang et al. (2007). "

**Comment [Bock25]:** Some clairification added as per Ref 1. Comment: "-From lines 17-20, I understand that the sensitivity-based classification is further refined using a more 'conventional' approach that looks at proximity and topographic divides. How does this refinement step works? Does it introduce significant changes in the classification? This is important to know in order to understand the value of the proposed sensitivity-based classification versus proximity or topography-based classification."

**Comment [Bock26]:** Added clarification to caption based on ref. 2 comments: "-p. 10058 and 10066, Figure 6 and Figure 14 are not that clear. It is understandable after reading the text, but it could be much improved"

362	the Geospatial Fabric, 1,864 were identified as having reference-quality data with at least 20
363	years of record. These reference quality streamgages were judged to be largely free of human
364	alterations to flow (Falcone et al., 2010). In the current study, reference quality was not
365	considered in the initial streamgage selection because the 20 years of record was considered too
366	restrictive. Therefore a subset of the 8,410 streamgages was selected for initial testing in the
367	MWBM calibration procedures based on the following criteria:
368	(1) Remove streamgages with less than 10 years of total measured streamflow (120 months)
369	within the time period $1950 - 2010$ .
370	(2) Remove streamgages with a drainage area defined by the Geospatial Fabric that are not
371	within 5% of the USGS National Water Information System (NWIS) reported drainage
372	area (U.S. Geological Survey, 2014). This eliminated many of the streamgages with
373	smaller drainage areas due to the resolution of the Geospatial Fabric.
374	(3) Remove streamgages that did not have at least 75% of its drainage area contained within
375	a single calibration region.
376	These criteria resulted in 5,457 potential streamgages for testing in the MWBM calibration
377	procedures (Fig. 7). Streamflow at these streamgages was aggregated and converted from daily
378	(cubic feet/second) to a monthly runoff depth (mm) (streamflow per unit area).
379	Figure 7. Streamgages tested in the study. GF notes geospatial fabric for national hydrologic
380	modeling (Viger and Bock, 2014).
381	3.4 Monthly Water Balance Model calibration
382	Two automated calibration procedures were implemented to produce an 'optimal' set of MWBM
383	parameters for each calibration region. The first procedure, Individual Streamgage Calibration,
384	calibrated each of the 5,457 streamgages individually. Results from the individual calibrations
385	were used to further filter the streamgages within the second procedure, Grouped Streamgage

386 Calibration, which calibrated selected streamgages together by calibration region.

### 387 3.4.1 Individual streamgage calibration

388	The first calibration procedure was an automated process that individually calibrated each of the
389	5,457 streamgages from the initial streamgage selection with measured streamflow (U.S.
390	Geological Survey, 2014). Results from these individual streamgage calibrations quantified the
391	'best' performance of the MWBM at each gage, providing a 'baseline' measure for evaluation.
392	The Shuffled Complex Evolution (SCE) global-search optimization algorithm (Duan et al., 1993)
393	has been frequently used as an optimization algorithm in hydrologic studies (Hay et al., 2006;
394	Blasone et al. 2007; Arnold et al., 2012), including previous studies with the MWBM (Hay and
395	McCabe, 2010). Further details can be found in Duan et al. (1993). SCE was used to maximize a
396	combined objective function based on: (1) Nash-Sutcliffe Efficiency (NSE) coefficient using
397	measured and simulated monthly runoff and (2) NSE using natural log-transformed measured
398	and simulated runoff (logNSE), using the entire period of record for each streamgage. The NSE
399	measures the predictive power of the MWBM in matching the magnitude and variability of the
400	measured and simulated runoff (Nash and Sutcliffe, 1970). The NSE coefficient ranges from $-\infty$
401	to 1, with 1 indicating a perfect fit, and values less than 0 indicating that measured mean runoff
402	is a better predictor than model simulations. The NSE has been shown to give more weight to
403	the larger values in a time series (peak flows) at the expense of lower values (low flows)
404	(Legates and McCabe, 1999), so the logNSE was incorporated into the objective function to give
405	weight to lowflow periods (Tekleab et al., 2011).

