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2 Title page

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5 **sediment discharging into the sea in response to human activities**

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21 **Variations in quantity, composition and grain size of Changjiang**
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31
32 Abstract: In order to evaluate the impact of human activities (mainly dam building) on the
33 Changjiang River sediment discharging into the sea, the spatial-temporal variations in the
34 sediment load of different tributaries of the river was analyzed to reveal the quantity, grain size
35 and composition patterns of the sediment entering the sea. The results show that the timing of
36 reduction in the sediment load of the main stream of the Changjiang was different from those
37 associated with downstream and upstream sections, indicating the influences of the
38 sub-catchments. Four step-wise reduction periods were identified, i.e., 1956-1969, 1970-1985,
39 1986-2002, and 2003-2010. The proportion of the sediment load originating from the Jinsha River
40 continuously increased before 2003; after 2003, channel erosion in the main stream provided a
41 major source of the sediment discharging into the sea. In addition, in response to dam construction,
42 although mean grain size of the suspended sediment entering the sea did not change greatly with

43 these different periods, the inter-annual variability for sediment composition or the relative
44 contributions from the various tributaries changed considerably. Before 2003, the clay, silt and
45 sand fractions of the river load were supplied directly by the upstream parts of the Changjiang;
46 after 2003, although the clay component may still be originated mainly from the upstream areas,
47 the source of the silt and sand components have been shifted to a large extent to the river bed
48 erosion of the middle reach of the river. These observations imply that the load, grain size and
49 sediment composition deposited over the coastal and shelf water adjacent to the river mouth may
50 have changed rapidly recently, in response to the catchment changes.

51

52 Keywords: grain size, sediment composition, sediment load, reservoir emplacement, Changjiang

53 River

54

55 1. Introduction

56 Recently, the global sediment flux into the sea has drastically decreased under the influence
57 of human activities (Vörösmarty et al., 2003; Walling, 2006), resulting in considerable changes in
58 the geomorphology and eco-environment of estuarine, coastal and continental shelf regions
59 (Syvitski et al., 2005; Gao and Wang, 2008; Gao et al., 2011). Thus, the source-sink processes and
60 products of the catchment-coast system, including those associated with sediment transport
61 pathways from catchment to continental margins under the impact of climate change and human
62 activities, have received increasing attention (Driscoll and Nittrouer, 2002; Gao, 2006).

63 Because marine deposits consist of the materials from different sub-catchments, variations in
64 the sediment characteristics at the deposition site should result from both sediment load reduction

65 and alterations in sediment grain size, as well as the proportion of the sedimentary materials from
66 different tributaries (which will be referred to as “sediment composition” in the present study)..
67 With regard to the sediment load reduction, there have been studies about the impact of human
68 activities (particularly large hydrologic projects), analyzing long-term variation trends for a
69 number of representative rivers (e.g., Milliman, 1997; Syvitski, 2003; Syvitski and Saito, 2007;
70 Milliman and Farnsworth, 2011; Yang et al., 2011). However, less attention has been paid to the
71 variations in the grain size and sediment composition in response to human activities, and to their
72 sedimentological and environmental consequences. The importance of these two aspects lies in
73 that they reflect the sediment contribution of different sub-catchments to the marine deposits and
74 determine the geochemical and sediment dynamic characteristics (Gao, 2007). Therefore,
75 knowledge about the catchment sediment characteristics for different periods is critical for an
76 accurate analysis of the sediment source/distribution pattern over the estuary and coast-continental
77 shelf regions, and for an improved prediction of the response of the marine sedimentary system to
78 climate change, sea level change, and human activities.

79 The Changjiang is one of the largest rivers in the world. A part of the sediment from the
80 Changjiang catchment has formed a large sub-aqueous delta system of around 10,000 km²
81 (Milliman et al., 1985); and the remainder escapes from the delta, being transported to the Yellow
82 Sea, the East China Sea, and the Okinawa Trough, thereby exerting a considerable impact on the
83 sedimentological and biochemical conditions of these areas (Liu et al., 2007; Dou et al., 2010).
84 Recently, the sediment load of the Changjiang into the sea was reduced considerably in response
85 to dam emplacement and soil water conservation projects (Yang et al., 2002). Dai et al. (2008)
86 estimated that the contribution of dam construction and the water and soil conservative measures

87 accounted for ~88% and $15 \pm 5\%$ of the decline in sediment influx, respectively. The Changjiang
88 catchment consists of numerous tributaries, which are characterized by different
89 rock properties and climate conditions. On the other hand, the intensity and duration of human
90 activities of these tributaries are also varied, which leads to different spatial-temporal patterns of
91 the sediment yield of the sub-catchments (Lu et al., 2003). Thus, the sediment supply of each
92 tributary to the main stream of the Changjiang also changed with time. Furthermore, dam
93 construction and land cover changes also have an important impact on changes in sediment grain
94 size for the tributaries and main stream of the Changjiang (Zhang and Wen, 2004). Therefore, the
95 sediment contribution made by the different tributaries to the sea-going sediment load, the grain
96 size and sediment composition may vary at the same time for the decrease in the total sediment
97 load of the Changjiang River.

98 In order to evaluate the impact of human activities (mainly dam construction) on the quantity,
99 composition and grain size of the Changjiang sediment discharging into the sea, we attempt to: (1)
100 analyze the effect of dam emplacement on the sediment load of different tributaries; (2) identify
101 the spatial-temporal variation patterns of sediment load within the main stream of the Changjiang
102 associated with dam emplacement; (3) reveal the quantity, grain size and composition features of
103 the sea-going sediment during different periods; and (4) delineate the variations in sediment load
104 originating from the tributaries of the Changjiang for different historical times.

