**Dear Editor**,

The English has been improved using AJE services as you requested. The cover letter from AJE is produced in this document and the modifications we did are highlighted in yellow in the following pages after the AJE letter.

These edits are described below. We performed all the requested modifications proposed by reviewers #2 and #1. We are indicating the way it has been answered (in blue below). We changed the title as requested, but it has been modified by the AJE review.

We tried also to rework a bit the abstract as requested by reviewer #2. We made several modifications as requested by the reviewers.

I hope this paper is now acceptable for your journal.

**Sincerely yours** 

**Michel** 

# 1 Answer to Anonymous Referee #1

Submitted on 31 Oct 2014

Suggestions for revision or reasons for rejection (will be published if the paper is accepted for final publication)

Although the authors rewrote more than 50% of the paper, it still seems to be a somewhat sloppy work. It is a pity for an exciting experiment, but the results are too preliminary. The authors merely touch the basic elements of erosion processes (for instance soil physical properties) and do not discuss this in an appropriate way.

Although "this paper has been corrected by experienced article writers", this writers did not a good job. The manuscript still contains a huge number of errors and untidiness. (for example: three different citations: Oostwood and Ergenzinger; Oostwoud and Ergenzinger; Oostwould and Ergenzinger ???).

## Corrected

There are two identical Tables 1 .... The use of English language is still somewhat inappropriate.

## We have submitted to AJE

Therefore all in all the manuscript should be rejected.

# 2 Answer to Anonymous Referee #2

Submitted on 21 Dec 2014

The research described in this article is very interesting because it addresses the question Suggestions for revision or reasons for rejection (will be published if the paper is accepted for final publication)

One of the major issue in all three reviews of the initial manuscript was the weak analysis and discussion of the results with regard to the (not really explicitly) formulated problem. By

including several scientific references in the discussion as well as pointing out some specific processes such as the role of crack closing this aspect has improved. However, the paper still lacks the specific problem that should be resolved and the hypotheses to be checked. This is a continuing weakness of the manuscript, that probably is difficult to rectify in this state of the research. The chapter « data processing » has undergone a substantial improvement. This applies also for the data documentation in the results part. The figures are now much more informative and better explained. Concerning the language the specific terminology is generally correctly used although there are several minor spelling errors.

Here some special issues :

The abstract should be written more explicitely. It should be more than a description of what has been done.

# We tried to slightly restructure the abstract and underlined the implication for erosion and also infiltration

The soil profile must be described properly, at least texture must be given (see review #1)

We have now described the soil profile with soil terminology even if there is no vegetation and soil development, we indicate the fact that no vegetation cover exists for the considered sample.

TLS data aquisition is time consuming : What was the actual time resolution? In table 1 and 2 has to be given the total scan time on the sample box and not only the starting (or ending point). In the present form the column caption « aquisition time » is not correct. When data acquisition is more than 10 minutes, this would include an important unsharpness in the data. In this case it is crucial, to know whether the scan was horizontally (then we have a bias in the data) or vertically (then the long sampling time only adds noise to the data)

We change the table headers and indicate the scan duration in minutes. In addition, the time of acquisition per cm. from left to right is now given.

The paragraph beginning at line 86 cannot be understood. What is « about 1.3 litres fell on the box »? This paragraph contains several spelling errors.

## We modified the paragraph by adding the total amount of rain 5.5 mm.

## Final suggestion:

Under the precondition that the mentioned points are resolved, I can recommend the publication of the manuscript in HESS, mainly because the presented methodology in erosion research is still upcoming and because the transfer of a bulk micro hillslope to the lab is somewhat original.



February 10, 2015

Dear Michel Jaboyedoff,

Thank you for choosing American Journal Experts. This manuscript, "Erosion processes in black marl soils at the millimetre scale: preliminary insights from an analogous model," demonstrates the ever-increasing utility of laser scanning technology in our understanding of natural processes by examining the minute processes occurring on a sloped soil surface. The first editor and I have revised the paper for grammar, phrasing, punctuation, and diction. A number of our changes are addressed directly in the text, and I have outlined several of the major changes below.

The choice of the correct preposition is a difficult aspect of the English language. In your manuscript, prepositions were edited to reflect standard English usage. For example, "changes of the soil surface" was changed to "changes in the soil surface", and "present at the surface of the sample" was edited to "present on the surface of the sample".

