

1 **Nitrogen Dynamics and Associated Greenhouse Gases Emissions in Constructed**  
2 **Wetlands: A Review**

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## 34 **Abstract**

35 **Wastewater, derived from various sources, contains levels of pollutants that exceed water quality**  
36 **guidelines. Constructed wetlands (CWs) are a rising technology globally used to remove these**  
37 **pollutants (e.g. carbon- C and nitrogen- N) from wastewater.** The N removal efficiency of CWs is  
38 very inconsistent and does not in itself reveal whether the removal processes are due to physical  
39 attenuation or whether the different species have been transformed to other reactive forms  
40 causing pollution swapping. **There are many pathways by which the removed N can contribute to**  
41 **water and air pollution:** accumulation and adsorption in soils, leaching to groundwater (nitrate-  
42  $\text{NO}_3^-$  and ammonium- $\text{NH}_4^+$ ), emission to atmosphere via nitrous oxide-  $\text{N}_2\text{O}$  and ammonia and/or  
43 conversion to **di-nitrogen ( $\text{N}_2$ )** gas. The kinetics of these pathways/processes varies with CW  
44 management and therefore needs to be studied quantitatively, for their sustainable use. For  
45 example, the quality of groundwater underlying CWs with regards to the reactive N (Nr) species  
46 is largely unknown. Equally, there is a dearth of information on the extent of Nr accumulation in  
47 soils and discharge to surface waters and air. Moreover, CWs are rich in dissolved organic carbon  
48 (DOC) and produce substantial amounts of **carbon dioxide ( $\text{CO}_2$ ) and methane ( $\text{CH}_4$ ).** These  
49 dissolved C species discharge to ground and surface waters and are emitted to the atmosphere.  
50 The dynamics of dissolved  $\text{N}_2\text{O}$ ,  $\text{CO}_2$  and  $\text{CH}_4$  in CWs is a key “loop hole” in global greenhouse  
51 gas budgets. In this review we provide an overview of the current knowledge and discussion  
52 about the dynamics of C and N in CWs and their likely impacts on aquatic and atmospheric  
53 environments. We recommend that the fate of various N species in CWs and their surface  
54 emissions and subsurface export need to be evaluated in a holistic way to better understand their  
55 potential for pollution swapping. Research on process based N removal and balancing of end  
56 products into reactive and benign forms is critical when assessing the environmental performance  
57 of CWs. Thus we strongly suggest further detailed examinations of *in situ* N transformations and  
58 the fate of such transformational products with respect to pollution swapping.

59

60 **Keywords:** Carbon, nitrogen, constructed wetlands, pollution swapping, nitrous oxide, methane

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## 62 **1. Introduction**

63 Increasing anthropogenic loading of reactive N (Nr) along the **nitrogen (N)** cascade in the  
64 environment raises many critical concerns for human health, drinking water quality (Gray, 2008),  
65 coastal and marine water degradation causing eutrophication, as well as algal blooms and hypoxia  
66 (Conley et al., 2009; Rabalais et al., 2010). **There are natural sinks for Nr along the N cascade**

67 (Galloway et al, 2003; Tanner et al., 2005) but other sinks that are rehabilitated or artificial, may  
68 also be introduced and evaluated at key locations e.g. constructed wetlands (CWs) (Gold et al.,  
69 2013) or permeable reactive interceptors (Fenton et al., 2014). Many investigations target single  
70 contaminant remediation, whilst disregarding the reality of mixed contaminants entering and  
71 leaving such systems. They do not consider the dynamic of pollution swapping (the increase in  
72 one pollutant as a result of a measure introduced to reduce a different pollutant) driven by  
73 transformational processes within and around the system. This review focuses on CWs which are  
74 now an established technology used globally as a mitigation option for nutrient retention at the  
75 delivery end of the transfer continuum (Dunne, et al., 2005; Sjøvik et al., 2006; Tanner and  
76 Kadlec, 2013; Tanner and Sukias, 2011). They have been successfully used for treating domestic  
77 sewage, urban runoff and storm water, industrial and agricultural wastewater, and leachate  
78 (Scholz and Lee, 2005; Gill et al., 2014). In addition, CWs provide multi-faceted ecosystem  
79 services that include both socio-economic and ecological aspects (Turner et al., 2008; Milder et  
80 al., 2010; Lei et al., 2011). Although CWs have a proven potential for the removal of organic  
81 matter and N, studies have shown that removal efficiencies have been highly variable, due to  
82 inadequate observations of N transformations and their removal mechanisms (Seitzinger et al.,  
83 2002) and lack of improved adoption and placement (Gold et al., 2013). The removal efficiencies  
84 can vary with the types of CWs (i.e. surface flow and subsurface flow) and with types of plants  
85 grown. CWs are complex bioreactors which facilitate a number of physical, chemical and  
86 biological processes but often continue to be evaluated as a 'black box' in terms of process  
87 understanding (Langergraber, 2008). The removal efficiencies can vary with the types of CWs  
88 (i.e. surface flow and subsurface flow) and with types of plants grown.

