1	The Role of Retrospective Weather Forecasts in Developing Daily Forecasts of
2	Nutrient Loadings over the Southeast US
3	
4	
5	Jeseung Oh <sup>1</sup> , Tushar Sinha <sup>2,*</sup> and A. Sankarasubramanian <sup>2</sup>
6 7 8 9 10	<ul> <li><sup>1-</sup>Department of Geological Sciences University of North Carolina, Chapel Hill, NC 27599-3315</li> <li><sup>2-</sup> Department of Civil, Construction and Environmental Engineering North Carolina State University, Raleigh, NC 27695-5908</li> </ul>
11	
12	(Published under Hess- Discussion and currently under review for full article)
13 14	
15	* Corresponding author: (email: tsinha@ncsu.edu, phone: +1-919-513-1804)
16 17	
18	Keywords: Daily streamflow, Nonparametric forecast, Weather forecast, Water quality
19	
20	
21	

### 22 Abstract

It is well-known in the hydrometeorology literature that developing real-time daily 23 streamflow forecasts in a given season significantly depend on the skill of daily precipitation 24 forecasts over the watershed. Similarly, it is widely known that streamflow is the most important 25 26 predictor in estimating nutrient loadings and the associated concentration. The intent of this study 27 is to bridge these two findings so that daily nutrient loadings and the associated concentration could be predicted using daily precipitation forecasts and previously observed streamflow as 28 surrogates of antecedent land surface conditions. By selecting 18 relatively undeveloped basins in 29 30 the Southeast US (SEUS), we evaluate the skill in predicting observed total nitrogen (TN) loadings in the Water Quality Network (WQN) by first developing the daily streamflow forecasts using the 31 retrospective weather forecasts based on K-nearest neighbor (K-NN) resampling approach and 32 then forcing the forecasted streamflow with a nutrient load estimation (LOADEST) model to 33 obtain daily TN forecasts. Skill in developing forecasts of streamflow, TN loadings and the 34 35 associated concentration were computed using rank correlation and RMSE, by comparing the respective forecast values with the WQN observations for the selected 18 Hydro-Climatic Data 36 Network (HCDN) stations. The forecasted daily streamflow and TN loadings and their 37 38 concentration have statistically significant skill in predicting the respective daily observations in the WQN database at all the 18 stations over the SEUS. Only two stations showed statistically 39 insignificant relationship in predicting the observed nitrogen concentration. We also found that the 40 41 skill in predicting the observed TN loadings increase with increase in drainage area which indicates that the large-scale precipitation reforecasts correlate better with precipitation and streamflow over 42 large watersheds. To overcome the limited samplings of TN in the WQN data, we extended the 43 44 analyses by developing retrospective daily streamflow forecasts over the period 1979-2012 using

reforecasts based on the K-NN resampling approach. Based on the coefficient of determination 45  $(R_{Q-daily}^2)$  of the daily streamflow forecasts, we computed the potential skill  $(R_{TN-daily}^2)$  in developing 46 daily nutrient forecasts based on the R<sup>2</sup> of the LOADEST model for each station. The analyses 47 showed that the forecasting skills of TN loadings are relatively better in winter and spring months 48 49 while skills are inferior during summer months. Despite these limitations, there is potential in utilizing the daily streamflow forecasts derived from real-time weather forecasts for developing 50 daily nutrient forecasts, which could be employed for various adaptive nutrient management 51 52 strategies for ensuring better water quality.

## 53 1. Introduction

Anthropogenic interventions of biogeochemical cycles have resulted in increased nutrient 54 loadings in streams over the past few decades (Galloway et al., 1995; Caraco and Cole, 1999). 55 Continuous concerns about water quality degradation have resulted in the development of active 56 57 water quality management programs such as total maximum daily load allocation (TMDL) as well 58 as establishment of policy instruments related to water quality trading. Of particularly interest is the total nitrogen (TN) loadings whose contribution from the land surface to the North Atlantic 59 Ocean, has increased from 5 to 20 folds in comparison to the pre-industrial/natural level (Howarth 60 61 et al., 1996). Nitrate levels have tripled in major rivers over the northeastern US since 1900's while nitrate concentration doubled in the Mississippi River basin since 1965 (Turner and Rabalais, 1991; 62 Howarth et al., 1996; Vitousek et al., 1997; Goolsby and Battaglin, 2001). 63

Excess nitrogen results in overproduction of phytoplanktons, which in turn causes anoxic 64 conditions and eutrophication in lakes and coastal regions (Vitousek et al., 1997; Pinckney et al., 65 66 1999). Such eutrophication, due to natural and anthropogenic nitrogen sources, is an important water quality degradation issue, which ranges from small streams (Duff et al., 2008) to large water 67 bodies such as Gulf of Mexico (e.g. Bricker et al., 1999; Alexander et al., 2000; Rabalais et al., 68 69 2002; Alexander and Smith, 2006). There have been several efforts to reduce nitrogen loadings to streams but such programs are often too costly. For example, North Carolina Department of 70 Energy and Natural Resources (DENR) have spent several billion dollars in nutrient management 71 72 of Falls Lake in the Neuse River to control total nitrogen loadings under permissible range (North Carolina DENR). But, availability of data on total nitrogen is limited with concentration being 73 74 measured on a non-continuous basis. Studies have tried to overcome these limitations by using the 75 long available records of streamflow, since both instream nutrient concentration and loadings

76 primarily depend on streamflow variability (Borsuk et al., 2004, Paerl et al., 2006, Lin et al., 2007) and antecedent flow conditions (Vecchia, 2003, Alexander and Smith, 2006). Various nutrient 77 simulation models have been developed to estimate loadings using semi-distributed hydrologic 78 79 models (e.g., WASP, HSPF, SWAT, GWLF) or statistical models (e.g., LOADEST). Both these types of models are typically implemented under a simulation mode by using observed 80 81 meteorological forcings to estimate nitrogen loadings. Similarly, considerable progress has been made in developing daily streamflow forecasts using statistical models, e.g. parametric models 82 (Rajagopalan and Lall 1999; Anderson et al., 2002; Salas and Lee, 2010), and semi-distributed 83 84 watershed models (e.g., Clark and Hay, 2004; Mcenery et al., 2005, Georgakakos et al., 2010). Developing daily streamflow forecasts over a large region using semi-distributed models require 85 intensive spatial data (e.g. topography, land cover, soils) and computational resources, hence, we 86 employed a semi-parametric approach in this study. In particular, we employed the K Nearest 87 Neighbor (K-NN) semi-parametric scheme to develop daily streamflow forecasts contingent on 88 updated climate forecasts since it can capture nonlinear relationships that are typically observed 89 in daily streamflow data (Salas and Lee, 2010). The K-NN scheme has been widely used in 90 hydrologic studies (Rajagopalan and Lall 1999; Prairie et al. 2006; Sharif and Burn 2006). 91 92 Although daily streamflow forecasts could be developed with reasonable skill, there is a gap in linking those forecasts to develop daily nutrient loading forecasts. Furthermore, several studies 93 have utilized antecedent streamflows as surrogates of initial catchment conditions (e.g. Chiew and 94 95 McMohan, 2002; Piechota et al., 2001; Wang et al., 2009) including 3-day average streamflow conditions (Majumdar and Kumar, 1990; Srinivas and Srinivasan, 2000; Cigizoglu, 2003). In our 96 analysis, the maximum auto-correlation between observed streamflow and the previous-day 97 streamflow occurred with a lag of 3 days for the selected sites (figure not shown). Given that 98

99 skillful forecasts of daily nutrient loadings could be utilized in improving in-stream water quality, 100 we intend to investigate the potential in forecasting daily nutrient loadings conditional on daily 101 precipitation forecasts and previously observed streamflow as surrogates of antecedent moisture 102 conditions for 18 watersheds that are minimally affected by anthropogenic interventions over the 103 Southeast US (SEUS).