In certain cases, sentences were restructured to eliminate weak or imprecise phrasing. For example, the sentence "It is clear from Figures 6 and 7 that the material has expanded..." was changed to "Figures 6 and 7 clearly show that the material has expanded..." Where possible, sentences should be constructed such that they contain nouns or pronouns other than "it" as their subject unless "it" refers to a specific noun.

Formal language is preferred in scientific writing. For this reason, the following edits were made in your manuscript: "about 10 cm" was changed to "approximately 10 cm" and "several artefacts" was changed to "certain artefacts".

Comments were left in several places where further clarification would be helpful or where confirmation of the meaning of the text is necessary. Please review these comments and all our changes carefully to ensure that the final version of the manuscript is fully accurate.

Thank you again for using our editing services; we wish you the best of luck with your submission.

Best Regards,

Matthew D. Senior Editor American Journal Experts

## 1 Erosion processes in black -marl soilss at the millimetreer

2 scale:, the preliminary insights from an analogical

## 3 analogous model

## 4

- 5 J. Bechet<sup>1</sup>, J. Duc<sup>1</sup>, M. Jaboyedoff<sup>1</sup>, A. Loye<sup>1</sup> and N. Mathys<sup>2</sup>
- 6 [1]{University of Lausanne, Risk-group ISTE Institute of Earth Sciences, Lausanne,7 Switzerland}
- 8 [2] { IRSTEA Grenoble, Unité de recherche Erosion Torrentielle, Neige et Avalanches, BP 76,
- 9 38402 Saint Martin d'Hères, France }
- 10 Correspondence to: M. Jaboyedoff (michel.jaboyedoff@unil.ch)

#### 11 Abstract

- 12 In order <u>t</u>To investigate some of the millimetre-millimetre-scale surface processes caused by
- 13 natural rainfall, an undisturbed sample of badlands soil (1 m long, 0.5 m wide and 0.15 m
- 14 thick) has been was carefully extracted in situ. The sample is composed of black\_-marls soil;
- 15 coming-from a badlands area of the Draix Observatory (SE of France). After it thoughtful
- 16 extraction, the undisturbed sample has been was placed with at the same slope angle (45°) as
- 17 its was originally original orientation placed set\_in situ\_and. This portion of soil \_was then
- 18 monitored for several processes by via a terrestrial laser scanner (TLS) with a millimetre
- 19 <u>millimetre-scale accuracy and resolution. This experiment allowed the identificationying</u>
- 20 ofidentified several surface processes interpreted as micro-landslides, swelling of the black\_-
- 21 marls material and lateral expansion elosing that closed desiccation cracks. These micro-
- 22 processes illustrate the complexity of the surface micro-topography changes processes that
- 23 that are controlling erosion and infiltration rates over times by surface micro topography
- 24 <mark>changes</mark>.
- 25

## 26 **1** Introduction

Small black\_-marls watersheds have strong responses to climate forcing (Malet et al., 2007).
This study aims to better understand the erosional behaviour and the micro-morphological

29 evolution on-of these materials caused by micro-scale mass -movements as-in response to 30 precipitation. The goal of this experiment was to extract an undisturbed sample of black marls 31 and expose this material to a natural rainfall event, in order to monitor the sample's surface 32 evolution using a 3D laser scanner. This monitoring has permitted us to study in high detail 33 the micro-topographic surface deformation and erosion processes with high detail which that may also-have also-an impact on infiltration rates during rainfall events (Mitchell and van 34 35 Genuchten, 1993; Römkens and Prasad, 2006). In addition, such-these surface changes may 36 also have an impact onplay a role in the triggering factor of landslides (Galeandro et al., 37 2014). Unlike as first intended Contrary to the original intention, tThe duration of the 38 experiment and the rainfall intensity does did not permit to investigatione of the splash 39 erosion as first intended, which can be important when rainfall is intense (Selby, 1993), 40 because its effect is was below the resolution of our acquisition and because of the rather 41 lowthe intensity of the rainfall, was rather lowwhich can be important when rainfall is intense (Selby, 1993). 42 43 In the past, several studies of artificial -rain simulations have been performed on the site of at 44 the Draix Observatory (ORE Draix), which is an observatory dedicated to the the research of 45 onof mountain hydrology and erosional processes. Previous works-studies focused on the 46 measurement of sediment transport (Oostwoud Wijdenes and Ergenzinger, 1998) and runoff 47 in as a function of precipitations (Mathys et al., 2005). The badlands ground surface ground 48 evolution has already been monitored during rainfall simulations by using a pin-type microrelief\_-meter and photographyies (Torri et al., 1999; Mathys et al., 2005), but these tools were 49 50 not able to reach the same precision as a laser scanner. Recent studies showed have shown the 51 potentiality of for observing and characterizing erosional processes at the level of micro-52 topography using a laser scanner (Schmid et al., 2004; Barneveld et al., 2013). 53 The soil sample was extracted from the Draix experimental site of Draix (ORE Draix, 54 IRSTEA), a badlands area near the city of Digne-les-Bains in the southern of French Alps, -

