



A plot-scale rainfall simulator for runoff and parameter estimation

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Development and testing of a large, transportable rainfall simulator for plot-scale runoff and parameter estimation

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Abstract

There is increased interest in the interplay between vegetation conditions and overland flow generation. The literature is unclear on this relationship and there is little quantitative guidance for modeling efforts. Therefore, experimental efforts are needed and these call for a lightweight transportable plot-scale ($> 10 \text{ m}^2$) rainfall simulator that can be deployed quickly and quickly redeployed over various vegetation cover conditions. Accordingly, a variable intensity rainfall simulator and collection system was designed and tested in the laboratory and in the field. The system was tested with three configurations of common pressure washing nozzles producing rainfall intensities of 62, 43, and 32 mm h^{-1} with uniformity coefficients of 76, 65, and 62, respectively, over a plot of 15.12 m^2 . Field tests were carried out in on a grassy field with silt-loam soil in Orroli, Sardinia in July and August 2010, and rainfall, soil moisture, and runoff data were collected. The two-term Philip infiltration model was used to find optimal values for the saturated hydraulic conductivity of the soil surface and bulk soil, soil water retention curve slope, and air entry suction head. Optimized hydraulic conductivity values were comparable to both the measured final infiltration rate and literature values for saturated hydraulic conductivity. This inexpensive rainfall simulator can therefore be used to identify field parameters needed for hydrologic modeling.

1 Introduction

Rainfall simulators have been used extensively to gather runoff, infiltration, and erosion data in both laboratory and field experiments. The results of these experiments are typically used for the purposes of understanding processes such as runoff and infiltration mechanisms, water routing, and sediment generation and transport at scales ranging from point to hillslope, with emphasis on how surface characteristics such as slope, aspect, soil properties, fire, vegetation, and microtopography affect these processes

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(e.g., Dunne et al., 1991; Bhardwaj and Singh, 1992; Foster et al., 2000; Stone et al., 2008; Fernandez-Galvez et al., 2008).

More recently, research has started to focus on the effects that changes in surface properties such as land cover or land use can have on the hydrologic cycle (e.g., García-Ruiz et al., 2005; Genxu et al., 2012; Lana-Renault et al., 2011; Maetens et al., 2012; Morán-Tejeda et al., 2010; Romero et al., 2011); these studies often assess how runoff and erosion at the plot and hillslope scale change as vegetation recovers from fire, agriculture, or other disturbances. Therefore, an emerging research need is the further understanding of the interplay between vegetation, additional surface properties, and runoff generation on scales from hillslope to watershed. To begin to answer this question, there is a need for broad experiments that evaluate characteristic land surfaces throughout a watershed and during the course of the year. Therefore, in order to improve the accuracy of a watershed hydrologic model, more sample locations are needed in order to properly characterize the surface runoff response. In practice, this requires a large rainfall simulator that is capable of producing high rainfall intensities and can easily be transported between field sites.

There are several different types of rainfall simulators, each with its own application, benefits, and shortcomings. The literature on rainfall simulators is extensive; for a more complete review see Battany and Grismer (2000). The standard small or laboratory scale (1 m² or smaller) rainfall simulator is a drip tank (Foster et al., 2000; Fernandez-Galvez et al., 2008; Cerdà et al., 1997, and many others). These devices are tanks with a uniform arrangement of holes. They can produce a wide range of rainfall intensities and typically provide spatially and temporally uniform coverage. The 0.5 m² simulator of Foster et al. (2000) produced intensities ranging from 7.74 to 28.57 mm h⁻¹ with temporal coefficients of variation for intensity ranging from 5.04 to 11.55 %, and the 0.24 m² rainfall simulator of Cerdà et al. (1997) produced 55 mm h⁻¹ with a spatial uniformity coefficient of 93 %. Larger intensities are also possible; for example Fernandez-Galvez et al. (2008) used a simulator with a range of 0 to 120 mm h⁻¹ with an intentionally heterogeneous distribution. The drop size and rainfall intensity in drip tank rainfall

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simulators are controlled by the diameter of the holes and the pressure in the tank. Another type of small-plot rainfall simulator can use nozzles with a rotating disk that directs water to the plot (originally Grierson and Oades, 1977). These devices also attain uniformity coefficients of at least 75 % for a full range of intensities. In general, small-plot rainfall simulators are easily portable due to their small size. However, their limited size (approx. 1 m²) makes them not well suited to capturing plot-scale heterogeneity in surface properties.

For field plots (up to approximately 10 m per side), two standard choices for rainfall simulators are sweeping sprinklers and rows or arrays of nozzles. One of the early simulators of this type, developed by Moore et al. (1983), used oscillating nozzles to obtain intensities of 3.5 to 185 mm h⁻¹ with uniformity coefficients between 80.2 and 83.7 over a 4.6 by 6.1 m plot. Similarly, the “EMIRE” rainfall simulator of Esteves et al. (2000) has a base unit that irrigates a 5 m by 5 m area with mean intensities from 60 to 76 mm h⁻¹ and a mean uniformity coefficient of 80.2 %. Fister et al. (2012) developed a rainfall simulator for a 2.2 m² plot that achieves 85 to 96 mm h⁻¹ with a mean uniformity coefficient of 60 %; the emphasis was placed on reproducibility rather than uniformity in this case. One of the main advantages of these types of devices is that they can achieve higher drop velocities than drip tanks due to the water pressure in the nozzles. These devices also tend to be expandable to larger areas by reproducing the base unit. Typical disadvantages of the sweeping or oscillating sprinkler rainfall simulators is that the design can include intricate parts and may need to be run by computer, adding to the system expense and complexity. Also, the spray may be intermittent instead of constant as the nozzles sweep back and forth across or rotate around a plot. A typical disadvantage of rows or arrays of spray nozzles is that they tend to have lower uniformity coefficients due to stationary nozzle patterns. The design of Esteves et al. (2000) is close to meeting our needs due to its simplicity and expandability, but the stand pipes are secured with guy wires, making the unit not freestanding and potentially difficult to move or install in desired locations.

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Finally, for larger plots on the hillslope scale, the standard design is a rotating boom rainfall simulator, first developed by Swanson (1965). A single unit of this simulator irrigates a by 3.7 m by 10.7 m area at up to 120 mm h^{-1} intensity, with two units used to cover plots up to about 23 m long. In addition to covering large areas, these systems can be mounted on a trailer, simplifying transportation between sites. However, trailer-mounted systems cannot be used on steep hillslopes or in other areas that vehicles cannot reach, such as forests, which does not make them well suited for use in the wide range of soil and vegetation combinations needed for the present research. Additionally, these systems are usually more complicated than drip tanks or stationary nozzle systems, and due to the circular spray pattern, these systems are less efficient for covering square or rectangular plots.

Overall, lacking standard designs for rainfall simulators, individual researchers develop devices that suit their particular needs. In this case, the need is for an inexpensive, mechanically simple rainfall simulator system that achieves reasonable uniformity and can be used to accurately estimate soil characteristics such as the saturated hydraulic conductivity in a wide variety of vegetated locations. Such a device should be easily constructed in areas where access to specialized components is limited, easily repaired should any components of the system break, and easily transported from one field plot to the next. An additional requirement is that the plot is significantly larger than a single vegetation patch in order to capture the net runoff response that results from the reinfiltration of surface runoff.

To this end, this paper describes the design and testing of a rainfall simulator using primarily components that can be found in any hardware store. The system irrigates an area of approximately 15 m^2 with a range of rainfall intensity rates. In field tests of the system in Sardinia, Italy, data collected included soil moisture throughout the plot and runoff collected using a tipping bucket flow gage.

The intended use of this system is to estimate soil hydraulic properties, primarily the saturated hydraulic conductivity (K_s). There is a long history of using field data and numerical inversion to estimate soil parameters (e.g., Kool et al., 1985; Russo et al.,

1991; Simunek and Genuchten, 1996; Lassabatere et al., 2006; Ramos et al., 2006; Xu et al., 2012; Rienzner and Gandolfi, 2014). However, these experiments typically use point measurements using instruments such as single ring permeameters or tension disk infiltrometers and therefore do not capture the effective response at the plot scale.