The manuscript is organized as follows: Section 2 details the data sources for daily streamflow, observed daily total nitrogen samplings and retrospective daily precipitation forecasts that were utilized in the study. Following that, we describe the methodology behind the development of daily streamflow and nutrient loadings forecasts. Section 4 provides the results on the skill in predicting the observed nutrient loadings over the selected 18 watersheds. Finally, in Section 5, we summarize the salient findings and conclusions arising from the study.

110

### 111 **2. Data Description**

112 This section outlines the streamflow, Water Quality Network (WQN), retrospective 113 weather forecasts associated with the development of total nitrogen forecasts over the SEUS.

## 114 2.1 HCDN Streamflow Database

Given the intent of the study is to associate daily nutrient loadings with daily precipitation forecasts, we focus our analysis on 18 undeveloped basins over the SEUS from the Hydro-Climatic Data Network (HCDN) database (Slack et al., 1993). Figure 1 shows the location of 18 HCDN stations and Table 1 provides the list of the 18 stations considered in this study along with their drainage areas. Daily streamflow records in the HCDN basins is purported to be relatively free of anthropogenic influences such as upstream storage and groundwater pumping and the accuracy ratings of these records are at least 'good' according to United States Geological Survey (USGS) standards. Since the streamflow data (Q) in the HCDN database is available only up to 1988, we
extended records up to 2009 based on the USGS historical daily streamflow database.

## 124 **2.2 Weather Forecasts Database**

We employed retrospective weather forecasts from the National Oceanic and Atmospheric 125 Administration (NOAA) to forecast daily streamflow at multiple sites in the SEUS (Hamill et al, 126 2004; Hamill et al., 2006). NOAA's Earth System Research Laboratory/Physical Science Division 127 (ESRL/PSD) reforecast project provides daily precipitation forecasts from the Global Forecast 128 System (GFS) model, which was formerly called the Medium-Range Forecast Model (MRF). The 129 130 GFS forecasts model has 28 sigma (pressure) levels and a T62 spatial resolution (~200 km grid size), which represents physical processes to estimate atmospheric forcings such as winds, 131 temperature, precipitation, geopotential heights at different pressure levels (Hamill et al., 2006). 132 15 ensemble forecasts are obtained by initializing different atmospheric states of the GFS model 133 every day. The control run is initialized by the National Center for Environmental Prediction 134 (NCEP)-National Center for Atmospheric Research (NCAR) reanalysis data (Kalnay et al., 1996) 135 136 while the other 14 ensemble members use a set of seven bred pairs of initial conditions (Toth and Kalnay, 1997), which are re-centered each day on the reanalysis initial condition. In this study, we 137 138 make use of daily precipitation reforecasts from the GFS model consists of 15 ensemble members, up to 15 days in advance, starting from 1979 to till date. We considered the ensemble mean of 139 daily precipitation forecasts to forecast daily streamflow and daily total nitrogen loadings for the 140 141 selected watersheds.

# 142 2.3 Water Quality Monitoring Network (WQN) Database

USGS provides national and regional descriptions of stream water quality conditions in
Water Quality monitoring Network (WQN) across the nation (Alexander et al., 1998). The WQN

145 database comprises of water quality data from the USGS monitoring networks for large watersheds (National Stream Quality Accounting Network, NASQAN) as well as watersheds that are 146 minimally developed (Hydrologic Benchmark Network, HBN). We used the observed daily 147 concentrations of Total Nitrogen (TN) for the 18 stations in the SEUS from the WON database. 148 By selecting watersheds from the HCDN database, we basically ensure that both the streamflow 149 is minimally affected by anthropogenic influences. Bu, water quality data is influenced by the land 150 use type. Based on the USGS National Land Use Classification Data (NLCD) data of 2001, we 151 calculated the percentage area under agriculture and urban (Table 1) land use. From Table 1, we 152 153 can see the distribution with seven, six and five watersheds having 20%-30%, 10%-20% and 0%-10% of area under agriculture respectively. On the other hand, the urban land use is less than 10% 154 with the exception being station #3(23%). TN loadings for these stations are available over a 155 period of 12-23 years with samplings being available on average 5 to 6 times per year (Table 1). 156 For additional details about the WQN database, see Alexander et al. (1998). We next provide 157 details on the methodologies behind the development of streamflow and total nitrogen loadings 158 159 forecasts for the selected watersheds.

## 160 3. Streamflow and Total Nitrogen Forecasting Models

161 The overall schematic diagram of the daily streamflow and nutrients forecasting methodology 162 is shown in Figure 2. Daily streamflow forecasts using weather forecasts have been pursued by 163 many studies (Day, 1985; Wang et al., 2011; Yang et al., 2014), but efforts to use those streamflow 164 forecasts to develop nutrient forecasts are very limited.

## 165 **3.1 Daily Streamflow Forecasts**

To develop daily streamflow forecasts, we first identified the grid points of large scale 1-dayahead forecasted precipitation (referred as FP, hereafter) that exhibit significant correlation with

daily streamflow from the HCDN database (Table 2). The correlation was considered statistically significant when it was greater than 1.96/sqrt(n-3), where 'n' denotes the number of days over 170 1979 to 2009 period. This helps us to identify the neighboring grid points that modulate the 171 streamflow of a particular watershed.

Given that the selected large scale FP grids are inter-correlated with each other, Principal 172 Component Analysis (PCA) was applied to select first few principal components which explained 173 over 90% variability in the precipitation data. PCA, also known as empirical orthogonal function 174 (EOF) analysis, transforms the correlated variables to orthogonal uncorrelated principal 175 176 components (See details in Oh and Sankarasubramanian, 2013). The number of principal components varies from 2 to 5 among different sites (Table 2). These principal components (PCs) 177 of 1-day ahead FP as well as daily streamflow over the previous three days, prior to the forecasting 178 179 date, were selected as the predictor for the semi-parametric statistical model. For example, to forecast streamflow on a particular day WQN data was observed, say March 14 in a given year, 180 predictors were the 1-day ahead FP issued on Mar 13 and the 1-day average daily streamflow from 181 182 March 11 to March 13 in that year. Thus, the number of predictors varies from 3 to 6 (i.e., the number of selected PCs shown in Table 2 and one predictor for the 1-day average streamflow). 183 184 The streamflow over the previous three days could be considered as a surrogate for antecedent soil moisture conditions. Then, the nearest neighbor resampling method was employed to predict daily 185 streamflow for that particular day in which WQN was observed. 186

## 187 *K* – *Nearest Neighbor (K-NN) resampling approach*

After obtaining the predictors, PCs of the FP grids and 1-day average streamflow, we utilized the K - Nearest Neighbor (K-NN) resampling method proposed by Lall and Sharma (1996). Similar application of K-NN resampling approach was employed for developing monthly streamflow forecasts conditional on climatic predictors (Souza et al., 2003; Devineni et al., 2008).
The K-NN approach resamples daily (or monthly data) from historical data to generate values that
were observed in the past. Typically, the K nearest neighbors are identified between predicted time
series and the historical series based on the Euclidean distance. Then a weighing function (e.g. Lall
and Sharma, 1996) is generally assigned such that more weights are given to the nearest neighbors
while less weights are given the farthest neighbors to estimate the predicted time series. Finally,
multiple ensembles are generated to estimate the conditional mean of the time series.