badlands area near the city of Digne les Bains. The experiment has been was performed in 55 56 Lausanne (Switzerland), where the precipitations can be considered is similar to the Alpine 57 region. For the monitoring of the sample surface, a terrestrial laser scanner (TLS) TLS has 58 beenwas used to (1) identify part some of the millimetre-millimetre-scale processes that are

- 59
- control<del>ling</del> erosion and infiltration by (2) quantifying the swelling of the material in the marks
- 60 during rainfall events and (3) mapping all the possible modifications of the terrain surface.

Formatted: Highlight

Formatted: Highlight

### 61 2 Geological settings

62 The badlands near the village of Draix are composed of weathered black -marls of Middle Jurassic age (Callovo-Oxfordian). Theis black marl formation is more than 2,000 m thick in 63 some certain places (Antoine et al., 1995). The studied study site has no vegetal cover. The 64 and the regolith is usually about approximately 40 cm to 1 m thick (Maquaire et al., 2002; 65 Antoine et al., 1995). About<u>The upper approximately</u> 10 cm of the top of the regolith is 66 67 constituted of by-loose detrital material made composed of local clasts and platelets produced 68 locally (Maquaire et al., 2002)., which This regolith corresponds to a sandy loam when it has been exposed during a time for long enough to be disaggregated by weathering (Antoine et al., 69 70 1995). When this-the upper regolith layer is fresh, it can be considered as-loamy sand or sand. 71 Below the regolith is constituted of a layer of plate-plate-like unstructured rock. Finally, a 72 compacted regolith of 10 to -20 cm thick lies in contact with the bedrock- (Maquaire et al., 73 2002).- The clay size fraction content-of the black -marls located in this part of France has been measured atfound to be 35±5% (Caris et Van Asch, 1991), but the clay mineral content 74 75 is of approximately 10% and is, mainly primarily illite and with traces of smectite and 76 interstratified clay minerals (Antoine et al., 1995). During rainfall events, water infiltrates the 77 ground, and the material ean-swells because of the fine-behaviour of the fine-grained material 78 behaviour and chemical reactions (Antoine et al., 1995). The loose upper-upper-layer detrital 79 material is very sensitive to erosion and is a good candidate for experiments of oninvolving 80 surface processes imaging.

#### 81 3 Methods

#### 82 3.1 Samples

83 The sample of soil used for this experiment was extracted from a marl outcrop with a  $45^{\circ}$ slope, and <u>T</u>the bedding crossing the surface of the outcrop is close to a perpendicular (Fig. 84 ure 1a). Nevertheless, the loose detrital material at the surface does not display any 85 86 identifiable bedding structure. The sample is 1 m long, 0.5 m wide and 0.15 m thick. To 87 extract this sample of soil, a metal case has been was designed to keep undisturbed the soil 88 structure undisturbed (Figure 1). The extractionis was performed by pressing the bottomless 89 box into the ground and inserting the bottom plate via, tapping with a hammer to slide the bottom plate of the box in order to isolate a sample of the bedrocksoil. This sample has 90 91 beenwas stored in the laboratory in dry conditions, which ean bewere similar to natural

Formatted: Highlight

**Comment [SE1]:** Please clarify what the bedding is perpendicular to. Perhaps "normal to the face of the outcrop" would provide more information.

Formatted: Highlight

**Comment [SE2]:** Please ensure that the intended meaning has been maintained in this edit.

<u>conditions</u>, <u>during-for</u> 3 months before the experiment started, <u>which can be similar to natural</u>
 <u>conditions</u>.

94 3.2 Experiment settingssetupdesign

95 On the 31<sup>st</sup> May, 2011, the soil sample was exposed to a natural rainfall from 11h01 to 17h47.

96 During the experiment, the soil sample was located kept in its extractioning metal extraction

97 casing, which and has been was tilted by at 45° in order to get obtain the same inclination as

98 its in situ conditions.

