

1 **Variability of natural low flow magnitudes in the Upper Colorado River Basin: Identifying**
2 **monotonic and periodic trends, and relative role of large-scale climate dynamics**

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20 **Abstract**

21 Low flow magnitude in a head water basin is important for planners because minimum available
22 amount of water in a given time period often leads to serious repercussions, in both up and the
23 downstream regions. This concern is common in the arid territory like Colorado River basin
24 located in the southwestern United States. Low flow variability in Colorado River is due to
25 complex interactions between several natural and anthropogenic factors but here we aim to
26 identify trends and systematic variability of low flows, and the relative role of climate at
27 different spatial locations of the basin. The research questions we aim to answer are: *How*
28 *variable are the low flow conditions in the headwater basin of Colorado River? Did location-*
29 *specific low flow change in the past years? How are low flows linked with synoptic ocean-*
30 *atmospheric conditions?* Towards that aim we select 17 stream gauge locations, which are
31 identified as “undisturbed” stream gauges meaning that these stations represent near-natural river
32 flow regimes in the headwater area of Colorado River providing a useful resource for assessment
33 of climate and hydrology associations without the confounding factor of major direct (e.g. water
34 abstraction) or indirect (e.g. land-use change) human modification of flows. A detailed
35 diagnostic analysis gave us fair understanding on the variability and changes in low flow
36 magnitude that is explained by climate. *Most notably, eastern and western sides of Upper*
37 *Colorado River Basin (UCRB) indicated opposite trending patterns of low flows, the west (east)*
38 *showing drier (wetter) conditions, and the low flow magnitudes were specifically found to be*
39 *having multi-decadal variability revealing the close associations with Interdecadal Pacific*
40 *Oscillation or Pacific Decadal Oscillation (PDO) patterns.*

41 **Key words:** low flow, variability, Colorado River

42

43 1. Introduction

44 Variability and change in stream flow can directly influence water supply (both quantity and
45 quality) for domestic, [agricultural](#), industrial, ecological, and other needs. Palmer et al. (2008)
46 indicated that river discharge in every inhabited river basin in the world would face changes;
47 some will have large increases while others will likely face the water scarcity. Understanding
48 variability of the volume of stream flow is important because very high flows can cause
49 damaging floods and erosion, while very low flows can fail to provide adequate water supply,
50 diminish water quality, and affect important ecological services (Smakhtin, 2001). Existing
51 evidence suggest that water scarcity due to low river flow could be one of the main drivers of
52 societal and cross-boundary conflicts (Gleick and Palaniappan, 2010; Gleick, 2010, 2014). Thus,
53 anticipating the magnitudes of seasonal and annual minimum flow in the headwater locations of
54 a river is important for up and downstream water management purposes. Intricate connections
55 between human and natural processes influence the water supply from the basin headwater and
56 as such minimum river flow is a result of complex interactions between human and biophysical
57 features and thus, differing from one region to another (Jones et al., 2012). Hence, characterizing
58 lower tail of river flow distribution demands more attention than it has received so far.

59 Water resources in the southwestern United States, are especially scarce and climatic changes
60 may cause significant alterations in water availability, quality, and demand. The hydrology of the
61 southwest is already characterized by strong variability on seasonal to multiannual time scales,
62 reflecting its sensitivity to fluctuations in large-scale atmospheric circulation patterns from the
63 Pacific Ocean, the Gulf of California, and the Gulf of Mexico (Seager et al., 2007). Amongst
64 major river basins, Colorado River is the critical source of water for 7 states in the arid
65 southwestern United States (especially for high [aggregated demand met in the municipal](#),

66 | agricultural, and industrial sectors), and that, this river has a history of going under low flow
67 | conditions (i.e. flow going under a minimum threshold condition) in the past (USGS, 2004;
68 | Meko et al., 2007; Ellis et al., 2010; Gleick, 2010). In addition, population growth, agricultural,
69 | urban, and industrial expansions within the past decades enhanced this effect. It has been
70 | reported in the scientific literature that the Colorado River flow is expected to reduce further
71 | under future warming scenarios due to a combination of strong temperature-induced runoff
72 | curtailment, reduced annual precipitation, and increased (potential) evapo-transpiration (Milly et
73 | al., 2005; Christensen and Lettenmeier, 2007; Seager et al., 2007), and consequently the seasonal
74 | distribution of flow will also change due to changing ratio of snow to total precipitation as well
75 | as changing timing of the snow melt (Fritze et al., 2011). Another risk to Colorado River stream
76 | flow is multi-decadal droughts, which is also expected to change under climate change (IPCC
77 | AR5, 2013). Therefore, impacts of drought conditions on the river flows, especially in the driest
78 | time, are also expected to change. But, little is known regarding how the low flow characteristics
79 | changed over time in response to changes in the climate.

80 | There is complexity and heterogeneity of low flow dynamics in a river basin. Therefore, it is
81 | difficult to generalize characterization of low flow. Low flow, defined in many different ways
82 | (section II in SI), could be a sole or combination of multiple factors in different seasons. Such
83 | factors may include slowly flowing ground water discharge, surface discharge from lakes,
84 | marshes, snow-pack dynamics, melting glaciers, basin precipitation, basin temperature and
85 | evaporation rates, basin soil, topography, geology and vegetation, river channel characteristics,
86 | and various man-induced effects (Smakhtin, 2001). For instance, in the summer time (July
87 | through October), low flows of most part of the United States, are usually derived by base flow
88 | (Reilly and Kroll, 2003; Flynn, 2003). On the other hand, in cold or mountainous regions, low

89 | flows are subject to the spatial influences of ice, snow or glacier melting in addition to the usual
90 | basin parameters (Smakhtin, 2001; Reilly and Kroll, 2003; Miller and Piechota, 2011; Curran et
91 | al., 2012; EPA, 2012). Therefore, we hypothesize that climate is linked with low flow variability
92 | and those links differ with locations because of variable physiographic parameters.

93 | This hypothesis leads to the science questions: *How variable are the low flow conditions in*
94 | *the headwater basin of Colorado River? Did location-specific low flow change in the past years?*
95 | *How are low flows linked with synoptic ocean-atmospheric conditions?*

96 | Through the Colorado River Compact, the Upper Colorado River Basin (UCRB) supplies
97 | water and hydropower for much of the southwestern United States and hence low flow dynamics
98 | of UCRB has large influence on both the up and the downstream water supply. These scientific
99 | questions will enable us to understand the statistical characteristics of regional low flow
100 | variability in this important river basin as well as capture their physical connections to large-
101 | scale ocean-atmospheric systems. Our research findings will support scientists and engineers to
102 | develop prediction tools that assist in climate informed and timely water management decisions
103 | during potential crises, as well as maintaining the minimum flow conditions in the river to
104 | sustain ecosystem services.

105 | The paper proceeds as follows. Section 2 describes the datasets used, section 3 summarizes
106 | calculation of low flow statistics, section 3 explains results and related discussions, and section 4
107 | summarizes the findings.

108 | **2. Data**

109 | We selected 17 “undisturbed” stream gauges in UCRB, which primarily contribute to the
110 | largest amount of total Colorado River stream flow (McCabe et al., 2007; Gao et al., 2011).

111 Consideration of the undisturbed stations minimizes the human induced effects on the natural
112 flow and captures natural variability and changes. We downloaded the daily river flow data from
113 USGS Hydro-Climatic Data Network 2009 (Lins, 2012). The detailed description of the data,
114 including the homogeneity/quality, is found in the supplementary information (SI) (section I).
115 Table 1-A lists the stations' information and Figure 1 shows the geographic locations of those
116 stations as well as the length of the data ranging from 25 to 61 years. As evident in Figure 1, the
117 Upper Colorado River Basin (UCRB) differs in topographical features, notably the eastern
118 stream gauges are located in the higher elevation areas than the western locations, also indicated
119 in Table 1.

120 To study large-scale climatological patterns, we used global Surface Temperature (ST) and
121 Mean Sea Level Pressure (MSLP) data from 1949-2011 from NCEP/NCAR reanalysis V1.0
122 monthly diagnostic products (Kalnay et al., 1996). We downloaded these datasets in ready to
123 analyze format from the International Research Institute for Climate and Society Data Library
124 (<http://iridl.ldeo.columbia.edu>).

125

126 **3. Calculation of low flow statistics**

127 To calculate the low flow statistics, we considered climate years that extends from April 1–
128 March 31, as suggested by previous research (Ries and Friesz, 2000; Flynn, 2003; Reilly and
129 Kroll, 2003; Pyrce, 2004; Risley et al., 2008; Martin and Arihood, 2010; Curran et al., 2012;
130 EPA, 2012). Daily mean flows for all complete climatic years of record are used to determine
131 low-flow statistics for all 17 stream-gaging stations. Low-flows in streams can be characterized
132 in many ways but in the United States, the 7-day low flow—annual or seasonal series of the

133 smallest values of mean discharge over any 7-consecutive days (q7)—is a common method for
134 determining the low flow magnitude (Ries and Friesz, 2000; Smakhtin, 2001; Flynn, 2003;
135 Pyrcce, 2004; Reilly and Kroll, 2003; Risley et al., 2008; Martin and Arihood, 2010; Curran et al.,
136 2012; EPA, 2012). We followed this approach in this research. Annual q7 generally occurs in the
137 driest season, mainly in the beginning of spring and/or summer for UCRB. But, different stream
138 gauges experience q7 in different months (Table S1). Because the summer time low flow
139 conditions is generally driven by the base flow in most part of the United States, and in winter,
140 that subjects to the influences of ice, snow or glacier melting, it is also crucial to study the
141 variations of low flows in different seasons in addition to annual low flows. In this research we
142 also considered four traditional seasons, namely Dec-Jan-Feb (DJF), Mar-Apr-May (MAM),
143 Jun-Jul-Aug (JJA), and Sep-Oct-Nov (SON).

144

145 **4. Results and Discussions**

146 This section presents an in depth descriptive analysis of low flow statistics at UCRB
147 locations (4.1). That helps to detect the seasonal and annual variability patterns and trends of q7,
148 thus answering the first two research questions. Following that we report the results of a
149 correlational investigation that identify relationships between low flow statistics and large-scale
150 climatic patterns, answering the third research question (4.2).

151 **4.1. Low flow variability and trends**

152 Figure 2 shows variation of annual q7 magnitude—annual smallest values of mean discharge
153 over any 7-consecutive days, and those for the traditional seasons are shown in Figure S1. Figure
154 2 (and Figure S1) indicates that annual (and seasonal) low flow magnitudes within UCRB have
155 high spatio-temporal variability. This is expected because the river basin characteristics,

156 especially the topographical features play a major role for q7 variation. Particularly, the
157 variability differs in the east and west sides of the basin (separated by -107.5° long) where the
158 elevations differ, as also indicated in Figure S2. In addition, spatial variability of q7 is quantified
159 in Table 1-B for annual and Tables S2-S5 for the traditional seasons.

160 Of the major assumptions in a correlational study is normal distribution idea. Since q7 falls at
161 an extreme tail of the daily mean flow distribution, non-normal behavior can be expected.
162 Therefore, to detect the non-normal behavior in q7 time series for different stream gauge
163 locations, we estimated skewness and kurtosis values and reported them in the summary statistics
164 tables. As those values indicate, annual and seasonal q7 distributions are normal in general,
165 because, as a rule of thumb, they have an absolute skewness value less than 3 and an absolute
166 kurtosis value less than 10.

167 Cross-correlation analysis: To determine co-variability, we conducted a cross correlation
168 analysis between q7 time series for different stream gauge stations. Tables 2 and 3 list the annual
169 and JJA cross-correlation analysis results and Tables S6-S8 list the other results. These tables
170 indicate that q7 magnitudes in most stations are positively correlated with each other; which is
171 most prominent and statistically significant for the cases in the summer (JJA), as in Table 3, and
172 followed by SON (Table S8). This finding indicates that variability of low flow at multiple
173 locations might be linked with common external factors, for JJA in particular. Generally, stations
174 close to each other yield highest positive significant correlations.