In the K-NN scheme, we used the Mahalanobis distance instead of the Euclidean distance, 198 since the selected predictors – PCs of the principal components and the streamflow over the past 199 observations- could be correlated. Therefore, for forecasting the streamflow for a given day 200 observed in the WQN data, all the neighbors were chosen based on the historical time series of 1-201 day ahead FP and previous 1-day average streamflow for that day over the period 1979 to 2009, 202 203 leaving out the daily predictors and predictands over the entire forecasting year (i.e., 365 days). 204 This implies that in order to forecast streamflow for a given day, 30 historical years are available (excluding the forecast year) for identifying similar conditions. Since this is a small sample size 205 for identifying neighbors, we also considered daily streamflow over the three previous days, 206 resulting in a total 120 neighbors, to develop streamflow forecasts for a given day. Mahalanobis 207 distance for all these 120 neighbors were estimated using equation (1) (Mahalanobis, 1936): 208

209 
$$D_{i,j} = \sqrt{(X_i - X_j)^T S^{-1} (X_i - X_j)}$$
 (1)

where  $X_i$  and  $X_j = (X_1, X_2, \dots, X_{120})$  are the multivariate vectors containing predictor variables at the conditioning time step *i* and *j* denote the rest of the time periods that are considered for identifying the neighbors with *T* representing the transpose operation and  $S^{-1}$  denoting the inverse of the predictor ( $X_j$ ) covariance matrix. The matrix,  $X_i = \begin{bmatrix} x_1, & x_2 \end{bmatrix}$  denotes the multivariate vector with  $x_1$ , and  $x_2$ , denoting the 1-day averaged streamflow before the forecasted day and the PCs of 1-day ahead FP. The first 50 nearest neighbors and their corresponding daily streamflow values were selected based on Mahalanobis distance,  $D_{i,j}$ , to develop ensemble of daily streamflow forecasts. These daily streamflow values from the 50 neighbors were used to draw 500 ensembles that represent the conditional distribution with the density/weight represented by each member, *j*, by the kernel in equation (2):

220 
$$w_j = \frac{1/j}{\sum_{k=1}^{K} 1/k}$$
,  $i = 1, 2, \dots, K$  (2)

where K = 50 (the number of neighbors),  $W_i$  represents the probability with which neighbor is resampled in constituting the 500-member ensemble. Finally, the forecasted streamflow for each day is calculated as the conditional mean of these 500 realizations obtained from the 50 neighbors. The ensemble mean of daily streamflow forecasts are specifically obtained for the days on which WQN data is available, so that the ensemble mean of daily streamflow forecasts could be used for developing forecasts of total nitrogen loadings, whose details are described in the next section.

### 227 **3.2 Daily Nitrogen Loadings and Concentration Forecasts Development**

Daily nitrogen loadings forecasts are developed by forcing the daily streamflow forecasts with the Load Estimation (LOADEST) program. The LOADEST model can be employed with the observed or predicted daily streamflow time series at any given site. Streamflow forecast developed using large-scale precipitation forecasts and previous 3-day average streamflow using the non-parametric model is forced with the LOADEST model to develop nutrient forecasts. LOADEST is a statistical model that estimates daily loadings based on the observed daily
streamflow and the centered time (*dtime*) of the year of the observation (Runkel et al., 2004).

235 
$$\ln(L_j) = a0 + a1 \ln(Q_j) + a2 \ln Q_j^2 + a3 \sin(2\pi dtime) + a4 \cos(2\pi dtime) + \hat{\varepsilon}_j \dots (3)$$

where  $L_j$  denotes the observed daily loadings from the WQN database with 'j' denoting the day of observation,  $Q_j$  is the observed daily flow and *dtime* is the centered time which is a function of the observation's number of days (from January 1) in the calendar year, *a0-a4* denote the model coefficients and  $\hat{\mathcal{E}}_j$  is the estimated residual for the model. The expression *dtime* is centered to avoid multi-collinearity and *dtime* also represents the seasonality in loadings pattern. For a detailed expression on *dtime*, see Cohn et al. (1992).

The LOADEST model allows the user to select the best-fitting regression model from eleven 242 predefined regression models using the Akaike Information Criterion (AIC) (Akaike, 1981). Five 243 regression models that include a linear time trend are not appropriate, since we are employing 244 observed streamflow to estimate simulated loadings for HCDN watersheds. Therefore, the 245 246 simulated nutrient loadings based on the remaining regression models (i.e., model forms: 1, 2, 4 247 and 6 as defined in Runkel et al., 2004) in the LOADEST program do not have any time trend. Equation (3) represents the model form 6. Model form 1 (2) considers only the first two (three) 248 249 terms in the right hand side (RHS) of equation (3), whereas model form 3 considers all the terms 250 except the third term in the RHS of equation (3). For further details on model forms, see Runkel et al. (2004). Table 3 shows the "goodness of fit" statistics (coefficient of determination (R<sup>2</sup>) and 251 252 AIC) in predicting the observed daily loadings in the WQN database (Table 1) and the coefficients of the best fitting regression model for total nitrogen for the selected18 stations. 253

From Table 3, we infer that R<sup>2</sup> ranges from 0.83-0.97 indicating good fit of the observed daily loadings over 18 stations. Using these parameters, we next estimate the forecasts of daily loadings 256 using the ensemble mean of daily streamflow forecasts developed using the retrospective weather 257 forecasts. Strictly speaking, these parameters should have been obtained by leaving out the observed WQN loadings on the day of the forecasting. Since we have more than 50 observations 258 259 at each site (Table 1), the regression coefficients and model forms did not change substantially. Hence, we used the parameters of the regression coefficients given in Table 3 to estimate the 260 forecasted total nitrogen loadings. These forecasted loadings are divided by the forecasted 261 streamflow to estimate the forecasts of total nitrogen concentrations for the 18 selected watersheds. 262 The forecasted daily streamflow and total nitrogen loadings and concentration are respectively 263 compared with the observed streamflow and the observed WQN daily loadings based on Spearman 264 rank correlation and Root Mean Square Error (RMSE) in predicting the observed information. 265

266

#### 267 **4. Results and Analysis**

In this section, we present skill in predicting variability (rank correlation) and accuracy (RMSE) of observed streamflow and WQN loadings using the forecasted daily streamflow obtained using the K-NN approach.

#### **4.1 Skill in Forecasting Daily Streamflow**

We first summarize the performance of the daily streamflow forecasts for only those days when TN loadings are measured (Figure 4a). Based on that, we infer that all the stations show statistically significant correlations with 8 sites showing correlations greater than 0.8 (Figure 3a). Similarly, RMSE (in cfs per unit area) is lesser than 1 for all states except stations #11, 17 and 18 (Figure 3b). These errors primarily occur due to the inability of the model to predict high values as indicated by very high residuals. For instance, RMSE for station #17 drops from 3.56 to 0.55 by excluding only one extreme observation recorded on 2/10/1981 (not shown here). The RMSE for station #11 (#18) are adjusted to 0.73 (0.29) by dropping one (two) high flow value(s). Although this conditional bias is not observed at all the stations, we infer that the daily streamflow forecast model has poor skill in predicting high flow values. We defer this issue for further discussion at the end of this section. Given this evaluation in predicting observed streamflow on days with WQN data, we next evaluate the performance of daily nitrogen loadings and concentration forecasts for the 18 stations.