99 The soil sample was scanned every 30 minutes using a ground-based TLS Leica ScanStation

100 II, that which produces point\_-clouds (x, y, and z data) in a three-3-dimensional  $\frac{1}{2}$ 

101 space (Figure 1c). The direction of the laser pulses' line-of-sight and the recorded time-of-

102 flight determine the position of the measured points. The rainfall has been was measured by

103 the weather station of at the University of Lausanne (PluvioMADD2 from MADD

104 Technology), located at 500 m from the experiment placelocation.

105 Since tThe sample was dry at the beginning of the experiment. The total precipitation during

106 the experiment was of 5.5 mm, which corresponds to about approximately 1.3 litres that

107 enteringed the metal case. Most of the rain-fall was absorbed by the sediment, which

108 limitedlimiting the transport of sediment by runoff. As a cConsequencetly, the sample did not

109 reach full saturation. We suspect that a small quantity of evaporation may have occurred

110 during the experiment, but we did not weight the box before and after the experiment, and this

111 minor influence has been neglected in our study.

## 112 **3.3 Data acquisition**

113 The surface evolution has been was monitored for 6 hours and 46 minutes by 12 successive

- laser scan acquisitions. The first scan and the last scan (acquired at <u>11h01</u><u>11:01</u> and 17h47,
- 115 respectively) have had approximatively a point spacing of approximately 0.001 m with a
- 116 duration time of 8 min. All the other scans have had a 0.002 m point spacing for a duration of
- 117 2 min. From 0 to 50 m, 50% of the laser beam is in-at a diameter of 4 mm diameter at full
- 118 width half height (FWHH) (Leica Geosystems AG, 2007). The scans are-were acquired
- 119 vertically from the left to the right of the box, which means that one centimetre in width is
- 120 scanned in less than 2.4 see. for the 2 min. scans and less than 9.6 see. for the 8 min. scans.
- 121 The scan distance was 2 m. The instrument was not moved throughout the experiment,

Formatted: Font: Not Italic

**Comment [QCE3]:** Please include both the name of the supplier/manufacturer and their location (including city, state, and country) for all specialized equipment, software, and reagents.

Formatted: Highlight

Formatted: Highlight
Formatted: Highlight

Formatted: Highlight

- 122 meaning that the position and the orientation of the scans is are identical for all of the
- 123 acquisitions. As a consequence Therefore, no scan alignment was necessary to compare the
- 124 obtained point clouds.

#### 125 3.4 Data processing

The point clouds were "manually" cleaned from of the points that are were not imaging the surface of the sample, i.e., the metal casing sides, the background of the scenery and some several artefacts, including points away from not located on the terrain's surface, such as rain drops. Only the surface of the sample was kept. Every cleaned scan has a very high point density: more than  $400_{2}$ -000 points for the first and last acquisitions and a minimum of  $110_{2}$ -000 points for the rest of theseother scans (Table 1).

132 Each TLS point cloud was firstly rotated by 45° to obtain on average an approximately 133 horizontal point cloud in order to be able to interpolate in 2.5-dimensions. The interpolations 134 were performed using an inverse distance method with a power of 1 (Shepard, 1968) with via 135 Surfer 8.0 software (GoldenSoftware). The point clouds were transformed into a regular 136 squared grid of 1 mm for the first and last scans and 2 mm for the other datasets. This 137 provided high-high-resolution digital elevation models (DEM) of altitude z above the mean 138 horizontal surface. The search radius for DEM generation was defined in-as\_1.5 times higher 139 than the pixel size, i.e., 1.5 mm for scans with a point spacing of 1 mm and 3 mm for the 140 scans with a point spacing of 2 mm (Figure 2). Although different values of DEM cell size 141 were tested, this value was chosen such as the most adequate satisfactory compromise 142 between accuracy and/ resolution. The so-generated DEMs were compared to quantify and 143 map surface changes, i.e., mass movements and, erosion/deposition processes. Each DEM 144 was subtracted from the initial (or reference) DEM (DeRose et al., 1998). The resulting z 145 difference grids have negative value pixels for the "erosion" and positive value pixels for the 146 "deposition". To limit the 'noise' of the measurements, absolute value differences in absolute 147 value inferior less than to 0.0015 m were ignored; this threshold was obtained by a trial and 148 error procedure.