175 Monotonic and periodic trends assessment: Impacts of global change and large-scale natural
176 climate variability is felt locally. Therefore, it is imperative to look at whether there have been
177 any significant monotonic trends and sub/multi decadal patterns in seasonal and annual low flow
178 statistics and how they compare amongst locations. Periodicity is the indicator for small-scale

179 hydrological system response to large-scale circulation patterns, such as El-Nino Southern
180 Oscillation (ENSO) or Pacific Decadal Oscillation (PDO). We considered non-parametric Mann-
181 Kendall trend tests to detect monotonic trends and wavelet analysis to identify periodicity or
182 multi-decadal patterns. The Mann–Kendall test is applicable to the detection of a monotonic
183 trend in a time series with no seasonal or other cycle. Mann (1945) formulated the non-
184 parametric test for monotonic trend detection, and Kendall (1975) derived the test statistic
185 distribution for testing non-linear trend and turning point. This method allows us to ignore high-
186 frequency (i.e. multiple change point) variations. Since there are chances of outliers in the low
187 flow data, non-parametric Mann–Kendall test is useful because its statistic is based on the sign of
188 differences, not directly on the values of the random variable, and therefore, the trends
189 determined are less affected by the outliers. On the other hand, Wavelet analysis has been widely
190 used to analyze time series data with localized variations of power. This method decomposes a
191 one dimensional time series (or frequency spectrum) into two dimensional time-frequency spaces
192 in order to analyze signals of the data containing non-stationary power at many different
193 frequencies and creates the time-scaled output signal (Torrence and Compo, 1998). We used an
194 interactive web toolkit developed by C. Torrence and G. Compo that uses Mortlet wavelet basis
195 function, incorporates the edge effects due to finite-length time series in a cone of influence, and
196 includes a statistical significance testing using specific theoretical wavelet spectra for both white
197 noise and red-noise processes (<http://ion.researchsystems.com>).

198 Figure 3 depicts the monotonic trend results for different time intervals (indicated in Table 1-
199 A) and Figure S3 indicates the same but for those stations which have more than 30-years of data
200 permitting more statistical power, both of which reveal identical behavior where the trends in
201 low-flow magnitudes exhibit variable nature in different seasons and annually. Stations, which

202 are located close to each other, are generally showing homogeneity in significant trends. More
203 specifically, we notice a clear distinction in monotonic trends in the eastern and the western sides
204 of -107.5° longitude. Annual, DJF, MAM, and SON q7 trends are negative on the west side and
205 positive on the east side. JJA q7 trends are usually negative everywhere, except some non-
206 significant positive trends. Negative trending patterns in the western part of UCRB, are
207 consistent with some of the previous studies, which indicated a general trend of low flow states
208 toward permanently drier conditions in the southwestern US due to a projected decrease in runoff
209 and soil moisture in the headwaters of Colorado river arising from a projected increase in
210 (potential) evap-transpiration (USGS, 2004; McCabe et al., 2007; Gleick, 2010; Seager et al.,
211 2007, 2013). However, positive trend patterns on the eastern part of UCRB, which has higher
212 elevation, does not follow the idea that “dry will get drier and wet will get wetter”. This
213 monotonic trend study indicates the importance of locally based studies needing further
214 investigations to detect the causes for the differences in trends other than diverse physiographic
215 characters of the basin (Figure 1).

216 Hydro-climatic analysis have also indicated that there is considerable non-stationarity in
217 measured and reconstructed stream flow estimates for the Colorado River basin, which may be
218 linked with inter-decadal, decadal, multi-decadal and even secular variations in ocean
219 temperatures (Cook et al., 2004; Gray et al., 2004; Hidalgo, 2004; McCabe et al., 2004). To
220 detect periodic trends in low flow magnitude, wavelet analysis was performed. The
221 decomposition of time series into time-frequency space permits the identification of the
222 dominant modes of variability and determining how these modes vary in time. A few examples
223 of periodicity of q7 magnitudes are shown in Figure 4, all of which indicate close to 10-16 years
224 periods. Rest of the test results for other stations is presented in the Tables S9-S10, which also

225 confirms recurrent 10-16 years periodicity of q7 data. Though the exact cause of these multi-
226 decadal variations is not fully understood yet, which will require longer datasets, we hypothesize
227 that this dominant periodicity might be closely related to Interdecadal Pacific Oscillation (IPO)
228 or Pacific Decadal Oscillation (PDO) patterns of northern Pacific (Zhang et al., 1997; Folland et
229 al., 2002; Dai, 2013), as also discussed in the following section.

230 **4.2. Linkage with Large-Scale Climate Patterns**

231 This section summarizes Pearson correlation patterns between low flow magnitudes (q7) and
232 large-scale climate variables. We've considered global surface temperature (ST) anomalies and
233 mean sea level pressure (MSLP) anomalies to determine variability of q7 dictated by climate. To
234 do that, first, we conducted the Principal Component Analysis (PCA) to determine the
235 orthogonal time series that are orthogonal to each other and explaining the common variance of
236 q7 across the stations within UCRB, both for traditional seasons and annually. This analysis
237 requires the data having equal length, thus, we considered only those stations having more than
238 30-years length. Thus, 14 stations having data from 1976-2011 were considered for PCA.

239 Figure S4 indicates the variance explained by the Principal Components (PCs) of annual and
240 seasonal q7. PC1 explains around 40-60% of the variance in the UCRB data, and then the
241 explained variance drops gradually by the other PCs. Figure 5 (Figure S5) shows the correlation
242 between annual (seasonal) PC1 and the global average climate in northern summer (Apr-May-
243 Jun-Jul-Aug-Sept/AMJJAS) and northern winter (Oct-Nov-Dec-Jan-Feb-Mar/ONDJFM). We
244 considered two distinct northern hemisphere seasons to check which timing of the year indicates
245 recognized ocean-atmospheric signals. Figure 5 indicates that PC1 of annual q7 has distinct
246 associations with the summer season climate, most notably, a positively correlated ENSO-like
247 surface temperature pattern extending from the coast of China to the Central Northern Pacific

248 surrounded by a negatively correlated horseshoe type pattern in the summer and a negatively
249 correlated ENSO-like surface temperature region in the tropical Pacific is visible (also indicated
250 by the positively correlated MSLP in the same region and season). DJF, MAM, and SON seasons
251 also yielded similar results as annual with little different patterns for JJA PC1. These prominent
252 sea surface temperature patterns in the northern Pacific (especially pole ward of 20° N) indicate
253 Pacific Decadal Oscillation (PDO) type behavior. The influence of distinct phases of PDO on
254 dominant mode of UCRB low flows is consistent with McCabe et al. (2004) who indicated the
255 effect of this multi-decadal ocean-atmospheric patterns on the drought frequency across the US,
256 especially over the Southwestern United States. A direct correlation analysis yielded statistically
257 significant associations between the northern hemispheric summer season PDO indices and the
258 annual (-0.40), DJF (-0.41), MAM (-0.48), JJA (-0.31) and SON (-0.40) time series of PC1
259 respectively. PDO shifts phases on at least inter-decadal time scale, usually about 20 to 30 years.
260 The Interdecadal Pacific Oscillation (IPO), on the other hand, displays similar sea-surface
261 temperature and sea-level pressure (SLP) patterns with the PDO but with a cycle of about 15–30
262 years. This, along with the wavelet analysis results above, indicate that IPO might be associated
263 closely with the prominent variability of the annual low flows and associated droughts; i.e. at the
264 positive (negative) phase of PDO/IPO the magnitude of q7 would be decreased (increased),
265 indicating a greater (lesser) frequency or magnitude of droughts, which supports McCabe et al.
266 (2004) and Dai (2013)'s findings.

267 Many studies also indicated persistent La Nina-like cold SST anomalies in the tropical
268 central and eastern Pacific Ocean leading to below-normal precipitation and often droughts over
269 Southwestern North America (e.g., Seager et al., 2005; Mo et al., 2009; Wang et al., 2010).
270 Because El-Nino Southern Oscillation (ENSO) associates with the PDO, an ENSO type pattern

271 is evidently allied with the low flow variability in Figure 5 and S5. Therefore, this study re-
272 establishes the connections between the variability of annual and seasonal droughts, but now via
273 low flow magnitudes, and the northern hemispheric summer time ocean-atmospheric patterns
274 such as ENSO/PDO/IPO.

275 **5. Summary**

276 Low-flow statistics for the streams is important for water supply planning and design, waste-load
277 allocation, reservoir storage design, and maintenance of quantity and quality of water for
278 irrigation, recreation, and wildlife conservation. Colorado River is the lifeline for many states in
279 the arid southwestern US. Water availability in the headwater basin matters a great deal for these
280 states. In this study we aim to understand the variability and changes in low flow conditions
281 during different seasons and annually as well as to detect what role the large-scale ocean-
282 atmospheric features play to modulate them. Since low flow is due to a complex mixture of
283 many local physiographic factors and climatic mechanisms, it has been hard historically to
284 generalize the low flow for the entire basin. However, this study indicates that significant
285 monotonic and periodic trends are existent for annual and seasonal low flow magnitudes but
286 differing in the eastern and the western parts of the basin due to variant topographical conditions
287 (east having higher elevation). Furthermore, the first Principal Component of annual and
288 seasonal low flow magnitudes (q7) across the Upper Colorado River Basin (UCRB) indicates
289 more than 40-60% variability and shows clear connections with the Pacific ocean patterns
290 (PDO/IPO) in northern hemispheric summer season, yielding 10-16 years periodicity. This
291 indicates a greater possibility of statistical predictions of low flow magnitudes using climate
292 indices, which forms our next step of research. A skillful and timely prediction of location
293 specific low flow statistics is important and necessary for environmental, industrial and

294 agricultural sectors, which has an aim to keep up with the water demand for human and
295 ecological systems during the time of water scarcity. This scientific research takes a step forward
296 to contribute to that reason.

297

298 **Author contribution:** I.P. designed the experiments and M.P.P carried out the analysis. I.P. and
299 M.P.P prepared the manuscript.

300

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304

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Table 1-A: Description of HCDN-2009 streamflow gaging stations in UCRB (Lins, 2012).

Station ID	Station Name	Drainage area in Sq-km	LAT_GAGE	LONG_GAGE	Altitude above NGVD29 (m)	Data available for years	Data length (years)
9066000	Black Gore Creek Near Minturn, CO. (east)	32.409	39.596	-106.265	2788.9	1965-2011	47
9034900	Bobrail Creek Near Jones Pass, CO. (east)	15.649	39.760	-105.906	3179.1	1967-2011	36
9066200	Booth Creek Near Minturn, CO. (east)	16.097	39.648	-106.323	2537.5	1966-2011	46
9306242	Corral Gulch Near Rangely, CO. (west)	81.986	39.920	-108.473	2005.6	1976-2011	36
9081600	Crystal River Ab Avalanche C, Near Redstone, CO. (east)	432.893	39.232	-107.227	2104.6	1957-2011	55
9035800	Darling Creek Near Leal, CO. (east)	22.944	39.801	-106.026	2724.9	1967-2009	43
9065500	Gore Creek At Upper Station, Near Minturn, CO. (east)	37.776	39.626	-106.278	2621.3	1965-2011	47
9047700	Keystone Gulch Near Dillon, CO. (east)	23.570	39.594	-105.973	2849.9	1959-2011	53
9066300	Middle Creek Near Minturn, CO. (east)	15.522	39.646	-106.382	2499.4	1966-2011	46
9035900	South Fork Of Williams Fork Near Leal, CO. (east)	72.842	39.796	-106.031	2728.0	1967-2011	45
9107000	Taylor River At Taylor Park, CO. (east)	331.619	38.860	-106.567	2846.8	1989-2011	23
9352900	Vallecito Creek Near Bayfield, CO. (west)	188.151	37.478	-107.544	2409.8	1964-2011	66
9183500	Mill Creek at Sheley Tunnel, Near Moab, UT. (west)	74.302	38.483	-109.404	1676.4	1989-2011	23
9210500	Fontenelle C Nr Herschler Ranch, Nr Fontenelle, WY. (west)	398.309	42.096	-110.417	2118.4	1953-2011	59
9223000	Hams Fork Below Pole Creek, Near Frontier, WY. (west)	333.153	42.110	-110.710	2272.3	1954-2011	58
9312600	White River Bl Tabbyvune C Near Soldier Summit, UT. (west)	195.295	39.876	-111.037	2203.7	1969-2011	43
9378170	South Creek Above Reservoir Near Monticello, UT. (west)	21.898	37.847	-109.370	2185.4	1987-2011	25

Note: CO = Colorado; UT = Utah; WY = Wyoming; LAT_GAGE = Latitude of a streamgauge; LONG_GAGE = Longitude of a streamgauge; NGVD29 = National Geodetic Vertical Datum of 1929

Table 1-B: Summary statistics table for annual q7 time series for each stream gauge station in UCRB.