# 285 4.2 Skill in Forecasting Total Nitrogen Loadings and Concentration

Using the ensemble mean of the daily forecasted streamflows as a predictor in the LOADEST 286 287 model, we estimate the forecasted TN loadings for those days in which measurements are available in the WQN database. Figure 4a (4b) shows rank correlation (RMSE) between forecasted daily 288 loadings and observed loadings for the 18 stations. Daily loadings of TN forecasts exhibit 289 290 statistically significant relationship between observation and forecasts at all the stations with correlation coefficients being greater than 0.8 in nine stations. We also infer the correlation is 291 higher in coastal regions as opposed to the inland watersheds. Similar to the skill of daily 292 293 streamflow forecasts, loadings forecasts produce high RMSE in some stations despite their ability to predict the observed variability. This failure in forecasting TN loadings is primarily due to the 294 295 inability in estimating high flow events as discussed in Section 4.1.

Further extending our analysis, we estimated TN concentration from the LOADEST model utilizing the forecasted streamflow and loadings and then compared the forecasted TN concentration with the observed concentrations available in the WQN database (Figure 5). Though the forecasted concentration is smaller compared to correlation reported for streamflow and loadings, the correlation is statistically significant at all stations except stations #6 and #18. Given that concentration is the ratio of loadings to the streamflow, the error in predicting both loadings and streamflow result in reduced skill. We are not reporting the RMSE since the trend is similarto Figures 4 and 5.

## **4.3 Factors affecting the skill in forecasting TN loadings**

In order to understand what factors control the skill in forecasting the TN loadings utilizing 305 the weather forecasts, we plotted the rank correlation against basin area (Figure 6). Rank 306 307 correlation in forecasting streamflow (Figure 6a) and TN loadings (Figure 6b) are statistically significant for all the stations and the skill increases as the drainage area increases, which is 308 consistent with previous findings (Bloeschl and Sivapalan, 2013). This is primarily due to the fact 309 310 that retrospective weather forecasts being available over large spatial scales, the developed streamflow and TN loadings forecasts modulate better with the observed streamflow and WQN 311 loadings. 312

To gain further understanding on how the developed model estimate the observed streamflow 313 and nutrients, we present the performance of daily streamflow and TN forecasts for the sites that 314 have the best and worst skill under each case (Figure 7). For quantifying the performance of 315 316 streamflow, we considered the continuous daily streamflow records available from USGS instead of comparing the performance on the days with WQN data. From streamflow forecasts for the site 317 318 with best skill (Figure 7a), we understand that overall performance is good, but the K-NN resampling approach based daily streamflow forecasts consistently underestimate high flow events. 319 This underestimation/error in streamflow forecasts partially arises from the errors in the 320 321 precipitation forecasts also. We discuss this issue in detail in the next section. From Figure 7b, the site performs poorly in forecasting flows above 8000 cfs. It is important to note that for the same 322 323 site we observed significant correlation in predicting both the streamflow and the loadings on those 324 days with WQN data being present. Thus, evaluating the performance of K-NN resampling model

325 over the entire time series of observed records provide a more confirmatory evaluation of the 326 model. The primary reason the K-NN resampling model performs poorly at site 3 (Rocky River near Norwood, NC) is due to the limited correlation between the observed precipitation and the 327 328 forecasted precipitation during the summer months (figure not shown). Thus, the error resulting from K-NN resampling arises from both errors in the precipitation forecasts and in estimating the 329 initial conditions as well as from the model itself. Even if one uses physically-based distributed 330 models (e.g., Sacramento model), the skill of streamflow forecasts is heavily dependent on the 331 skill of precipitation forecasts as well as the season of forecasting. 332

333 Figure 7c shows the performance of the TN loadings forecasts obtained using the streamflow forecasts with the LOADEST model. Even here, the same issue is highlighted with the limited 334 ability of the forecasts in predicting the nutrients on days with high flows resulting in 335 underestimated TN loadings. But, the model estimates the variability of the observed nutrients 336 very well. Figure 7d shows the performance of TN loadings for a station with the worst skill. The 337 skill of the streamflow forecasts resulting from the K-NN resampling approach in predicting the 338 339 observed daily streamflow recorded at USGS stations is marginal with an average daily correlation of 0.6. Given that the  $R^2$  of the LOADEST model is 0.912 for the selected station (Table 3), the 340 341 poor performance primarily results from the inability of the streamflow forecasts which partly arises from the resampling model as well as from the skill of the precipitation forecasts. Thus, to 342 develop nutrient forecasts, it is important that the skill of daily precipitation and streamflow 343 344 forecasts should be good and also the load estimation model should have very high skill in predicting the observed nutrients. Given that these basins are virgin, it could be argued the 345 346 predominant source of nutrient loadings arise from the nonpoint sources whose primary transport 347 is the streamflow. Thus, for developing a broader understanding of what could be achieved in forecasting daily nutrients in virgin basins, one could look at the skill in predicting daily streamflow forecasts using the retrospective weather forecasts for the selected 18 stations. We summarize this information under the discussion in the next section by summarizing the skill of daily streamflow forecasts under each month for the selected 18 stations.

#### 352 **4.4 Discussion**

The intent of this study is to develop daily forecasts of total nitrogen (TN) loadings and its 353 concentration in 18 HCDN watersheds that are minimally impacted by anthropogenic influences 354 over the southeastern US. Given that these watersheds experience virgin flow, our hypothesis is 355 356 that most of the nutrient transport at daily time scales could be explained based on observed 357 streamflow. For this purpose, we related the observed daily streamflow and loadings using the LOADEST model (Table 3), which showed significant skill in predicting the daily variability in 358 359 TN loadings purely based on observed streamflow. Given that the predominant driver of streamflow in watersheds under rainfall-runoff regime is precipitation, we utilized the 360 retrospective 1-day ahead precipitation forecasts from the reforecasts database of Hamill et al., 361 (2004) and daily streamflow over previous three days (as a surrogate for soil moisture storage) for 362 developing daily streamflow forecasts on the days with recorded WQN observations. The 363 364 forecasted ensemble average of the streamflow obtained using the K-NN resampling model was used within the LOADEST model to estimate the forecasted daily TN loadings and the 365 concentrations. We observed the correlation between observed TN loadings and the forecasted TN 366 367 loadings being significant in almost all the stations. But, the forecasted concentration showed reduced skill, since it accounted for the errors in both loadings and streamflow. Though one could 368 improve the streamflow forecasts developed using the K-NN resampling approach by considering 369 370 physically distributed hydrologic models and by explicitly considering additional input variables

(e.g., temperature forecasts, humidity), we certainly captured the first-order information on the daily streamflow variability by utilizing the retrospective precipitation forecasts and employed that for assessing the potential in developing nutrient forecasts. Another advantage with the streamflow forecasts using K-NN approach is in specifying the conditional distribution of flows. Thus, one could use the conditional distribution of streamflows with the LOADEST model to develop the conditional distribution of loadings, which could be used to estimate the probability of violating the concentration at the daily time scale.