#### 149 4 Results

The precipitations started at<u>lasted from</u> 13h30 <u>until the end of the experiment, with</u> <u>and</u>, then it stopped for<u>a</u> 30-minute hiatus starting <u>-30 minutes around at approximately</u>  $16h00_{-}$  and then continued up to the end of the experiment. The maximum rainfall intensity (5.2 mm h<sup>-1</sup>) was reachedoccurred at 16h30. A-<u>The cumulative</u> precipitation <u>amountquantity was of 5.5 mm</u>
 was recorded at the end of the experiment. Three different surface processes of topographic
 changes were identified <u>and discriminated</u>. <u>The fF</u>igure 3 <u>displays shows</u> the difference

between the initial and final scans (11h01 and 17h47, respectively). The rising of <u>Changes in</u>

157 the surface <u>elevation</u> (z) appears in blue <u>colours(increased elevation)</u>, and the decrease in z

158 appears in red colours (decreased elevation; (Figure 3). We considered only the changes that

159 are fully visible in the oblique view of the experiment (<u>i.e., the</u> box with a slope of  $45^{\circ}$ ).

The micro-topographicy changes <u>that</u> occurred <u>at-along location-transect 1-2</u> <u>is-are</u> shown in Figures 3 and 4. As <u>can be observed in Ffigure 4c shows</u>, a depression <u>(in red)</u> is-formed at the top of a small ridge <u>(in red)</u>, and <u>a the surface elevation of a</u> depression below this ridge is filled inincreased (in blue). This <u>change</u> can be interpreted as a downward mass\_-movement at the millimetre scale. The small particles moved down, <u>fillingward and filled</u> a desiccation crack. This phenomenon occurred <u>right\_immediately</u> after the highest intensity rain event, between 17h00 and 17h30 (Figure 5).

When the rain intensity reached 1 mm h<sup>-1</sup>, the entire surface of the soil started to rise up all 167 over the surface (it means perpendicularly to the surface of the  $45^{\circ}$  slope at  $45^{\circ}$ ), appearing 168 169 which appears as a pale blue layer on in Figures 3 and is illustrated in Efigure 6. This process 170 affecting the entire surface has been measured from resulted in an overall rise of 1.5 to 3 mm 171 for-over the course of the entire experiment, i.e., for-a cumulativeed 5.5 mm of rain (Figure 172 5). It starts The surface elevation increased slowly after the first rain. At 16h30, the surface 173 subsided, and this the process was momentarily slowed down. After that Subsequently, the 174 processes re-accelerated following the new-peak in rain intensity (Figure 5). Note that no 175 significant rise of in the topographic surface has occurred at the top of the sample, which was 176 protected from the rain (Figure 3).

Another observed process is linked to changes <u>of in</u> the soil surface by lateral expansion
(Figures 6 and 7). This process <u>has been observed occurred</u> continuously through the closing

179 of the desiccation  $\operatorname{cracks}_{a,\overline{s}}$  which were present at <u>on</u> the surface of the sample.

The <u>last final observed</u> process was the stripping of soil particles by the kinetic energy of the
 raindrops. Although these changes were observed by the authors during the experiment<sub>a</sub>.
 Unfortunately its their magnitude was <u>unfortunately</u> too small to be significantly monitored

183 by the <u>utilised TLS used in this study</u>.

**Comment [SE4]:** Please note that the highest intensity rainfall was previously said to have occurred at 16h30. Please check whether one of these statements is a typing error.

### 184 **5** Discussion and conclusions

200

185 The geometric-topographic changes observed at locationalong transect 1-2 (Figures 3 and 4) occurred almost instantaneously, which excludes the rain splashing process.; they These 186 changes were are interpreted as a micro-landslide. Looking at Based on the duration of the 187 188 scans and the size of the transported mass<del>-moving</del>, we <del>can</del>-assume that it is <del>quite</del>-highly 189 unlikely that a scan had just crossed the region of interest when the movements occurred 190 within it (0.05 m scanned in around-approximately 50 see.). In addition, the two profiles 191 before and after the occurrence of the changes mimic the profile changes usually observed in 192 real landslides. Such a The observed process can be is likely related to the initiation of 193 miniature debris flows (MDFs) observed by Oostwoud Wijdenes and Ergenzinger (1998), but 194 the latter needed heavy rainfall to be transformed in MDFs. 195 The observed rise of in the surface, which reacts fullyoccurreds with rainfall with atfollowing 196 a 30--minutes delay\_after the initiation of rainfall-with rainfall, is certainly linked to the 197 swelling of the material. The soil sample experienced swelling after the first rainfall intensity 198 peak, contraction 30 minutes after the stop cessation of rainfall and new renewed rising of the 199 topographic surface once the rain started again. This measured cyclic behaviour demonstrates