5 Some studies have focused on vertical variability of K_s (Olyphant, 2003; Segal et al., 2008), but few have worked at the plot or hillslope scale. Ram et al. (2012) performed experiments on 4 m wide border strips and used nonlinear optimization to estimate the van Genuchten parameters α and n , but fixed K_s based on Guelph permeameter measurements. Also, Verbist et al. (2009) carried out rainfall simulation experiments
10 on 6 m \times 2 m plots and determined that the values of K_s resulting from fitting the two-term Philip infiltration model to measured data were comparable to the values obtained using full nonlinear optimization. Following this approach, we seek to use this new rainfall simulator and simplified parameter estimation methods to calculate physically reasonable values of the saturated hydraulic conductivity of the soil surface and soil
15 below the surface.

2 Rainfall simulator system design

Generally, there are ten important criteria for ideal rainfall simulators (Moore et al., 1983):

1. drop size distribution similar to that of natural rainfall
- 20 2. drop velocity similar to that of natural rainfall
3. uniform intensity and random drop size distribution over the plot
4. continuous application over the plot
5. nearly vertical impact angle
6. reproducible storm durations and intensities

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7. ability to perform in conditions such as high temperatures and wind
8. sufficient areal coverage to meet needs of experiment
9. plot-to-plot and site-to-site portability
10. low cost.

5 Typically, there is some compromise among the criteria as it is difficult to meet all of them. In addition to the primary objective of ease of portability, the choice was made to emphasize low cost, areal coverage, and reproducibility, with the added objective of being able to meet these criteria at at least two different rainfall intensities.

2.1 Simulator components

10 The water delivery system of this rainfall simulator consists of four lines of either 11 or 12 common pressure washing nozzles (Fig. 1a) mounted on a 4 m × 4 m × 2 m metal frame (Fig. 2). The nozzles point upwards at alternating angles of 48° and 54° from horizontal (Fig. 1b) and the entire system is connected using 1.5 cm inner diameter PVC pipe. Each of the four lines is operated independently with a valve and pressure gauge to regulate the flow. Additionally, a plastic mesh suspended under the nozzles served the purpose of helping to randomize the spray pattern (Foster et al., 2000). Supporting components of the system are the metal supports for the nozzles and piping, a 2 m³ plastic tank for use in the field, a submersible pump, and the tubing and valves necessary to connect the pump to the four inlet lines.

2.2 Instrumentation and data collection

20 An in-line volumetric flow meter on the tubing between the pump and the nozzle lines measured the total amount of water delivered to the system. The reading on this flow meter was recorded at the beginning and end of each experiment. Soil moisture probes (Campbell Scientific CS616), inserted at approximately a 30° angle from horizontal,

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measured soil moisture in the top 15 cm of the soil. A data logger recorded readings from the probes at one second intervals. To measure the amount of water delivered to the plot, two 10 cm simple rain gages were placed inside the irrigated area.

Runoff from the plot was collected using a tipping bucket flow gage (Fig. 3), based on the design of Chow (1976). The metal box is open on the sides and supports two plexiglas buckets that hold approximately two liters each. The recording mechanism is the magnetic switch component taken from a tipping bucket rain gage. Water flows into the bucket from a hose attached to the back of the structure (see Fig. 3b), and the data logger records the time of the pulse created by each tip.

Since the tipping of the bucket is not instantaneous, the amount of water in the bucket can vary with the flow rate. Specifically, while the mass of water needed to make the bucket tip is constant, water continues to collect in the bucket as it tips. Therefore, a higher flow rate can result in a higher volume of water in each bucket. To calibrate the flow gage for this effect, water with a known flow rate q was directed into the flow gage and the instrument was video recorded for the amount of time needed for several tips to occur. The volume in each of the tip of the bucket was inferred from the flow rate and the amount of time the water stream spent in the bucket, $V = q(t_{\text{exit}} - t_{\text{enter}})$. t_{enter} and t_{exit} were gathered from the time stamps on the corresponding frames of the video. This process was repeated several times at different known flow rates. Figure 4 shows the data points and relationship of the bucket volume as a function of the tipping frequency as tested in the lab, namely

$$V = 0.0144f + 2.15, \quad (1)$$

where f is the frequency in tips per minute and V is the bucket volume in liters. The volume is nearly constant, only ranging between 2.1 and 2.3 L per tip, but the data do show a weak dependence on tipping rate. For each pulse recorded during the experiments, the corresponding number of tips per minute was calculated as $f = 1/(t_j - t_{j-1})$, where t is measured in minutes, then Eq. (1) was used to calculate the expected volume.

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3 Rainfall simulator testing

Before using the rainfall simulator in the field, laboratory tests were performed to determine the uniformity of the depth of water in the plot with several configurations. Figure 5 shows three potential configurations for this particular system. If a nozzle was not used in a given configuration, it was removed and replaced with a short plugged length of PVC pipe. The first configuration (Fig. 5b) has 46 nozzles, with 33.3 cm between the center of each nozzle. In the second configuration, which has 31 nozzles, every third nozzle is removed, resulting in nozzle spacings of 33.3 and 66.6 cm. The third configuration has 24 nozzles with 66.6 cm between all nozzles.

For each configuration, the simulator was run for 30 min with water collected in 63 containers arrayed beneath the sprinklers. The 10 cm diameter collection containers were placed in an array with 9 containers in the direction of the ground slope with and 7 containers in the direction perpendicular to the ground slope, both with 50 cm spacing. The uniformity of the applied water was calculated using the Christiansen Coefficient of Uniformity (Christiansen, 1942):

$$CU = 100 \times \left(1 - \frac{\sum |x - \bar{x}|}{n\bar{x}} \right), \quad (2)$$

where \bar{x} is the average of all of the measurements, $\sum |x - \bar{x}|$ is the sum of the individual deviations from the mean, and n is the number of measurements taken. Figure 6 shows the rainfall intensity for the three nozzle configurations. The three configurations, with 46, 31, and 24 nozzles, had average rainfall intensities of 61.6, 43.3, and 31.8 mm h⁻¹, respectively. The standard deviations of the three configurations of 18.2, 18.9, and 14.9 mm h⁻¹ are similar to those of the basic unit of the “EMIRE” rainfall simulator, which covers a similar area using stationary nozzles (Esteves et al., 2000). Since each configuration has an operating pressure of 8×10^3 Pa, the average intensity decreases with a decreasing number of nozzles. However, the coefficient of uniformity also decreases with fewer nozzles due to less spatial coverage of the plot. Ideally, one

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seeks a CU near 80 % (e.g. Neff, 1979); however, in the present application there is a tradeoff between uniformity and other objectives such as size, cost, and portability. Accordingly, while the CU values of 75, 65, and 62 % are not ideal, they are a reasonable compromise for the objective of emphasizing the applicability and reproducibility of our results over the uniformity (Fister et al., 2012). Additionally, rainfall simulators with the highest CU values are typically the smallest (1 m² range); therefore we consider our CU values to be acceptable based on the size of the plot under consideration.

Drop size distributions (DSD) were measured with a disdrometer for several locations within the irrigation area for both configuration 1 (61.6 mm h⁻¹) and configuration 3 (31.8 mm h⁻¹). In each case, the DSD for the number of drops in 300 s ranged from 0.25 to 3.3 mm drop diameter. The maximum peak frequency occurred at 1 mm with additional peaks at 2 and 2.6 mm. These results are similar to those documented by Sauvageot and Lacaux (1995). The full details of the disdrometer analysis are shown in Corona (2013).