105

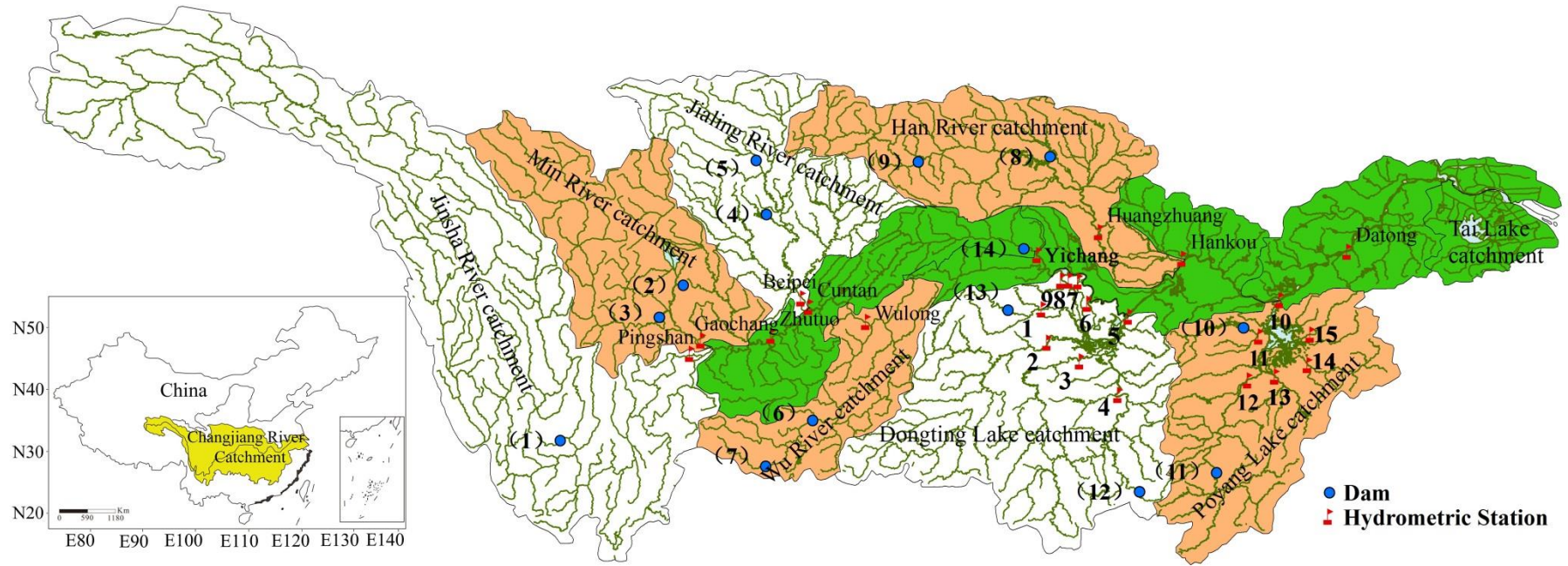
106 2. Regional setting

107 The Changjiang, with a drainage basin area of approximately $1.80 \times 10^6 \text{ km}^2$, originates in the
108 Qinghai-Tibet Plateau and flows 6,300 km eastward toward the East China Sea. The upper reach

109 of the river, from the upstream source to the Yichang gauging station (Fig.1), is the major
110 sediment-yielding area of the entire catchment (Shi, 2008). The main upstream river has four
111 major tributaries, i.e., the Jinsha, Min, Jialing, and Wu Rivers. The upper reach region is typically
112 mountainous, with an elevation exceeding 1,000 m above sea level (Chen et al., 2001). The
113 mid-lower reach extends from Yichang to the Datong gauging station, with three large inputs
114 joining the main stream in this section: the Dongting Lake drainage basin, the Hanjiang River, and
115 the Poyang Lake drainage basin. The catchment area of this section mainly comprises alluvial
116 plains and low hills with elevations of less than 200 m (Yin et al., 2007). The Dongting Lake is the
117 second largest freshwater lake in China, and part of the main river flow enters Dongting Lake via
118 five different entrances. Four tributaries enter Dongting Lake from the south and southwest, and
119 water from Dongting Lake flows into the Changjiang main river channel at the Chenglingji
120 gauging station (Dai et al., 2008). The Dongting Lake was a major sink of the upstream sediment
121 of the Changjiang and, due to sediment decrease from the upstream Changjiang, has become a
122 weak sediment source to its downstream sections (Dai and Liu, 2013). Poyang Lake is the largest
123 freshwater lake in China, and it directly exchanges and interacts with the river. Poyang Lake
124 receives runoff from 5 smaller tributaries (the Gan, Fu, Xin, Rao, and Xiu Rivers) and discharges
125 freshwater into the Changjiang at Hukou (Shankman et al., 2006). The estuarine reach of the
126 Changjiang extends from Datong (tidal limit) to the river mouth. The Datong gauging station is
127 the last station along the Changjiang before going to the sea, and its hydrological records are often
128 used to derive a represent sediment flux of the Changjiang into the adjacent East China Sea.

129 Due to intensified human activities, the catchment forest vegetation degenerated continuously,
130 with a large-scale reduction of forest cover in the Changjiang catchment (Xu, 2000), resulting

131 serious deterioration of the ecological environment (Lu and Higgitt, 2000). Starting from the late
132 1980s, a major soil conservation campaign was implemented in high sediment yielding regions of
133 the upper Changjiang catchment. However, due to the highly variable natural conditions of the
134 tributaries, the effect of this campaign was different in every upstream tributary. For example,
135 most of the Jialing River catchment is characterized by hills areas, with potential severe slope
136 erosion (Zhang and Wen, 2004), yet its vegetation restoration rate is quite high due to the humid
137 climate; hence, the effect of vegetation recovery on the reduction of slope erosion is prominent
138 (Lei et al., 2006). As such, the sediment yield of these parts of the Jialing River catchment has
139 rapidly decreased since the soil conservation campaign carried out in the 1980s (BSWC, 2011).
140 However, the downstream Jinsha River has a different situation. This section, 782 km in length, is
141 the main sediment yield area; although its area only accounts for 7.8% of the upstream Changjiang,
142 the average annual sediment load reaches 35.50% of the quantity at the Yichang station (Zhang
143 and Wen, 2004). The high and steep mountains here are characterized by landslides and debris
144 flow, reducing the effect of vegetation restoration (Lei and Huang, 1991; Yang, 2004). Therefore,
145 the water and soil erosion prevention scheme does not work in the Jinsha River as well as in the
146 Jialing River (BSWC, 2011); in the former case, reservoir interception is still the dominant factor
147 of the sediment load reduction.



1 Shimen 2 Taoyuan 3 Taojiang 4 Xiangtan 5 Chenglingji 6 Ouchi 7 Mituosi 8 Shadaoguan
 9 Xinjiangkou 10 Hukou 11 Wanjiabu 12 Waizhou 13 Lijiadu 14 Meigang 15 Hushan

149
 150 **Figure 1.** The Changjiang catchment and location of the hydrologic stations for the Changjiang catchment. The numeric symbols in the figure denote some important reservoir sites,
 151 including: (1) Er'tan; (2) Heilongtan; (3) Tongjiezi; (4) Shengzhong; (5) Baozhushi; (6) Wujiangdu; (7) Puding; (8) Danjiangkou; (9) Ankang; (10) Zhelin; (11) Wan'an; (12) Dongjiang; (13)
 152 Jiangya; and (14) Three Gorges Dam).