### 406 **3.4.2 Grouped streamgage calibration**

- 407 The second calibration procedure was an automated process that calibrated groups of
- 408 streamgages together for each calibration region to derive a single set of MWBM parameters
- 409 (Table 1) for each calibration region (Fig. 6). The NSE and logNSE values from the individual
- 410 streamgage calibrations (described in the previous section) were used to identify streamgages
- 411 that should not be used for grouped streamgage calibration. If the individual streamgage
- 412 calibration was not 'satisfactory', then it was felt that it would not provide useful information for
- 413 the grouped streamgage calibration procedure.
- 414 Satisfactory individual streamgage calibrations were identified with the following procedure:

415 (	(1)	Eliminate all streamgages with NSE values $< 0.3$	
<b>TIJ</b> (	1/	$\Box$	

415	(1) Eliminate all streamgages with INSE values < 0.3.		
416	(2) If the number of remaining streamgages for a given calibration region is $> 10$ , then		
417	eliminate all streamgages with NSE $< 0.5$ .		
410	(2) If the number of states $\alpha$ are a given calibration ratio is $25$ then aliminate all		
418	(5) If the number of streamgages for a given canoration region is $> 25$ , then emininate an streamgages with NSElog $< 0$		
419	sucanigages with NSElog < 0.		
420	(4) If the number of remaining streamgages for a calibration region is $< 5$ , check to see if any		
421	of the eliminated streamgages were reference streamgages (as defined in Falcone et al., 2010),		
422	then add the reference streamgages back in if the NSE value $> 0.0$ . Reference streamgages are		
423	USGS streamgages deemed to be largely free of anthropogenic impacts and flow modifications		
424	(Falcone et al., 2010; Kiang et al., 2013).		
425	These criteria, while somewhat arbitrary, were chosen so that no calibration region had less than		
426	5 streamgages for the grouped streamgage calibration. Using the above criterion, of the 5,457		
427	streamgages individually calibrated, 3,125 remained as candidates for the grouped streamgage		
428	calibration procedure.		
429	The grouped streamgage calibration procedure used the SCE global-search optimization		
430	algorithm with a multi-term objective function (Eq. 1). Measured and simulated values for		
431	selected streamgages contained within a calibration region were scaled toby Z-scores to remove		
432	differences in magnitudes between streamgages (Eq. 2). The multi-term objective function		
433	minimized the sum of the absolute differences between Z-scores from four measured and		
434	simulated time series: <u>-{mean monthly runoff (MMO,MMS),</u> monthly runoff (MO,MS), mean		
435	monthly runoff, annual runoff (AO,AS) (U.S. Geological Survey, 2014), and monthly snow		
436	water equivalent (SO,SSSWE)) for all selected streamgages within a given calibration region:		
437	- The observed and simulated Z scores (Z) were calculated at each streamgage as:		
438	$min\sum_{i=1}^{n}[3 MMO_{i} - MMS_{i}  +  MO_{i} - MS_{i}  +  AO_{i} - AS_{i}  + 0.5 SO_{i} - SS_{i} ] $ (Eq.1)		
439			

**Comment [Bock27]:** Added details on reference gages at request of ref. 1: "-p. 10037, line 17: Please give a very brief definition of a reference streamgage."