89  
90 Denitrification is the main mechanism in CWs to reduce nitrate ( $\text{NO}_3^-$ -N) but its intermediate  
91 product, nitrous oxide ( $\text{N}_2\text{O}$ ), is a potent greenhouse gas (GHG), causing 6% of anthropogenic  
92 global warming as well as being the leading cause of stratospheric ozone depletion (Beaulieu et  
93 al., 2011). It is known that CWs significantly contribute to atmospheric  $\text{N}_2\text{O}$  emissions either  
94 directly to the atmosphere from the surface of the wetland as its byproduct (IPCC, 2014; Sjøvik et  
95 al., 2006; Ström et al., 2007; Elberling et al., 20011; van der Zaag et al., 2010) or indirectly via  
96 dissolved  $\text{N}_2\text{O}$  in the effluent or groundwater upon discharge to surface waters. The global  
97 importance of this  $\text{N}_2\text{O}$  emission has been increasingly recognised (IPCC, 2014). Indirect  $\text{N}_2\text{O}$   
98 emissions from CWs effluents that discharge directly to aquatic environments have a default  
99 emission factor (EF) of 0.005 but with a wide range (0.0005-0.25). The wide range of EF is  
100 highlighting the large uncertainty, which warrants further refinement. Indeed, the production and

101 reduction processes of N<sub>2</sub>O in the environment are not yet fully understood (Butterbatch-Bahl et  
102 al., 2014).

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104 In CWs soil organic N is degraded releasing ammoniacal-N which further undergoes  
105 transformation to other forms. Ammonium can accumulate in CWs due to limited oxidation of  
106 NH<sub>4</sub><sup>+</sup> in the reduced environment. For example, questions arise with respect to NH<sub>4</sub><sup>+</sup>  
107 concentrations in groundwater underlying the CWs (often higher in groundwater than the effluent  
108 e.g. Harrington et al., 2007; Dzakpasu et al., 2012). Dissimilatory NO<sub>3</sub><sup>-</sup> reduction to ammonium  
109 (DNRA) occurring in wetland soils can contribute to ammonium (NH<sub>4</sub><sup>+</sup>) pollution (Burgin et al.,  
110 2013). Mass balance analysis of the different components of the N cycle and kinetics of the N  
111 transformation processes occurring within the treatment cells using the isotope-tracing <sup>15</sup>N  
112 technique can provide mechanistic information for N transformation products (Lee et al., 2009;  
113 O’Luanaigh et al., 2010) and therefore can be used to start to answer such questions. Huygens, et  
114 al. (2013) described two complementary stable isotope methods that can be used to determine N  
115 cycle processes in CWs: i) the isotope pairing and revised isotope pairing techniques which  
116 calculate total N<sub>2</sub> production by denitrification and anaerobic ammonium oxidation (anammox)  
117 and ii) isotope dilution and tracing techniques which assess gross N transformation in wetland  
118 soils. Denitrification in porewater samples can be measured by analysing samples for dissolved  
119 N<sub>2</sub> in a membrane inlet mass spectrometer (MIMS; Kana et al., 1994) and N<sub>2</sub>O in gas  
120 chromatograph (GC; Jahangir et al., 2012). The studies of natural abundance of <sup>15</sup>N and <sup>18</sup>O (δ<sup>15</sup>N  
121 and δ<sup>18</sup>O) in NO<sub>3</sub><sup>-</sup> is an insightful tool for the investigation of the sources, fate and  
122 transformational processes of N in a system (e.g. in shallow groundwater, Baily et al., 2011).  
123 However, this analysis cannot in itself explain whether the denitrification process has occurred *in*  
124 *situ* or if the gases have been transported from other soil zones to groundwater. A <sup>15</sup>N based  
125 method could be developed to measure *in situ* N cycle processes and to understand the movement  
126 (both upward and downward) of the N cycle products in wetland sediments. This would involve  
127 incubation of intact laboratory soil columns from the site in question enriched with <sup>15</sup>NH<sub>4</sub><sup>+</sup><sup>14</sup>NO<sub>3</sub><sup>-</sup>  
128 and <sup>14</sup>NH<sub>4</sub><sup>+</sup><sup>15</sup>NO<sub>3</sub><sup>-</sup> kept at average site temperature (Huygens et al., 2013). The *in situ*  
129 denitrification, DNRA and C dynamics of groundwater below the CWs can be elucidated in the  
130 field with shallow piezometers or boreholes using push-pull investigations (Addy et al., 2002;  
131 Jahangir et al., 2013; Burgin et al., 2013). These methods in combination will be able to provide a  
132 comparative analysis of the rates of N transformation processes occurring in CWs. For example,  
133 *in situ* quantification of the rates of NH<sub>4</sub><sup>+</sup> oxidation and denitrification in CWs will give insights  
134 into their role in delivering NO<sub>3</sub><sup>-</sup> and or NH<sub>4</sub><sup>+</sup> to groundwater and N<sub>2</sub>O to the atmosphere.

135 CWs can either be sources or sinks of C (Whiting and Chanton, 2001) but only a few studies have  
136 considered carbon dioxide (CO<sub>2</sub>) efflux from CWs (Mander et al., 2008). CWs are characterised  
137 by anaerobic conditions with low organic matter turnover (Augustin et al., 1998) but they also  
138 release C back into the atmosphere (Clair et al., 2002). Wetlands can contribute to the dissolved  
139 organic carbon (DOC) load transfer to ground and surface waters and can exchange substantial  
140 amounts of CO<sub>2</sub> and methane (CH<sub>4</sub>) with the atmosphere (Clair et al., 2002; Elberling et al.,  
141 2011). CO<sub>2</sub> is lost by plant respiration and aerobic organic matter decomposition, whilst CH<sub>4</sub> is  
142 lost by methanogenic activity during anaerobic decomposition of DOC. Methane is emitted from  
143 wetlands due to anoxic or suboxic conditions that occur in the soils (Ström et al., 2007; Søvik et  
144 al., 2006; Pangala et al., 2010). In addition to the positive aspects highlighted previously, CWs  
145 can have negative climatic impacts (Ström et al., 2007), enhance pollution swapping (Reay,  
146 2004) and act as an active source of GHGs (IPCC, 2014; Clair et al., 2002; Mander et al., 2008;  
147 Mitsch and Gosselink, 2000). For example, van der Zaag et al. (2010) measured CH<sub>4</sub>-C emissions  
148 as 0.2- 27% of the total C removed and N<sub>2</sub>O emissions as 0.1- 1.16% of the total N removed in  
149 CWs. Estimation of GHG emissions from different CWs management systems, especially in  
150 response to a changing climate, is poorly known. Consequently, the improved understanding of  
151 integrated C and N dynamics is required to improve the CW management to optimise their  
152 environmental benefits.