153 3. Material and method

154 3.1 Data sources

155 3.1.1 Water discharge and sediment load data

156 A long-term discharge and sediment monitoring program for the entire catchment has been
157 implemented since the 1950s, by the Changjiang Water Resource Commission (CWRC) under the
158 supervision of the Ministry of Water Resources, China (MWRC). These monitoring data of each
159 station include field survey and measurement of water discharge, suspended sediment
160 concentration, suspended sediment load, and suspended sediment grain size, in accordance with
161 China's national data standards (Ministry of Water Conservancy and Electric Power, 1962, 1975):
162 10-30 vertical profiles within the water column were established for the measurements of each
163 river cross-section, with the number of profiles varying with the width of the river. For each
164 profile, the flow velocity are measured (using a direct reading current meter) at different depths
165 (normally at surface, 0.2H, 0.6H, 0.8H and the bottom, where H is the water depth). Meanwhile,
166 the water mass of these layers is sampled for the measurements of the suspended sediment
167 concentration (using filtration) and grain size (using the suspension settling method). Such
168 measurements are repeated daily at each station. The homogeneity and reliability of the
169 hydrological data, with an estimated daily error of 16% (Wang et al., 2007), has been strictly
170 examined by the CWRC before release. The data for the period of 1956-2001 were either
171 published in the Yangtze River Hydrological Annals or provided directly by the CWRC. After
172 2002, these hydrological data were reported in the Bulletin of China River Sediment published by
173 the Ministry of Water Resources, China (BCRS, 2002-2010; available at:
174 <http://www.mwr.gov.cn/zwzc/hygb/zghlnsgb/>).

175 We acquired the annual sediment load data from 26 hydrological stations distributed in the
176 main reach and seven of the tributaries (for the location of these stations, see Fig.1). The dataset
177 for these gauging stations covers a 55-year period (i.e., 1956-2010).

178

179 3.1.2 Dam data

180 In the present study, the reservoirs with a storage capacity of greater than 0.01 km^3 (i.e., “large
181 and medium sized reservoirs” according to the MWRC) are considered. Data on reservoir emplacement
182 during 1949-2001 were obtained from the MWRC (2001), and those built during 2002-2007 were
183 obtained from the annual reports published by the MWRC (<http://www.mwr.gov.cn/zwzc/hygb/slbgb/>).
184 In total, we count 1,132 large and medium sized reservoirs located within the Changjiang catchment, of
185 which 1,037 reservoirs are situated upstream of the Datong station (Fig.1b). The database includes
186 information on reservoir storage capacity, construction and impoundment time.

187 In the present study, the reservoir storage capacity index (RSCI) is defined as the ratio of the
188 reservoir storage capacity to the annual average water discharge of the contributed catchment; thus, the
189 total RSCI of a catchment is the ratio of total capacity of reservoir to the annual average water
190 discharge.

191

192 3.2 Analytical methods

193 The Mann-Kendall test (M-K test) is a nonparametric method, which has been used to
194 analyze long-term hydro-meteorological temporal series (Mann, 1945; Kendall, 1955). This test
195 does not assume any distribution form for the data and is as powerful as its parametric competitors
196 (Serrano et al., 1999). Trend analysis of the sediment load changes was conducted based on this

197 method. Before using the M-K test, the autocorrelation and partial autocorrelation functions were
 198 used to examine the autocorrelation of all the hydrological data. The results indicated that there
 199 was no significant autocorrelation in the data. The modified M-K method was used to analyze
 200 variations in the sediment load data: $X_t = (x_1, x_2, x_3 \dots x_n)$, where the accumulative number m_i for
 201 samples for which $x_i > x_j$ ($1 \leq j \leq i$) was calculated, and the normally distributed statistical variable
 202 d_k was expressed as (Hamed and Rao, 1998)

$$203 \quad d_k = \sum_{i=1}^k m_i \quad 2 \leq k \leq n \quad (1)$$

204 The mean and variance of the normally distributed statistic d_k were defined as

$$205 \quad E[d_k] = \frac{k(k-1)}{4} \quad (2)$$

$$206 \quad Var[d_k] = \frac{k(k-1)(2k+5)}{72} \quad 2 \leq k \leq n \quad (3)$$

207 Then, the normalized variable statistical parameter UF_k was calculated as

$$208 \quad UF_k = \frac{d_k - E[d_k]}{\sqrt{Var[d_k]}} \quad k = 1, 2, 3, \dots, n \quad (4)$$

209 where UF_k is the forward sequence. The backward sequence UB_k was obtained using the same
 210 equation but with a retrograde sample. The C values calculated with progressive and retrograde
 211 series were named as C_1 and C_2 . The intersection point of the two lines, C_1 and C_2 ($k=1, 2, \dots, n$)
 212 was located within the confidence interval, providing the beginning of the step change point
 213 within the time series. Assuming a normal distribution at the significant level of $P=0.05$, a positive
 214 Man-Kendal statistics C larger than 1.96 indicates a significant increasing trend, whilst a negative
 215 C value with an absolute value of lower than 1.96 indicates a significant decreasing trend.