	where $\begin{cases}  SO_i - SS_i  if SS_i < SO_i^{0.75} \\  SO_i - SS_i  SS_i > SO_i^{1.25} \end{cases}$			
440	The measured and simulated Z-scores were calculated as:		<b>Comment [Bock28]:</b> Added equation fo the	
441	$Z = (x-u)/\sigma$ (Eq. 24)		Ref 1: "-p. 10038, lines 1-4: The definition of the multi-term objective function is unclear. Are the four terms summed up? Why considering heth mean monthly rungif and partial.	
442 443	where x is the time-series value, u is the mean, and $\sigma$ the standard deviation of the measured and simulated variablestreamflow.		runoff (I would imagine that they convey the same information, the former being equal to the latter divided by 12)? I think inserting an equation with the mathematical definition	
444	'Measured' SWE was determined for each HRU from the Snow Data Assimilation System	$\left( \right)$	of the objective function would help here." <b>Comment [Bock29]:</b> Ref 2: -p. 10037, lines 24-25: The multi-term objective function is unclear. Inserting an equa flow with the	
445	(SNODAS; National Operational Hydrologic Remote Sensing Center, 2004) and included a +/-		mathematical definition of the objective function would help here. Though	
446 447	25% error bound. The unconstrained automated calibration (without a restriction on SWE) led to unrealistic sources of snowmelt in the summer that enhanced the low-flow simulations. The 25%		NSE, logNSE are mentioned in the manuscript, it is not clear what the authors try to represent the term multi-term objective function. What about SQRT NSE?	
448	error bound is arbitrary; calibrating to the actual SNODAS SWE values was found to be too		<b>Comment [Bock30]:</b> Both reviewers wanted	
449	restrictive, but adding this error bound to the SWE values resulted in better overall runoff		Comment [Bock31]: Changed at the	
450 451	simulations. The absolute difference of the simulated SWE Z-scores within +/- 25% of the measured SWE Z-score were designated as 0. Otherwise, the absolute difference was computed		suggestion of ref. 1: "-p. 10038, line 9: "simulated streamflow" should be "simulated variable" (since one of the four is SWE and not runoff)"	
452	between the simulated SWE Z-score and either the upper or lower bounds (Eq. 1).		Comment [Bock32]: Clarified SWE Z-scores	
453	The grouped calibration procedure was run for all 110 calibration regions. For each calibration		At request or ref.1: -p. 10038, lines 14-16: Please clarify how the error bounds were taken into account. Did you modify the definition of the Z score for the SWE?"	
455	analysis section 3 1) were calibrated; parameters deemed not sensitive (narameter sensitivity			
456	< 5% of total variance) were set to their default values (see Table 1). The entire period of the		Comment [Bock33]: Added clarifying	
457	streamflow record for each streamgage was split by alternating years. After calibration, mean		statement at the recommendation of ref. 1: "Recall here that a parameter is deemed	
458	monthly measured and simulated Z-scores for runoff at all selected streamgages within a		insensitive if sensitivity index is below 5%"	
459	calibration region were compared on a mean monthly basis.		Comment [Bock34]: redundant, deleted at	
460	Figure 8 shows an example of the graphic used to evaluate the measured and simulated mean	recommendation of ref. 1: "p. 10038, lines 21- 23: "on a mean monthly based". Unclear. Possibly it might just be dropped, since it was already said that monthly variables are used to compute the		
461	monthly Z-scores for 21 streamgages selected for the region located in the Tennessee River		Zscores."	
462	calibration region (part of NHDPlus Region R06 in Fig. 2); the orange, red, and black dots			
463	indicate calibration, evaluation, and the entire period of record, respectively. A tight grouping			
464	around the one-to-one line indicates good correspondence between measured and simulated Z- 17			

 $(0 \text{ if } 0.75 < SO_i - SS_i < 1.25)$ 

465 scores. Points closer to the upper right corner of each plot represent high-flow periods. Points 466 closer to the lower left corner of the plot represent low-flow periods. Streamgages within a 467 calibration region were assigned the same parameter values; therefore streamgages that plotted 468 outside (two standard deviations) of the one-to-one line were considered to not be representative 469 of the calibration region, and the calibration procedure for that calibration region was repeated 470 without those streamgages.

Figure 8. Measured versus simulated mean monthly Z-scores for the Tennessee River calibration
region (see Fig. 10b for location). Orange is calibration, red is evaluation, and black is all
years.

The goal of the second calibration procedure was to find a single parameter set for each 474 calibration region. Past applications of the MWBM (Wolock and McCabe, 1999, McCabe and 475 Wolock, 2011a) used a single set of fixed MWBM parameters for the entire CONUS. Many of 476 the streamgages included in the second calibration procedure could be affected by significant 477 anthropogenic effects; the seasonal adjustment factors, calibrated at each individual streamgage, 478 could account for these effects and result in satisfactory NSE values. Streamgages that were 479 480 removed due to poor performance in the second calibration were assumed to have anthropogenic 481 effects not consistent with the streamgages that plotted along the one-to-one line. Poor 482 performance may result because the MWBM fails to reliably simulate runoff for a watershed 483 because of model limitations (i.e. not including all important hydrologic processes), but the calibration regions are assumed to be homogeneous based on the FAST analysis. Therefore it is 484 assumed that if some of the streamgages within a region have satisfactory results, then the 485 MWBM is able to simulate runoff in that region. 486