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## 154 2. Removal Efficiency, Hydraulic Loading and Retention Time

155 In CWs used for wastewater treatment, the efficiency of OM removal is often satisfactory, but N  
156 removal is generally limited. The limited removal can be caused by the typically reduced  
157 environment that enhances NH<sub>4</sub><sup>+</sup> accumulation and limits NH<sub>4</sub><sup>+</sup> oxidation, in addition to the lower  
158 denitrification rates due to lack of NO<sub>3</sub><sup>-</sup> and N recycling by vegetation. On average, it appears  
159 that 50% of the influent N can be removed but the removal rates are very inconsistent. A  
160 summary of influent composition and their removal efficiencies is given in Table 1. In European  
161 systems, for example, typical removal percentages of ammoniacal-N in long-term operation are c.  
162 35% but can be enhanced if some pre-treatment procedures are followed (Verhoeven, and  
163 Meuleman, 1999; Luederitz et al., 2001). O’Luanaigh et al. (2010) in a subsurface CW showed  
164 41% N removal across a tertiary treatment system (receiving secondary treated effluent) with  
165 little evidence of any seasonal change. In a single house domestic wastewater treatment using  
166 CWs, Gill et al. (2012) has shown mean removal efficiencies of 33-36% of total N (TN) but did  
167 not investigate the N cycling processes. To enhance such a low removal, the nature of N  
168 transformations and the fate of the removed N within the CWs should be investigated. However,

169 to our knowledge, no study is available that provides a holistic evaluation of N attenuation and  
170 transformations to other reactive N as losses to the atmosphere ( $\text{N}_2\text{O}$ ,  $\text{NH}_3$ ,  $\text{N}_2$ ) and receiving  
171 surface waters or groundwater ( $\text{NO}_3^-$ ,  $\text{NH}_4^+$ ).

172

173 The removal efficiency of pollutants in CWs depends on hydraulic loading rates (HLR) and  
174 hydraulic retention time (HRT) (Toet et al., 2005). **The effects of HLR and HRT can vary with  
175 the nature of the use of CWs e.g. whether they are used for treating single or mixed contaminants.  
176 Therefore, to have the least amount of Nr delivery to the receiving waters or to the atmosphere,  
177 CWs need to be optimally designed with respect to HLR and HRT.** Investigation into the effects  
178 of fluctuating hydraulic loadings (hydraulic pulsing) on N removal efficiency and its  
179 transformation products will provide information about the potential of pollution swapping for  
180  $\text{NO}_3^-$ ,  $\text{NO}_2^-$ ,  $\text{NH}_3$ ,  $\text{NH}_4^+$  and  $\text{N}_2\text{O}$ . If the dominant product is  $\text{N}_2$ , the system will be more benign  
181 to the environment than many of the other potential end products. If N is transformed into  $\text{NH}_4^+$  it  
182 should be fixed in the soils and subsoils but there is also the potential for transport to ground and  
183 surface waters connected to CWs if the cation exchange sites become saturated. **However,  
184 contrasting results on the impacts of hydraulic loadings on nutrients removal efficiency are  
185 available in literature. Luo et al. (2005) reported that low HLR results in incomplete  
186 denitrification, whereas Zhang et al. (2006) argued that low HLR increases  $\text{NH}_4^+$  and chemical  
187 oxygen demand (COD) removals by 25 and 11%, respectively. Several authors have used a wide  
188 range of HLR and HRT to measure nutrient removal efficiency but experimental evidence is still  
189 scarce (Toet et al., 2005). For example, Zhou and Hosomi (2008) constructed a surface flow  
190 wetland with a HLR of  $0.20 \text{ m}^3 \text{ m}^{-2} \text{ d}^{-1}$  and an average water depth of 0.056 m. The authors  
191 estimated the N removal efficiency in terms of a comparison between N inputs and outputs but  
192 did not quantify how this rate corresponded to the different forms of the removed N. Similarly,  
193 Song et al. (2011) reported an average HRT of 2.4 days during wetland operation periods for the  
194 removal of  $\text{NO}_3^-$  and  $\text{NH}_4^+$  without documented design optima or pollution swapping. The way in  
195 which a CW performance is assessed can lead to different conclusions regards removal of Nr. For  
196 future studies, evaluation of systems in a holistic manner, which include pollution swapping at  
197 different HLRs and HRTs is important, particularly within the context of the changing  
198 hydrological cycle in a changing climate. Also local legislative targets should be considered and  
199 weighting factors (are GHGs more important than water quality targets) developed for different  
200 species within the system to come up with overall balances for the CW. Thus far, estimates of  
201 nutrient removal efficiencies are based on the subtraction of the concentrations of nutrients in the  
202 effluent from the influent, but the transformation kinetics of the removed nutrients are unknown.**