216

217 4. Results

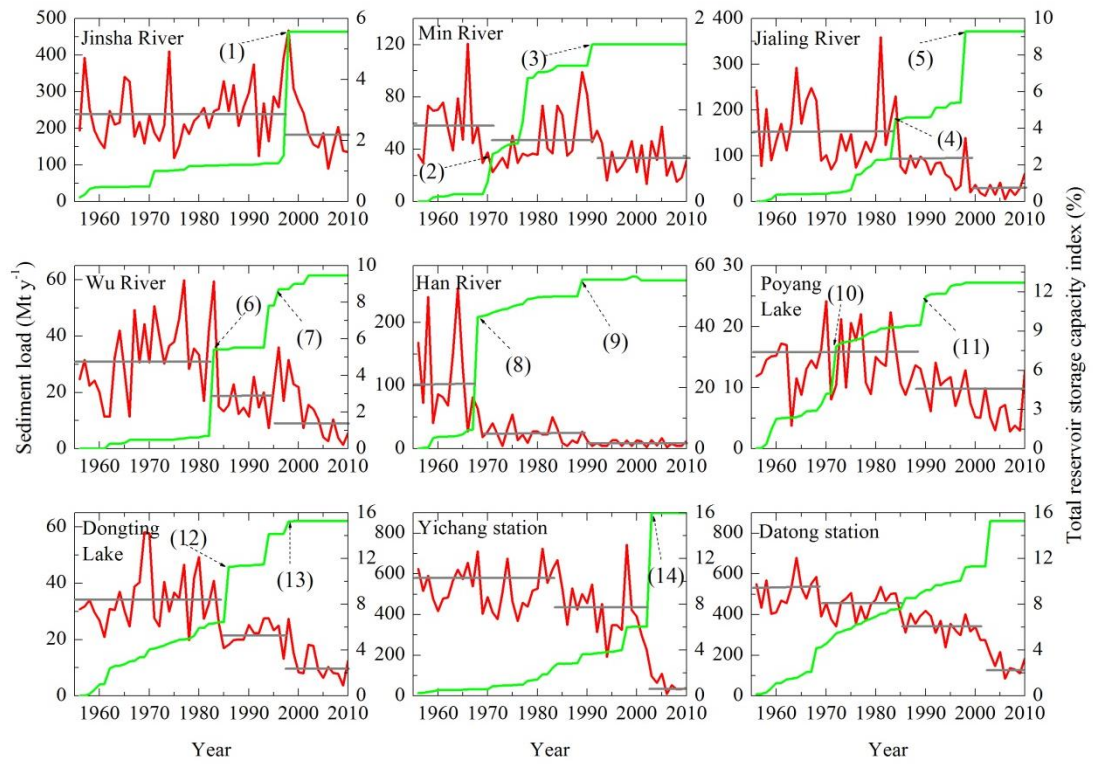
218 4.1 Stepwise variations in the reservoir storage capacity of the

219 **tributaries**

220 The total RSCI values of the seven tributaries and the main stream of the Changjiang reveal
221 stepwise increasing trends (Fig. 2). The variations in reservoir storage capacity of the four
222 upstream tributaries indicated that the total RSCI of the Min River catchment is low (1.72% in
223 2010) and those of the Jialing and Wu Rivers rapidly increased in 1985. In response to the
224 construction of the Er'tan reservoir, the total RSCI of the Jinsha River also rose considerably in
225 1998. As a result of rising in the reservoir storage capacity of these four rivers, the total RSCI of
226 the Changjiang catchment, upstream of the Yichang station with increases by 2.8% in 1985 and
227 16.0% in 2003, also showed the stepwise patterns. The middle reaches of the Changjiang
228 catchment consisted of three major tributaries, namely, the Han River, Dongting Lake and Poyang
229 Lake. The total RSCI of the Han River began to increase in 1966, and greatly rose in 1968. In
230 addition, the rapid increment in the total RSCI of the Poyang and Dongting Lakes were also
231 present for 1972 and 1985, respectively. Generally, as a consequence of dam construction, the total
232 RSCI of the Changjiang upstream of the Datong greatly increased in 1969 and 2003, respectively.

233 The changes of the total RSCI and sediment load of the tributaries and the whole Changjiang
234 catchment indicate that the stepwise decrease of sediment load is apparently related to the
235 significant increase of the total RSCI. In the case of the Yichang and Datong stations, over the last
236 few decades, there is significant negative correlation between average sediment load and total
237 RSCI at both the Yichang and Datong stations (Fig.3), which reflected the impact dams have on
238 the sediment load.

239



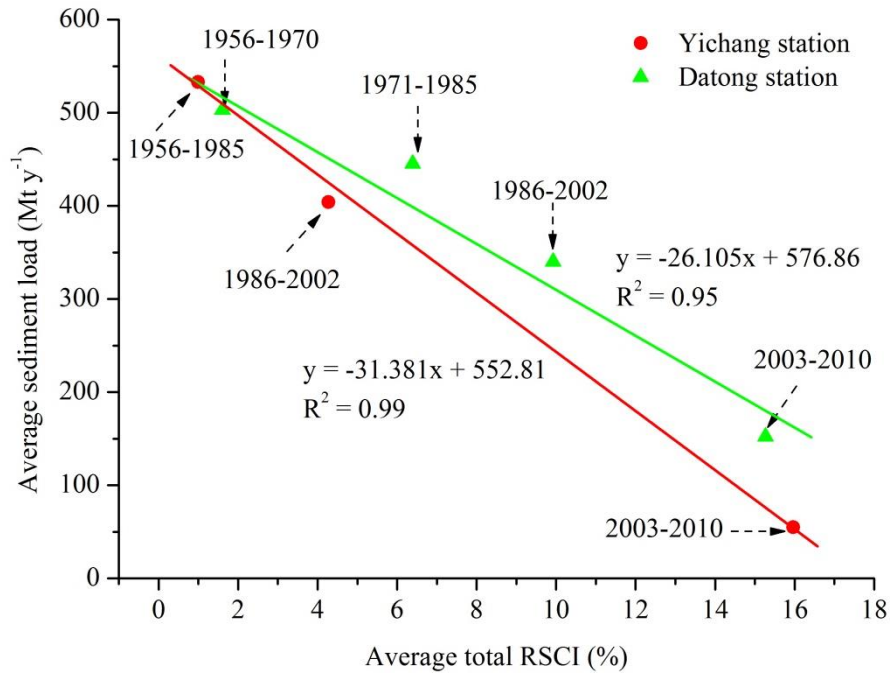
240

241 **Figure 2.** Relationship between the reduction in sediment load (red line) and the total reservoir storage

242 capacity index (green line) in the tributaries and the main stream. Numeric symbols represent reservoirs listed in

243 Figure 1.

244



245

246 **Figure 3.** The relationship between average sediment load and total RSCI of different periods at Yichang and

247 Datong station.

248

249 4.2 Spatial-temporal sediment load variations within the catchment

250 The trends, derived on the basis of the M-K method, of sediment load of the seven tributaries

251 (Figs. 4 and 5) indicate that the downstream sediment load began to decrease earlier than the

252 upstream sediment load. Due to the sediment load reduction of the Jialing and Wu Rivers, the total

253 sediment load of the four upstream rivers began to decrease in 1984. In the middle stream of the

254 Changjiang, the sediment load for the Han River and Dongting - Poyang Lakes began to reduce in

255 1966, 1984 and 1985, respectively; the M-K trends of sediment load of the three sub-catchments