### 487 **4 MWBM calibration region results**

### 488 **4.1 Individual streamgage calibration results**

- 489 The individual streamgage calibrations provided information regarding: (1) the potential
- 490 suitability of a given streamgage for inclusion in a grouped calibration, and (2) a 'baseline'
- 491 measure for evaluation of the grouped calibration results. Reference and non-reference
- 492 streamgages were considered in this application; if the runoff at a streamgage could not be

493	calibrated individually to a 'satisfactory' level (based on criterion outlined in section 3.4.2), then
494	it was felt that it would not provide useful information for the grouped streamgage calibration
495	procedure. Figure 9 shows the NSE (Fig. 9a) and logNSE (Fig. 9b) coefficients from the
496	individual streamgage calibrations for the CONUS. Scattered throughout the CONUS are NSE
497	and logNSE values less than 0.0 (triangles in Fig. 9). These poor results are likely streamgages
498	with poor streamflow records, either due to measurement error or anthropogenic effects (dams,
499	water use, etc.).

# Figure 9. Individual streamgage calibration results: (a) Nash-Sutcliffe Efficiency (NSE) coefficient and (b) log of the NSE (logNSE).

### 502 4.2 Grouped streamgage calibration results

### 503 4.2.1 Mean monthly z-scores

Figure 10a shows a scatterplot of measured versus simulated mean monthly Z-scores for runoff, 504 similar to Figure 8, but based on all available years (the black dots in Fig. 8) for all the final 505 calibration streamgages (1,575 streamgages). Four regions are highlighted to illustrate the 506 507 monthly variability in MWBM results across the CONUS (see Fig. 10b for locations). The four regions are: New England (67 streamgages, red); Tennessee River basin (21 streamgages, 508 509 orange); Platte Headwaters (15 streamgages, blue); and Pacific Northwest (33 streamgages, green) (Fig. 10b). 510 Figure 10. (a) Measured versus simulated mean monthly Z-scores for runoff at all streamgages 511 and (b) location of highlighted streamgages for four calibration regions: New England (67 512 streamgages, red); Tennessee River (21 streamgages, orange); Platte Headwaters (15 513 streamgages, blue); and Pacific Northwest (33 streamgages, green). 514 In Fig. 10a, three of the regions (New England, Tennessee River, and Pacific Northwest), show 515

simulated Z-scores that correspond favorably to measured Z-scores for each of the twelve

517 months, including periods of low and high runoff. These regions represent marine or humid

climates with homogenous physio-climatic conditions and an even spatial distribution of

519 streamgages, where models should be expected to perform well (see Fig. 9) There is a higher

520 variability in model results for the high-flow months (May - June) for streamgages within the

- 521 Platte Headwaters (Fig. 10a; blue dots) than for low-flow months. This variability may be
- related to factors controlling the magnitude and timing of snow melt runoff (Fig. 9).

For each calibration streamgage, a set of four months were identified that represent different
parts of the measured mean monthly hydrograph (highest- and lowest- flow month and the two

- median-flow months). The measured and simulated mean monthly streamflow Z scores
- 526 corresponding to the four months are plotted as cumulative frequencies (Fig. 11) to compare how

well the simulated Z scores matched measured Z scores for different parts of the hydrograph

528 over the entire set of calibration gages. For the highest-flow, there is an under-estimation of

runoff, with the greatest divergence between the two distributions in the middle to lower half of

the distribution (Fig. 11a). For the median-flow, the measured and simulated Z scores are well

- 531 matched. For the 10 lowest-flow, simulated Z scores are greater than measured Z scores, with the 532 greatest divergence between the two distributions in the middle to upper half of the distribution
- 533 (Fig. 11c).
- Figure 11. Z-score cumulative frequency for (a) highest-, (b) median-, and (c) lowest-flow
  months.

The median Z-score errors (simulated - measured) by region for the (a) highest-, (b) median-, and (c) lowest-flows are shown in Figure 12. The largest errors are for the highest-flows (Fig. 12a). The MWBM simulations under-estimate the highest flows for much of the CONUS. The errors for median-flows are fairly uniform and consistent across the CONUS (Fig. 12b), with a median error close to 0. For the lowest-flow months the MWBM over-estimates low flows for a large portion of the Midwest (Fig. 12c).

Figure 12. Z-score error (simulated - measured) for (a) highest-, (b) median-, and (c) lowestflow months.