4 Field experiments

4.1 Field site

The experiments were carried out at an existing field site in Orroli, Italy, on the island of Sardinia (39°41′12.57″ N, 9°16′30.34″ E) in July and August 2010. The site was accessible by a path leading downhill from the road for approximately 200 m. The simulator components were carried from the road to the site and assembled around the chosen plot. The water tank was carried to the site while empty and then filled from a water truck located on the road. During the experiments, a generator powered a submersible pump in the water tank to deliver water to the rainfall simulator.

In the chosen field plot, tall grasses covered approximately 95 % of the ground surface, and the land has a gentle slope of approximately 4.6°. The soil is a silt loam with bulk density of 1.38 g cm⁻³ (Montaldo et al., 2008). Around the frame of the rainfall

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simulator, thin sheet metal was inserted approximately 10 cm into the ground to help prevent lateral flow of water across the plot boundaries. To collect the surface runoff, a shallow trench was dug at the bottom edge of the plot, and a plastic ledge was inserted into the ground just below the surface to direct the water onto plastic sheeting.

From here, the water flowed by gravity into a tube that was connected to the tipping bucket flow gage. As shown in Fig. 7, the soil moisture probes were placed at 1 m intervals in the irrigated section of the plot. Additionally, to verify rainfall intensity, two rain gages were placed inside the plot. A slight rill already existed in the plot due to an animal footpath, and as a result the outlet of the plot was approximately at the location of soil moisture probe 14.

4.2 Experiments

Three experiments were conducted over the course of several days. Each experiment had a different combination of initial soil moisture (θ_0), experiment duration (T , min), and rainfall intensity (p , mm h^{-1}). For each experiment period, the rainfall was allowed to continue until the soil moisture and runoff reached a constant rate. The test conditions are described in Table 2.

5 Method of soil parameter estimation

The intended use of this rainfall simulator is to glean surface properties from rainfall experiments. Alberts et al. (1995) demonstrated a sufficient fit of an infiltration model to observed data can be found by running the model with a range of values for the saturated hydraulic conductivity and selecting the value that minimizes model error with respect to the observed infiltration. This method can be applied here to obtain $K_{s, \text{bulk}}$ and $K_{s, \text{surf}}$ from the experiment data in a similar manner as Verbist et al. (2009). Here, we define $K_{s, \text{surf}}$ and $K_{s, \text{bulk}}$ to be the saturated hydraulic conductivities of the surface layer and remainder of the soil profile, respectively. We use these two values

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rather than one K_s value to separate the behavior of the soil surface, which determines the amount of water that infiltrates, from that of the bulk soil, which determines overall soil water content. Based on previous work for this site (Montaldo et al., 2008), the estimated saturated hydraulic conductivity for this soil is approximately 18 mm h^{-1} . If K_s values obtained from model fitting are close to this previous estimate, then the rainfall simulator indeed performs as intended.

A common model for infiltration is the two-parameter Philip equation (Philip, 1957):

$$i_t = \frac{1}{2} S_t t^{-1/2} + A_t, \quad (3)$$

where i [L T^{-1}] is the infiltration rate at time t , S_t [$\text{L T}^{-1/2}$] is the sorptivity, and A_t [L T^{-1}] is a parameter similar to, but not necessarily equal to the saturated hydraulic conductivity K_s . It is used in this analysis because of its simplicity and success in short time scales. In this implementation, the sorptivity expression of Sivapalan et al. (1987) is used:

$$S_t = \left[\left(2K_{s, \text{surf}} (\theta_s - \theta_t)^2 \left(-\frac{\psi_{ae}}{\theta_s - \theta_h} \right) \right) \left(\frac{1}{(2b+3) + \frac{1}{2b} - 1} + \frac{\theta_s - \theta_h}{\theta_s - \theta_t} \right) \right]^{1/2}, \quad (4)$$

where $K_{s, \text{surf}}$ is the saturated hydraulic conductivity of the surface layer [L T^{-1}], θ_s is the saturated volumetric soil moisture (porosity), θ_t is the volumetric soil moisture at time t , θ_h is the hydrostatic (minimum) volumetric soil moisture, ψ_{ae} is the air entry suction head [L], and b is the soil water retention curve parameter (Clapp and Hornberger, 1978). Eagleson (1978) showed the application of Philip's (1960) work on the diffusion equation at short times as

$$A = \frac{1}{2} (K_s + K_0). \quad (5)$$

Using the Campbell (1974) definition of hydraulic conductivity for the surface layer as $K_{\text{surf}} = K_{\text{s, surf}}(\theta_t/\theta_s)^{2b+3}$, Eq. (5) may be written as a dynamic parameter, namely

$$A_t = \frac{1}{2}K_{\text{s, surf}} \left(1 + (\theta_t/\theta_s)^{2b+3}\right). \quad (6)$$

The unsaturated hydraulic conductivity of the bulk soil is also calculated using the definition of Campbell (1974), namely

$$K_{\text{bulk}} = K_{\text{s, bulk}} \left(\frac{\theta_t}{\theta_s}\right)^{2b+3} \quad (7)$$

where $K_{\text{s, bulk}}$ [L T^{-1}] is the saturated hydraulic conductivity of the bulk soil rather than the soil surface.

The calculation of soil moisture comes from a one-dimensional water balance on a soil layer with thickness dz , beginning with $\Delta S = Q_{\text{in}} - Q_{\text{out}}$, where S is the water storage in the layer, Q_{in} is the water entering the layer, and Q_{out} is the water exiting the layer. S can be represented as $\theta \Delta z$, so $\Delta S = \Delta z(\theta_t - \theta_{t-\Delta t})$. Q_{in} is the water infiltrating from above, so during a small time step Δt , $Q_{\text{in}} = i \Delta t$. Neglecting evapotranspiration, Q_{out} is the drainage of water through the bottom of the layer. Use of Darcy's Law, $q = -K \frac{dh}{dz}$, with a unit gradient yields $Q_{\text{out}} = K \Delta t$. Therefore, the soil moisture prediction equation is

$$\frac{\theta_t - \theta_{t-\Delta t}}{\Delta t} z = i_t - K_{\text{bulk}}, \text{ or}$$

$$\theta_t = \theta_{t-\Delta t} + (i_t - K_{\text{bulk}}) \frac{\Delta t}{z} \quad (8)$$

where z [L] is the soil depth being considered and i_t [L T^{-1}] is the infiltration rate, which is the smaller of the maximum infiltration in Eq. (3) and the precipitation rate. See Table 3 for the parameters used in this analysis.

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Infiltration is modeled using the above equations with a time step of $\Delta t = 1$ min. First, A and S are calculated according to Eqs. (6) and (4). Then, the potential infiltration i_t^* is calculated using Eq. (3). The actual infiltration i_t is $\min(i_t^*, \rho)$. The current bulk soil hydraulic conductivity is calculated with Eq. (7), and then the bulk soil moisture θ_t is calculated using Eq. (8).

After the first set of calculations, an adjustment is made for the time before ponding since the original Philip equation is for ponded conditions only. Following Dingman (2004), t_s is defined as the first time in the original calculations where $\rho > i_t^*$. The total infiltrated volume before ponding is

$$I_p = \sum_{t=0}^{t_s} i_t^* \quad (9)$$

The time to ponding is

$$t_p = \frac{I_p}{\rho} \quad (10)$$

and the condensed ponding time is

$$t_{cp} = t_p - t_s \quad (11)$$

Then, the calculations for $i_{\hat{t}}$, $K_{\hat{t}}$, and $\theta_{\hat{t}}$ are repeated using $\hat{t} = t - t_{cp}$. Finally, the modeled runoff is calculated as

$$f_{\hat{t}} = \max(\rho - i_{\hat{t}}, 0) \quad (12)$$

Equations (3), (4), and (6) show that the infiltration of water through the surface, and accordingly the runoff, depend strongly on $K_{s, \text{surf}}$ but not on $K_{s, \text{bulk}}$. Likewise, Eqs. (8) and (7) show that $K_{s, \text{bulk}}$ affects θ_t but not i_t . Therefore,

$$e_{Q_t} = |Q_{t, \text{mod}} - Q_{t, \text{obs}}| \quad (13)$$

and

$$e_{\theta} = \left(\frac{\sum_{t=0}^T (\theta_{t,\text{mod}} - \theta_{t,\text{obs}})^2}{T} \right)^{1/2} \quad (14)$$

can be used as measures of error that, when minimized, indicate the optimal values of $K_{s, \text{surf}}$ and $K_{s, \text{bulk}}$, respectively. e_{Q_f} uses the final Q value to capture the overall behavior of the plot in producing runoff, and e_{θ} uses the time series of θ to capture the evolution of soil moisture during the experiment.