It is important to note that all the skill reported in Figures 3-6 consider the ability to predicting 378 379 those days when the WQN observations are available. The primary difficulty in assessing the potential for developing nutrient forecasts at daily time scale is the discontinuous nutrient 380 samplings recorded in the WQN database. Oh and Sankarasubramanian (2012) addressed this issue 381 by computing the coefficient of determination  $(R^2)$  of the winter TN loadings forecasts as a product 382 of the  $R^2$  in forecasting the seasonal streamflow and the  $R^2$  of the LOADEST model for the winter 383 season. Similarly, we express the skill of TN forecasts (equation 4) at daily time scale as a product 384 of the R<sup>2</sup> of streamflow (Q) forecasts developed from the K-NN approach for each day in the 385 calendar year and the  $R^2$  of the LOADEST model reported in Table 3. 386

$$R_{TN-daily}^2 = R_{(LOADEST)}^2 * R_{Q-daily}^2 \qquad \dots (4)$$

 $R_{Q-daily}^{2}$  is computed between the observed daily streamflow over the period 1979-2010 and the computed ensemble mean of the streamflow forecast from the K-NN resampling approach. Since the skill of daily streamflow forecasts differ substantially depending on the season, we plot the  $R_{TN-daily}^{2}$  as a box-plot for each month (Figure 8). Basically, Figure 8 pools the daily correlation,  $R_{TN-daily}^{2}$ , for a given month across the 18 stations. For instance, in January, we expect 31\*18 daily

correlations and the box-plot simply summarizes the skill in predicting daily TN for that month 393 over the Southeast US. About 75% of the  $R_{IN-daily}^2$  at daily level are statistically significant level 394 over the period January to May and also from November and December (Figure 8). Daily TN 395 forecasts show relatively better skills in predicting observed TN variability during winter and 396 spring. On the other hand, the skill of  $R_{TN-daily}^2$  is poor during the summer and fall seasons. It has 397 been well-known that retrospective precipitation forecasts have lower skill during the warm-season 398 (Hamill et al., 2004). One of the possible reasons of relatively poor skill during summer and fall 399 is that weather phenomena during these seasons depend greatly on local scale processes while 400 large-scale models do not have the ability to capture it (Hamill et al., 2006). Thus, the poor skill 401 of  $R_{TN-daily}^2$  primarily arises from the skill in forecasting precipitation during the summer and fall 402 seasons. Additionally, the role of temperature during the summer season is also much higher with 403 404 enhanced evapotranspiration. However, considering temperature as additional predictor did not result in substantial increase in the  $R_{Q-daily}^2$  for the summer season. Perhaps, if one considers a 405 406 physically-based hydrologic model, the skill in predicting daily streamflow could improve during the summer season. We plan to investigate this as a future work in assessing the potential for 407 developing nutrient forecasts with streamflow forecasts being derived from a physically-based 408 distributed hydrologic model. Thus, the potential skill  $(R_{TN-daily}^2)$  in predicting daily nutrients is 409 statistically significant for the winter and spring season in almost all the stations. One could utilize 410 411 this to develop adaptive nutrient management strategies for controlling the point sources (e.g., 412 waste water treatment plants) so that the downstream TN concentration does not exceed the 413 desired/ EPA standards.

414 Given that consideration of both forecasted precipitation and 3-day average streamflow prior to the forecasting day exhibit significant skill in predicting the observed TN loadings from the 415 WON database, we investigated the role of each predictor in contributing to the overall skill 416 reported in Figures 3-5. This analyses will also provide information on the role of basin storage, 417 3-day average streamflow, in contributing to the forecast skill. For this purpose, we developed the 418 419 streamflow forecasts using only one predictor and then used that streamflow forecast to estimate 420 the TN loadings. Figure 9 quantifies the role of each predictor, 3-day average streamflow prior to forecasting day (Q) and 1-day ahead precipitation forecasts (FP), in contributing to the skill, 421 422 correlation and RMSE, in forecasting TN loadings for all the 18 sites. It is important to note that 423 the correlation and RMSE were obtained by forecasting for the actual day for which the samples are available in a given site in the WQN database. Figure 9 clearly indicates that the combination 424 of both 3-day average streamflow and 1-day ahead precipitation forecasts as predictors result in 425 improved correlation and reduced RMSE in estimating daily TN loadings at all the sites. 426 Comparing the skill obtained using only one predictor, 3-day average streamflow or forecasted 427 428 precipitation, we infer that for most of the watersheds, the skill obtained using 3-day average 429 streamflow (prior to the forecasting day) alone as a predictor provides better skill in comparison 430 to the skill obtained using forecasted precipitation alone as a predictor with the exception being stations 6, 8 and 18. On an average, in most of the basins, 3-day average streamflow prior to the 431 forecasts alone can explain around 25% (average correlation across all the sites is 0.52) of the 432 433 variability in the observed nutrients. Several studies have shown that antecedent moisture/flow conditions also play a critical role in influencing the nutrient loadings from the watershed (Vecchia 434 435 2003, Alexander and Smith 2006). This analyses further confirms the critical role of basin storage, 436 both streamflow and nutrients, in influencing the forecast skill. On the other hand, forecasted

precipitation alone, can explain on average 20% (average correlation across all the sites is 0.45) of
the variability in the observed TN loadings in the WQN database. Thus, including both of them
as a predictors in the proposed modeling framework results in overall improvement.

We also investigated how the type of land use influence the skill in forecasting TN loadings. 440 Figure 10 shows the scatter plot between the forecast skill, correlation coefficient between the 441 observed TN loadings and the forecasted TN loadings, and the percentage area under agriculture 442 for each watershed. This indicates basins with higher percentage of agricultural land exhibits 443 higher skill in forecasting the TN loadings. Basin with increased agricultural activity could 444 445 potentially experience increased fertilization application, which could increase the streamflowinduced transport. This indicates the role of basin nutrient storage in influencing the forecast skill. 446 Similar analyses on urban land use did not reveal any relationship with the skill. Thus, analyses 447 from Figures 9 and 10 show that both antecedent moisture conditions and in-basin nutrient storage 448 influence the forecast skill for the selected 18 stations over the SEUS. 449

450

#### 451 **5. Summary and Conclusions**

We developed a semi-parametric statistical model, which utilizes 1-day ahead precipitation 452 453 forecasts from the reforecasts from the NOAA GFS climate model (Hamill et al., 2004) and daily streamflow over the previous three days as predictors, to develop daily streamflow forecasts, which 454 in turn was used to implement a load estimation model, LOADEST, for estimating daily nutrients. 455 456 For each day, conditioned on previous day's streamflow and 1-day ahead forecasted precipitation, 50 nearest neighbors over a three-day window were selected based on the Mahalanobis distance 457 and then observed daily streamflow corresponding to those 50 neighbors were resampled to 458 459 constitute 500 ensemble members to develop a daily streamflow forecast. It is important to note

460 that to develop a forecast for a given day in a year, the entire year's predictors and predictand were left out for identifying the 50 nearest neighbors. Finally, the conditional mean of these daily 461 streamflow ensemble was forced in the LOADEST model to obtain daily forecasts of TN loadings 462 and concentration for days with recorded WQN observations. Skill in developing forecasts of 463 streamflow, TN loadings and the associated concentration were computed using rank correlation 464 465 and RMSE, by comparing the respective forecast values with the WQN observations for the selected 18 HCDN stations. The forecasted daily streamflow and TN loadings and their 466 concentration exhibit statistically significant skill in predicting the respective daily observations 467 468 in the WQN database at all the 18 stations over the SEUS.