### 544 4.2.2 Nash-Sutcliffe efficiency

- 545 Figure 13 compares the NSE from the individual streamgage calibrations (gageNSE) with the
- 546 grouped calibrations (groupNSE) for all final streamgages used in the second calibration
- 547 procedure. NSE values > 0.75 (dashed line) and > 0.5 (solid line) indicate very good and

548	satisfactory results (Moriasi et al., 2007). Overall, most NSE values fall above the 0.5 NSE		
549	threshold of satisfactory performance (median of gageNSE and groupNSE = $0.76$ ). The gageNSE		
550	values are used here as a 'baseline' for evaluation of the groupNSE results. The groupNSE		
551	values were not expected to be greater than the gageNSE values since (1) NSE was not used as		
552	an objective function in the grouped calibration, and (2) grouped calibrations found the 'best'		
553	parameter set for a set of streamgages versus an individual streamgage. Figure 13 shows an equal		
554	distribution of NSE values around the one-to-one line, indicating that the grouped calibration		
555	provided additional information over the individual streamgage calibrations (cases where		
556	groupNSE are greater than gageNSE in Fig. 13). The difference between the gageNSE and		
557	groupNSE becomes larger as the NSE values decrease, reflecting the increasing uncertainty in		
558	the grouped calibrations in areas with lower gageNSE values.		
559	Figure 13 Nash Sutcliffe Efficiency from individual (gageNSE) and grouped (groupNSE)		
560	calibration Calibration regions in New England (67 streamgages red): Tennessee River (21		
500	stream second or no soly. Diotta Handwatars (15 stream second hive), and Davifia Northwest (22		
501	streamgages, orange); Platte Headwaters (15 streamgages, blue); and Pacific Northwest (33		
562	streamgages, green) are highlighted (see Fig. 10b for location).		
563	Four regions are highlighted in Fig. 13 to illustrate the variability of NSE across the CONUS		
564	(see Fig. 10b for locations). The highlighted regions in New England (red), Tennessee River		
565	(orange), and Pacific Northwest (green), show good groupNSE and gageNSE results. Four of		
566	the 15 streamgages in the Platte Headwaters (blue) have groupNSE values $\leq 0.5$ . This is		
567	probably related to simulation error during the snowmelt period (May - June, Fig. 10a).		
5.60	Figure 14 shows the modion group NSE by solibustion assign for the CONUS. The nottern is your		
568	Figure 14 shows the median group NSE by canoration region for the CONOS. The pattern is very		
569	similar to that shown for the individual streamgage calibration results in Fig. 9a and highlights		
570	the problem areas shown in Fig. 12.		
571	Figure 14. Median Nash Sutcliffe Efficiency (NSE) by calibration region. Median Nash Sutcliffe		
572	Efficiency (NSE) by calibration region of streamgages used for calibration.		
573			

**Comment [Bock35]:** Added clarification to caption based on ref. 2 comments: "-p. 10058 and 10066, Figure 6 and Figure 14 are not that clear. It is understandable after reading the text, but it could be much improved"

### 574 **5 Discussion**

575 This study presented a parameter regionalization procedure for calibration of the MWBM,

- resulting in an application that can be used for simulation of hydrologic variables for both gaged
- 577 and ungaged areas in the CONUS. The regionalization procedure grouped HRUs on the basis of
- 578 similar sensitivity to five model parameters. Parameter values and model uncertainty
- information within a group was then passed from gaged to ungaged areas within that group.

### 580 5.1 Regionalized parameters

Results from this study indicate that regionalized parameters can be used to produce satisfactory 581 MWBM simulations in most parts of the CONUS (Fig. 13). Despite the differences between the 582 individual streamgage calibration and grouped calibration, Figure 13 illustrates that the grouped 583 calibration strategy, which focused only on only sensitive parameters, can provide just as much 584 information as the individual streamgage calibration with no constraints on the parameters 585 optimization other than the default ranges. The MWBM is a simple hydrologic model as it has 586 minimal parameters, which are conceptual in nature (not physically based). It may be that this 587 type of model is best for regionalization when parameter sensitivity can be identified and HRU 588 behavior can be classified by a small number of clearly defined spatial groups. More 589 complicated models with many more interactive parameters may not respond as well to this 590 simple type of regionalization; more parameters may lead to more parameter interaction and 591 situations of equifinality which might confuse the analysis. 592 The adjustments of precipitation and temperature parameters for the individual streamgage 593

- calibrations accounted for local errors such as rain\_gage under catch of precipitation. In addition
- these climate adjustments also account for local anthropogenic effects on streamflow (e.g. dams,
- 596 diversions) since streamgages were not screened for these effects prior to individual streamgage
- 597 calibration. In the grouped streamgage calibrations, the same precipitation and temperature
- <sup>598</sup> adjustments are applied at every streamgage within the calibration region, making these climate
- adjustments more of a regional adjustment and producing more of a 'reference' condition for
- 600 each calibration region.