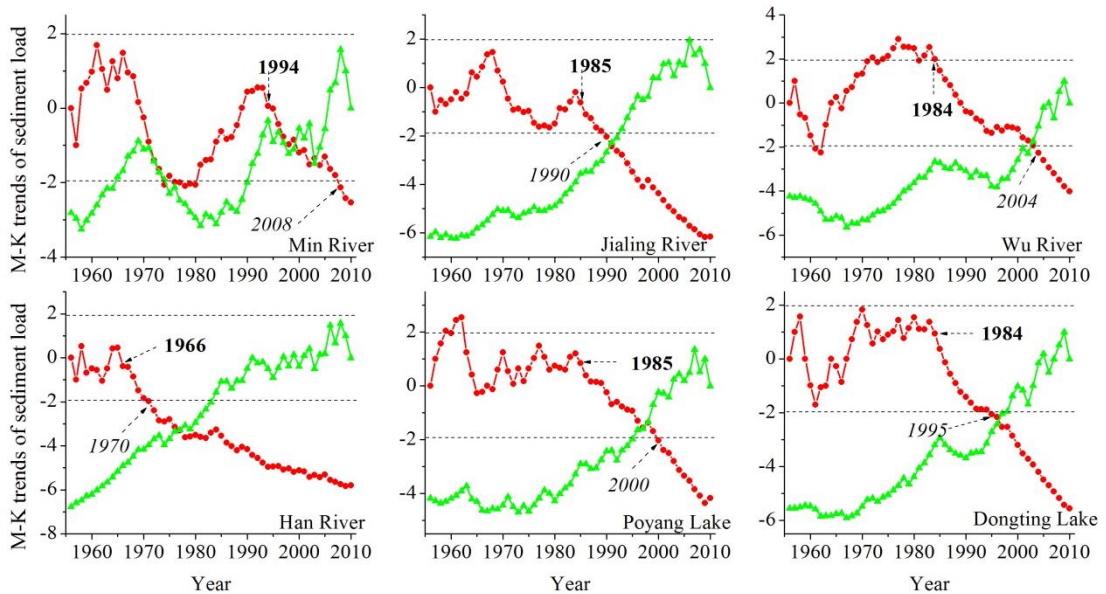
256 exhibited significant trends of decrease (at the 95% confidence level) in 1970, 1995 and 2000,

257 respectively.

258 Due to the different patterns of sediment load variations of the seven sub-catchments, there

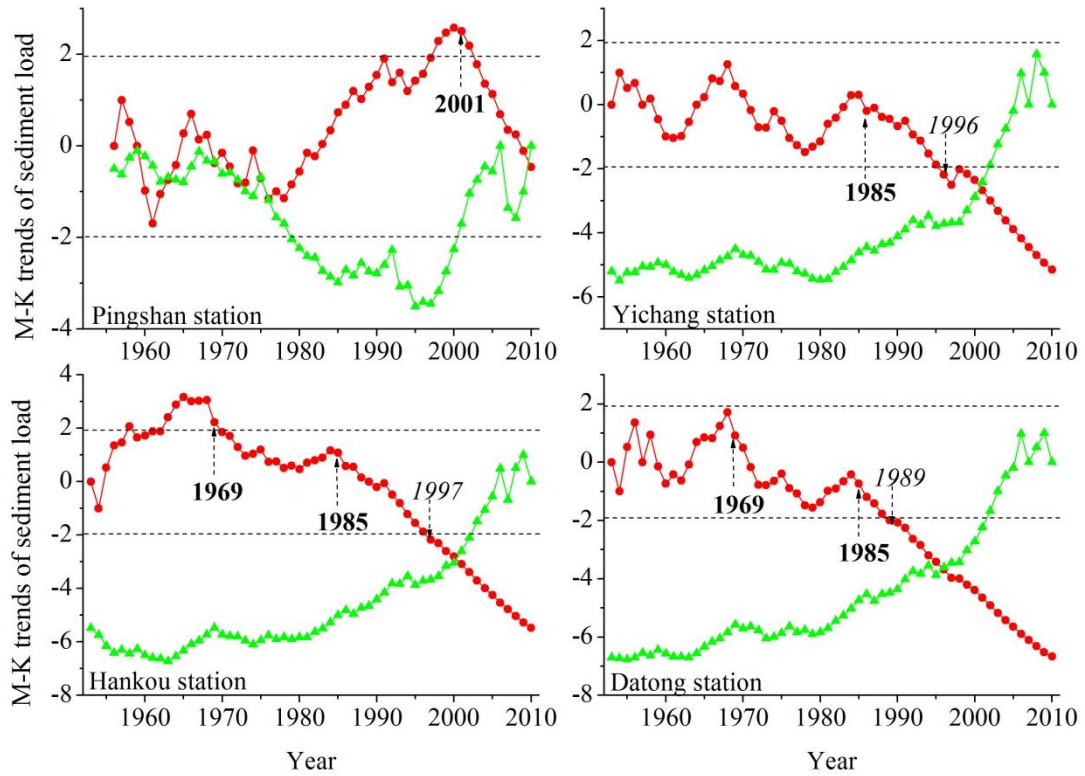
259 were significant spatial-temporal differences in the sediment load variations of the mainstream

260 Changjiang: the sediment load began to decrease later at upstream locations than at downstream
 261 locations. The sediment load upstream of the Yichang station began to reduce in 1985, with a 95%
 262 confidence level for the year of 1996. Impacted by sediment load decreasing of the Han River,
 263 beginning from 1966, the sediment load reduction trends of the middle reach (Hankou - Datong
 264 stations) were observed in 1969. Furthermore, as a result of sediment load reducing of upstream
 265 and middle reach tributaries in 1985, the sediment load of the middle reach of the Changjiang
 266 began to further decrease in 1985. The M-K trends of sediment load of the Datong, Hankou and
 267 Yichang stations are associated with a 95% confidence level in 1989, 1997 and 1996, respectively.



268
 269 **Figure 4.** M-K trends of the sediment load for the Jinsha, Min, Jialing, Wu, and Han Rivers and the Poyang
 270 and Dongting Lake system. The symbol \bullet and \blacktriangle denotes C_1 and C_2 , respectively. The bold is the beginning time
 271 of sediment load decreasing; and the italic is the time when the M-K trends of the sediment load pass the 95%
 272 confidence test.