Station ID	Station Name	Mean (cms)	Median (cms)	Standard deviation (cms)	Skewness	Kurtosis
9066000	Black Gore Creek Near Minturn, CO. (east)	0.67	0.65	0.12	0.22	-0.83
9034900	Bobtail Creek Near Jones Pass, CO. (east)	1.73	1.76	0.31	0.08	-0.73
9066200	Booth Creek Near Minturn, CO. (east)	6.73	6.95	1.52	-0.43	-0.11
9306242	Corral Gulch Near Rangely, CO. (west)	1.78	1.79	0.33	0.19	0.04
9081600	Crystal River Ab Avalanche C, Near Redstone, CO. (east)	2.53	2.31	1.01	2.91	9.08
9035800	Darling Creek Near Leal, CO. (east)	2.09	1.89	0.98	2.90	9.19
9065500	Gore Creek At Upper Station, Near Minturn, CO. (east)	0.76	0.71	0.37	1.84	4.30
9047700	Keystone Gulch Near Dillon, CO. (east)	0.23	0.20	0.18	1.70	4.36
9066300	Middle Creek Near Minturn, CO. (east)	42.70	41.93	8.86	1.03	1.76
9035900	South Fork Of Williams Fork Near Leal, CO. (east)	29.15	28.93	3.90	-0.34	-0.58
9107000	Taylor River At Taylor Park, CO. (east)	4.76	4.61	0.94	-0.25	0.25
9352900	Vallecio Creek Near Bayfield, CO. (west)	17.85	18.00	5.35	0.37	0.32
9183500	Mill Creek at Sheley Tunnel, Near Moab, UT. (west)	10.24	10.86	3.94	-0.26	-0.22
9210500	Fontenelle C Nr Herschler Ranch, Nr Fontenelle, WY. (west)	0.43	0.35	0.32	1.05	0.45
9223000	Hams Fork Below Pole Creek, Near Frontier, WY. (west)	2.55	2.47	1.58	0.29	-0.51
9312600	White River Bl Tabbyune C Near Soldier Summit, UT.	17.06	16.43	5.39	1.20	3.73
9378170	South Creek Above Reservoir Near Monticello, UT. (west)	0.05	0.04	0.06	1.24	0.86

Note: CO = Colorado; UT = Utah; WY = Wyoming

Table 2: Pearson correlation coefficients between annual q7 time series of different stations (1989-2011). 90% statistically significant estimates are shown.

Station ID	9034900	9066200	9306242	9081600	9035800	9065500	9047700	9066300	9035900	9107000	9352900	9183500	9210500	9223000	9312600
9034900															
9066200															
9306242															
9081600															
9035800															
9065500		0.41													
9047700			0.39												
9066300		0.62		0.45			0.38								
9035900		-0.35		0.39											
9107000	0.41						0.50	0.51							
9352900		0.39		0.68				0.37							
9183500			0.56												
9210500			0.40	0.40			0.41		0.44						
9223000			0.38	0.37			0.44						0.66		
9312600				0.41			0.62	0.36	0.37			0.45	0.55	0.64	
9378170			0.64									0.75			
9066000				0.50			0.43				0.41		0.72	0.46	0.43

Table 3: Pearson correlation coefficients between JJA q7 time series of different stations (1989-2011). 90% statistically significant estimates are shown.

Station ID	9034900	9066200	9306242	9081600	9035800	9065500	9047700	9066300	9035900	9107000	9352900	9183500	9210500	9223000	9312600	9378170
9034900	0.73															
9066200	0.81	0.78														
9306242	0.64	0.47	0.47													
9081600	0.74	0.48	0.78	0.58												
9035800	0.61	0.61	0.70	0.58	0.76											
9065500	0.87	0.71	0.86	0.61	0.89	0.78										
9047700	0.78	0.73	0.78	0.41	0.59	0.56	0.76									
9066300	0.83	0.77	0.85	0.48	0.61	0.65	0.74	0.72								
9035900	0.87	0.74	0.80	0.54	0.75	0.62	0.83	0.80	0.83							
9107000	0.82	0.48	0.78	0.61	0.92	0.62	0.90	0.67	0.65	0.77						
9352900	0.68	0.57	0.60	0.55	0.62	0.57	0.69	0.53	0.57	0.47	0.69					
9183500	0.73	0.61	0.77	0.63	0.80	0.77	0.79	0.53	0.60	0.60	0.71	0.67				
9210500	0.63	0.44	0.47	0.58	0.37		0.45	0.58	0.49	0.61	0.49	0.42	0.50			
9223000	0.68	0.49	0.50	0.76	0.54	0.56	0.59	0.60	0.55	0.65	0.56	0.54	0.60	0.87		
9312600	0.70	0.44	0.71	0.72	0.81	0.67	0.72	0.62	0.68	0.74	0.80	0.52	0.72	0.68	0.74	
9378170	0.72	0.70	0.73	0.48	0.69	0.62	0.78	0.60	0.49	0.58	0.62	0.60	0.86	0.43	0.49	0.52

(A)



(B)

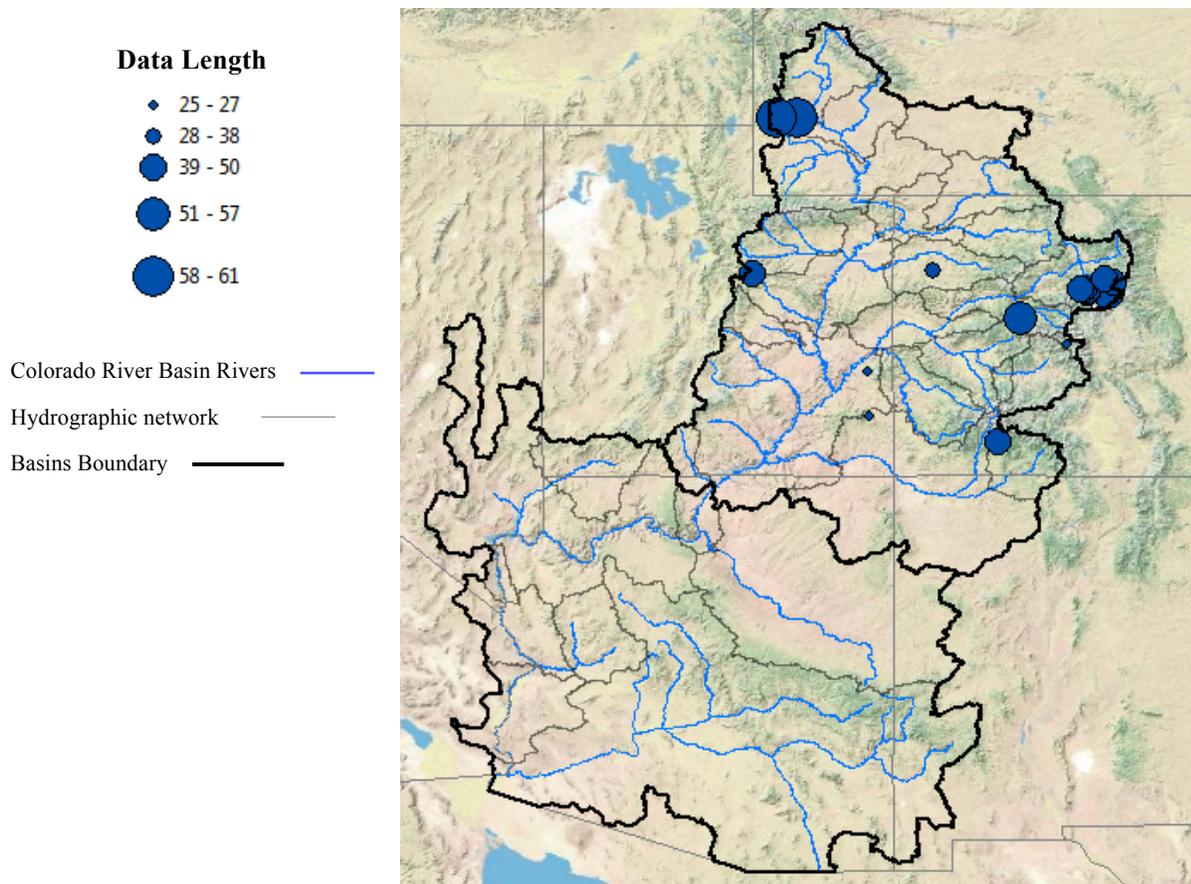


Figure 1: (A) Colorado River basin Location (upper and lower) **(B)** Hydrographic network, major rivers and tributaries, stream gauge stations' locations (blue bubbles), and data length displayed as proportional to the blue bubble diameters.

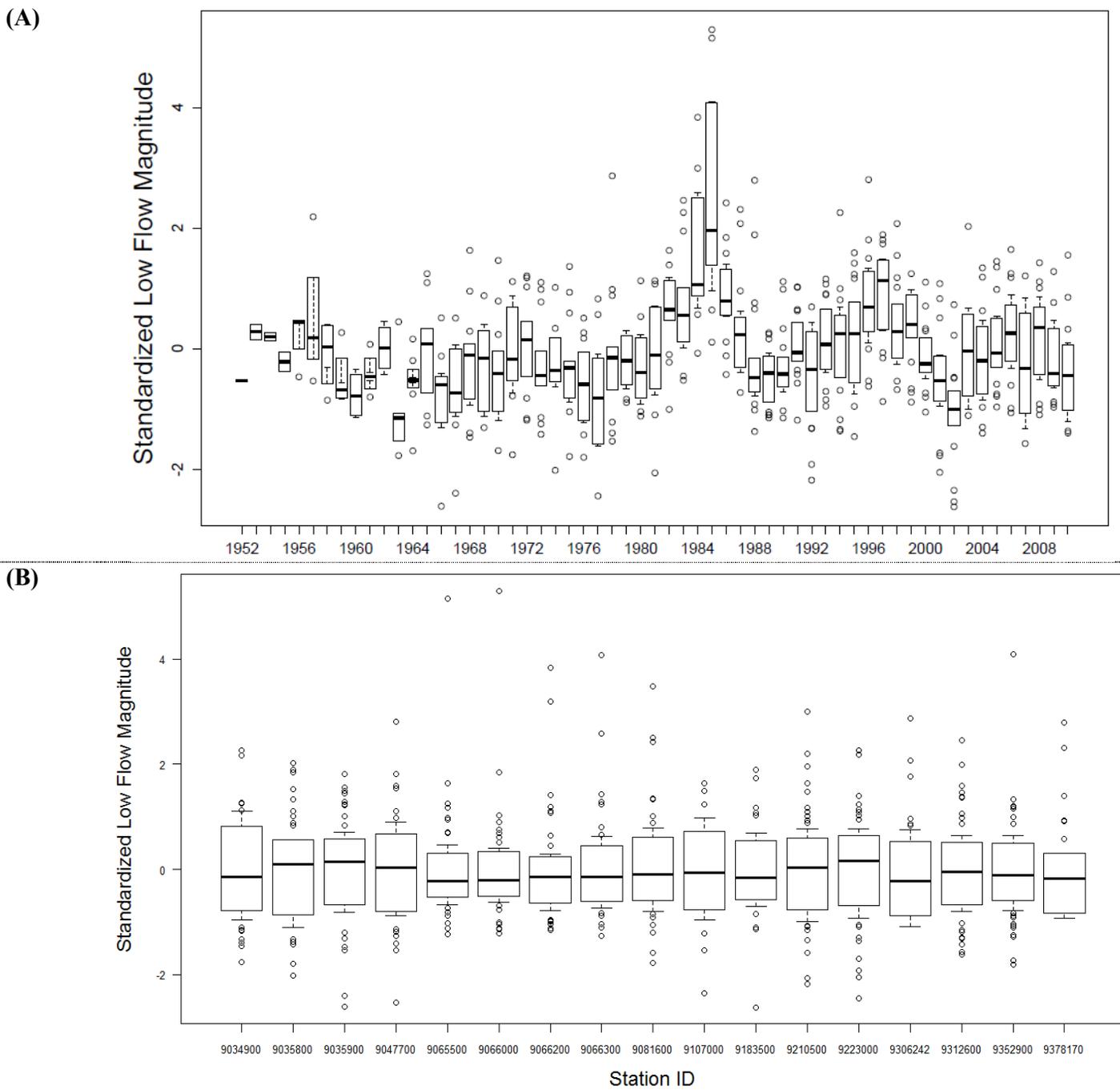
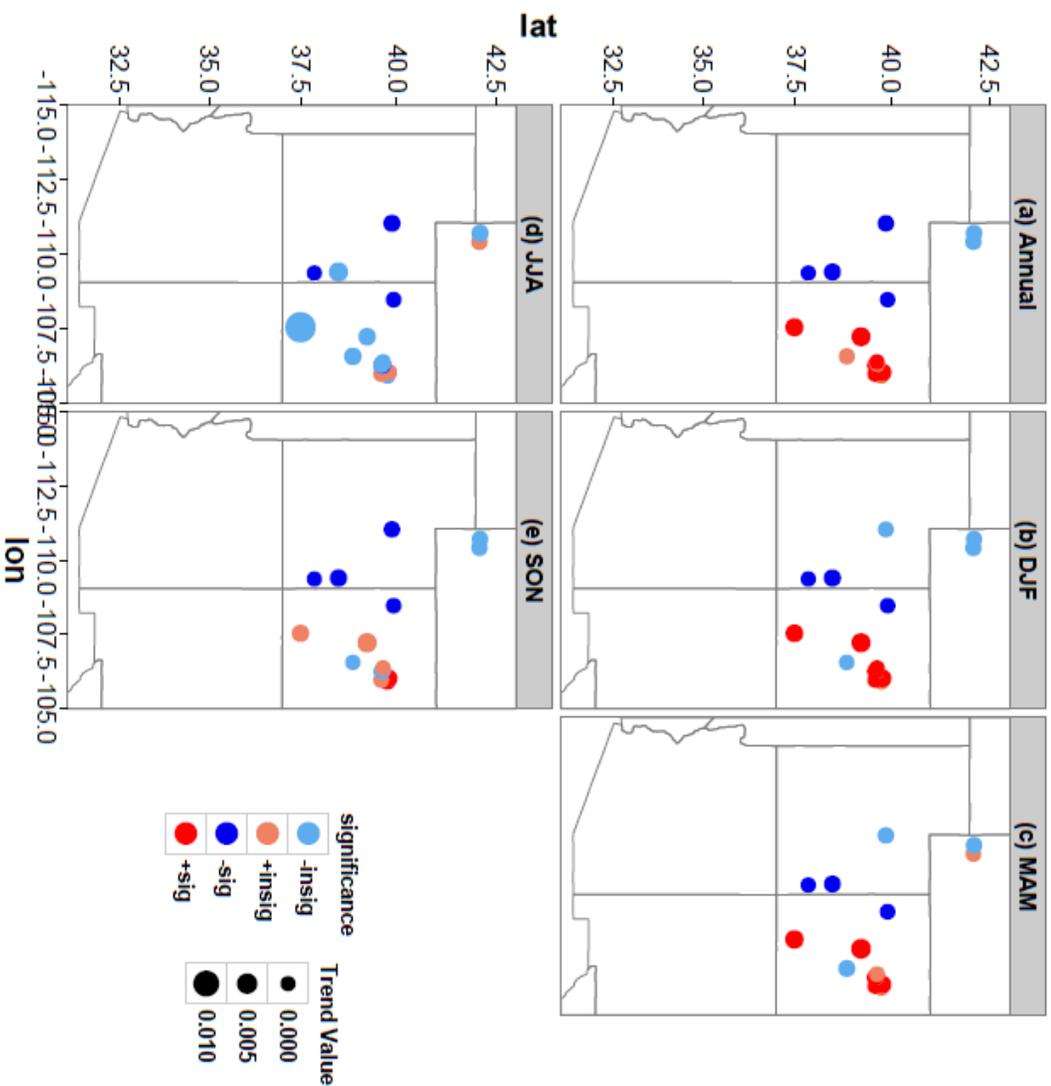