**Comment [Bock36]:** Clairifcation to ref. 1 comment: "-Also, the result of Figure 13 is very puzzling. As the authors note on page 10041 (line 27 onward), the groupNSE values are expected

to be lower than the gageNSE values. Figure 13 instead shows many cases were groupNSE is much larger than gageNSE. I really struggle to believe that NSE can be increased so much and so often when using a model calibrated with a different objective function."

### **5.2 Parameter sensitivities and dominant process**

The MWBM parameter sensitivities varied by hydroclimatic index (RR and RV) and across the 602 603 CONUS (Fig. 4). The parameter sensitivity patterns give an indication of dominant hydrologic processes based on MWBM. The dominant process can be seasonal and MWBM performance 604 may be enhanced by extending the use of SA along the temporal domain to identify and 605 606 temporally vary the parameters that are seasonally important to the MWBM. For example, error in peak flow months is the primary cause for poor model performance in the Platte Headwaters 607 (Fig. 10). For the Platte Headwaters, the final parameter set performed well for simulated Z-608 scores for the regionalized low- and median-flow conditions (Fig. 10a, July through April), but 609 was not able to replicate measured mean monthly flows for May and June. In this case, the 610 dominant processes controlling hydrologic behavior change with season and the parameters 611 controlling the dominant response may have to change accordingly (Gupta et al., 2008; Reusser 612

613 et al., 2011).

### 614 5.3 Model accuracy

615 The pattern of MWBM accuracies shown in Fig. 9 and 14 are similar to those shown by Newman

et al. (2015; Fig. 5a) in which a daily time-step hydrologic model was calibrated for 671 basins

across the CONUS. Our study and the Newman et al. (2015) study both indicate the same

<sup>618</sup> 'problem areas' with the poorest performing basins generally being located in the high plains and

619 desert southwest. Newman et al. (2015) attributed variation in model performance by region to

spatial variations in aridity and precipitation intermittency, contribution of snowmelt, and runoffseasonality.

622 The inferior MWBM results in the 'problem areas' can be attributed to multiple factors which

623 likely include inadequate hydrologic process representation and errors in forcing data (e.g.

624 climate data), and/or measured streamflow. Archfield et al. (2015) state that the performance of

625 continental-domain hydrologic models is considerably constrained by inadequate model

representation of dominant hydrologic processes. For example, the simplicity of the MWBM

627 presents limitations on the representation of deeper groundwater reservoirs, gaining and losing

stream reaches, simplistic AET, and the effects of surface processes (infiltration and overlandflow) that need to be represented at finer time steps than monthly.

The dominant hydrologic processes in the 'problem areas' appear to be poorly represented at the daily (Newman et al., 2015) and monthly time steps. This may be due to inadequate forcing <u>datas</u>, the quality of which 'is paramount in hydrologic modeling efforts' (Archfield et al., 2015) and/or the lack of 'good' reference streamflow data for calibration and evaluation. Both surely play a role and emphasize the need for incorporation of additional datasets so that calibration and evaluation of intermediate states in the hydrologic cycle are examined.