273



274

275 **Figure 5.** M-K trends of the sediment load for different gauging stations of the Changjiang main river. The
 276 symbol ● and ▲ denotes C_1 and C_2 , respectively. The bold is the beginning time of sediment load decreasing;
 277 and the number in italics denotes the time when the M-K trends of the sediment load pass the 95% confidence test.

278

279 4.3 Stepwise reduction of the sediment load entering the sea

280 The M-K trends of sediment load variation at the Datong station show that 1969 and 1985 are
 281 two critical temporal nodes, reflecting the beginning time of sediment load decrease. While the
 282 M-K trends of the sediment load passes the 95% confidence test for 1989, another important time
 283 node (2003) is not shown in the M-K trends of sediment load of the Datong station. Taking into
 284 account the great impact of the Three Gorges Dam on the sediment load decrease of the main
 285 stream Changjiang (Hu et al., 2011), the variations of the sediment load entering the sea could be
 286 divided into four stepwise reduction stages, namely, 1956-1969, 1970-1985, 1986-2002, and

287 2003-2010.

288 The variations of sediment load discharging into the sea, measured at the Datong station,
289 indicated that although the sediment load of the Datong station, 503 Mt y⁻¹ on average, exhibited
290 fluctuations from 1956 to 1969, with the quantity generally remaining at a high level (Tab.1). The
291 Han River was once the most important sediment source of the middle reach Changjiang (Yin et
292 al., 2007); however, because the annual sediment load supplied by the Han River decreased by 95
293 Mt, the sediment load of the Datong station was reduced to 445 Mt in the period of 1970-1985.
294 Previous studies elsewhere have suggested that the sediment load from the Changjiang entering
295 the sea began to decrease in the 1980s (Yang et al., 2002); however, we would propose that such a
296 decreasing trend already occurred in as early as 1970, and the impact of the reduced sediment load
297 of the Han River on the overall sediment flux of the Changjiang was neglected in these previous
298 studies. The sediment load for the upstream Changjiang had a decreasing trend starting from 1985;
299 in term of the quantity, there was a reduction from 533 Mt y⁻¹ during 1956-1985 to 404 Mt y⁻¹
300 during 1986-2002. The sea-going sediment load of the Changjiang became less than 340 Mt y⁻¹
301 during this period. With the emplacement of the Three Gorges Dam in 2003, the sediment load
302 upstream of the Changjiang decreased to 55 Mt y⁻¹ during 2003-2010, with the sediment discharge
303 into the sea being around 152 Mt y⁻¹ (Tab. 1).

304

305 Table 1. The mean value of sediment load of the Changjiang main river during different period

Time	Pingshan station Mt y ⁻¹	Yichang station Mt y ⁻¹	Hankou station Mt y ⁻¹	Datong station Mt y ⁻¹
1956-1969	232	547	461	503
1970-1985	226	521	426	445
1986-2002	275	404	331	340
2003-2010	151	55	118	152

306

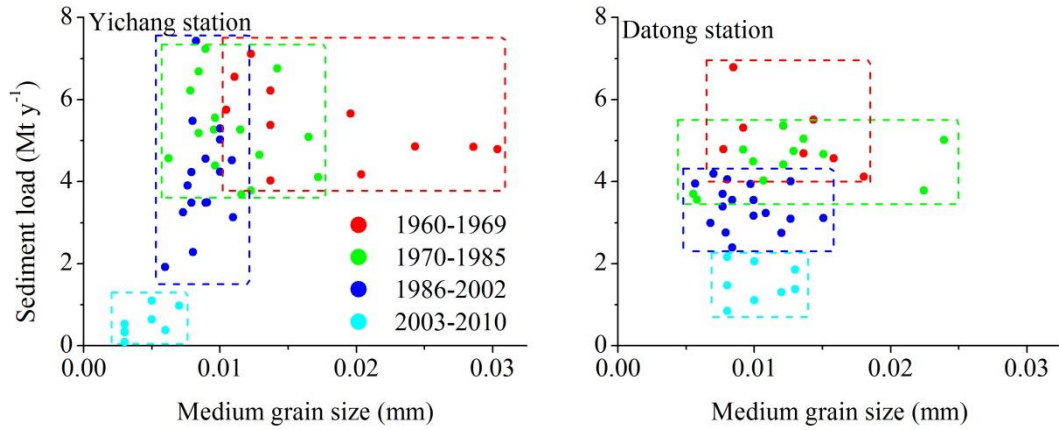
307 Overall, four stepwise reduction stage periods of the sea-going sediment load were observed,
308 namely, 1956-1969, 1970-1985, 1986-2002, and 2003-2010. Further, the sediment load decrease
309 may be related to sediment load decrease of different tributaries: the reduction during 1970-1985
310 was correlated with the Han River, whilst the upstream tributaries (mainly the Jialing and Wu
311 Rivers), together with the sub-catchment of the middle reach (mainly the Poyang Lake), were
312 responsible for the decrease during 1970-1985. The sediment load decrease during 2003-2010
313 resulted mainly from the emplacement of the Three Gorges Dam.

314

315 4.4 Variations in the grain size of the sediment entering the sea

316 Because most of the coarse-grained sediment is intercepted by reservoirs, the sediment grain
317 size downstream of the reservoirs becomes significantly finer (Xu, 2005). The variation in the
318 medium grain size (D_{50}) of suspended sediment at the Yichang station (Fig. 6) indicates that the
319 average value of D_{50} was 0.017 mm in 1960-1969, 0.012 mm in 1970-1985, 0.009 mm in
320 1986-2002, and 0.004 mm in 2003-2010, suggesting that the sediment grain size from the
321 upstream Changjiang exhibited a continuous decreasing trend. In contrast, the decreasing trend of
322 D_{50} for the Datong station was not as significant as that for the Yichang station during these four
323 stages: the average D_{50} in 1960-1969 (0.12 mm) was similar to that in 1970-1985 (0.13 mm), with
324 a slight decrease for the year of 2002 (0.09 mm) and for the period of 2003-2010 (0.10 mm).

325



326

327 **Figure 6.** Relationship between the medium grain size of suspended sediments and the sediment load during

328 different periods at the Yichang and Datong stations. Data are not available for the Datong station in 1968-1970,

329 1972-1973, and 1975.