203 A holistic assessment of ecologically engineered systems has been thoroughly outlined in Healy  
204 et al. (2011) and developed further in Fenton et al. (2014). Such assessments can be applied in  
205 evaluating nutrient dynamics in CWs. In a subsurface CW, O’Luanaigh et al. (2010) used <sup>15</sup>N  
206 isotopes to show that N spiralling occurred between NH<sub>4</sub><sup>+</sup> to organic N and back to NH<sub>4</sub><sup>+</sup> within  
207 the CW. So, studying the dynamics of N in CW systems is crucial because the forms of removed  
208 N are particularly of concern with respect to their potential for pollution swapping; global  
209 warming and water pollution. In addition to the estimation of nutrient removal rates,  
210 investigations into the effect of HLR and HRT on the different forms of nutrients in the final  
211 effluents and their fate in the natural environment are needed to address the pollution swapping  
212 potential of CWs. In addition, the Nr delivered to groundwater can be transformed *in situ* to other  
213 reactive or benign forms. Hence, quantification of such Nr loadings to groundwater and their *in*  
214 *situ* consumption (e.g. N<sub>2</sub>O to N<sub>2</sub>) is required to understand their environmental consequences.

215

### 216 3. Accumulation of C and N in CWs Soils

217 The soil in CWs is a major sink for N. Data on the N accumulation (dissolved organic N- DON,  
218 TN, NH<sub>4</sub><sup>+</sup> or NO<sub>3</sub><sup>-</sup>-N) within the soil profile of various CWs are scarce as most studies focus on  
219 N balances between influent and effluent N loads. The accumulation of N could be a substantial  
220 part of the total N added to CWs; 30-40% (Shamir et al., 2001), 39% (Harrington et al., 2007),  
221 9% (Mander et al., 2008) and 2.5% (Obarska-Pempkowiak and Gajewska, 2003). The wider  
222 range of N accumulation could be due to variations in CW types and their management. The  
223 accumulated species of N are rather reactive unless they are transformed to N<sub>2</sub> by biogeochemical  
224 processes. For example, Shamir et al. (2001) measured about 80% of the total accumulated N as  
225 organic and the remainder as NH<sub>4</sub><sup>+</sup>. The organic N could be mineralised to NH<sub>4</sub><sup>+</sup> and NO<sub>3</sub><sup>-</sup>,  
226 reactive forms of N, depending on the physico-chemical properties of soil. The Nr could be  
227 assimilated by plants and microbes living in CWs, which are recycled in a soil-plant-soil  
228 continuum. Typically, N accumulation has been found to decrease with soil depth: for example,  
229 Shamir et al. (2001) reported mean NH<sub>4</sub><sup>+</sup> concentrations in 0-0.15, 0.15-0.30 and 0.30-0.60 cm  
230 depths of 156, 151, and 28 mg N kg<sup>-1</sup>, respectively. However, dissolved nutrients can be  
231 preferentially leached down into deeper soil layers via different pathways e.g. root channels. In  
232 terms of the input-output balance, these are considered as removed N, but they remain in such a  
233 biogeochemically active system. In addition to N, organic C accumulation occurs in CW soils. In  
234 a gravel-bed CW, Nguyen (2000) measured 17.5, 16.2 and 3.6% of total C (TC) accumulation at  
235 surface, 0-10 and 10-40 cm depth. As such, CWs represent organic C and Nr rich systems where  
236 the products of the continuously occurring biogeochemical processes can be transported to fresh

237 waters and to the atmosphere. The  $\text{NO}_3^-$  could be denitrified to  $\text{N}_2\text{O}$  or  $\text{N}_2$  and  $\text{NH}_4^+$  could be  
238 mineralised to  $\text{NO}_3^-$  or fixed in soil matrix. Estimation of the reaction kinetics of these processes  
239 is required to accurately measure the fate of the added nutrients in CWs. With methodological  
240 advancement, measurement of these processes is now possible. For example, application of  $^{15}\text{N}$   
241 tracer technique ( $^{15}\text{NO}_3^-$ ) can give insights into the  $\text{NH}_4^+$  oxidation and fixation capacity of  
242 sediments below CWs. Estimating the rates of nutrient accumulation in soil and subsoils and their  
243 *in situ* transformation kinetics in various types of CWs is important. Elucidating the fate of these  
244 added nutrients will help to reduce their potential for pollution swapping which has been ignored  
245 so far in managing CWs. The stability of the accumulated C and N under changing climatic  
246 scenarios also needs to be addressed to consider the long term sustainability of CWs.