The study also found that the skill in predicting the observed TN loadings is higher for large 469 watersheds indicating the large-scale precipitation forecasts from the reforecast database better 470 correlate with precipitation and streamflow over large watersheds. Analyses also showed that 471 compared to the forecast precipitation, the 3-day average streamflow prior to the forecasting period 472 played a dominant role in contributing to the skill of the forecast. We also observed the skill in 473 forecasting TN loadings is higher for basins having higher percentage of the area under agriculture. 474 These findings confirm that basin storage, both streamflow and nutrients, play a critical role in 475 476 influencing the skill of the forecast. Further, to overcome the limited samplings of TN in the WQN data, we extended the analyses by developing retrospective daily streamflow forecasts over the 477 period 1979-2012 using reforecasts based on the K-NN resampling approach. Based on the 478 coefficient of determination ( $R_{Q-daily}^2$ ) of the daily streamflow forecasts, we computed the potential 479 skill ( $R_{IN-daily}^2$ ) in developing daily nutrient forecasts based on the R<sup>2</sup> of the LOADEST model for 480 each station. The analyses showed that the forecasting skills of TN loadings are relatively better 481 in winter and spring months while skills are inferior during summer months. These findings are 482

483 consistent with other studies (Devineni and Sankarasubramanian, 2010; Sinha and 484 Sankarasubramanan, 2013) which show that large-scale precipitation forecasts derive their skill from ENSO climatic modes in the SEUS. One possible reason for this poor skill in summer is due 485 to the dominance of local-scale processes during the summer season. Other possible reasons could 486 be due to the limitations in the methodology. We resampled neighbors to develop daily streamflow 487 ensemble, which of course will not have members beyond the maximum observation over the 488 selected 50 neighbors. Further, air temperature can play a dominant role during the summer and 489 fall seasons, resulting in enhanced evapotranspiration and reduced baseflow from the watershed. 490 491 Despite these limitations, there is potential in utilizing the daily streamflow forecasts for developing daily nutrient forecasts, which could be employed for various adaptive nutrient 492 management strategies for ensuring better water quality. 493

Though the watersheds considered under this study have experienced moderate agricultural 494 activity, extending the above modeling framework for basins experiencing significant urbanization 495 will require additional information. For instance, as the basin gets urbanized, it is natural to expect 496 497 the point TN loadings from waste water treatment (WWT) plants to influence the downstream loadings and concentration. Under such situation, it would be useful to consider the discharges 498 from the WWT plants as predictors in developing the model. One could also use the TN forecast 499 to control point loadings so that the downstream TN concentration is within the prescribed standard. 500 For basins experiencing significant non-point pollution from agriculture, one could use 501 information from remote sensing satellites that quantify the chlorophyll concentration could be 502 also be considered as nutrient storage in the river reach and water bodies (Jones et al., 2005). Thus, 503 adequate monitoring of changes in basin land use and nutrient conditions could provide additional 504

information in developing a TN forecasting model for watersheds experiencing significant humaninterference.

508	Acknowledgments: The first author's PhD dissertation research was partially supported by the
509	U.S. National Science Foundation CAREER grant CBET-0954405. Any opinions, findings, and
510	conclusions or recommendations expressed in this paper are those of the authors and do not reflect
511	the views of the NSF. Authors also wish to thank Dr. Rob Runkel of USGS for his support in
512	setting up the LOADEST model.
513	
514	
515	References
516	Akaike, H., 1981. Likelihood of a model and information criteria. J. Econometrics, 16, 3–14.
517	Alexander, R.B., Slack, J.R., Ludtke, A.S., Fitzgerald, K.K., Schertz, T.L., 1998. Data from
518	selected US Geological Survey national stream water quality monitoring networks. Water
519	Resour. Res., 34, 2401-2405.
520	Alexander, R.B., Smith, R.A., Schwarz, G.E., 2000, Effect of stream channel size on the delivery
521	of nitrogen to the Gulf of Mexico, Nature, 403, 758-761.
522	Alexander, R.B., Smith, R.A., 2006. Trends in the nutrient enrichment of U.S. rivers during the
523	late 20th century and their relation to changes in probable stream trophic conditions. Limnol.
524	Oceanogr., 51, 639-654.
525	Anderson, M.L., Chen, Z-Q., Kavva, M.L., Feldman, A., 2002. Coupling HEC-HMS with
526	atmospheric models for prediction of watershed Runoff. Journal of Hydrologic Engineering,
527	7(3), 312-318.

- Bloeschl G., M. Sivapalan, et al., 2013. Runoff prediction in ungauged basins: synthesis across
  processes, places and scales. Cambridge University Press.
- Borsuk, M.E., Stow, C.A., Reckhow, K.H., 2004. Confounding effect of flow on estuarine
  response to nitrogen loading. J. Environ. Eng., 130, 605-614.
- 532 Bricker, S.B., Clement, C.G., Pirhalla, D.E., Orlando, S.P., Farrow, D.R.G., 1999. National
- 533 Estuarine Eutrophication Assessment. Effects of Nutrient Enrichment in the Nation's
- 534 Estuaries, NOAA—NOS Special Projects Office.
- 535 Caraco, N.F., Cole, J.J., 1999. Regional-scale export of C, N, P and sediment: what river data tell
- us about key controlling variables, in: Tenhunen, J.D. Kabat, P.(Eds.), Integrating Hydrology
- Ecosystem Dynamics and Biogeochemistry in Complex Landscapes. John Wiley, New York,
  pp. 239–253.
- 539 Chiew, F. H. S., McMahon, T. A., 2002. Global ENSOstreamflow teleconnection, streamflow

540 forecasting and interannual variability. Hydrol. Sci. J., 47, 505–522.

- 541 Cigizoglu, H.K., 2003.Estimation, forecasting and extrapolation of river flows by artificial neural
  542 networks. Hydrological Sciences Journal, 48.3, 349-361.
- 543 Clark, M.P., Hay, L.E., 2004. Use of medium-range numerical weather prediction model output
  544 to produce forecasts of streamflow. J. Hydrometeorol.,5(1), 15-32.
- 545 Cohn, T. A., Caulder, D. L., Gilroy, E. J., Zynjuk, L. D., Summers, R. M., 1992. The validity of
- a simple statistical model for estimating fluvial constituent loads: An Empirical study
- 547 involving nutrient loads entering Chesapeake Bay, Water Resour. Res., 28(9), 2353–2363.
- 548 Day, G., 1985. Extended Streamflow Forecasting Using NWSRFS. Journal of Water Resources
- 549 Planning and Management, 111(2), 157-170.

- Devineni N., Sankarasubramanian, A., 2010. Improved categorical winter precipitation forecasts
  through multimodel combinations of coupled GCMs. Geophys. Res. Lett., 37(24), L24704.
- 552 Devineni, N., A.Sankarasubramanian, and S.Ghosh, Multi-model Ensembling of Probabilistic
- 553 Streamflow Forecasts: Role of Predictor State Space in skill evaluation, *Water Resources*
- 554 *Research*,44, W09404, doi:10.1029/2006WR005855,2008.
- Duff, J.H., Tesoriero, A.J., Richardson, W.B., Strauss, E.A., Munn, M.D., 2008. Whole stream
  response to nitrate loading in three streams draining agricultural landscapes. J. Environ.
- 557 Qual., 37, 1133–1144.
- 558 Galloway, J.N., Schlesinger, W.H., Levy I, H., Michaels, A., Schnoor, J.L., 1995. Nitrogen
- fixation: Anthropogenic enhancement-environmental response. Global Biogeochem. Cy.,
  9(2), 235-252.
- 561 Georgakakos, A.P., Yao, H., Georgakakos, K.P., 2010. Upstream regulation adjustments to
- ensemble streamflow predictions, HRC Technical Report 7, Hydrologic Research Center,
- 563 San Diego, CA, 30 June, 2010 (NA08NWS4620023), 63pp.
- Goolsby, D.A., Battaglin, W.A., 2001, Long-term changes in concentrations and flux of nitrogen
  in the Mississippi River basin, USA. Hydrol. Process., 15(7), 1209-1226.
- Hamill, T.M., J. S. Whitaker and S. L. Mullen, 2006. Reforecasts, an important dataset for
  improving weather predictions. *Bull. Amer. Meteor. Soc*, 87, 33-46.
- 568 Hamill, T.M., J. S. Whitaker, and X. Wei, 2004. Ensemble reforecasting: Improving medium-
- range forecast skill using retrospective forecasts. *Mon. Wea. Rev.*, 132, 1434–1447.
- 570 Howarth, R.W., Billen, G., Swaney, D., Townsend, A., Jaworski, N., Lajtha, K., Downing, A.,
- 571 Elmgreen, R., Caraco, N., Jordan, T., Berendse, F., Freney, J., Kudeyarov, V., Murdoch, P.,

