

# 1 **Supplementary information “Water balance and its** 2 **intra-annual variability in a permafrost catchment:** 3 **hydrological interactions between catchment, lake** 4 **and talik”**

## 7 **1.1 Meteorological data**

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9 Results from the local Automatic Weather Station (AWS) at TBL show that almost 60% of the 2012  
10 annual P was recorded during the active period (May to September) with maximum monthly P in  
11 September, 84 mm. This is more than three times the recorded P for the active period in 2011. The  
12 corrected annual P 2012 in Kangerlussuaq was 346 mm, which is significantly higher than the long  
13 term mean of 173 mm. On annual basis, the P is slightly higher at TBL compared to Kangerlussuaq  
14 and the short time series from the TBL catchment is considered to be comparable with the long term  
15 mean values from Kangerlussuaq.

16 Based on the 36 y time series from the DMI station in Kangerlussuaq a rain of 24 h  
17 duration with a return period of one year was defined to be approximately 13 mm/day. The largest 24-  
18 hour rain for the 36-year period was 30 mm /day. In the TBL catchment, two rain events with a rain  
19 intensity higher than 13 mm/day were recorded during the studied hydrological year, indicating that  
20 rain events with high intensity at TBL are in the same order of magnitude as high intensity rains in  
21 Kangerlussuaq. The rain events at the TBL catchment are characterized by relative long duration and  
22 low intensity.

23 Data from the AWS on shortwave in- and out- going radiation, long wave in- and  
24 outgoing radiation, albedo, air temperature, relative humidity and wind speed was used to calculate the  
25 potential evapotranspiration (PET) of the site, using the Penman-Montieth equation (Penman 1947).  
26 The total calculated PET for the hydrological year is 413 mm and the mean daily PET during the  
27 active period is 2.5 mm/day in 2011 and 2.4 mm/day in 2012.

28                   The mean hourly wind speed was 2.1 m/s with a maximum mean hourly wind speed of  
29 18.6 m/s. During winter, the threshold wind speed for snow drift is exceeded 106 times, indicating  
30 snow drift might be an important process and component of the water balance of the catchment area.  
31 Snow accumulation and snow melt was calculated with a degree day method using air temperature and  
32 precipitation data from the AWS. The snow model was verified with photos from the time lapse  
33 camera. The threshold temperature for snow melt/accumulation was set to 0°C and the degree day  
34 coefficient was set to 2 mm/day/°C. Snow depth measurements in three transects, one in a valley along  
35 the lake, one on the lake ice and one on a hillside, were performed in April 2011. The mean snow  
36 depth in the three transects ranged from 11 cm on the hillside to 24 cm in the valley. The mean snow  
37 depth on the lake was 17 cm.

38                   In April 2013, sublimation was measured at three sites in the TBL catchment  
39 during a period with temperatures between -14 to -2 °C, moderate winds in the range from 2-5  
40 m/s and clear sky. At each site, 5 boxes were filled with snow, weighted and left out for three  
41 consecutive days. The weight loss in each box was observed every 24 h. The measurements stopped  
42 after three days due to snow fall. The measured sublimation averaged to 2.75 mm/day ranging  
43 from 1.8 to 4.2 mm/day depending on the location of the measurements.

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## 46 **1.2 Active layer and lake ice**

47 The thickness of the active layer was measured with a soil probe at a number of locations in 2008,  
48 2010 and 2011 (Clarhäll 2011). The survey resulted in a mean active layer thickness of 0.69 m. The  
49 maximum and minimum depths were found to be 0.94 m and 0.55 m respectively. In the stream  
50 depression at the lake outlet the active layer was measured to 0.90 m. Soil temperature is monitored  
51 every 0.25 m down to 2 m depth at one location in the catchment. During both monitored seasons,  
52 2010 and 2011, the uppermost sensor at 0.25 m depth has been freezing in October and no positive  
53 temperatures have been recorded between November first and mid-May, Figure 3 (lower left). The

54 uppermost TDR sensors, placed at 5 cm and 10 cm depth, indicate that the uppermost 5-10 cm of soil  
55 freezes in early October and thaws in the first half of May.

### 56 **1.3 Lake level measurements**

57 Data from the pressure transducers (Level TROLL 700) was corrected for barometric effects using air  
58 pressure data from the AWS. The total amplitude of lake pressure for the observed period is  
59 approximately 440 mm with lowest values during the summer of 2011. In 2010 the lake level was  
60 measured to be 180 mm below the lake threshold for surface water runoff. Measurements were  
61 performed with a leveling instrument. The threshold for groundwater outflow in the active layer, i.e.  
62 the depth of the active layer in the lake outlet, was in the summer of 2010 situated 90 cm below  
63 ground surface. During the studied hydrological year the net change in lake water pressure was an  
64 increase with 73 mm. In general, the amplitude of pressure changes are 30 mm during the frozen  
65 period when ice and snow covers the lake and approximately one order of magnitudes higher in the  
66 active period. For both seasons 2011 and 2012 a pressure increase was observed as response to snow  
67 melt with larger increase in 2012.