247

#### 248 4. C and N Dynamics and Greenhouse Gas Emissions

249 Processes involved in N removal and N transformations in wetlands include sedimentation of  
250 particulates (Koskiaho, 2003); nitrification, denitrification and DNRA (Poach et al., 2003; Burgin  
251 et al., 2013), microbial assimilation and plant uptake-release (Findlay et al., 2003), anammox and  
252 DEAMOX (DEnitrifying AMmonium OXidation). Müller et al. (2014) developed a  $^{15}\text{N}$  tracing  
253 model, which is able to identify four different pathways of  $\text{NO}_2^-$  reduction to  $\text{N}_2\text{O}$ : i) reduction of  
254  $\text{NO}_2^-$  associated with nitrification, ii) reduction of  $\text{NO}_2^-$  associated with denitrification, iii)  
255 reduction of  $\text{NO}_2^-$  associated with organic N oxidation, and iv) co-denitrification, a hybrid  
256 reaction where one N atom in  $\text{NO}_2^-$  originates from organic N and the other from  $\text{NO}_2^-$  reduction  
257 via denitrification. Most of these pathways transfer Nr (mainly  $\text{NH}_4^+$  and  $\text{N}_2\text{O}$ ) to the  
258 environment, however, some pathways can convert Nr to  $\text{N}_2$  (e.g. denitrification, anammox and  
259 DEAMOX). During denitrification,  $\text{NO}_3^-$  is used as a terminal electron acceptor to produce  $\text{N}_2$  or  
260  $\text{N}_2\text{O}$  (Starr and Gillham, 1993). Anammox can remove  $\text{NO}_2^-$  and  $\text{NH}_4^+$  as  $\text{N}_2$  in CWs when the  
261 existing environment is hypoxic. The DEAMOX can remove  $\text{NO}_3^-$  and  $\text{NH}_4^+$  as  $\text{N}_2$  where  $\text{NO}_3^-$  is  
262 converted to  $\text{NO}_2^-$  by autotrophic denitrification with sulphide (Kalyuznyi et al., 2006).  
263 Denitrification has been estimated to be a significant N removal process but actual quantification  
264 data are scarce. Limited studies have estimated N losses by denitrification e.g. 19% (Mander et  
265 al., 2008) and 86% (Obarska-Pempkowiak and Gajewska, 2003) of the total N input based on the  
266 mass balance study. To our knowledge, no data are available on denitrification measurements in  
267 soil/subsoils of surface flow CWs.

268



269 In addition, the two other processes that can remove Nr from the CWs (anammox and  
270 DEAMOX) are not well understood and so it is crucial to identify which of the processes are  
271 occurring in a specific type of CW and at what rate they occur. Once a process that provides N<sub>2</sub>  
272 as the end product is determined in a specific system then the CW management could be directed  
273 towards enhancement of that process. Hence, quantifying the rates of these processes under  
274 various CW types is required for improved N management towards lowering Nr in the  
275 environment. There is large uncertainty about the processes involved and their magnitude. Many  
276 studies focus on the mass balance approach and use the difference as the rate of N attenuation.  
277 N<sub>2</sub>O emissions from CWs have been reported to be high (EF 0.005; IPCC, 2014) but there is  
278 large uncertainty around these values due to the limited scope of the research that has been  
279 carried out to quantify such processes. The enhanced reduction of N<sub>2</sub>O to N<sub>2</sub> needs further  
280 elucidation.

281

282 Similarly, C transformations involve respiration, fermentation, methanogenesis, CH<sub>4</sub> oxidation  
283 and reduction of S, Fe and NO<sub>3</sub><sup>-</sup>. Anaerobic methane oxidation coupled with denitrification, a  
284 recently proposed pathway of the C cycle (á Norði and Thamdrup, 2014; Haroon et al., 2013;  
285 Islas-Lima et al., 2004), can reduce CH<sub>4</sub> emissions in CWs. The C removal processes are  
286 sedimentation, microbial assimilation, gaseous emissions, dissolved C losses through water to  
287 ground and surface water bodies and chemical fixation (bonding with chemical ions). As CWs  
288 are designed to remove pollutants in an anaerobic/suboxic environment, they change the C and N  
289 dynamics and contribute significantly to CH<sub>4</sub> and N<sub>2</sub>O emissions (Johansson et al., 2002, 2003;  
290 Mander et al., 2005, 2008; Stadmark and Leonardson, 2005; Liikanen et al., 2006). The  
291 hydrological, chemical and microbial processes of CWs are likely to be different from the  
292 processes occurring in more natural wetlands, because they receive nutrient rich waters from  
293 various sources. Increased nutrients and organics inputs will increase the productivity of wetland  
294 ecosystems and increase the production of GHGs. Søvic et al. (2006) measured N<sub>2</sub>O, CH<sub>4</sub> and  
295 CO<sub>2</sub> emissions in various CWs in different European countries and suggested that the potential  
296 atmospheric impacts of CWs should be examined as their development is increasing globally.  
297 Ström et al. (2007) recommended that CW management processes must consider the negative  
298 climatic aspects of increased emissions of GHGs in addition to their primary functions.  
299 Therefore, estimation of the contribution of CWs to global warming is required. Modelling or up-  
300 scaling of GHG emissions at national and regional scales to a global scale is important to improve  
301 global GHG budgets. The use of CWs to improve water quality is likely to increase to meet

302 policy drivers such as the European Union Water Framework Directive and so the uncertainty of  
303 GHG budgets needs to be reduced to prevent pollution swapping.

304

305 In this regard, measurement of spatial and temporal variations (seasonal and diurnal) of GHG  
306 emissions is required. Moreover, plant mediated GHG emissions could be an important  
307 component of total emissions but research in this area is very limited. The GHG from CWs can  
308 vary with vegetated and non-vegetated systems. Vegetation and its composition affect the  
309 nutrient dynamics and the production, consumption and transport of greenhouse gases and hence  
310 their exchange between wetlands and atmosphere (Ström et al., 2003, 2005; Sövic et al., 2006;  
311 Johansson et al., 2003). Emergent plants can transport atmospheric O<sub>2</sub> to rooting zone and  
312 contribute to C and N dynamics in wetland soils e.g. N<sub>2</sub>O and CO<sub>2</sub> production and CH<sub>4</sub>  
313 consumption (Brix, 1997). Vascular plants can exchange GHGs between the rooting zone and  
314 atmosphere (Yavitt and Knapp, 1998). Emissions of N<sub>2</sub>O and CH<sub>4</sub> in CWs can vary across CW  
315 typologies e.g. surface flow or subsurface flow (Van der Zaag et al., 2010). Therefore, the  
316 assessment of GHG emissions in various types of CW (surface flow, subsurface flow; vertical  
317 and horizontal), under different management systems (vegetated, nonvegetated, plant species  
318 composition) and usage (municipal waste water, agricultural runoff, landfill leachate) is  
319 necessary in light of the national and global GHG budgets and mitigation of GHG emissions. In  
320 addition, such measurements will help scientists, environmental managers and policy makers to  
321 adopt environmental friendly construction and management of CWs. Assessment of the reactive  
322 versus the benign forms of C and N transformation products in various types of CWs will give  
323 insights into their environmental friendly design and management.