that <u>such athis</u> process is linked to <u>the</u>-moderate rainfall intensit<u>iesy</u>, which allows <u>the</u>-water to infiltrate <u>the fine-grained material, causing and swell part of theswelling fine\_grained</u> material. Furthermore, t<u>T</u>his process is not necessarily caused by clay minerals, <u>since because</u> they are present <u>only</u> in small quantitiesy (<u>mostly-primarily</u> illite) in the study area (Antoine et al., 1995). The swelling <u>must</u>-dissipates rapidly for moderate rainfall events because of the diffusion of water when the rainfall stops, <u>which leadsing</u> to a decrease <u>of in</u> the <u>effect of the</u> water-<u>effect</u>.

that the swelling and contraction of the soil surface is a reversible process. We can assume

208 The crack-closing lateral expansion of the surface within cracks-is also certainly linked to 209 swelling, but does not retreat reverse when the rain stops. It is clear from the Ffigures 6 and 7 210 clearly show that the material has expanded, bBut and that it is not affected by the material as 211 not transported, since because no deposition was observed at the bottom. We do not findhave 212 not found any definitive explanation for this that accounts for the difference with between the 213 rising up-process and the lateral expansion process., but-However, it-the difference must be related to gravity, which increases the effect of swelling downward. Then Wwhen both the 214 215 two\_sides of the crack touch each othermake contact, the moistened zone doubles its Formatted: Highlight

**Comment [SE5]:** Please clarify. The word "latter" is generally only used after 2 or more objects/options have been listed. It is unclear what 'latter' refers to here. Perhaps you mean, "but MDFs require heavy rainfall."

Formatted: Highlight

- thickness, decreasing the water diffusion in the material. In addition, such these processes
- 217 must be part-components of the creeping (Selby, 1993), probably likely because probably the
- 218 retreat by drying will beis less effective for in the downward-downslope partportion, leading
- 219 to slow progressive downward movements-downward.

220 The above interpretations are also important for the understanding of the infiltration process.

221 It is clear that cCracks clearly play an important role in the infiltration rates (Mitchell and van

222 Genuchten, 1993; Römkens and Prasad, 2006) and subsequently consequently for in the

destabilisation of slopes (Stumpf et al., 2013; Galeandro et al., 2014). The above analysed

224 processes <u>analysed here are playingplay</u> a role for in the closure of cracks, as was showned in

- 225 <u>Ffigure</u>-7. In the present case study, we demonstrated how micro-scale infiltration can
- influence the degradation of <u>soil</u> surface of <u>soil</u> by inducing downward mass movements that are not reversible.
- 228 We have also shown here the great potential of high-high-resolution three-dimensional TLS or 229 photogrammetry point clouds of TLS or photogrammetry to analyse for the analysis of the processes that lead to erosion throughout surface mass movements at the millimetre scale. 230 231 Investigations about of erosional processes using point clouds are increasing in number. They 232 These studies use either <u>IL</u>aser scanners for either for micro-scale surface imaging 233 (Schmid et al., 2004; Barneveld et al., 2013) or measuring for cracks apertures (Sanchez et al., 234 2013). In addition, photogrammetry and , especially, sStructure from mMotion (SfM) 235 methods are now being developed to analyse soil surfaces (Snapir et al., 2014).
- 236 This paper shows that monitoring the changes at the millimetre scale to illustrate examine soil 237 surface changes and erosion is now possible. This development will help to designaid in 238 designing future experiments to analyse <u>certain single</u>-processes, such as swelling, crack 239 closure, micro-landslides, and initiation of MDFs. With heavier rainfall, those sediments will 240 start to be mobilised on for and transported across longer distances, enabling the study of like 241 MDFs and the formation of rills. It-Thise study also shows demonstrates also that material and 242 rain intensity must be sufficient-suitable to permit an-the efficient detection of rain splash processes and associated erosion; specifically, a rainfall intensity of more greater than 20 mm 243 244  $h^{-1}$  is necessary (Mathys et al., 2005).
- 245