Figure 2: Variability of annual q7 (standardized). **(A)** UCRB while all stations pulled together for every year; **(B)** Individual stations, all years pulled together.

A)



B)

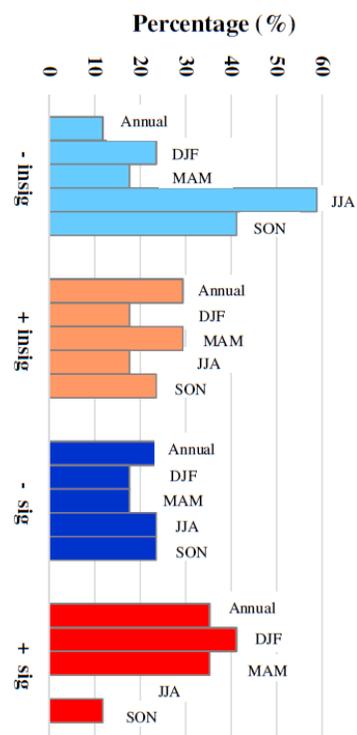


Figure 3: A) Monotonic trends for q7 at each stream-gauge location in cms/day/year. (a) Annual, (b) DJF, (c) MAM, (d) JJA, (e) SON. Color bubbles indicate location of each station, sign and significance of the trend estimates. 90% significant levels are used. The size of the bubble is proportional to the magnitude of the trend. B) The bar plots of different types of trends in annual and four different seasons.

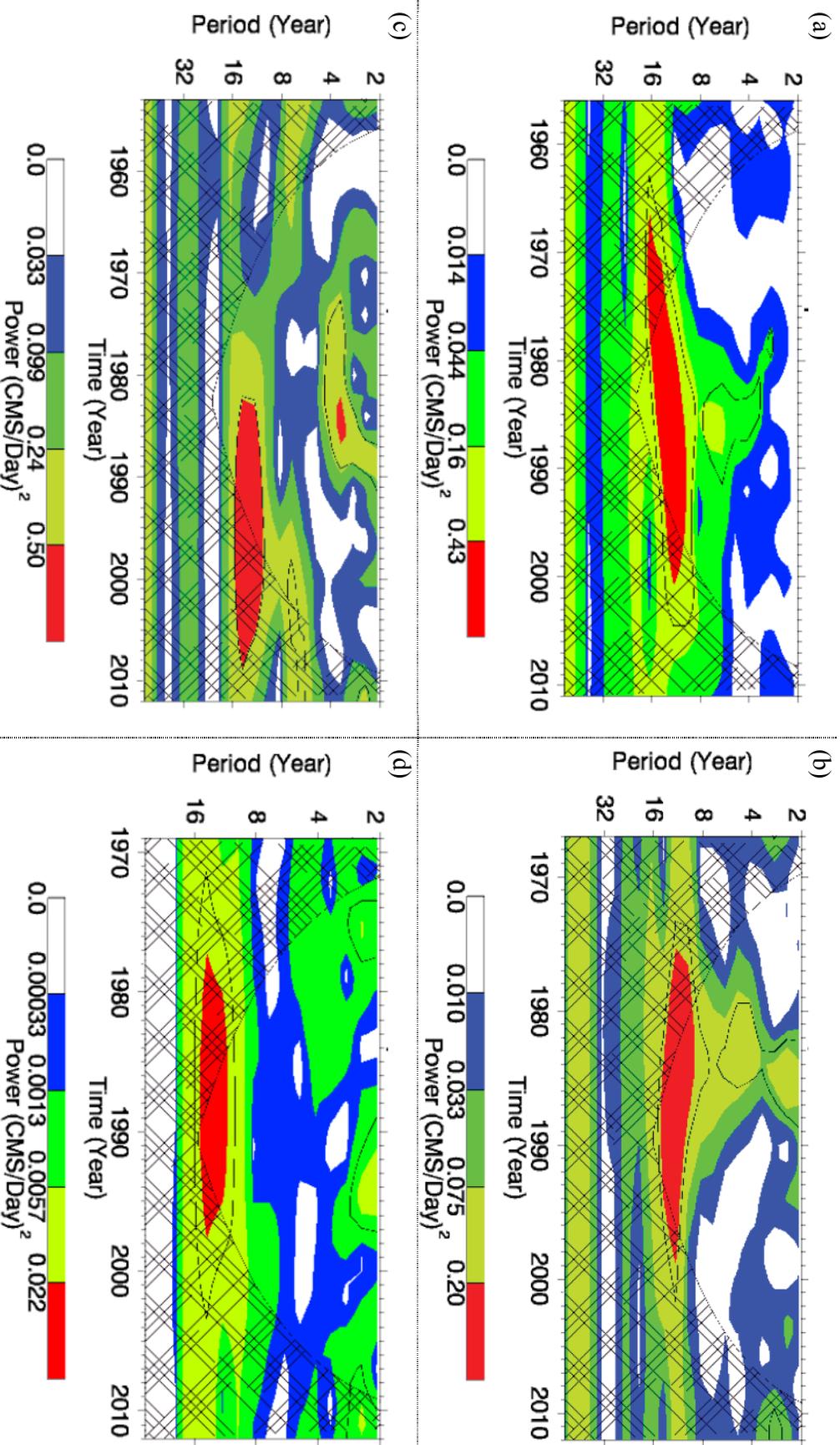


Figure 4: Example periodicity of low flow magnitudes. Wavelet power spectrum for (a) annual q7 time series for Crystal station (east), (b) JJA q7 time series for South Fork station (east), (c) JJA q7 time series for Fontenelle station (west), (d) SON q7 time series for White station (west). The black contours enclose regions of greater than 90% confidence for a red-noise process. Cross-hatched regions indicate the cone of influence (Torrence and Compo, 1998).

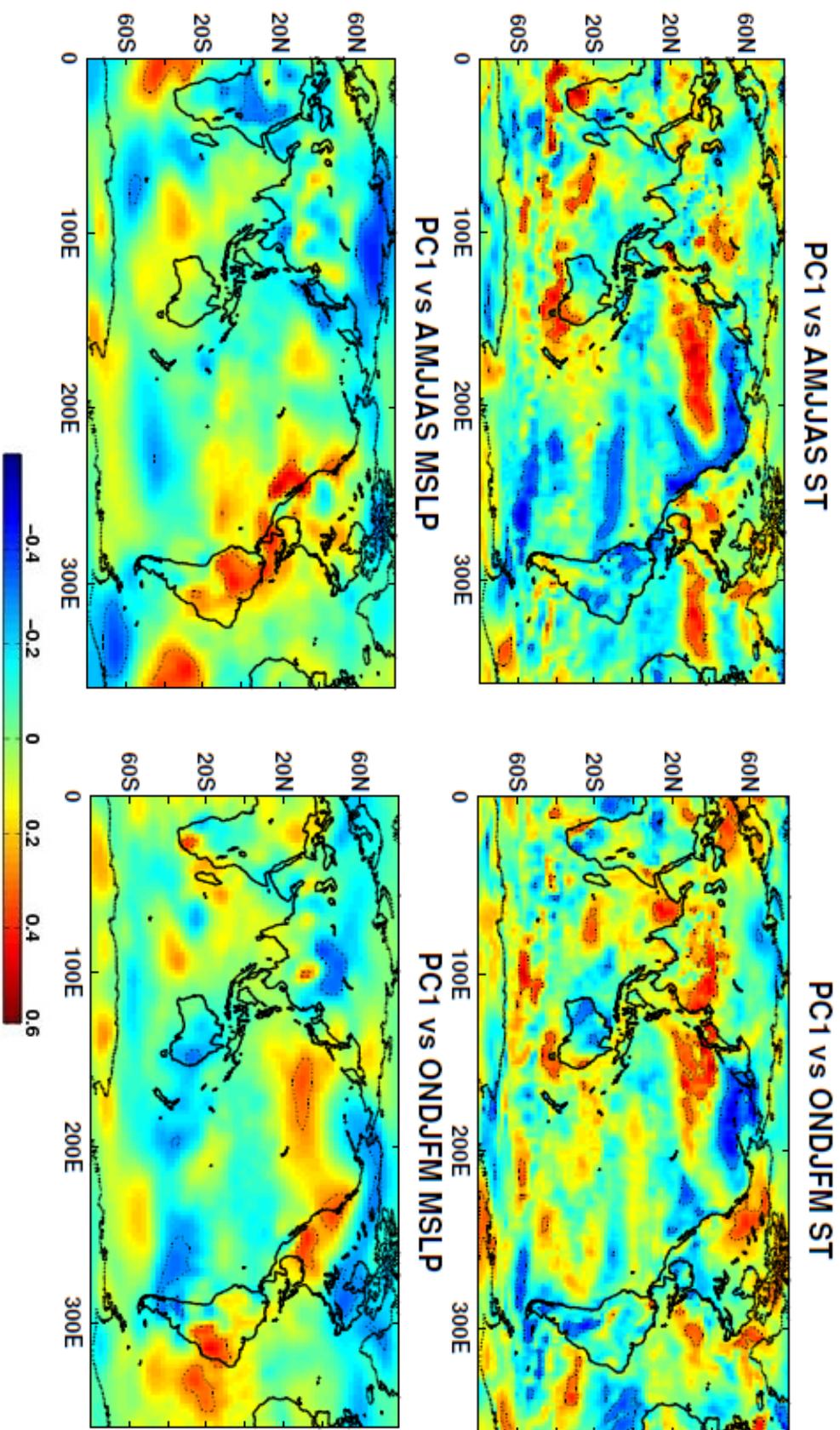


Figure 5: Pearson correlations between PC1 of annual q7 with northern summer (AMJJAS) and northern winter (ONDJFM) climate data (95% significant regions are marked by dotted contours).