324

## 325 5. Surface Emissions vs. Subsurface Export of C and N

326 Dissolved GHGs in porewater in wetland soils and subsoils can be emitted to the atmosphere by  
327 transpiration of vascular plants (from within the rooting zone) and via groundwater, upon  
328 discharge to the surface waters. Dissolved GHGs in groundwater can flow towards surface waters  
329 by advective transport and or by dispersion of groundwater. The GHGs produced in soils/subsoils  
330 in CWs can also be emitted to the atmosphere by ebullition and diffusion. Elberling et al. (2011)  
331 reported that in wetlands, the transport of subsurface soil gases occurs both via diffusive transport  
332 in the pores and through the vascular plants. In addition to measuring the surface emissions,  
333 Ström et al. (2007) also measured a considerable quantity of CH<sub>4</sub> in porewater and found a  
334 significant correlation ( $p = 0.021$ ) between the surface emissions and porewater CH<sub>4</sub>  
335 concentrations in vegetated wetlands. Measuring porewater GHGs and linking these to the

336 surface emissions and subsurface export to groundwater below CWs will help to estimate a better  
337 GHG balance from both a national and global context. Elberling et al. (2011) linked subsurface  
338 gas concentrations in wetlands to the surface fluxes using a diffusion model which has  
339 demonstrated the need for future studies on subsurface GHG production, consumption and net  
340 GHG emissions in CWs ecosystem in a climate change context. It is important to characterise  
341 soils and subsoils physical and hydraulic properties and to assess their potential to percolate  
342 dissolved nutrients and gases to the underlying groundwater. To our knowledge, this indirect  
343 pathway of GHG emissions from CWs has never been reported despite the fact that this would  
344 appear to have a high biogeochemical potential to produce and exchange GHGs. The balance  
345 between N and C input and output flows between CWs and aquatic and atmospheric  
346 environments together with the direct and indirect emissions of C and N species could be an  
347 important input to global C and N budgets.

348

## 349 6. Hydrogeochemistry below CWs

350 CWs can be designed with or without a clay liner or a compacted soil bed at the base, which can  
351 lead to large differences in permeability of the underlying layers. The variation in permeability of  
352 a CW soil bed will affect solute, nutrient and GHG flows and their interactions with the  
353 underlying groundwater (Dzakpasu et al., 2012; 2014). Groundwater hydrogeochemistry below  
354 CWs can therefore provide a unique insight into such interactions. An example of such  
355 interactions would be between nutrient rich water discharging from CW cells mixing with  
356 laterally moving regional groundwater. It should be noted that groundwater can also discharge  
357 into CWs depending on the hydraulic gradients. This necessitates that multiple multi-level  
358 piezometers or boreholes are installed at such sites to elucidate groundwater flow direction,  
359 hydraulic gradients and hydraulic permeability. Such monitoring networks allow water samples  
360 to be taken and then sources of nutrients in the groundwater body below CWs can be identified.  
361 For example, natural abundance of N ( $\delta^{15}\text{N}$ ) and oxygen ( $\delta^{18}\text{O}$ ) can be used to identify the  
362 sources of  $\text{NO}_3^-$  and infer transformational processes responsible for concentrations in  
363 groundwater (Baily et al., 2011). The local site hydrology (precipitation, groundwater table  
364 fluctuations and evapotranspiration) has greater impacts on the pollutant removal by physical  
365 attenuation and by biochemical transformations. Hydrogeochemical studies on an accurate spatial  
366 and temporal resolution should explain the effects of precipitation on nutrient removal by dilution  
367 as well in situ nutrient turnover. Despite conditions within CWs being saturated throughout the  
368 year, changes in the water table elevation affect the oxygen ( $\text{O}_2$ ) concentrations in CWs which is

369 key parameter for the biogeochemistry of soils and subsoils. For example, lowering the water  
370 level increases the O<sub>2</sub> concentrations and accelerates soil OM decomposition, thereby increasing  
371 CH<sub>4</sub> oxidation and CO<sub>2</sub> emissions. Highly contrasting results on gas emissions with fluctuating  
372 water levels have been reported and the controlling mechanisms are unclear (Elberling et al.,  
373 2011). Effective CW management requires an understanding of the effects of wetland hydrology  
374 on the physical and biochemical attenuation of nutrients in order to assess their impacts on the  
375 surface emissions and subsurface export of nutrients and GHGs. Data on the species of N in  
376 groundwater below the CWs are required to provide an in-depth understanding of wetland  
377 ecosystem services, particularly if CW systems have the potential to leak pollutants down into the  
378 groundwater (Dzakpasu et al., 2014). A study on a clay lined CW system reported that the  
379 nutrient content in the underlying groundwater to be NH<sub>4</sub><sup>+</sup>-N 4.0 mg L<sup>-1</sup>, NO<sub>3</sub><sup>-</sup>-N 0.2 mg L<sup>-1</sup> and  
380 molybdenum reactive phosphorus (MRP) <0.01mg L<sup>-1</sup>. The elevated NH<sub>4</sub><sup>+</sup>-N indicates transport  
381 of nutrient from the CW to groundwater through leakage or *in situ* N transformations such as  
382 DNRA. However, the impact of CWs on the local underlying groundwater quality has been  
383 seldom assessed. Linking geochemistry of groundwater below CWs to site hydrology, water table  
384 fluctuations and soil/subsoils physico-chemical properties is required to elucidate the major  
385 environmental drivers of C and N removal and or pollution swapping.