246 Acknowledgements

Formatted: Highlight

Formatted: Highlight

- 247 <u>Thanks-The authors would like to thank to the Observatoire de Recherche en Environnement</u>
- 248 (ORE) of Draix, for letting us to-sample soil into a protected area. Thanks are also due to We
- 249 would also like to thank the UNIBAT service of the University of Lausanne (UNIL); for
- 250 providing meteorological data from their weather station. We also thank greatly appreciate the
- 251 comments and suggestions of also-our colleagues Dr. M.-H. Derron, Benjamin Rudaz and
- 252 Antonio Abellán-for their comments and suggestions. We thank American Journal Experts for
- 253 the improvements of the English language of this paper.
- 254

**Comment [SE6]:** If each of these three colleagues has a doctorate, please use the abbreviation "Drs."

#### 255 References

- Antoine, P., Giraud, A., Meunier, M. and Van Asch, T.: Geological and geotechnical
  properties of the 'terres noires' in southeastern France: Weathering, erosion, solid
  transport and instability, Engineering Geology, 40, 223-234, 1995.
- Barneveld, R.-J., Seeger, M. and Maalen-Johansen, I.: Assessment of terrestrial laser scanning
  technology for obtaining high-resolution DEMs of soils. Earth Surf. Process.
  Landforms, 38, 1096-9837, 2013.
- Caris, J.P.T. and Van Asch, T.: Geophysical, geotechnical and hydrological investigations of
   a small landslide in the French Alps, Engineering Geology, 31, 249-276, 1991.
- DeRose, R. C., Gomez B., Marden, M. and Trustrum, N. A.: Gully erosion in Mangatu Forest,
   New Zealand, estimated from digital elevation Models, Earth Surface Processes and
   Landforms, 23, 1045-1053, 1998.
- 267 Galeandro A., Doglioni A., Simeone V. and Šimůnek J. S.: Analysis of infiltration processes
  268 into fractured and swelling soils as triggering factors of landslides. Environ. Earth Sci.
  269 71, 2911–2923, 2014.
- 270 Leica Geosystem AG: Leica ScanStation 2 technical note, Heerbrugg, Switzerland, Vl.07,
  271 2007.
- Malet, J.-P., Durand, Y., Remaître, A., Maquaire, O., Etchevers, P., Guyomarc'h, G., Déqué,
  M. and van Beek, L.P.H. Assessing the influence of climate change on the activity of
  landslides in the Ubaye Valley. In: McInnes, R., Jakeways, J., Fairbank, H., Mathie, E.
  (Eds): Proceedings of the International Conference on Landslides and Climate Change Challenges and Solutions, Taylor & Francis, London, pp. 195-205, 2007.
- Maquaire O., Ritzenthaler A., Fabre D., Ambroise B., Thiery Y., Truchet E., Truchet E.,
  Malet J.-P. Caractérisation des profils de formations superficielles par pénétrometrie
  dynamique à énergie variable: application aux marnes noires de Draix (Alpes-de-HauteProvence, France). Comptes Rendus Géosciences, 334, 835–841, 2002.
- Mathys N., Klotz S., Esteves M., Descroix L. and Lapetite J.-M.: Runoff and erosion in the
  Black Marls of the French Alps: observations and measurements at the plot scale.
  Catena, 63, 261-281, 2005.

- 284 Mitchell AR AND van Genuchten MTh. Flood irrigation of a cracked soil. Soil Science
   285 Society of America Journal. 57, 490–497, 1993.
- 286 Oostwoud Wijdenes, D.J. and Ergenzinger, P.: Erosion and sediment transport on steep marly
  287 hillslopes, Draix, Haute-Provence, France: an early experimental field study. Catena,
  288 33, 179-200, 1998.
- 289 Römkens M.J. M. and Prasad, Rain Infiltration into swelling/shrinking/cracking soils.
   290 Agricultual Water Management 86, 96 205, 2006.
- Sanchez M., Atique A., Kim S., Romero E. and Zielinski M. Exploring desiccation cracks in
   soils using a 2D profile laser device. Acta Geotechnica, 8, 583–596, 2013.
- Schmid, T., Schack-Kirchner, H. and Hildebrand, E.: A Case Study of Terrestrial LaserScanning in Erosion Research: Calculation of Roughness Indices and Volume Balance
  at a Logged Forest Site. Eds: M. Thies, B. Koch, H. Spiecker and H. Weinacker: LaserScanners for Forest and Landscape Assessment. ISPRS Archives V. XXXVI-8/W2 WG VIII/2, 114-118, 2004.
- Selby, M.J., Hillslope materials and processes: Oxford, UK, Oxford University Press, 252–
  258, 1993.
- Shepard, D., A two dimensional interpolation function for irregularly-spaced data, New York,
   USA. Proceeding ACM '68 Proceedings of the 1968 23<sup>rd</sup> ACM national conference,
   517-524 1968.
- Snapir B., Hobbs S. and Waine T.W. Roughness measurements over an agricultural soil
  surface with Structure from Motion. ISPRS J. of Photogrammetry and Remote Sensing,
  96, 210–223, 2014.
- Stumpf, A., Malet, J.-P., Kerle, N., Niethammer, U., & Rothmund, S. Image-based mapping
  of surface fissures for the investigation of landslide dynamics. Geomorphology, 186,
  12-27, 2013.
- Torri, D., Regüés, D., Pellegrini, S. and Bazzoffi P.: Within-storm surface dynamics and
  erosive effects of rainstorms. Catena, 38, 131-150, 1999.
- 311