Supplementary Information (SI)

Section I: Description of stream flow data

This study uses UCRB stream flow data derived by the Hydro-Climatic Data Network 2009 (HCDN)—a data set developed by the US Geological Survey (USGS) (Lins, 2012). When originally published, the network was composed of 1,659 stations (Slack and Landwehr, 1992) for which the years of primarily “natural” flow were identified (Lins, 2012). The HCDN data set is useful for studying surface water and was specifically developed for examining the effects of climate change on hydrologic conditions. The stream gauge stations selected for inclusion in the HCDN are from locations that are not affected by “artificial diversions, storage, or other human-made works in or on the natural stream channels or watersheds” and has been employed in numerous other stream flow studies (e.g., Douglas et al., 2000; Martin and Arihood, 2010). 17 stations from this network, with drainage basin areas between 10 km² to 432 km², throughout the UCRB are examined in this research. The names and detailed descriptions of those stream-gaging stations (17 stations) of UCRB are shown in Table 1-A.

Section II: Definition of “low flow”

Low flow has different definitions. Many define this as “the actual flows in a river occurring during the dry season of the year”, others define as “the length of time and the conditions occurring between flood events”, or “the changes in the total flow regime of a river on sustainable water yield or riverine and riparian ecology” (Smakhtin, 2001). On the other hand, international glossary of hydrology (WMO, 1974) defines low flow as “flow of water in a stream during prolonged dry weather”. This definition does not make a clear distinction between low

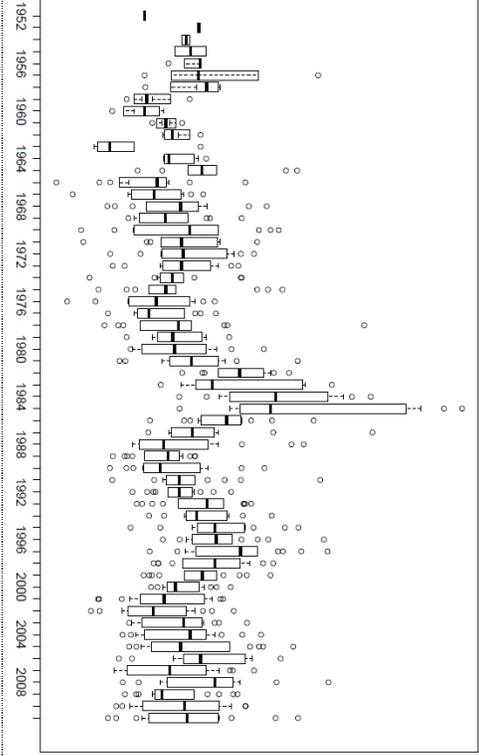
flows and droughts though, which is a natural event resulting from a less than normal precipitation for an extended period of time (EPA, 2012). Hence, low flow is a seasonal phenomenon, and an integral component of a flow regime of any river and drought is a more general phenomenon that includes low-flow periods, but a continuous seasonal low-flow event might not necessarily constitute a drought (Smakhtin, 2001).

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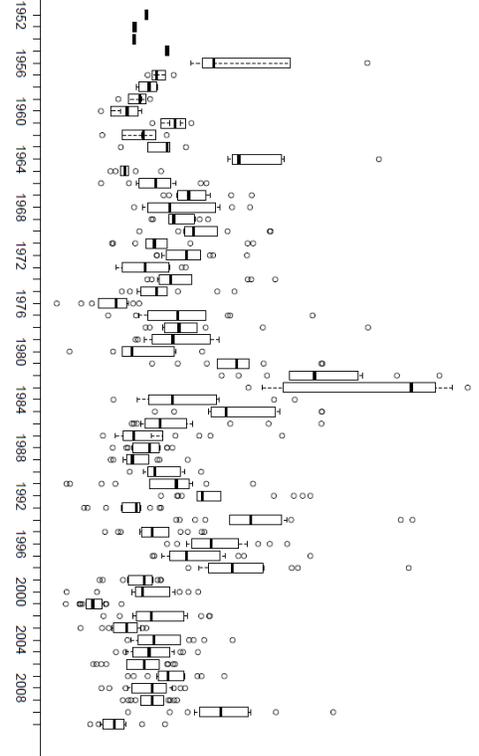
(A)

Standardized Low Flow Magnitude



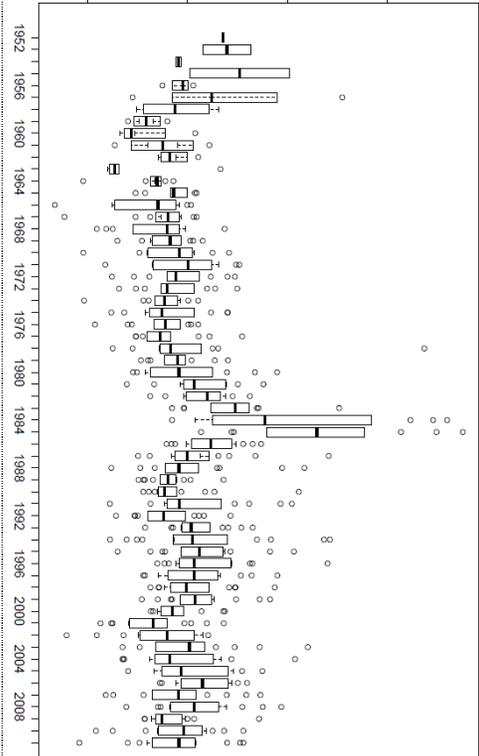
DJF

Standardized Low Flow Magnitude



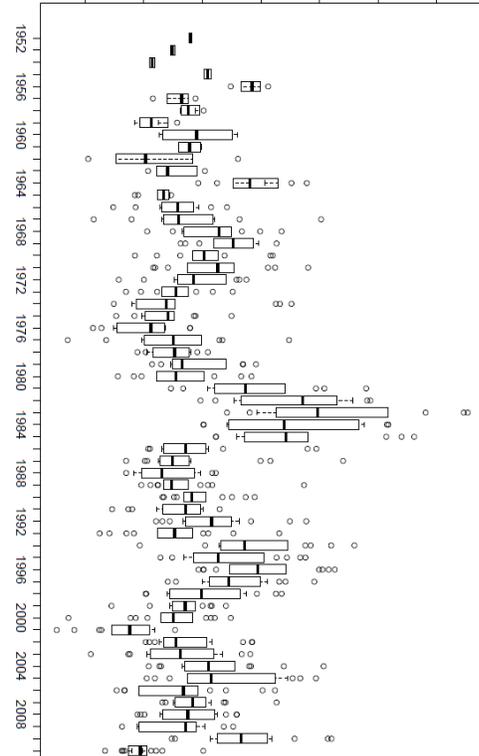
JJA

Standardized Low Flow Magnitude



MAM

Standardized Low Flow Magnitude



SON

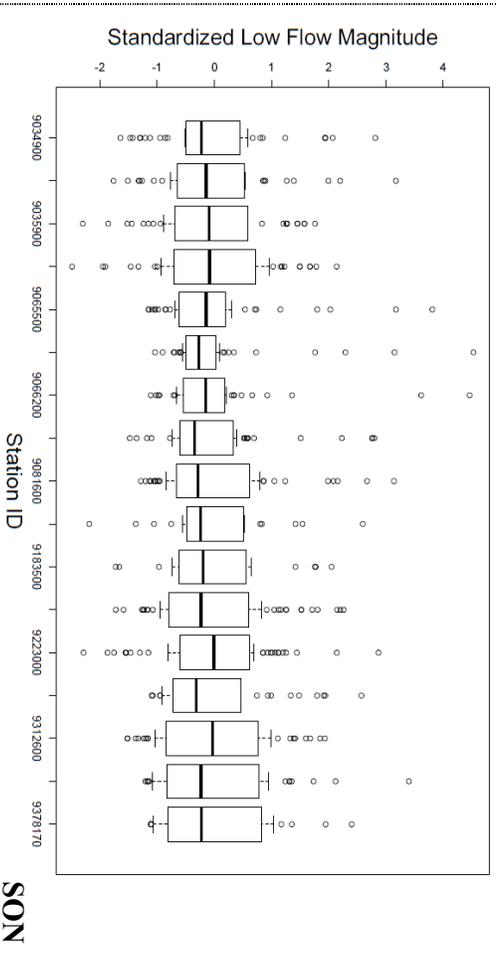
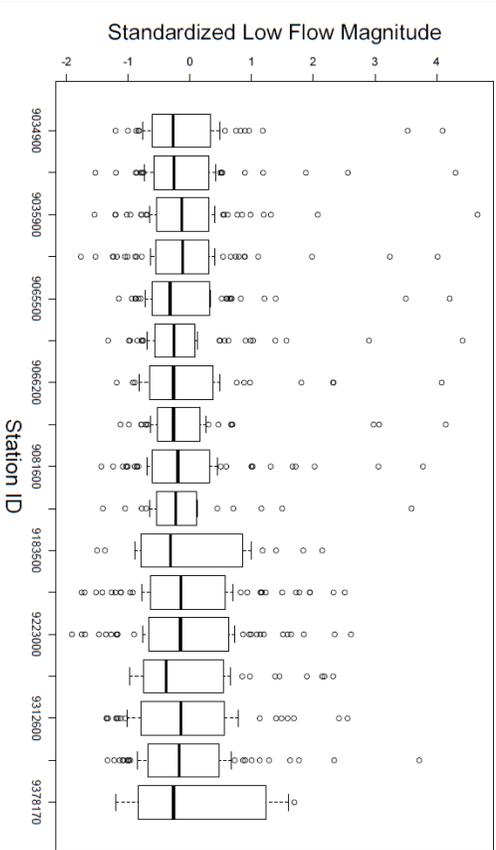
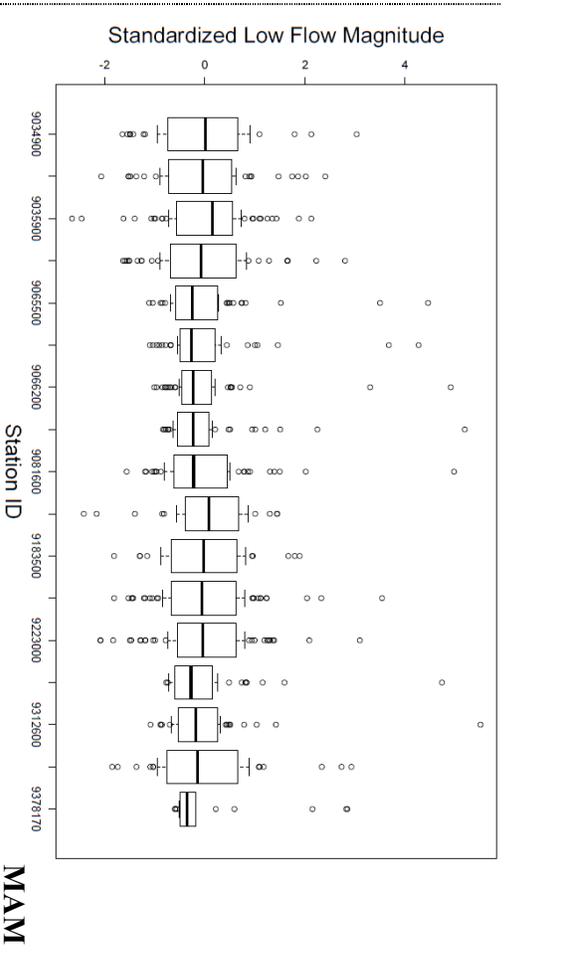
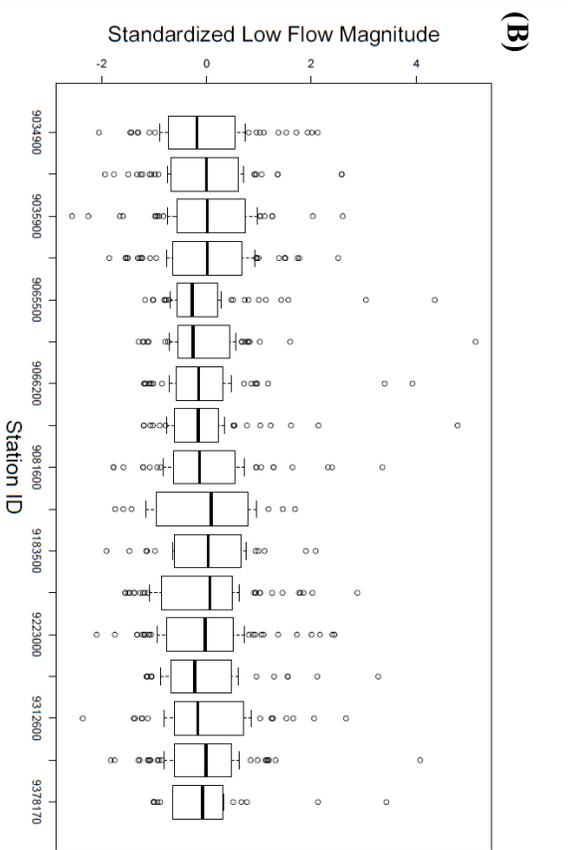
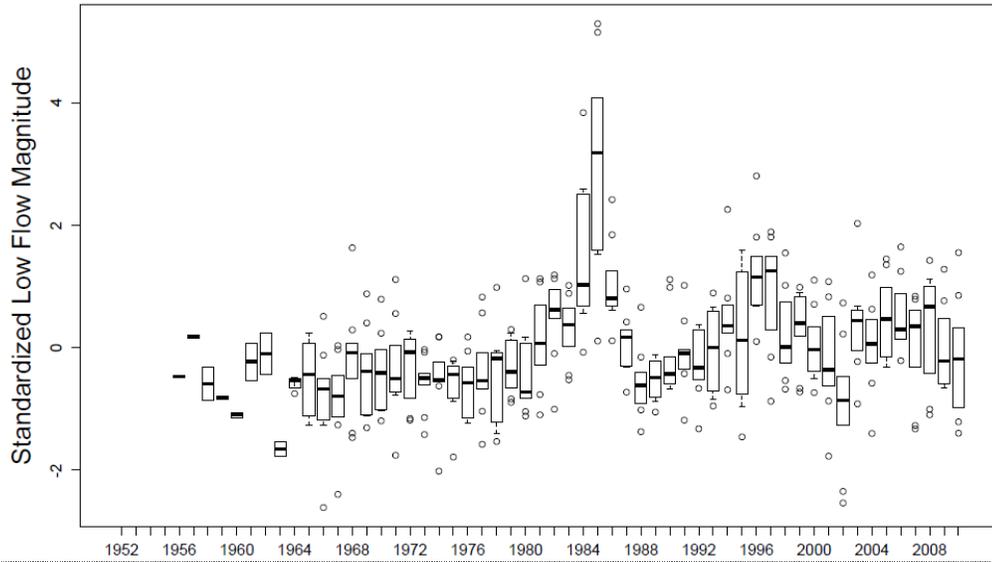