386

## 387 7. Methodological Developments

388 Quantification of the occurrence and magnitude of N transformations and hydrochemical  
389 properties is crucial to improve the assessment of CW ecosystem services and to minimize their  
390 potential for pollution swapping. With the recent advancement of isotope pairing and dilution  
391 techniques, single or simultaneously occurring N transformation processes in CWs can be  
392 quantified in laboratory or *in situ* conditions (Huygens et al., 2013; Müller et al., 2014). The  
393 isotope technique relies on the introduction of a known amount of <sup>15</sup>N into the CW system and  
394 then quantification of N concentrations and isotopic compositions through different N pools after  
395 incubation for a specific period. Laboratory methods involve collection of intact soil/sediment  
396 cores, with subsequent incubation in the laboratory. The *in situ* field techniques involve release of  
397 <sup>15</sup>N solution *in situ* in the CWs soils. The *in situ* stable isotope techniques may be an appropriate  
398 tool for the determination of simultaneously occurring N transformation processes in wetland  
399 soils (Huygens et al, 2013).

400

401 Details of the application of isotope pairing/ revised isotope pairing techniques and isotope  
402 dilution/ tracing techniques have been presented elsewhere (Huygens et al., 2013). In brief,

403 incubation of intact soil cores with differentially labelled  $^{15}\text{NH}_4^{14}\text{NO}_3$  and  $^{14}\text{NH}_4^{15}\text{NO}_3$  can be  
404 used to quantify the rates of different N transformation processes ((Rütting and Müller, 2008).  
405 The quantification of simultaneously occurring N transformation rates rely on the analysis with  
406 appropriate  $^{15}\text{N}$  tracing models. Development in the recent years in  $^{15}\text{N}$  tracing techniques is now  
407 available, which are able to identify process specific  $\text{NO}_2^-$  pools (Rütting and Müller, 2008),  
408 pathways specific  $\text{N}_2\text{O}$  production and emission as well as  $\text{N}_2\text{O}/\text{N}_2$  ratios (Müller et al., 2014).  
409 Traditional techniques for investigation of gross N dynamics in sediments (Blackburn, 1979)  
410 should be combined with the latest  $^{15}\text{N}$  tracing techniques where all N transformation rates are  
411 included that are important in wetlands and under anoxic condition (Huygens et al., 2013). Thus,  
412 current models should consider processes such as anammox and/or DEAMOX and then be tested  
413 in CW environments under various conditions. The *in situ*  $\text{NO}_3^-$  push-pull method has been used  
414 to determine denitrification in shallow groundwater (<3 m) in riparian wetlands (Addy et al.,  
415 2002; Kellogg et al., 2005) and in deep groundwater in arable/grassland (Jahangir et al., 2013).  
416 This method can be used in CW sites using piezometers or boreholes screened at different  
417 intervals to investigate *in situ* C and N dynamics. In this method the parameters of importance  
418 to be analysed are  $^{15}\text{N}_2\text{O}$ ,  $^{15}\text{N}_2$ ,  $^{15}\text{NO}_3^-$ , and  $^{15}\text{NH}_4^+$ . In addition measurements of DOC and gases  
419 ( $\text{CO}_2$  and  $\text{CH}_4$ ) will provide insights into the C consumption and transformation associated with  
420 the N transformations.

421

422 Carbon and N dynamics are influenced by the interacting effects of soil conditions with microbial  
423 community structure and functioning. Occurrence of C and N cycling processes are controlled by  
424 the environment which involves transcription of genes, translation of messenger RNA and  
425 activity of enzymes (Firestone et al., 2012). As such, activities of microbial communities under  
426 various environmental conditions and how these contribute to C and N dynamics is a very  
427 important area of future research (Müller and Clough, 2013). Molecular approaches can be  
428 important tools to identifying and quantifying the genes that code for enzymes mediating C and N  
429 cycles (Peterson et al., 2012). These tools help assess the relationships among genes,  
430 environmental controllers and the rates of C and N processes.

431

## 432 **8. Conclusions and Recommendations**

433 The transformational processes on a mixture of contaminants within and below CWs can cause  
434 pollution swapping. A holistic assessment of C and N dynamics in CWs is needed to fully  
435 understand their removal, transport and impact on water quality and emissions to the atmosphere.

436 Mixed contaminants entering CWs and those formed within and underneath CWs during  
437 transformational processes must be considered in future studies. The overall balance of these  
438 constituents will decide on whether a CW is a pollution source or a sink. This will necessitate a  
439 higher degree of multi-level spatial and temporal monitoring and use of multi-disciplinary  
440 techniques both *in* and *ex situ* to fully characterise all pathways of C and N loss.