To optimize $K_{s, \text{surf}}$ and $K_{s, \text{bulk}}$, the infiltration model was run using all combinations of the two values ranging from 1 to 30 mm h^{-1} at a step of $\Delta K = 10^{-7} \text{ms}^{-1} = 0.36 \text{mm h}^{-1}$. e_{Q_f} and e_{θ} were calculated for each combination, and the combination with the smallest value of

$$e^* = e_{Q_f} + e_{\theta} \quad (15)$$

was selected as the optimal values of $K_{s, \text{surf}}$ and $K_{s, \text{bulk}}$.

6 Results and discussion

6.1 Soil moisture

While the ultimate goal is to accurately model the infiltration and runoff at the plot scale, soil moisture values can be used in the calculation of soil saturated hydraulic conductivity. Since the soil moisture values vary throughout the plot, spatial averaging of the measured data is employed. However, examining the spatial variability within the plot allows for a better understanding of each experiment.

According to previous field work at this site (Montaldo et al., 2008) and examination of the measured soil moisture at saturation, the soil has an average porosity of $\theta_s = 0.568$.

This seems to suggest a change in water input from solely local rainfall to rainfall and overland flow from upslope.

The third experiment, performed on 5 August (Fig. 8c), started with $0.55 \leq \hat{\theta}_{x,0} \leq 0.78$, and a mean value of 0.67, placing it in between Experiments 1 and 2 in terms of initial soil moisture. However, this experiment was performed with nozzle configuration 3, meaning a rainfall intensity of 31.8 mm h^{-1} . As a result, it took nearly three hours for the whole plot to reach saturation. Again, probe 12 takes the longest to reach saturation. There is also a sharp change in the slopes of the soil moisture time series for probes 4, 5, and 6; for each of these three probes, the change in slope occurs soon after the probes immediately downslope (8, 9, and 10, respectively) reach saturation. This feature again seems to indicate a change in the contribution from rainfall only to rainfall plus overland flow from upslope.

Overall, the time to saturation in each of the three experiments increased as the rainfall intensity and the initial soil moisture decreased, as expected. To investigate the relationship between the location of the soil moisture probes and the spatial variability of the measurements, the soil moisture results have been categorized by location. Probes 2, 3, 5, 6, 9, 10, 13, and 14 are considered interior probes, and the remaining are considered border probes. Figure 9 shows the comparison of the averages of the interior and border soil moisture probes compared to the overall average. In each case, the border soil moisture probes overall have lower readings than the interior probes. The difference is particularly pronounced in Experiment 2, when the difference between the interior and border probes in the middle of the experiment exceeds the difference in the initial values. This is also true, to a lesser degree, in Experiment 3. A possible explanation for the discrepancy is that the rainfall intensity distribution as shown in Fig. 6 generally appears to favor the interior of the plot over the edges.

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6.2 Runoff

The tips of the tipping bucket flow gage were converted to time series of runoff rate and cumulative runoff through Eq. (1). Specifically, with t_i being the time in minutes of the i th bucket tip and V_i is the volume of that tip in liters, the runoff rate q_i in mm h^{-1} is

$$q_i = \frac{V_i}{t_i - t_{i-1}} \times \alpha \quad (18)$$

where

$$\alpha = 60 \text{ min hr}^{-1} \cdot 0.001 \text{ m}^3 \text{ L}^{-1} \cdot 1000 \text{ mm m}^{-1} / 15.12 \text{ m}^2 = 3.968 \quad (19)$$

is the constant that converts L min^{-1} to mm h^{-1} . The infiltration rate that results from this runoff is

$$i_j = p - q_j. \quad (20)$$

The cumulative runoff Q_i is

$$Q_i = \sum_{j=0}^i V_j \times \beta \quad (21)$$

where

$$\beta = 0.001 \text{ m}^3 \text{ L}^{-1} \cdot 1000 \text{ mm m}^{-1} / 15.12 \text{ m}^2 = 0.06614 \quad (22)$$

converts L to mm. Since the tipping bucket flow gage causes q_i to oscillate around what should be a smooth time series, a 4-point moving average was calculated as

$$\hat{q}_i = \frac{1}{4} \sum_{j=i-3}^i q_j. \quad (23)$$

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Figure 10 shows both \hat{q}_i and Q_i , in addition to the times at which the soil moisture probes reached full saturation. These figures provide a visual check of whether or not the rainfall simulator provides reasonable results. In each of the three experiments, runoff begins approximately when the first probe reaches saturation. The first probe to reach saturation in all three experiments was probe 14, which is next to the outlet of the plot. With $t_{x,s}$ as the time when probe x reaches saturation, the average time to saturation for each experiment, \bar{t}_s , is defined as $\frac{1}{n} \sum_{x=1}^n t_{x,s}$ with $n = 15$. \bar{t}_s occurs while the runoff rate is increasing, and the runoff rate changes little after the final probe reaches saturation. The exception is Experiment 2, which did not continue beyond time when the final probe reached saturation.

It is also useful to compare the runoff rates between experiments. For both Experiments 1 and 2, the average of the runoff rates near the end of the experiment is approximately 47.6 mm h^{-1} . When compared to the applied rainfall rate of 61.6 mm h^{-1} , the final runoff coefficient ($C_f = q_f/p$) once the plot is fully saturated is 0.77. For Experiment 3, the average rate near the end of the experiment is 24.2 mm h^{-1} , or $C_f = 0.76$. The similar runoff coefficients emphasize the proportionality between rainfall and runoff rates.

The overall runoff coefficient is

$$C = Q_f/P \quad (24)$$

where Q_f is the cumulative runoff for the entire experiment and P is the total precipitation applied during duration T . While the C_f values are equivalent across the experiments, the C values differ, with values of 0.61, 0.46, and 0.51. This is expected due to the varying antecedent soil moisture states; as higher θ_0 values are correspond to higher C values. Therefore, this brief analysis of the measured data of soil moisture and runoff, the experiments show that the rainfall simulator produces physically reasonable results.

6.3 Parameter estimation

Using the estimation method shown in Sect. 5, the $K_{s, surf}$ values for the three experiments are 16.5, 8.56, and 4.24 mmh⁻¹, respectively, and the $K_{s, bulk}$ values are 11.8, 7.48, and 8.20 mmh⁻¹. Verbist et al. (2009) found a K_s range of approximately 17 to 55 mmh⁻¹ for a coarse loamy soil for four replicates at 120 mmh⁻¹ precipitation intensity. Since our experiments were performed at two different precipitation intensities, we consider the overall range of 4.24 to 16.5 mmh⁻¹ to be acceptable.