Figure S1: Variability of seasonal q7 (standardized). **(A)** UCRB while all stations pulled together for every year, **(B)** Individual stations, all years pulled together.

(A)



(B)

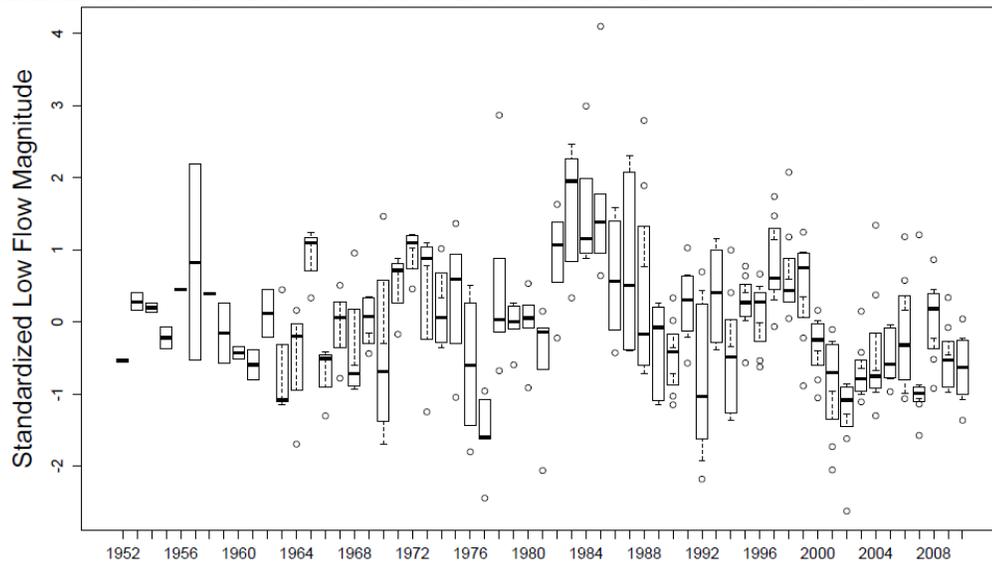
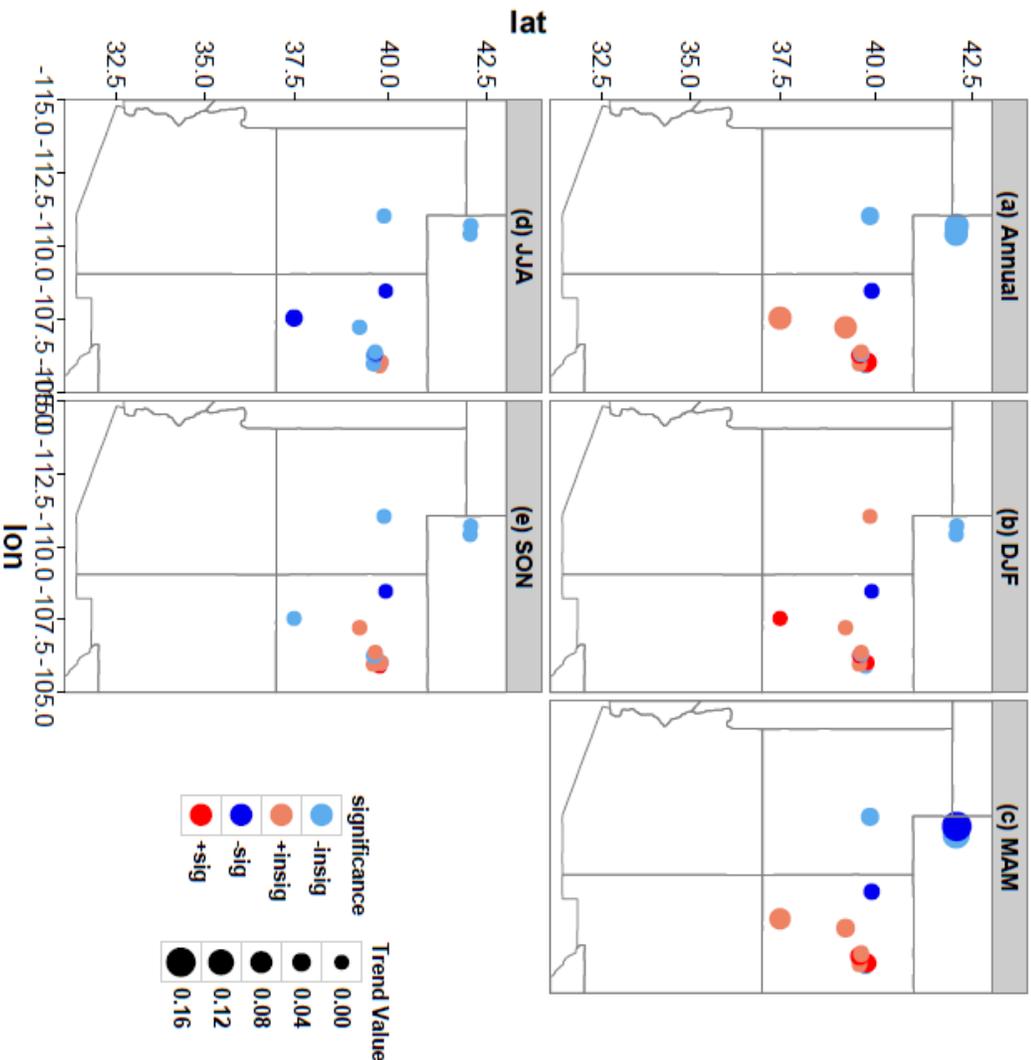


Figure S2: Variability of annual q7 (standardized) of UCRB. **(A)** All “eastern” stations pulled together for every year; **(B)** All “western” stations pulled together for every year.

A)



B)

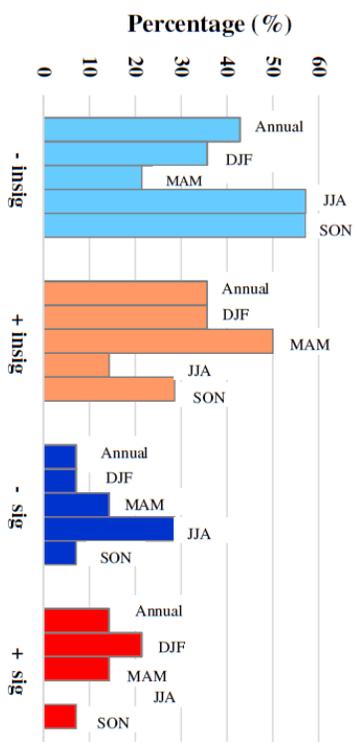


Figure S3: A) Monotonic trends for q_7 at each stream-gauge location having more than 30-years data in cms/day/year. (a) Annual, (b) DJF, (c) MAM, (d) JJA, (e) SON. Color bubbles indicate location of each station, sign and significance of the trend estimates. 90% significant levels are used. The size of the bubble is proportional to the magnitude of the trend. B) The bar plots of different types of trends in annual and four different seasons.

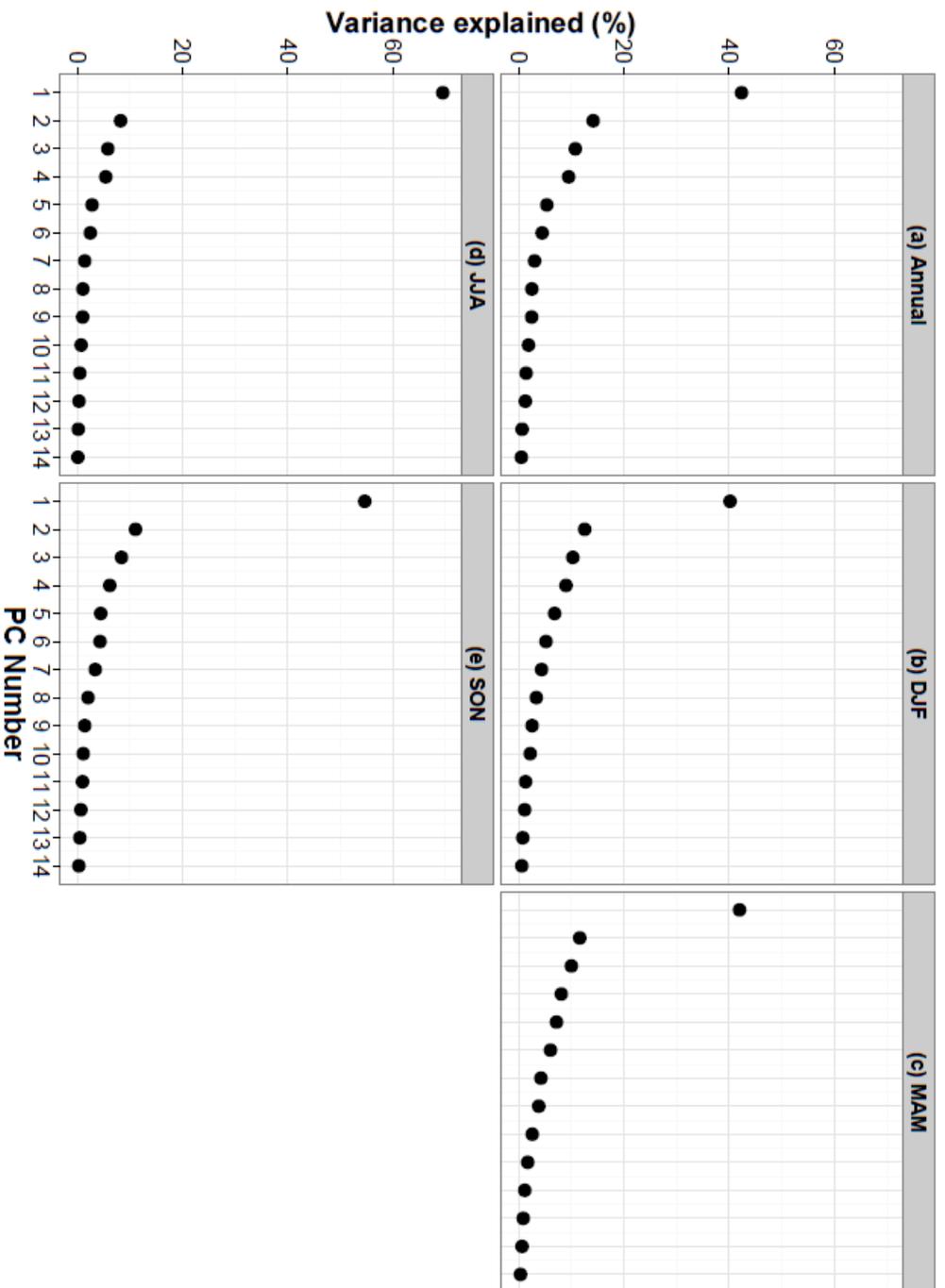


Figure S4: Variance explained by each Principal Component for (a) annual, (b) DJF, (c) MAM, (d) JJA, (e) SON.