441

442 Leakages of nutrients and gases from CWs to groundwater can only be elucidated through the  
443 techniques itemised herein and a better understanding of the transformation processes along a  
444 **vertical geochemical gradient** from the CW into the underlying aquifer. The reactive versus the  
445 benign forms of the N transformation products should be evaluated in various CWs. An  
446 understanding of how N removal occurs and how losses of N and associated gases impact water  
447 and air quality is required. Denitrification, DEAMOX and anammox are the processes which  
448 remove N to its benign form ( $N_2$ ) and all other processes produce only reactive forms of N. Data  
449 on **when, where and at what rates denitrification, DEAMOX and anammox occur in CWs are**  
450 **needed as well as what are the key factors that control such processes. The provenance of  $NH_4^+$  in**  
451 **groundwater below CW cells needs clarification and its impact on down gradient receptors.**

452

453 **This review shows that CWs have the potential to produce  $N_2O$ , DON, DOC, DIC,  $CO_2$  and  $CH_4$**   
454 **and that the GHGs produced in CWs can be emitted to the atmosphere. They can also be exported**  
455 **to fresh waters via groundwater and degassed upon discharge to surface waters. Moreover, the**  
456 **DOC and DIC transferred to the fresh water sediments (rivers and lakes) can produce GHGs that**  
457 **in turn get emitted to the atmosphere. Therefore, it is important to know the concentrations of**  
458 **dissolved C and N species and the rates of production and consumption of GHGs in groundwater**  
459 **below CWs. The amount of C and N exported from terrestrial ecosystem via the subsurface**  
460 **pathway to fresh waters has been the missing piece of our understanding of global C and N**  
461 **budgets. It is clear that data on the various C and N species along with the GHGs in various CWs**  
462 **systems are thus crucial to make a robust input-output balance of C and N in such a rising and**  
463 **engineered ecosystem. Spatial and temporal variation in  $N_2O$  emissions in CWs under different**  
464 **management systems is critical to get much more rigorous estimates of  $N_2O$  emission factors.**  
465 **These data will bring down the existing uncertainties in global C and N budgets.**

466

467 **Managing wetting and drying spells (pulsing hydrology) in CWs can enhance mineralisation of**  
468 **organic N and oxidation of  $NH_4^+$  to  $NO_2^-/NO_3^-$  and then denitrification/anammox could transform**  
469 **these species to benign  $N_2$ . This requires more research into the N cycle processes over the**

470 wetting drying spells which is now possible with the advancement in  $^{15}\text{N}$  tracing and modelling  
471 techniques. With the recent advancement of isotope pairing and dilution techniques single or  
472 simultaneously occurring N transformation processes can be quantified. The isotope based  
473 techniques can also be extended to other elements e.g., a  $^{33}\text{P}$  tracing model has been developed  
474 recently to study phosphorous cycle in soil (Müller and Bünemann, 2014). Further reducing the  
475 saturated hydraulic conductivity below the wetland bed will help reduce nutrients leaching to  
476 groundwater below the CWs cells. The selection of plant species is important to increase  
477 nutrients removal, sequester more C and decrease greenhouse gas emissions which is an area that  
478 still requires more research across types of CWs and countries. More research is also still needed  
479 to be done on the impacts of hydraulic retention time on nutrients dynamics and removal.  
480 Subsurface export of nutrients and GHGs to groundwater through leachates, preferential flow  
481 paths created by dead roots and holes of burrowing invertebrates and subsequent transport to  
482 surface water bodies should be accounted in CW management. Rates of nutrient accumulation or  
483 fixation in soils and their *in situ* transformation in CWs need to be quantified to know their  
484 contribution to C sequestration fate in the environment. Estimation of GHG production in CWs,  
485 their consumption and emissions to the atmosphere in a changing climate is urgent. The  
486 ecosystems services from CW will make them an attractive option for water pollution mitigation  
487 while providing an important aquatic habitat, but their integrated impact on environmental quality  
488 needs to be addressed in a holistic manner to avoid unwanted side effects.

489

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493

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735 Table 1 Summary of influent composition and their removal efficiencies being reported by  
 736 several authors

CWs type	NO <sub>3</sub> <sup>-</sup> (mg L <sup>-1</sup> )	NH <sub>4</sub> <sup>+</sup> (mg L <sup>-1</sup> )	BOD <sub>5</sub> (mg L <sup>-1</sup> )	MRP (mg L <sup>-1</sup> )	TN Removal (%)	References
SF	15.5	4.5	1.5 (DOC)	0.2	61	Song et al., 2011
SF	5	40			97- 98	Dzakpasu et al., 2011
SF	4.4	39			39-48	Vymazal, 2007
SF		36	170	9.6	39	Vymazal, 2010
SF	2.6	64	816	10	>90	DEHLG, 2010*
SF	6.4	34	854	4.3	>90	DEHLG, 2010**
SF	<2	196	277		35	Shamir et al., 2001
SF	<1.0	80	1200	25	>90	Harrington et al., 2007
SF	1.54	0.95	32 (COD)	1.43	45	Toet et al., 2005
SSF	3.9	74.9	514 (COD)	11.1	29	O’Luanaigh et al., 2010
SSF		87 (TN)	167		46-48	Mander et al., 2008

737 SF-surface flow; SSF-subsurface flow; NO<sub>3</sub><sup>-</sup>- nitrate; NH<sub>4</sub><sup>+</sup>- ammonium; BOD- biological oxygen  
 738 demand; DOC- dissolved organic carbon; MRP- molybdenum reactive phosphorus; TN-total N;  
 739 \*soiled water; \*\*wastewater

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