The comparison of the observed and modeled time series based on these hydraulic conductivity values are shown in Fig. 11. Additionally, the observed final infiltration rate for each experiment is defined as

$$i_f = \rho - \frac{\sum_{t=t_c}^T (q_t)}{T - t_c} \quad (25)$$

where t_c is the time at which the infiltration and runoff appear to become constant. The comparison of $K_{s, surf}$, $K_{s, bulk}$, and i_f is shown in Fig. 12. The parameter which shows the largest range over the three experiments is $K_{s, surf}$. One possible reason for the reduction in $K_{s, surf}$, particularly between the first two experiments, is surface sealing that can occur after water is applied to the surface. Another reason for the difference between $K_{s, surf}$ in Experiments 1 and 2 and that of 3 is the smaller rainfall intensity in the third experiment. The most stable values are for $K_{s, bulk}$. This is likely due to the behavior of the bulk soil being less affected by the conditions that vary between experiments. The literature value of saturated hydraulic conductivity for a silt loam with a mean clay fraction of 0.14 is approximately 25 mmh⁻¹ (Clapp and Hornberger, 1978). The optimized values here are smaller, despite having a clay fraction of 0.05. One possible reason for the disparity is the high silt fraction in this soil (0.76). The most apparent reason, however is that field values of saturated hydraulic conductivity are highly dependent on rainfall intensity (e.g., Bowyer-Bower, 1993; Stone et al., 2008; Langhans

et al., 2010). Larger experimental rainfall intensities are likely to lead to higher $K_{s, \text{surf}}$ and $K_{s, \text{bulk}}$ values of this soil.

Since the soil properties do not change between experiments, ideally the range of $K_{s, \text{surf}}$ and $K_{s, \text{bulk}}$ values should be smaller than is shown in Fig. 12. The remaining model parameters that have not been determined by field data are b and ψ_{ae} , so they can be included in the optimization. Values of b from 5 to 9 with a step of $\Delta b = 1$ and values of ψ_{ae} from 0.3 to 0.8 m with $\Delta\psi = 0.083$ m were incorporated into the parameter space so each model realization used a unique combination of the four parameters. Again, the parameter assignment that minimized e^* was selected as the optimal set for each equation. These results are shown in the second block of Table 4.

b and ψ_{ae} are constant for one plot, so the final step was to average the values of b and ψ_{ae} for the three experiments to get the optimal values $\hat{b} = 7$ and $\hat{\psi}_{ae} = 0.44$ m. The model was run using \hat{b} and $\hat{\psi}_{ae}$ along with the combinations of $K_{s, \text{surf}}$ and $K_{s, \text{bulk}}$ used in the first optimization. The values that minimize e^* are $K_{s, \text{surf}} = 11.8, 12.5,$ and 6.04 mm h^{-1} , along with $K_{s, \text{bulk}} = 18.6, 13.6,$ and 6.76 mm h^{-1} . The modeled Q and θ are shown in Fig. 13, and the rate comparison is in Fig. 14, with the overall results in Table 4. The $K_{s, \text{bulk}}$ values for the higher intensity experiments are close to the 18 mm h^{-1} that is expected for this soil, which indicates that despite the heterogeneity of rainfall intensity throughout the plot, this rainfall simulator can be used to generate reasonable estimates of the primary soil properties that govern infiltration and runoff.

While little difference can be seen between Figs. 11 and 13, a comparison of Figs. 12 and 14 emphasizes the changes in $K_{s, \text{surf}}$ and $K_{s, \text{bulk}}$ following the optimization process. The ranges of the rates for Experiments 2 and 3 are smaller than with the previously assumed values for b and ψ_{ae} . Additionally, $K_{s, \text{bulk}}$ is closer to the final infiltration rate, particularly for Experiment 2, and the $K_{s, \text{bulk}}$ values for Experiments 1 and 2 are more comparable, which is an improvement since these two experiments have the same water application rate. The reduction in $K_{s, \text{surf}}$, from Experiment 1 to Experiment 2 is still present, which is still likely a result of surface sealing in between experiments. Since the values of $K_{s, \text{surf}}$ and $K_{s, \text{bulk}}$ are more consistent within experiments and

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across experiments, it appears that this optimization method has improved the estimates for the saturated hydraulic conductivities of the soil in this plot.

7 Conclusions

This paper presented the design, construction, and testing of a low-cost rainfall simulator system. This system has several significant advantages. Primarily, it is comprised of readily available and inexpensive materials. Additionally, the 4 m by 4 m module can be replicated to expand the areal coverage of the simulator. It is fairly easy to change the configuration of the system to produce different intensities, and the low operating pressure means that the power requirements of the system are low as well.

There are some limitations of the system. The rainfall intensity pattern is not as uniform as typically desired for rainfall simulators. Additionally, since some water falls on the surface outside the effective irrigated area, the simulator is not as water efficient as it could be. Due to the low operating pressure, the system needs to be close to level to provide even distribution of water, so it cannot be used on steep slopes. Finally, since the water is sprayed upwards, the system does not work properly in windy conditions without wind shields that are at least 3 m tall in place around the plot frame.

After measuring soil moisture and runoff during three experimental runs of the system, analysis using the Philip infiltration model showed that this rainfall simulator can produce data that allow for the estimation of reasonable values for the saturated hydraulic conductivity of the soil. Specifically, we showed that performing a three-step parameter search optimization that first estimates K_s with assumed values of b and ψ_{ae} , then estimates b and ψ_{ae} using the K_s initial estimates, then uses b and ψ_{ae} to get new values for the K_s estimates allows for the estimation of a total of four soil hydraulic parameters. The implication of this result is the ability to use this rainfall simulator system to estimate parameters needed for hydrologic modeling with less reliance on prior knowledge.

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Continued work with this system will be to use the rainfall simulator on plots with different vegetation fractions, types, and patterns. The purpose of gathering the data in these different conditions will be to calibrate model parameters for a coupled hydrologic-vegetation dynamics model. This model will be used to investigate how changes in vegetation growth during the transition from the rainy season to the dry season in Sardinia will affect overall surface runoff volumes.

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Table 1. Summary of nozzle configurations and performance.

Configuration	#of nozzles	Pressure (Pa × 10 ³)	Mean Intensity (mm h ⁻¹)	Standard deviation (mm h ⁻¹)	Coefficient of Uniformity (%)
1	46	8	61.6	18.2	75.7
2	31	8	43.3	18.9	65.2
3	24	8	31.8	14.9	62.1

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Table 2. Field test conditions.

Exp #	Date	$\theta_0(-)$	ρ (mm h ⁻¹)	T (min)
1	29 Jul	0.50	61.6	37
2	3 Aug	0.32	61.6	90
3	5 Aug	0.38	31.8	169

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Table 3. Parameters used in Philip infiltration model.

Symbol	Value	Units	Source
θ_s	0.568	[-]	Field data
θ_h	0.08	[-]	Montaldo et al. (2008)
ψ_{ae}	0.79	[m]	Clapp and Hornberger (1978)
b	8	[-]	Clapp and Hornberger (1978)
z	0.15	[m]	Field data

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Table 4. Parameters for the model optimization (a) using assumed values of b and ψ_{ae} , (b) including all four parameters in the optimization, and (c) using $\hat{\psi}_{ae}$ and \hat{b} .

Experiment	$K_{s, \text{bulk}}$ mm h^{-1}	$K_{s, \text{surf}}$ mm h^{-1}	b [-]	ψ_{ae} m	e_{θ} [-]	e_{Q_i} mm
(a) Optimization with assumed ψ_{ae} and b						
1	11.8	16.5			2.84×10^{-3}	6.94×10^{-3}
2	7.48	8.56	8	0.79	2.45×10^{-2}	6.61×10^{-5}
3	8.20	4.24			8.21×10^{-3}	2.10×10^{-2}
(b) Optimization using all parameters						
1	11.2	18.0	5	0.63	2.69×10^{-3}	4.63×10^{-4}
2	13.7	14.8	9	0.38	2.10×10^{-2}	8.60×10^{-4}
3	6.52	8.68	7	0.30	5.50×10^{-3}	2.12×10^{-3}
(c) Optimization using $\hat{\psi}_{ae}$ and \hat{b}						
1	11.8	18.6			2.82×10^{-3}	2.05×10^{-3}
2	12.5	13.6	7	0.44	2.10×10^{-3}	6.94×10^{-3}
3	6.04	6.76			6.50×10^{-3}	5.56×10^{-2}