### 636 6 Conclusions

A parameter regionalization procedure was developed for the CONUS that transferred parameter 637 values and model uncertainty information from gaged to ungaged areas for a MWBM. The 638 FAST global-sensitivity algorithm was implemented on a MWBM to generate parameter 639 sensitivities on a set of 109,951 HRUs across the CONUS. The parameter sensitivities were 640 used to group the HRUs into 110 calibration regions. Streamgages within each calibration region 641 were used to calibrate the MWBM parameters to produce a regionalized set of parameters for 642 each calibration region. The regionalized MWBM parameter sets were used to simulate monthly 643 runoff for the entire CONUS. Results from this study indicate that regionalized parameters can 644 be used to produce satisfactory MWBM simulations in most parts of the CONUS. 645 The best MWBM results were achieved simulating low- and median-flows across the CONUS. 646 The high-flow months generally showed lower skill levels than the low- and median-flow 647 months, especially for regions with dominant seasonal cycles. The lowest MWBM skill levels 648

- 649 were found in the high plains and desert southwest and can be attributed to multiple factors
- 650 which likely include inadequate hydrologic process representation and errors in forcing data
- 651 and/or measured streamflow. Calibration and evaluation of intermediary fluxes and states in the
- 652 MWBM through additional measured datasets may help to improve MWBM representations of
- these model states by helping to constrain parameterization to measured values.

**Comment [Bock37]:** Response to Ref. 1 Minor Points: -p. 10024, lines 1-2 "to transfer .. model uncertainty information". What type of uncertainty information is transferred and how? This is mentioned here and in the conclusion but it is not clearly discussed throughout the paper.

# 654 **7** Acknowledgments

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**Comment [Bock39]:** Removed reference based on modifications to text (comment Bock 23)

**Comment [Bock40]:** Added reference at recommendation of ref. 2.

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Parameter	Definition	Range	Default
1. Drofac	Controls fraction of precipitation that becomes runoff	0, 0.10	0.05
2. Rfactor	Controls fraction of surplus that becomes runoff	0.10, 1.0	0.5
3. Tsnow	Threshold above which all precipitation is rain $(^{\circ}C)$	-10.0, -2.0	-4.0
4. Train	Threshold below which all precipitation is snow $(^{\circ}C)$	0.0, 10.0	7.0
5. Meltcoef	Proportion of snowpack that becomes runoff	0.0, 1.0	0.47
6. Ppt_adj	Seasonal adjustment factor for precipitation (%)	0.5, 2.0	1
7. Tav_adj	Seasonal adjustment for temperature (°C)	-3.0,_3.0	0

Table 1. Monthly Water Balance Model parameters and ranges.



Figure 1. Conceptual diagram of the Monthly Water Balance Model (McCabe and Markstrom

2007). <u>Processes influenced by m</u> odel parameters used in Fourier Amplitude Sensitivity Test
 (FAST) are identified by green arrow and numbered (Table 1).

**Comment [Bock42]:** Changed based on Ref1 suggstion : "-caption of Fig. 1: "model parameters used in..." Maybe better: "processes influenced by the model parameters used in..."





**Comment [Bock43]:** Remove Fig. 3 at suggestion of Ref. 1: "-Figure 3: maybe not needed. Anyway, if maintained, vertical axis should show units of measurements. Also, it would probably be better to show Drofac and Rfactor in a separate panel."

Figure 3. A portion (0 to 500 parameter sets) of the parameter sampling scheme for the Monthly Water Balance Model in the Fourier Amplitude Sensitivity Test (FAST). A total of 1000 parameter sets were generated for implementation in FAST.

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**Comment [Bock45]:** Old picture with FOPV as y-axis labels



**Comment [Bock46]:** New figure with Param. Sens. As y-axis labels







### Figure 7. Streamgages tested in the study. GF notes geospatial fabric for national hydrologic modeling (Viger and Bock, 2014).

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region (see Fig. 10b for location). Orange is calibration, red is evaluation, and black is all years.

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Figure 9. Individual streamgage calibration results: (a) Nash-Sutcliffe Efficiency (NSE)
coefficient and (b) log of the NSE (logNSE).



Figure 10. (a) Measured versus simulated mean monthly Z-scores for runoff at all streamgages
and (b) location of highlighted streamgages for four calibration regions: New England (67
streamgages, red); Tennessee River (21 streamgages, orange); Platte Headwaters (15
streamgages, blue); and Pacific Northwest (33 streamgages, green).





months.



Figure 12. Z-score error (simulated - measured) for (a) highest-, (b) median-, and (c) lowestflow months.



Figure 13. Nash Sutcliffe Efficiency from individual (gageNSE) and grouped (groupNSE)
calibration. Calibration regions in New England (67 streamgages, red); Tennessee River (21
streamgages, orange); Platte Headwaters (15 streamgages, blue); and Pacific Northwest (33
streamgages, green) are highlighted (see Fig. 10b for location).