330

331 In addition, the degree of inter-annual variation in the upstream sediment grain size

332 continuously decreased during the four periods at the Yichang station, i.e., the range of D_{50}

333 variations is gradually narrowed (with a continuously reduced standard deviation), and the

334 distribution range of the D_{50} data and sediment load moves towards the side of finer grain sizes;

335 however, such a change is not so significant at the Datong station (Fig. 6). The sediment grain size

336 variations of the two stations also indicated that the average value of D_{50} for the Yichang station

337 was greater than that for the Datong station in 1960-1969, but the two stations had similar values

338 in 1970-1985 and 1986-2002; after 2003, the average value of D_{50} of the Yichang station was

339 smaller than that of the Datong station. Furthermore, D_{50} ranged from 0.003-0.007 mm for the

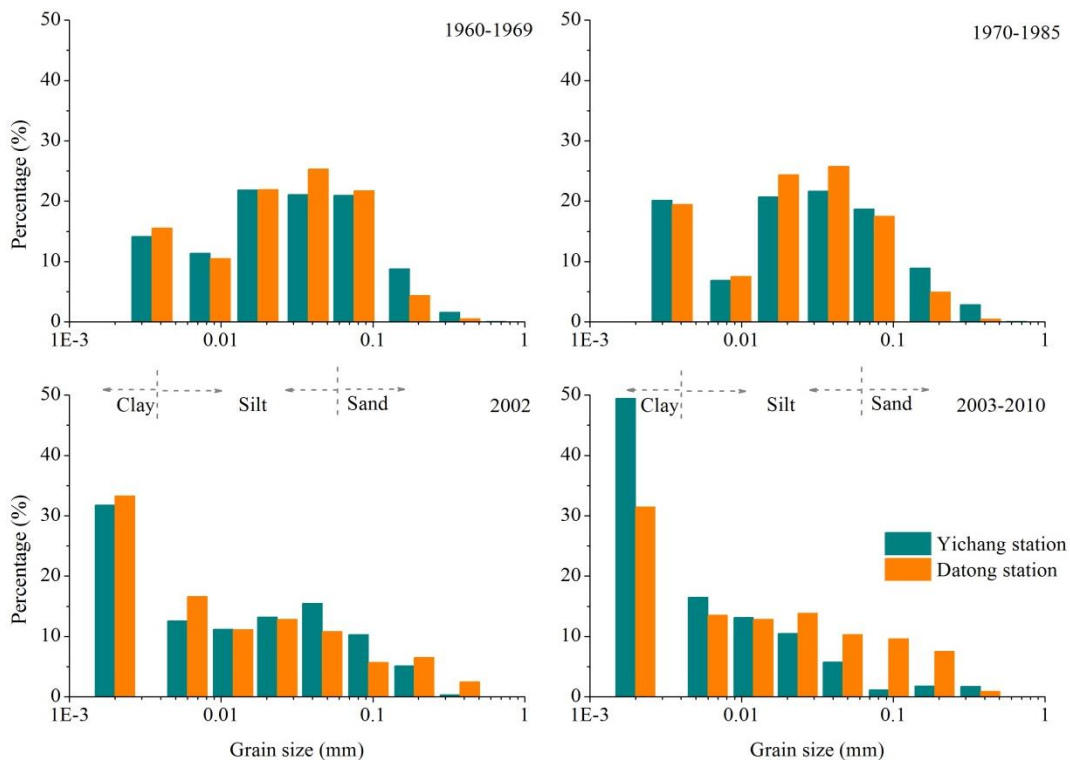
340 Yichang station and 0.008-0.013 mm for the Datong station, in 2003-2010, suggesting that the D_{50}

341 variation range of the two stations did not overlap after 2003.

342 Compared with previous periods, after 2003, the clay and silt contents at the Yichang station

343 greatly increased, and the sand fraction significantly decreased (Fig. 7). At the Datong station,
 344 however, although the sand fraction had no apparent variation trends, the clay content increased,
 345 and the silt content reduced. Furthermore, before 2003, the silt and clay contents did not differ
 346 much between the Yichang and Datong stations, and the sand content at the Yichang station was
 347 slightly greater than that at the Datong station; however, after 2003, the sand content at the Datong
 348 station became significantly greater than that at the Yichang station, and the clay content at the
 349 Datong station became lower than that at the Yichang station, implying that the sediment sources
 350 other than the seven tributaries supplied a sand fraction to the Yichang-Datong section of the
 351 Changjiang. This observation suggests that although the average value of the grain size of the
 352 sediment entering the sea during the different periods did not alter greatly, the inter-annual
 353 variation range, sediment components and the material sources changed considerably.

354



355

356 **Figure 7.** Distribution of the suspended sediment grain size of the Yichang and Datong stations in 1960-1969,
357 1970-1985, 2002, and 2003-2010.

358

359 5. Discussion

360 As outlined above, the Changjiang sediment load is influenced by mixing of
361 weathering products supplied by the different sub-catchments. The spatial-temporal differences
362 among the sub-catchments, in terms of sediment load variations, caused the sediment load
363 reduction and changes in the sediment composition. According to the concept of sediment budget
364 (Houben, 2012), the following equation may be used to calculate the sediment balance of the main
365 stream Changjiang:

$$366 \quad \sum S_{input} = \Delta S + S_{output} = S_{Jinsha} + S_{Min} + S_{Jialing} + S_{Wu} + S_{Han} + S_{Poyang} \quad (5)$$

367 where $\sum S_{input}$ is the contribution of the tributaries to the main stream sediment load, S_{output} is
368 the sediment load entering the sea (measured at the Datong station), ΔS is the quantity of
369 deposited (+) / eroded (-) sediment of the main stream Changjiang and the Dongting Lake. Thus,
370 the contribution of the different tributaries to the overall sediment load can be expressed by:

$$371 \quad \frac{S_{Jinsha}}{S_{output}} + \frac{S_{Min}}{S_{output}} + \frac{S_{Jialing}}{S_{output}} + \frac{S_{Wu}}{S_{output}} + \frac{S_{Han}}{S_{output}} + \frac{S_{Poyang}}{S_{output}} - \frac{\Delta S}{S_{output}} = 1 \quad (6)$$

372 The calculated results indicated that (Tab.2), in 1956-1969, the sediment load of the Datong
373 station was mainly originated from the Jinsha, Jialing and Han Rivers, with their contributions
374 being 35.0%, 24.3%, and 19.0%, respectively. As the sediment load of the Han River decreased,
375 the Jinsha and Jialing Rivers accounted for 46.7% and 27.6%, respectively, in the sediment load at
376 the Datong station during the 1970-1985 period; whereas the contribution from the Han River
377 decreased to 5.8%. During the 1986-2002 period, due to the reduced sediment yield in the Jialing

378 River, the contribution of the Jinsha River to the sediment load at the Datong station further
 379 increased to 64.2% and that of the Jialing River decreased to 15.0%. The sediment composition
 380 changed considerably during the 2003-2010 period due to the TGD emplacement: the sediment
 381 proportion due to channel erosion of the main stream reached 48.3% and the proportion of the
 382 Jinsha River decreased dramatically to 24.1%. Furthermore, both the Jialing and Han Rivers only
 383 contributed 5.3% to the sediment load at the Datong station.