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11, 4267–4310, 2014

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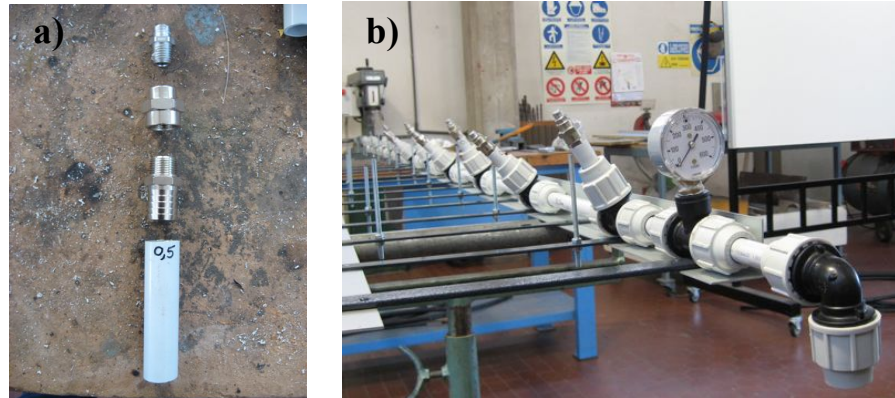


Fig. 1. Nozzle line setup. **(a)** Nozzle assembly. The threaded pressure washing nozzle was connected to a short length of PVC pipe using two connectors. **(b)** Alternating angles of nozzles and pressure gauge used to regulate flow.

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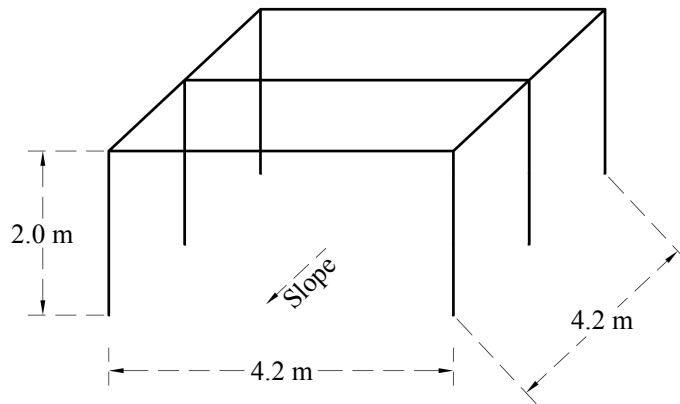


Fig. 2. Schematic of rainfall simulator frame.

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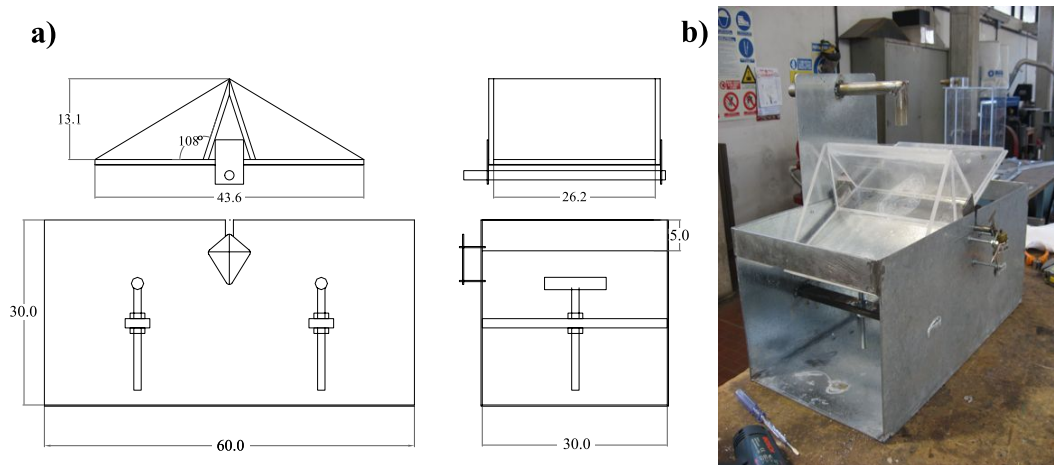
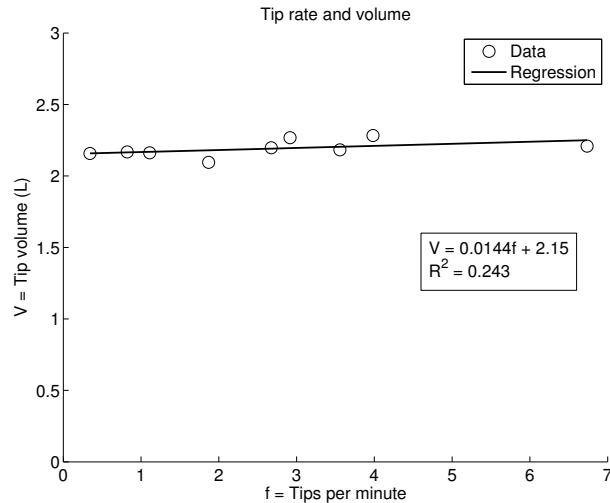


Fig. 3. Tipping bucket flow gage. **(a)** Dimensions of the flow gage. All dimensions in centimeters. **(b)** Completed flow gage constructed from sheet metal and plexiglass.

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**Fig. 4.** Data and regression for bucket volume as a function of tipping frequency.[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

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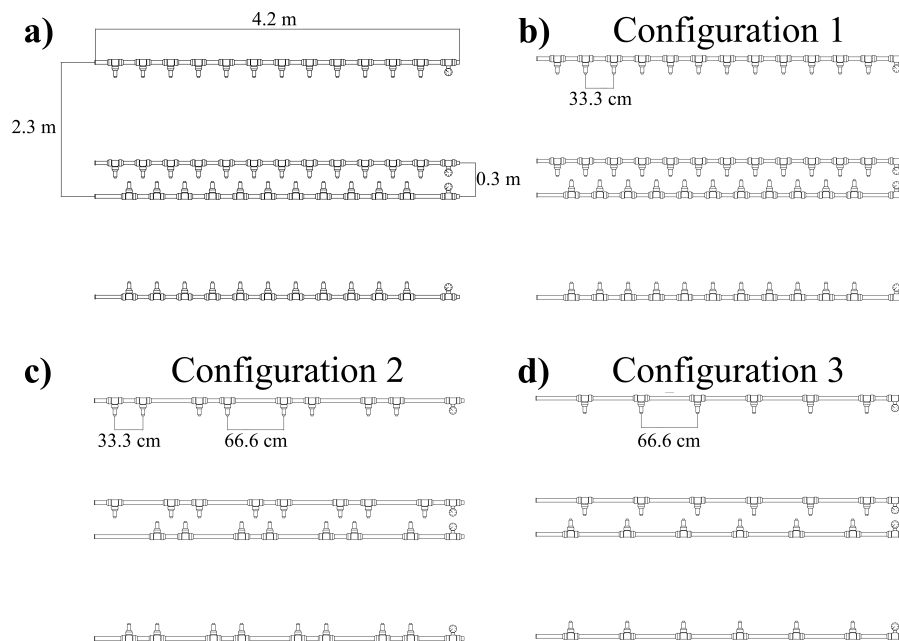


Fig. 5. Nozzle configurations for three intensities with nozzle spacing, resulting intensity, and experimental coefficient of uniformity of water application. **(a)** Spacing between nozzle lines. **(b)** 46 nozzles, 61.6 mm h^{-1} , $\text{CU} = 75.64$. **(c)** 31 nozzles, 43.3 mm h^{-1} , $\text{CU} = 65.16$. **(d)** 24 nozzles, 31.8 mm h^{-1} , $\text{CU} = 62.1$.