- 572 Zhao-liang, Z., 1996. Regional nitrogen budgets and riverine N & P fluxes for the drainages
- 573 to the North Atlantic Ocean: Natural and human influences. Biogeochemistry, 35, 181-226.
- 574 Jones, M.O., J. Kimball, S.W. Running, B.K. Ellis, and A.E. Klene, 2005. Application of
- 575 MODIS for monitoring water quality of a large oligotrophic lake. *Eos Trans. AGU*, 85(52),
- 576 B41A-0160.
- Kalnay, E., Kanamitsu, M., and co-authors, 1996. The NCEP/NCAR 40-Year Reanalysis Project.
  Bulletin of the American Meteorological Society Vol. 77, No. 3, pp. 437-472.
- Lall, U., Sharma, A., 1996. A nearest neighbor bootstrap for resampling hydrologic time series.
- 580 Water Resour. Res., 32(3), 679–693.
- Lin, J., Xie, L., Pietrafesa, L.J., Ramus, J.S., Paerl, H.W., 2007. Water quality gradients across
- Albemarle-Pamlico estuarine system: Seasonal variations and model applications. J. Coastal
  Res., 23, 213-229.
- Mahalanobis, P.C., 1936. On the generalised distance in statistics. Proceedings of the National
  Institute of Sciences of India 2, 1, 49–55.
- 586 Mujumdar, P.P., Kumar, D.N., 1990, Stochastic models of streamflow: some case studies,
- 587 Hydrological Sciences Journal 35 (4), 395-410.
- 588 Mcenery, J., Ingram, J., Duan, Q., Adams, T., Anderson, L., 2005. NOAA'S Advanced
- 589 Hydrologic Prediction Service: Building Pathways for Better Science in Water Forecasting.
- 590 Bull. Amer. Meteor. Soc., 86, 375–385.
- 591 North Carolina DENR Fiscal analysis for proposed nutrient strategy for Falls of Neuse Reservoir
- report: available at: http://portal.ncdenr.org/c/document\_library/get\_file?uuid=2a29f5a4-
- 593 3db1-4c63-bd63-cad51a5ac385&groupId=38364, 2010.

594	Oh, J., Sankarasubramanian, A., 2012. Interannual hydroclimatic variability and its influence on
595	winter nutrient loadings over the Southeast United States. Hydrol. Earth Syst. Sci., 16, 2285-
596	2298.

- 597 Paerl, H.W., Valdes, L.M., Peierls, B.L., Adolf, J.E., Harding, L.W., 2006. Anthropogenic and
- climatic influences on the eutrophication of large estuarine ecosystems. Limnol. Oceanogr.,51, 448-462.
- Piechota, T. C., Chiew, F. H. S., Dracup, J. A., McMahon, T. A., 2001. Development of
  exceedance probability streamflow forecast. J. Hydrol. Eng., 6, 20–28.
- 602 Pinckney, J.L., Paerl, H.W., Harrington, M.B., 1999. Responses of the phytoplankton community
- growth rate to nutrient pulses in variable estuarine environments. J. Phycol., 35, 1455–1463.
- Prairie, J., Rajagopalan, B., Fulp, T., Zagona, E., 2006. Modified K-NN model for stochastic
  streamflow simulation. Journal of Hydrologic Engineering, 11(4), 371-378.
- Rabalais, N.N., Turner, R.E., Scavia, D., 2002. Beyond science into policy: Gulf of Mexico
- hypoxia and the Mississippi River. Bioscience, 52(2), 129-144.
- Rajagopalan, B., Lall, U., 1999. A k-nearest-neighbor simulator for daily precipitation and other
  weather variables. Water Resour. Res., 35, 3089-3101.
- 610 Runkel, R.L., Crawford, C.G., Cohn, T.A., 2004, Load Estimator (LOADEST): A FORTRAN
- Program for Estimating Constituent Loads in Streams and Rivers. U.S. Geological SurveyReport.
- Salas, J., Lee, T., 2010. Nonparametric Simulation of Single-Site Seasonal Streamflows. J.
- 614 Hydrol. Eng., 15(4), 284–296.

- 615 Sinha, T., Sankarasubramanian, A., 2013. Role of initial soil moisture conditions and monthly
- updated climate forecasts in developing operational streamflow forecasts. Hydrol. Earth Syst.
  Sci., 17, 721-733.
- 618 Sharif, M., Burn, D., 2006. Simulating climate change scenarios using an improved K-nearest
- 619 neighbour model. J. Hydrol., 325, 179-196.
- 620 Slack, J.R., Lumb, A., Landwehr, J.M., 1993. Hydro-Climatic Data Network (HCDN)
- 621 Streamflow Data Set, 1874-1988. U.S. Geological Survey Report.
- 622 Souza Filho, F., Lall, U., 2003. Seasonal to interannual ensemble streamflow forecasts for Ceara,
- Brazil: applications of a multivariate, semi-parametric algorithm. Water Resour. Res., 39,
  1307–1325.
- Srinivas, V.V. and Srinivasan, K., 2000. Post-blackening approach for modeling dependent
  annual streamflows, Journal of Hydrology 230 (1), 86-126.
- Turner, R.E., Rabalais, N.N., 1991. Changes in Mississippi River Water Quality this Century.
  Bioscience, 41(3), 140-147.
- Vecchia, A. V., 2003. Relation Between Climate Variability and Stream Water Quality in the
  Continental United States, Hydrolog. Sci. Tech., 19, 77–98.
- 631 Vitousek, P.M., Aber, J.D., Howarth, R.W., Likens, G.E., Matson, P.A., Schindler, D.W.,
- 632 Schlesinger, W.H., Tilman, G.D., 1997. Human alteration of the global nitrogen cycle:
- 633 Sources and consequences. Ecol. Appl., 7, 737-750.
- 634 Wang, Q. J., Robertson, D. E., Chiew, F. H. S., 2009. A Bayesian joint probability modeling
- approach for seasonal forecasting of streamflows at multiple sites. Water Resour. Res., 45,
- 636 W05407, doi:10.1029/2008WR007355.
- 637

638	Wang, E., Y. Zhang, J. Luo, F. H. S. Chiew, and Q. J. Wang, 2011. Monthly and seasonal
639	streamflow forecasts using rainfall-runoff modeling and historical weather data, Water
640	Resources Research, 47(5).
641	Yang, L., F. Tian, Y. Sun, X. Yuan, and H. Hu., 2014. Attribution of hydrologic forecast
642	uncertainty within scalable forecast windows, Hydrol. Earth Syst. Sci., 18(2), 775-786.
643	
644	
645	
646	
647	
648	
649	

Table 1: Baseline information for the 18 selected stations. Percentage land use area under urban

and agriculture are calculated based on the 2001 USGS NLCD data. Values in the parentheses in

the last column show the total number of daily total nitrogen loadings and concentration

654 samplings available for each station.