- 312 Table 1: Summary of the TLS scans table of fromin the TLS campaign experiment for the
- 313 TLS campaign of the<u>on</u> 31st of June, of 2010. This table compiles the information concerning
- the scan time, the number of points before and after cleaning, the mean spacing between the
- 315 points and the density of points.
- 316

Start of acquisition	Number of points	Mean spacing	Distance from TLS [m]	Number of points	Density	
.11h01 (8)	601349	1	2	496391	84.84	
12h00 (2)	149951	2	2	123233	21.06	
14h00 (2)	149951	2	2	123477	21.10	
14h30 (2)	149951	2	2	123269	21.07	
15h00 (2)	149951	2	2	123496	21.11	
15h30 (2)	149951	2	2	122857	21.00	
16h00 (2)	149951	2	2	120162	20.54	
16h30 (2)	149951	2	2	118301	20.22	/
17h00 (2)	149951	2	2	117290	20.05	/
17h30 (2)	149951	2	2	116322	19.88	/
17h32 (8)	601349	1	2	467484	79.90	/
17h47 (8)	601349	1	2	469906	80.31	
						/

Formatted: Highlight
Formatted: Superscript, Highlight
Formatted: Highlight
Formatted: Highlight
Formatted: Highlight
Formatted: Highlight
Formatted: Highlight
Formatted: Highlight
Formatted: Highlight
Formatted: Highlight
Formatted: Highlight
Formatted: Highlight
Formatted: Highlight
Formatted: Highlight
Formatted: Highlight



Figure 1: a. <u>V</u>view of the sample collection process, ing and the metal case (1 m x 0.5 m x 0.20 m), and the approximate orientation of the bedding. showing tThe bottom plate is partially entered inserted and the approximate orientation of the bedding. b: View of the soil sample of soil inclined atwith an inclination of 45° during the rainfall event. c. Position of the TLS, which is protected from rainfall by a tent, relatively to the metal case-which is protected from rainfall by a tent. 







335 Figure 3: <u>A comparison of TLS scans-comparison. The</u>; <u>colours represent topographic</u>

- changes from between 11h01 to and 17h47. and The figure also includes a surface detailed
- 337 <u>view of the surfaceview</u>. The locations of **F**figures 4 to 6 are indicated.



339 Figure 4: Profile changes showing a-mass movement occurring between 17h00 and 17h30.-...



341  $\underline{\mathbf{T}}$  the black box<u>es</u> indicates the position of the profiles in c.



**Comment [SE7]:** Please change "Rise up" to "Rise".



338

343 Figure 5: Observed phenomenon according to the in relation relative to the precipitation data.

344 The total rain amount in mm, the cumulative precipitation and the rain intensity are presented.

345 The duration of the rise is shown in mming up is in mm, and and the lateral expansion with its

346 intensity and mass movement processes are presented in relative scales.





350 Figure 6: a. Profile changes showing <u>the</u>evolution of lateral expansion and <u>surface</u>rise <del>up</del>

frombetween 16h00 and 17h30. b. A 3D view of the zone at 11h01 and b. at 17h47. T the

352 black box<u>es</u> indicate<u>s</u> the position of the profiles in a.

353



354

355 Figure 7: 3D views changes showing of the evolution of lateral expansion, which showing

356 <u>closes a crack closure between from 11h01 (a) to and 17h47 (b).</u>