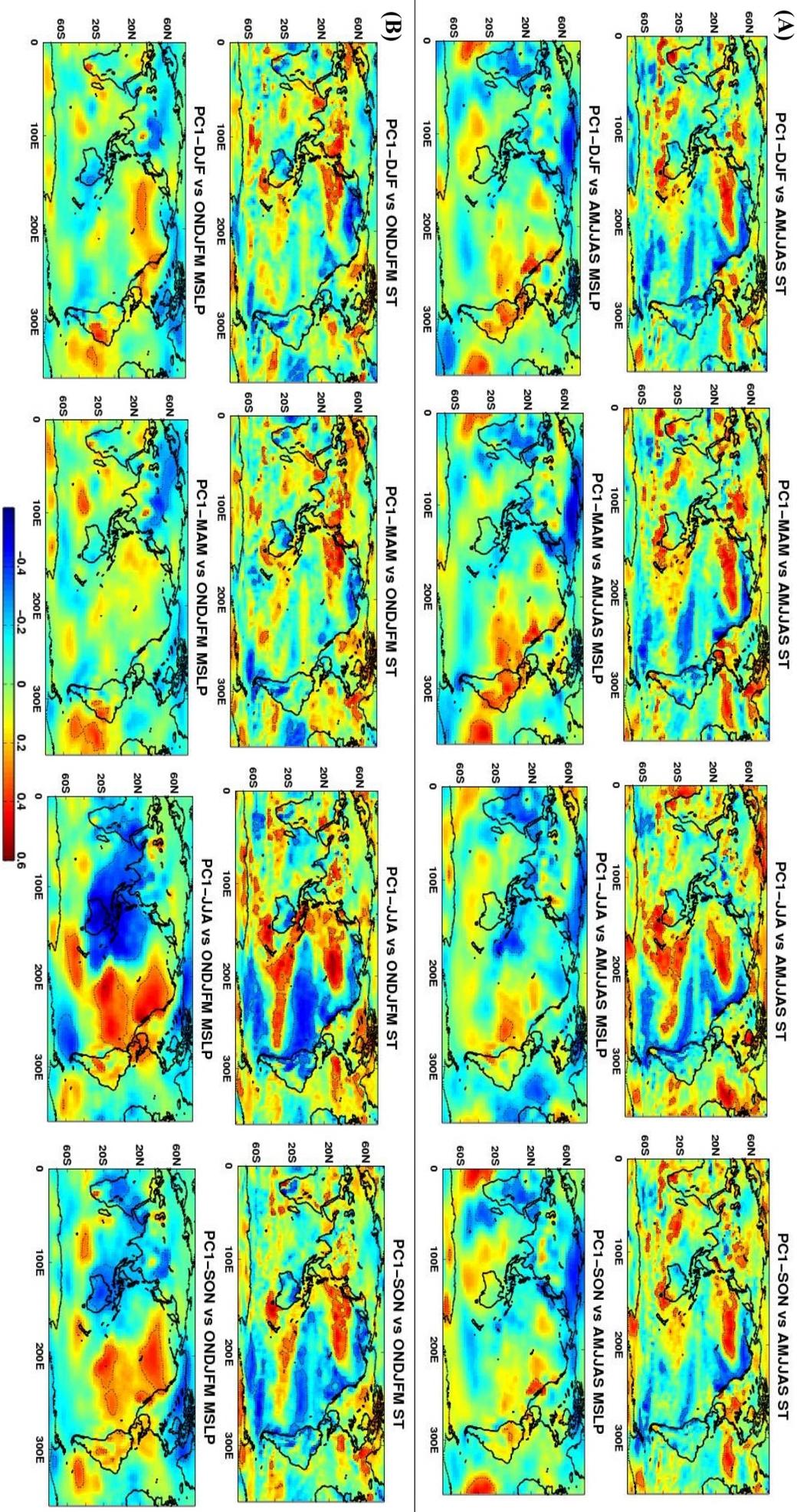


Figure S5: Pearson correlations between PC1 of seasonal q7 with climate data. (A) Traditional seasons q7 versus northern summer (AMJJAS) climate, (B) Traditional seasons q7 versus northern winter (ONDJFM). (95% significant regions are marked by dotted contours).

Table S1: Occurrence month(s) of annual low flow (q7) for different stream gauge stations in UCRB.

Station ID	Station Name	Annual q7 Occurrence Month(s)
9066000	Black Gore Creek Near Minturn, CO. <u>(east)</u>	Oct-Mar
9034900	Bobtail Creek Near Jones Pass, CO. <u>(east)</u>	Dec-May
9066200	Booth Creek Near Minturn, CO. <u>(east)</u>	Aug-Apr
9306242	Corral Gulch Near Rangely, CO. <u>(west)</u>	Oct-Mar
9081600	Crystal River Ab Avalanche C, Near Redstone, CO. <u>(east)</u>	Nov-Apr
9035800	Darling Creek Near Leal, CO. <u>(east)</u>	Nov-Apr
9065500	Gore Creek At Upper Station, Near Minturn, CO. <u>(east)</u>	Nov-Apr
9047700	Keystone Gulch Near Dillon, CO. <u>(east)</u>	Aug-Apr
9066300	Middle Creek Near Minturn, CO. <u>(east)</u>	Dec-Apr
9035900	South Fork Of Williams Fork Near Leal, CO. <u>(east)</u>	Oct-Mar
9107000	Taylor River At Taylor Park, CO. <u>(east)</u>	Oct-Mar
9352900	Vallecio Creek Near Bayfield, CO. <u>(west)</u>	Oct-Feb
9183500	Mill Creek at Sheley Tunnel, Near Moab, UT. <u>(west)</u>	Jul-Aug & Oct-Feb
9210500	Fontenelle C Nr Herschler Ranch, Nr Fontenelle, WY. <u>(west)</u>	Jul-Aug & Oct-Mar
9223000	Hams Fork Below Pole Creek, Near Frontier, WY. <u>(west)</u>	Jul-Aug & Oct-Feb
9312600	White River Bl Tabbyune C Near Soldier Summit, UT. <u>(west)</u>	Jul-Aug & Oct-Jan
9378170	South Creek Above Reservoir Near Monticello, UT. <u>(west)</u>	Jul-Aug & Oct-Mar

Note: CO = Colorado; UT = Utah; WY = Wyoming; LAT_GAGE = Latitude of a streamgauge; LONG_GAGE = Longitude of a streamgauge.

Table S2: Summary statistics table for DJF q7 time series for each stream gauge station in UCRB.

Station ID	Station Name	Mean (cms)	Median (cms)	Standard deviation (cms)	Skewness	Kurtosis
9066000	Black Gore Creek Near Minturn, CO. (east)	0.063	0.056	0.029	3.14	15.79
9034900	Bobtail Creek Near Jones Pass, CO. (east)	0.021	0.020	0.004	0.33	-0.46
9066200	Booth Creek Near Minturn, CO. (east)	0.023	0.021	0.012	2.09	6.29
9306242	Corral Gulch Near Rangely, CO. (west)	0.014	0.012	0.011	1.27	1.90
9081600	Crystal River Ab Avalanche C, Near Redstone, CO. (east)	1.222	1.189	0.259	0.95	1.63
9035800	Darling Creek Near Leal, CO. (east)	0.054	0.055	0.012	0.41	0.47
9065500	Gore Creek At Upper Station, Near Minturn, CO. (east)	0.076	0.068	0.033	2.48	7.85
9047700	Keystone Gulch Near Dillon, CO. (east)	0.055	0.055	0.010	0.25	-0.42
9066300	Middle Creek Near Minturn, CO. (east)	0.008	0.007	0.007	2.68	10.75
9035900	South Fork Of Williams Fork Near Leal, CO. (east)	0.210	0.209	0.046	0.14	0.78
9107000	Taylor River At Taylor Park, CO. (east)	0.845	0.858	0.103	-0.30	-0.91
9352900	Vallecito Creek Near Bayfield, CO. (west)	0.490	0.487	0.152	1.21	4.57
9183500	Mill Creek at Sheley Tunnel, Near Moab, UT. (west)	0.138	0.140	0.022	0.17	0.08
9210500	Fontenelle C Nr Herschler Ranch, Nr Fontenelle, WY. (west)	0.553	0.562	0.142	0.46	0.05
9223000	Hams Fork Below Pole Creek, Near Frontier, WY. (west)	0.322	0.320	0.103	0.49	0.14
9312600	White River Bl Tabbyune C Near Soldier Summit, UT. (west)	0.091	0.085	0.035	0.42	0.50
9378170	South Creek Above Reservoir Near Monticello, UT. (west)	0.002	0.002	0.002	2.04	5.32

Note: CO = Colorado; UT = Utah; WY = Wyoming. Highlighted skewness and kurtosis values indicate non-normality of the data.

Table S3: Summary statistics table for MAM q7 time series for each stream gauge station in UCRB.

Station ID	Station Name	Mean (cms)	Median (cms)	Standard deviation (cms)	Skewness	Kurtosis
9066000	Black Gore Creek Near Minturn, CO. (east)	0.076	0.065	0.048	2.91	10.05
9034900	Bobtail Creek Near Jones Pass, CO. (east)	0.020	0.020	0.004	0.48	0.89
9066200	Booth Creek Near Minturn, CO. (east)	0.027	0.024	0.015	3.45	14.36
9306242	Corral Gulch Near Rangely, CO. (west)	0.025	0.016	0.023	1.11	0.01
9081600	Crystal River Ab Avalanche C, Near Redstone, CO. (east)	1.374	1.278	0.377	2.38	10.00
9035800	Darling Creek Near Leal, CO. (east)	0.053	0.052	0.010	0.29	0.01
9065500	Gore Creek At Upper Station, Near Minturn, CO. (east)	0.086	0.075	0.045	2.85	10.07
9047700	Keystone Gulch Near Dillon, CO. (east)	0.054	0.054	0.010	0.47	0.11
9066300	Middle Creek Near Minturn, CO. (east)	0.009	0.007	0.008	3.49	15.65
9035900	South Fork Of Williams Fork Near Leal, CO. (east)	0.207	0.215	0.049	-0.36	0.60
9107000	Taylor River At Taylor Park, CO. (east)	0.906	0.906	0.101	-0.85	1.50
9352900	Vallecito Creek Near Bayfield, CO. (west)	0.605	0.568	0.208	0.92	1.38
9183500	Mill Creek at Shaley Tunnel, Near Moab, UT. (west)	0.145	0.143	0.026	0.30	-0.58
9210500	Fontenelle C Nr Herschler Ranch, Nr Fontenelle, WY. (west)	0.669	0.651	0.170	0.91	1.83
9223000	Hams Fork Below Pole Creek, Near Frontier, WY. (west)	0.402	0.396	0.146	0.26	0.66
9312600	White River Bl Tabbyune C Near Soldier Summit, UT. (west)	0.157	0.139	0.104	4.02	21.19
9378170	South Creek Above Reservoir Near Monticello, UT. (west)	0.010	0.004	0.016	2.27	4.01

Note: CO = Colorado; UT = Utah; WY = Wyoming. Highlighted skewness and kurtosis values indicate non-normality of the data.

Table S4: Summary statistics table for JJA q7 time series for each stream gauge station in UCRB.