384

385 **Table 2.** The sediment contribution proportion (%) of different tributaries to the sediment load entering the sea of
 386 the Changjiang.

River/Catchment	1956-1969	1970-1985	1986-2002	2003-2010
Jinsha River	35	46.7	64.2	24.1
Min River	8.8	8.6	10.1	6.1
Jialing River	24.3	27.6	15	5.3
Wu River	4.4	8.2	4.5	2.2
The total of the upstream four rivers	72.5	91.1	93.8	37.7
Han River	19	5.8	2.8	5.3
Channel erosion	6.1	0.9	1.1	48.3
Poyang Lake	2.4	2.2	2.3	8.7

387

388 The above analysis indicates that as the sediment load at the Datong station decreased,
 389 although the average sediment grain size did not display clearly-defined variations, the sediment
 390 composition changed considerably. Before 2003, the four rivers of the upstream Changjiang were
 391 the dominating sediment source to the sediment load entering the sea, and their total contribution
 392 was 72.5% during 1956-1969, 91.1% during 1970-1985, and 93.8% during 1986-2002. In addition,
 393 during these periods, the variations in the sediment composition were mainly determined by the
 394 changes in the sediment contributions of the Jinsha, Jialing, and Han Rivers, i.e., with the
 395 sequential reduction in the sediment loads of the Han and Jialing Rivers, the proportion of the

396 sediment load originating from the Jinsha River continuously increased. However, after 2003, the
397 sediment contribution of the upstream to the sediment load of the Datong station greatly decreased.
398 The middle reach of the Changjiang became one of the major sinks of the upstream sediment
399 (Yang et al., 2011); after 2003, channel erosion of the middle reach main stream became the most
400 important source of sediment load at the Datong station.

401 Apart from the dam interception effect, the soil conservation campaign starting from 1989
402 and implemented for the high sediment yielding regions of the upper Changjiang basin (Hu et al.,
403 2011) may be another factor accelerating the decreasing trend of the sediment grain size at the
404 Yichang station. The different grain sizes of the suspended sediment at the Yichang and Datong
405 stations indicate that the clay, silt and sand fluxes at the Yichang station were greater than those at
406 the Datong station during 1960-1969, 1970-1985 and 1986-2002 periods (Tab. 3), which implies
407 that the sediment fractions of clay, silt and sand entering the sea were mainly originated from the
408 upstream Changjiang, with weak sediment exchange between the water column and the riverbed.
409 After the emplacement of the TGD in 2003, the clay, silt and sand fractions originating from the
410 upstream Changjiang decreased dramatically. With regard to the amount of sediment originating
411 from the Poyang Lake and the Han River to the main stream Changjiang, we may still use the
412 sediment budget concept, to calculate the balance for the different sediment fractions for the
413 Yichang-Datong reach:

$$414 \quad S_{Yichang} + S_{Han} + S_{Poyang} = \Delta S + S_{datong} \quad (7)$$

415 The calculations show that the eroded sediment of the main river channel (between Yichang
416 and Datong) and the Dongting Lake contributed 13 Mt y⁻¹ of clay, 43 Mt y⁻¹ of silt and 20 Mt y⁻¹
417 of sand to the sediment load at the Datong station in 2003-2010, which accounted for 27.1%, 55.8%

418 and 74.1% of the corresponding sediment components of the Datong station. Taking into account
 419 the eroded sediment supply within the estuarine areas (Li, 2007), the percentages of the silt and
 420 sand fractions discharging into the sea, due to the material supply by the eroding main river
 421 channel, may exceed 55.8% and 74.1%, respectively. These data imply that the clay fraction at the
 422 Datong station should be originated mainly from the upstream Changjiang, and the silt and sand
 423 fractions largely consisted of the eroded sediment of the middle reach river channel.

424

425 **Table 3.** Annual quantities of clay, silt, and sand at the Yichang and Datong stations during different periods.

Time Period	Clay (Mt y ⁻¹)		Silt (Mt y ⁻¹)		Sand (Mt y ⁻¹)	
	Yichang	Datong	Yichang	Datong	Yichang	Datong
1960-1969	78	78	297	291	172	134
1970-1985	105	86	257	257	159	102
1986-2002	128	113	212	174	63	50
2003-2010	27	48	25	77	3	27

426

427 6. Conclusions

428 (1) The increment in reservoir storage capacity is significantly correlated with the decrease in
 429 the Changjiang sediment load, which reflected the impact of dams on the sediment load of the
 430 tributaries and the entire Changjiang catchment.

431 (2) The patterns of sediment delivery from the sub-catchments of the Changjiang River have
 432 been changed, with significant spatial-temporal differences in the sediment load variations of the
 433 main stream Changjiang: four stepwise reduction stages were identified, i.e., 1956-1969,
 434 1970-1985, 1986-2002 and 2003-2010. There was a lag of the decrease in the sediment load at
 435 upstream locations compared with those at downstream locations.

436 (3) Before 2003, the variations in the sediment composition in the marine areas were mainly

437 determined by the changes in the sediment contribution made by the Jinsha, Jialing, and Han
438 Rivers. However, after 2003, channel erosion of the main stream Changjiang supplied around 48.3%
439 of the sediment load into the sea.

440 (4) In response to dam construction, although mean grain size of the sediment entering the
441 sea during the different periods did not show clearly-defined variations, the inter-annual variations
442 in terms of the size range, sediment components and source areas changed considerably.

443 (5) Before 2003, the clay, silt and sand fractions entering the sea were mainly originated from
444 the upstream regions of the river. In contrast, after 2003, the origin of the clay component of the
445 sediment was dominated by the upstream areas, whilst the silt and sand component were mainly
446 supplied by the eroding bed of the middle-reach main channel of the Changjiang River.

447

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