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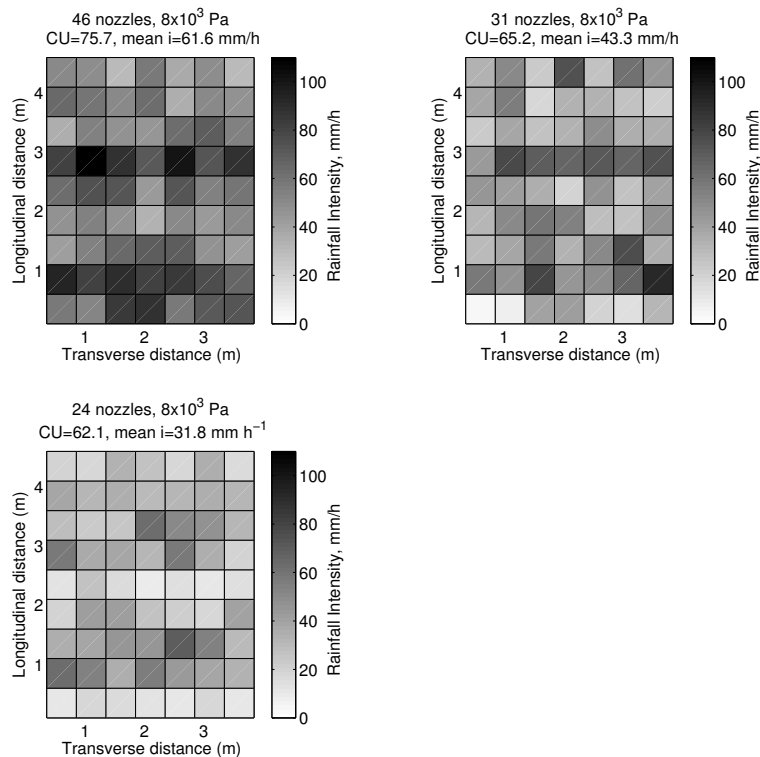


Fig. 6. Distribution of rainfall intensities and efficient of uniformity for the three nozzle configurations.

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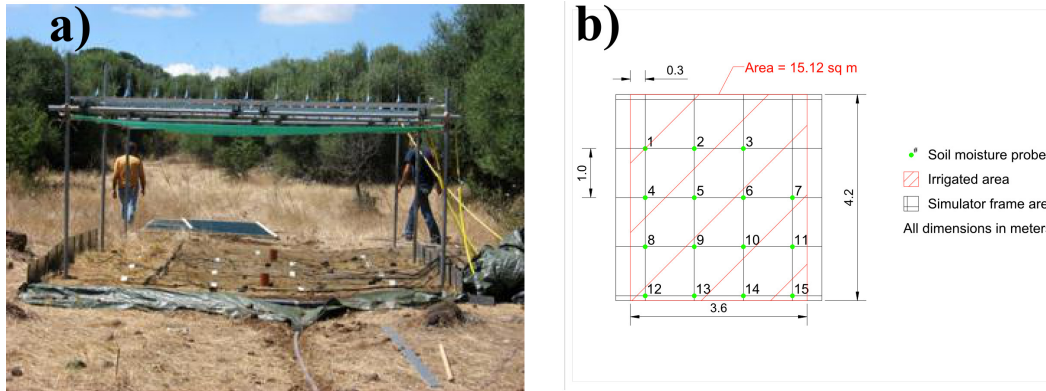


Fig. 7. Field experiment setup. **(a)** Field setup, including soil moisture probes, border around edges, rain gages, nozzles, distribution mesh, and collection system. **(b)** Details of the plot area, with soil moisture probe location and effective coverage area. The outlet of the plot is at probe 14.

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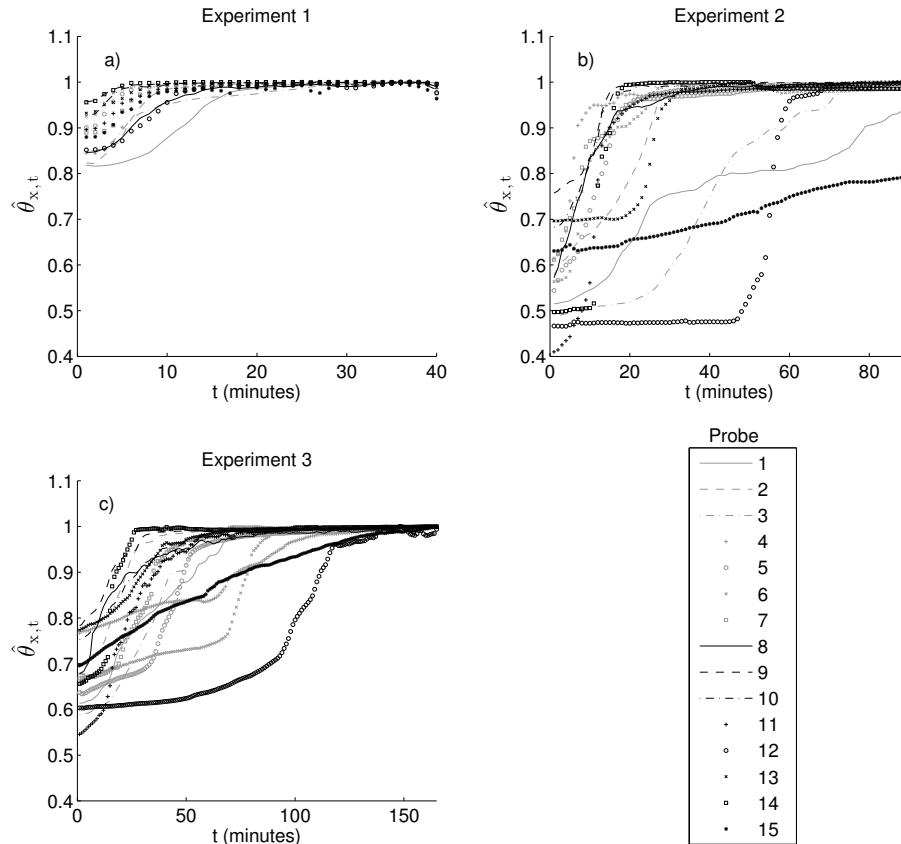


Fig. 8. Soil moisture as $\hat{\theta}_{x,t}$ for each TDR probe. **(a)** Experiment 1. The soil started near saturation due to previous test runs. **(b)** Experiment 2. **(c)** Experiment 3.

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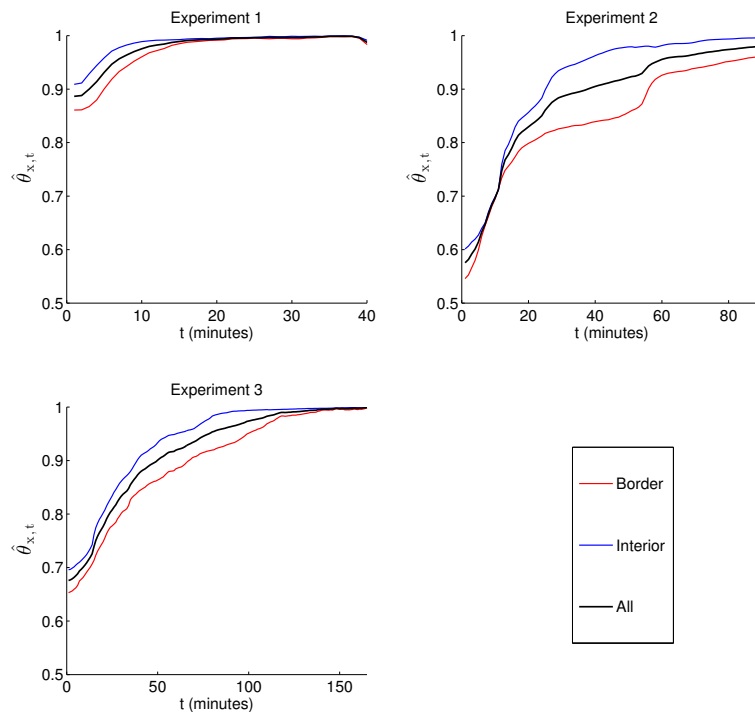


Fig. 9. Soil moisture averaged for all probes, border probes, and interior probes for the three experiments.