Station ID	Station Name	Mean (cms)	Median (cms)	Standard deviation (cms)	Skewness	Kurtosis
9066000	Black Gore Creek Near Minturn, CO. (east)	0.144	0.125	0.072	2.54	8.35
9034900	Bobtail Creek Near Jones Pass, CO. (east)	0.152	0.132	0.079	2.53	8.06
9066200	Booth Creek Near Minturn, CO. (east)	0.076	0.063	0.051	2.08	5.51
9306242	Corral Gulch Near Rangely, CO. (west)	0.025	0.016	0.023	1.11	0.01
9081600	Crystal River Ab Avalanche C, Near Redstone, CO. (east)	3.743	3.418	1.696	1.73	3.83
9035800	Darling Creek Near Leal, CO. (east)	0.141	0.125	0.065	2.31	7.56
9065500	Gore Creek At Upper Station, Near Minturn, CO. (east)	0.341	0.269	0.224	2.54	8.13
9047700	Keystone Gulch Near Dillon, CO. (east)	0.115	0.109	0.049	1.82	5.74
9066300	Middle Creek Near Minturn, CO. (east)	0.046	0.035	0.040	2.77	8.37
9035900	South Fork Of Williams Fork Near Leal, CO. (east)	0.541	0.510	0.207	2.49	10.04
9107000	Taylor River At Taylor Park, CO. (east)	1.867	1.642	0.861	2.22	6.83
9352900	Vallecito Creek Near Bayfield, CO. (west)	2.058	1.881	1.122	1.41	2.90
9183500	Mill Creek at Sheley Tunnel, Near Moab, UT. (west)	0.208	0.184	0.090	0.64	-0.52
9210500	Fontenelle C Nr Herschler Ranch, Nr Fontenelle, WY. (west)	0.842	0.785	0.382	0.65	-0.04
9223000	Hams Fork Below Pole Creek, Near Frontier, WY. (west)	0.549	0.508	0.274	0.41	0.04
9312600	White River Bl Tabbyune C Near Soldier Summit, UT. (west)	0.120	0.106	0.087	0.72	0.09
9378170	South Creek Above Reservoir Near Monticello, UT. (west)	0.003	0.002	0.003	0.50	-1.22

Note: CO = Colorado; UT = Utah; WY = Wyoming. Highlighted skewness and kurtosis values indicate non-normality of the data.

Table S5: Summary statistics table for SON q7 time series for each stream gauge station in UCRB.

Station ID	Station Name	Mean (cms)	Median (cms)	Standard deviation (cms)	Skewness	Kurtosis
9066000	Black Gore Creek Near Minturn, CO. (east)	0.082	0.071	0.040	3.14	10.63
9034900	Bobtail Creek Near Jones Pass, CO. (east)	0.039	0.037	0.009	0.83	0.63
9066200	Booth Creek Near Minturn, CO. (east)	0.036	0.033	0.022	3.04	11.14
9306242	Corral Gulch Near Rangely, CO. (west)	0.018	0.013	0.014	1.00	-0.05
9081600	Crystal River Ab Avalanche C, Near Redstone, CO. (east)	1.776	1.606	0.549	1.27	1.43
9035800	Darling Creek Near Leal, CO. (east)	0.074	0.072	0.020	0.90	1.37
9065500	Gore Creek At Upper Station, Near Minturn, CO. (east)	0.116	0.112	0.043	2.08	5.32
9047700	Keystone Gulch Near Dillon, CO. (east)	0.072	0.071	0.018	-0.07	-0.02
9066300	Middle Creek Near Minturn, CO. (east)	0.016	0.012	0.010	1.57	2.33
9035900	South Fork Of Williams Fork Near Leal, CO. (east)	0.287	0.280	0.065	-0.12	-0.57
9107000	Taylor River At Taylor Park, CO. (east)	1.165	1.096	0.267	0.42	1.47
9352900	Vallecito Creek Near Bayfield, CO. (west)	0.871	0.777	0.400	1.08	1.42
9183500	Mill Creek at Sheley Tunnel, Near Moab, UT. (west)	0.159	0.146	0.049	0.64	-0.01
9210500	Fontenelle C Nr Herschler Ranch, Nr Fontenelle, WY. (west)	0.645	0.599	0.214	0.56	-0.46
9223000	Hams Fork Below Pole Creek, Near Frontier, WY. (west)	0.408	0.409	0.151	0.05	0.35
9312600	White River Bl Tabbyune C Near Soldier Summit, UT. (west)	0.091	0.090	0.058	0.26	-0.95
9378170	South Creek Above Reservoir Near Monticello, UT. (west)	0.002	0.002	0.002	0.77	-0.16

Note: CO = Colorado; UT = Utah; WY = Wyoming. Highlighted skewness and kurtosis values indicate non-normality of the data.

Table S6: Pearson correlation coefficients between DJF q7 time series of different stations. 90% statistically significant estimates are shown.

Station ID	9034900	9066200	9306242	9081600	9035800	9065500	9047700	9066300	9035900	9107000	9352900	9183500	9210500	9223000	9312600	9378170	9034900
9066000																	
9034900																	
9066200																	
9306242																	
9081600	0.38																
9035800					0.42												
9065500		0.41	0.47														
9047700				0.45													
9066300			0.48		0.40												
9035900	0.36				0.41	0.49			0.47								
9107000						0.36											
9352900	0.42			-0.37	0.64												
9183500				0.41	0.41	0.41											
9210500	0.65				0.39			0.39		0.38							
9223000				0.38									0.58	0.39			
9312600			0.36		0.48									0.35			
9378170				0.44				0.36			0.38		0.63		0.36		

Table S7: Pearson correlation coefficients between MAM q7 time series of different stations. 90% statistically significant estimates are shown.

Station ID	9034900	9066200	9306242	9081600	9035800	9065500	9047700	9066300	9035900	9107000	9352900	9183500	9210500	9223000	9312600	9378170	9034900
9066000																	
9034900																	
9066200																	
9306242																	
9081600																	
9035800																	
9065500			0.43														
9047700				0.35													
9066300			0.46														
9035900									0.43								
9107000																	
9352900		-0.52				0.57											
9183500				0.52								0.51					
9210500																	
9223000													0.60	0.41			
9312600												0.41	0.37				
9378170		-0.48			0.52	-0.35				-0.43		0.65				0.40	

Table S8: Pearson correlation coefficients between SON q7 time series of different stations. 90% statistically significant estimates are shown.

Station ID	9034900	9066200	9306242	9081600	9035800	9065500	9047700	9066300	9035900	9107000	9352900	9183500	9210500	9223000	9312600	9378170	9034900
9066000																	
9034900																	
9066200																	
9306242																	
9081600	0.64																
9035800	0.49			0.39	0.74												
9065500	0.42		0.51		0.66	0.49											
9047700	0.65				0.50	0.44											
9066300			0.48		0.52	0.53	0.56	0.47									
9035900	0.80				0.70	0.53	0.57	0.59	0.62								
9107000	0.53		0.54	0.36	0.65	0.59	0.81	0.50	0.51	0.62							
9352900																	
9183500	0.49			0.68	0.57	0.68		0.37			0.53						
9210500	0.62			0.47	0.57	0.44		0.37		0.58	0.48		0.50				
9223000	0.70			0.37	0.55	0.48		0.42		0.63	0.39		0.43	0.86			
9312600	0.66			0.63	0.71	0.77	0.53	0.47	0.41	0.62	0.74		0.78	0.74	0.69		
9378170	0.45	-0.37		0.56	0.35	0.56					0.40		0.77	0.49	0.49	0.63	

Table S9: Periodicities identified (in number of years) for annual q7 magnitude for different stream gauge locations.

Station ID	Station Name	Active years identified	Periodicity in years (within the cone of influence)
9066000	Black Gore Creek Near Minturn, CO. (east)		
9034900	Bobtail Creek Near Jones Pass, CO. (east)		
9066200	Booth Creek Near Minturn, CO. (east)		
9306242	Corral Gulch Near Rangely, CO. (west)		
9081600	Crystal River Ab Avalanche C, Near Redstone, CO. (east)	1967-2003	12 - 16
9035800	Darling Creek Near Leal, CO. (east)	1978-1990 1980-2000	6 - 8 15
9065500	Gore Creek At Upper Station, Near Minturn, CO. (east)	1976-2000	10 - 15
9047700	Keystone Gulch Near Dillon, CO. (east)		
9066300	Middle Creek Near Minturn, CO. (east)		
9035900	South Fork Of Williams Fork Near Leal, CO. (east)		
9107000	Taylor River At Taylor Park, CO. (east)		
9352900	Vallecito Creek Near Bayfield, CO. (west)		
9183500	Mill Creek at Shelley Tunnel, Near Moab, UT. (west)	1996-2004	5 - 7
9210500	Fontenelle C Nr Herschler Ranch, Nr Fontenelle, WY. (west)	1974-2003	13 - 16
9223000	Hams Fork Below Pole Creek, Near Frontier, WY. (west)	1971 - 2003	10 - 16
9312600	White River BI Tabbyune C Near Soldier Summit, UT. (west)	1975-1995	12-16
9378170	South Creek Above Reservoir Near Monticello, UT. (west)		

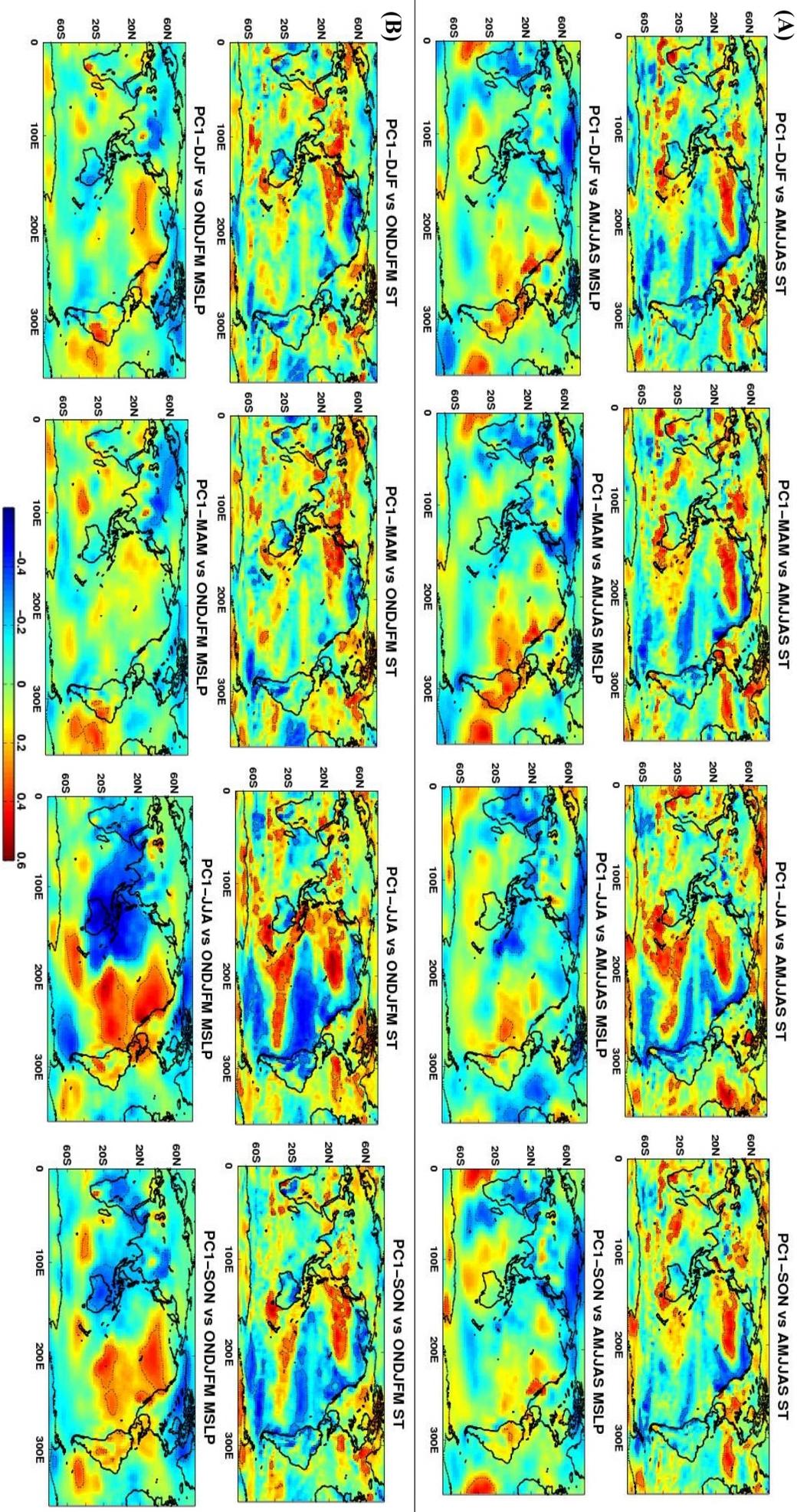


Figure S5: Pearson correlations between PC1 of seasonal q7 with climate data. (A) Traditional seasons q7 versus northern summer (AMJJAS) climate, (B) Traditional seasons q7 versus northern winter (ONDJFM). (95% significant regions are marked by dotted contours).