			Drainage	% Area	% Area		
Station	Station	Station Name	Area	under	under	Number of Years	
Index	Index Number					(# of daily Obs.)	
			( <b>km</b> <sup>2</sup> )	Agriculture	Urban		
1	2047000	Nottoway river near Sebrell, VA	3732.17	16.9	5.0	17 (95)	
2	2083500	Tar river at Tarboro, NC.	5653.94	28.9	8.0	22 (152)	
3	2126000	Rocky river near Norwood, NC	3553.46	28.7	22.8	14 (65)	
4	2176500	Coosawhatchie river near Hampton, SC	525.77	23.9	6.8	13 (100)	
5	2202500	Ogeechee river near Eden, GA	6863.47	23.5	5.0	20 (141)	
6	2212600	Falling creek near Juliette, GA	187.00	0.6	2.4	14 (56)	
7	2228000	Satilla river at Atkinson, GA	7226.07	20.4	7.6	20 (123)	
8	2231000	St. Marys river near Macclenny, FL	1812.99	3.8	5.9	14 (108)	
9	2321500	2321500Santa Fe river at Worthington springs, FL		12.3	6.4	21 (82)	
10	2324000 Steinhatchee river near Cross city, FL		906.50	0.8	4.8	19 (92)	
11	2327100         Sopchoppy river near Sopchoppy, FL		264.18	0.0	1.0	22 (125)	
12	2329000 Ochlockonee river near Havana, FL		2952.59	28.6	6.9	22 (133)	
13	<b>2358000</b> Apalachicola river at Chattahoochee, FL		44547.79	22.5	9.8	23 (152)	
14	2366500         Choctawhatchee river near Bruce, FL		11354.51	19.6	5.6	21 (119)	
15	2368000   Yellow river at Milligan, FL		1616.15	17.6	6.5	21 (123)	
16	2375500	2375500 Escambia river near Century, FL		12.5	4.8	22 (145)	
17	2479155	Cypress creek near Janice, MS	136.23	0	0.9	16 (54)	
18	2489500	Pearl river near Bogalusa, LA	17023.99	15.2	6.8	12 (57)	

Table 2: Station ID, number of selected PCs, and cumulative Eigen values for large scale

657 precipitation grids from the NOAA's GFS model, which provide 1-day ahead precipitation

658 forecasts (Grid numbers are shown in Figure 1)

Station		# of	Cumulative Eigen value		
ID	Selected Grids (total # of selected grids)	Selected PCs	of selected PCs		
1	5, 7, 12, 19-21 (6)	4	0.962		
2	5, 7, 12 (3)	2	0.905		
3	4-5, 11-13, 18-20 (8)	4	0.948		
4	11-13, 18-20, 27 (7)	3	0.909		
5	17-18, 24-27 (6)	3	0.922		
6	9-12, 16-19, 23-26 (12)	4	0.903		
7	24-26, 31-33 (6)	3	0.918		
8	17-19, 24-26, 31-33 (9)	4	0.918		
9	17-19, 24-26, 31-33 (9)	4	0.918		
10	16-19, 23-26, 30-33 (12)	5	0.921		
11	16-18, 23-25, 30-32 (9)	4	0.930		
12	16, 30-32 (4)	3	0.975		
13	23-25, 30-32 (6)	3	0.934		
14	18, 22-25, 30-32 (8)	4	0.938		
15	18, 22, 24 (3)	2	0.912		
16	22, 29-31 (4)	3	0.977		
17	15-17, 22-24, 29-31 (9)	4	0.929		
18	17, 22-24, 30-31 (6)	3	0.932		

659

Station	R <sup>2</sup> (Daily)	AIC (Daily)	Model No	Coefficients of selected LOADEST model				
Index				a0	a1	a2	a3	a4
1	0.948	0.892	4	6.768	1.114	-0.283	-0.069	
2	0.966	-0.131	4	8.122	0.980	0.108	-0.018	
3	0.966	0.496	4	8.863	1.066	-0.195	0.090	
4	0.956	0.905	6	4.446	1.013	0.026	0.238	-0.036
5	0.916	0.837	4	7.721	1.069	-0.084	-0.317	
6	0.853	2.094	1	2.647	1.095			
7	0.968	0.518	6	7.521	1.005	-0.025	-0.083	0.103
8	0.963	0.250	6	6.428	1.088	-0.075	-0.027	0.187
9	0.986	-0.219	6	5.690	1.086	-0.037	-0.078	0.059
10	0.979	0.279	6	5.549	1.241	-0.069	-0.096	0.071
11	0.979	0.516	6	4.351	1.139	-0.043	0.187	0.007
12	0.923	0.585	1	7.341	0.846			
13	0.902	0.193	4	10.563	0.981	0.074	0.165	
14	0.835	0.423	4	9.077	0.931	-0.145	-0.042	
15	0.834	1.085	6	7.238	1.123	-0.131	-0.004	0.176
16	0.873	0.758	4	8.868	1.039	0.147	0.032	
17	0.912	1.233	4	4.555	1.188	0.206	0.328	
18	0.899	0.853	1	10.193	1.047			

Table 3: Performance of LOADEST model in predicting the observed TN loadings from the
WQN database. Models with linear time components (Model No: 3,5, 7-9) are not considered.



666 Figure 1. Locations of 18 water quality monitoring stations and grids of forecasted precipitation

667 from NOAA's reforecast model.





Figure 2: Schematic diagram illustrating the overall approach to forecast daily streamflow and
total nitrogen loadings conditioned on the predictors (gray boxes), daily weather forecasts and
daily average streamflow values for previous 3 days, based on Kernel-Nearest Neighbor (K-NN)
resampling approach.



680



Figure 3: a) Rank correlation and b) RMSE (cfs per unit area) between observed daily

683 streamflow and forecasted daily streamflow for those days with TN loadings being available in

684 the WQN database.





Figure 4: a) Rank correlation and b) RMSE between observed TN loadings and forecasted TN
loadings for those days with TN loadings being available in the WQN database.



690 Figure 5: Correlation between observed TN concentration and forecasted TN concentration for





Figure 6: Role of basin scale, drainage area, in forecasting observed (a) streamflow and (b) TN
loadings provided in the WQN database for the 18 stations.



Figure 7: Comparison of observations and forecasts of streamflow (a and b) and TN loadings (c



- 712 Tarboro, NC) and worst forecasting skill (Streamflow: Rocky river near Norwood, NC, TN:
- 713 Cypress creek near Janice, MS).



Figure 8: Box plot showing rank correlations between observed daily streamflow and forecasted
streamflow aggregated over each month from 1979 to 2009 period. Each plot includes 558
correlations (18 stations × 31 years). The solid line represents the statistically significant (95%)
R<sup>2</sup> corresponding to the null hypothesis that R<sup>2</sup> being equal to zero.



Figure 9: The role of different predictors, 3-day average daily streamflow prior to forecasting
day (Q) and 1-day ahead precipitation forecasts (FP), in forecasting the observed TN loadings is
expressed as (a) correlation coefficient and (b) RMSE between the observed TN and the
forecasted TN loadings for the 18 selected sites.



760

Figure 10: Role of the type of land use, percentage area under agriculture, in influencing the

762 forecast skill which is expressed as the correlation coefficient between the observed TN loadings

and the forecasted TN loadings for the 18 selected watersheds.