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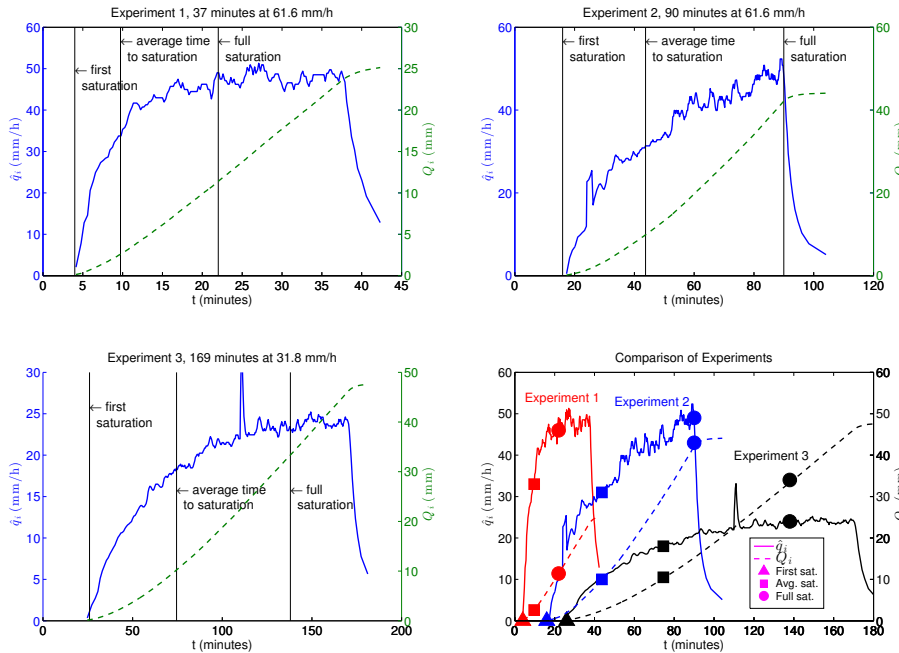


Fig. 10. Runoff rate and cumulative runoff for the three experiments. The vertical lines represent times when the soil moisture probes reached maximum values.

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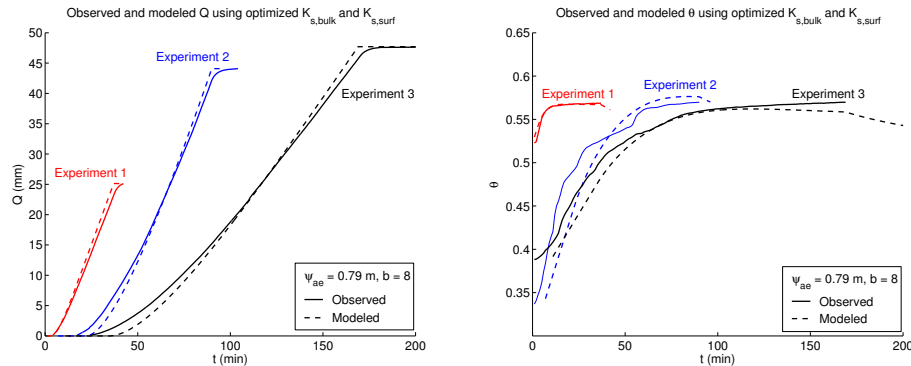


Fig. 11. Observed and modeled Q and θ based on optimized hydraulic conductivities for the three experiments with $\psi_{ae} = 0.79$ m and $b = 8$. Optimized hydraulic conductivity values are: Experiment 1: $K_{s, surf} = 16.5$ mm h⁻¹, $K_{s, bulk} = 11.8$ mm h⁻¹; Experiment 2: $K_{s, surf} = 8.56$ mm h⁻¹, $K_{s, bulk} = 7.48$ mm h⁻¹; Experiment 3: $K_{s, surf} = 4.24$ mm h⁻¹, $K_{s, bulk} = 8.2$ mm h⁻¹.

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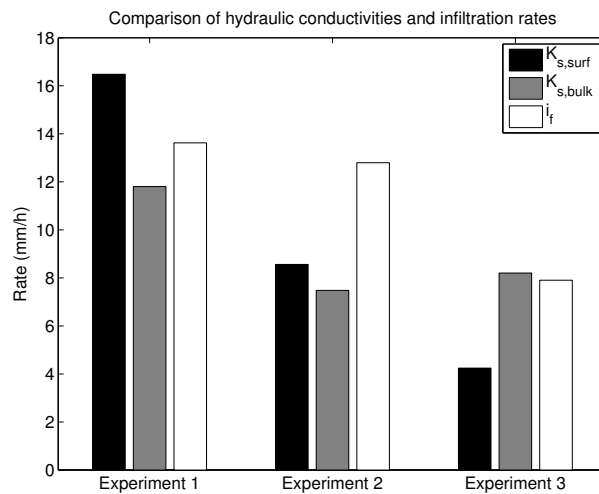


Fig. 12. Comparison of infiltration and conductivity rates for the three experiments.

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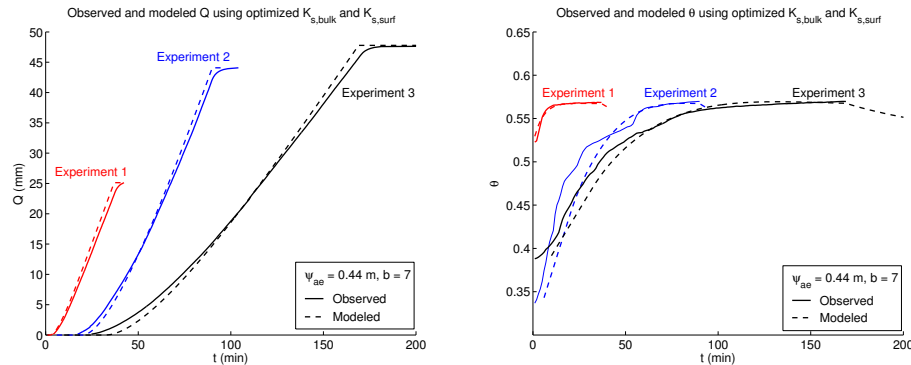


Fig. 13. Observed and modeled Q and θ using $\hat{b} = 7$ and $\hat{\psi}_{ae} = 0.44$ m. Optimized hydraulic conductivity values are: Experiment 1: $K_{s,surf} = 18.6 \text{ mm h}^{-1}$, $K_{s,bulk} = 11.8 \text{ mm h}^{-1}$; Experiment 2: $K_{s,surf} = 13.6 \text{ mm h}^{-1}$, $K_{s,bulk} = 12.5 \text{ mm h}^{-1}$; Experiment 3: $K_{s,surf} = 6.76 \text{ mm h}^{-1}$, $K_{s,bulk} = 6.04 \text{ mm h}^{-1}$.

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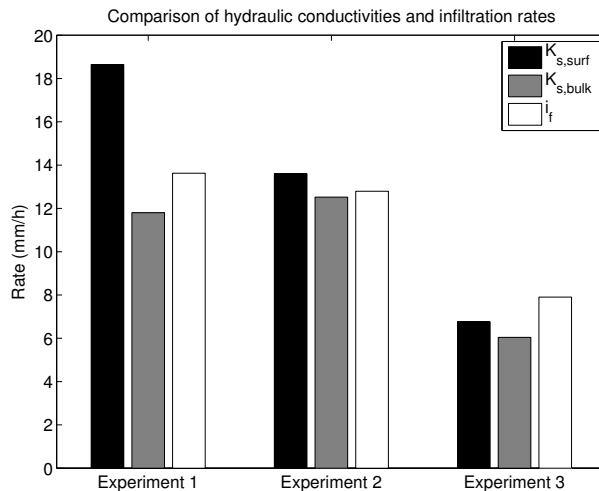


Fig. 14. Rates for the optimization using \hat{b} and $\hat{\psi}_{ae}$.

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