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### 69 **1.4 Soil moisture data and porosity**

70 In total 44 TDR sensors (Campbell CS616) measuring soil moisture content were installed in three  
71 clusters. Two clusters containing 4\*4 vertical profiles with TDR sensors equally distributed from the  
72 ground surface down to the bottom of the active layer. Additionally TDR sensors were placed in a  
73 transect at three sites, each site containing a vertical profile of 4 sensors evenly distributed from  
74 ground surface to the bottom of the active layer. TDR data were evaluated in order to determine the  
75 change in total water content in the active layer over the hydrological year, Table S1. For each cluster  
76 of sensors 1-3 (Figure 1) the mean value of the soil moisture content at each individual depth was  
77 calculated for the start and end of the hydrological year and the annual difference in water content for  
78 each depth was determined and expressed in mm of water, Table S1. All individual values were  
79 summed in order to determine the total change in water storage for each profile. TDR cluster number 2

80 and 3 are installed in the regolith type “peaty silt” and cluster number 3 is placed in the regolith type  
81 “Eolian silt-fine sand” (Figure 1). The TDR sensors are installed in relatively wet areas within the  
82 catchment. The regolith type “peaty silt”, covering 2% of the total catchment area, is only present in  
83 small areas close to the lake, i.e. in topographically low areas within the catchment where water is  
84 accumulating and the soil moisture is assumed as higher than the mean soil moisture content in the  
85 catchment area. “Eolian silt-fine sand” is the dominating type of regolith in the area covering 51% of  
86 the catchment land area and is present in areas ranging from the catchment boundary to the lake, but  
87 still in relatively low lying areas. It is assumed to represent the mean soil moisture of the catchment  
88 area. The high altitude areas of the catchment are covered by till or bedrock outcrops (Figure 1). The  
89 water content in these areas is very low and an eventual difference of soil water content over the year  
90 can be neglected.

91           A mean value of the change in soil water content was calculated for TDR 2 and 3, both  
92 representing the QD-class “Peaty silt” and the total storage change over the hydrological year in this  
93 QD-class was calculated by multiplying the storage change with the area of the QD-class resulting in a  
94 total storage change over the hydrological year of 2196 m<sup>3</sup>. The change in water content determined  
95 from TDR number 1 representing QD-class “Eolian silt-fine sand”, 15358 m<sup>3</sup>, was calculated in the  
96 same way. The total storage change over the hydrological year is 17555 m<sup>3</sup>.

## 97 **1.5 Borehole data**

98 The DH-GAP01 borehole is packed off and instrumented with a U-tube multi-sensor system (Freifeld  
99 2009) located at a vertical depth interval corresponding to 130-140 m. This system allows for water  
100 sampling and monitoring of in situ pressure (M), temperature (T) and electrical conductivity (EC). The  
101 M/T/EC sensors (AquaTROLL 200) are located below the inflatable packer at vertical depth of 138.57  
102 m (relative to the top of casing), and are connected to the surface via a non-vented cable. The  
103 amplitude of the measured head during the observed period is 1.4 m. A detailed description of the  
104 drilling and installation of the U-tube multi-sensor system is described in SKB (2010).

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106 The hydraulic gradient between the lake and the bedrock was calculated using pressure data from the  
107 DH-GAP01 borehole and the pressure transducer on the lake bottom. The height difference, dz,  
108 between the pressure transducer in the lake and the sensors in the bore hole is 128.1 m. The mean  
109 difference in hydraulic head over the year is 1.85 m, ranging from 1.21 m to 2.47 m. The mean  
110 hydraulic gradient was calculated to 0.014 ranging from 0.019 to 0.009. However, the gradient is  
111 continuously directed downwards. The hydraulic gradient between the lake and the borehole is  
112 increasing over the studied hydrological year, i.e. the net increase in lake pressure is reflected as an  
113 increased hydraulic gradient between the two domains.

## 114 **2 References supplementary material**

115 **Clarhäll A (editor). 2011.** SKB studies of the periglacial environment – report from field  
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117 AB.

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119 **Penman, H.L., 1948.** Natural evaporation from open water, bare soil and grass. Proceedings  
120 of the Royal Society of London 193, 120–145.

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## 125 **3 Tables Supplementary Material**

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127 **Table S1. Soil moisture data at start and end of the hydrological year, table 4A. Total**  
128 **change in water content in the catchment is presented in table 4B.**

**A**

<b>TDR1</b> <b>(Postglacial silt- fine sand)</b>	Water content, % 2011	Water content, % 2012	Difference, %	Difference, mm water*
0-10 cm	0.42	0.42	0.00	0.00
10-20 cm	0.37	0.38	0.01	0.01
20-30 cm	0.38	0.41	0.04	0.04
30-70 cm	0.36	0.41	0.05	0.14
SUM:				0.19
<b>TDR2 (Peaty silt)</b>				
0-5 cm	0.33	0.81	0.48	0.24
5-15 cm	0.45	0.68	0.23	0.23
15-20cm	0.52	0.66	0.14	0.07
20-25 cm	0.47	0.53	0.06	0.03
25-30cm	0.48	0.56	0.08	0.04
30-70 cm	0.50	0.57	0.06	0.22
SUM:				0.82
<b>TDR3 (Peaty silt)</b>				
0-10 cm	0.47	0.61	0.14	0.14
10-20cm	0.46	0.53	0.07	0.07
20-30cm	0.47	0.53	0.05	0.05
30-35cm	0.50	0.54	0.04	0.02
35-45cm	0.47	0.53	0.07	0.07
45-55cm	0.34	0.45	0.12	0.17
SUM:				0.52

\*mm water calculated as water content in %= mm of water in 10 cm of soil

**B**

	Mean difference of water, mm	Area, m <sup>2</sup>	Mean difference of water, m <sup>3</sup>
<b>Postglacial silt-fine sand</b>	19.00	808328	15358
<b>Peaty silt</b>	67.00	32781	2196
	SUM:		17555

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