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Impacts of climate variability on wetland salinization in the North American Prairies

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Abstract

The glaciated plains of the North American continent, also known as the "prairies", are a complex hydrological system characterized by hummocky terrain, where wetlands, containing seasonal or semi-permanent ponds, occupy the numerous topographic depressions. The prairie subsoil and many of its water bodies contain high salt concentrations, in particular sulfate salts, which are continuously cycled within the closed drainage basins. The period between 2000 and 2012 was characterized by an unusual degree of climatic variability, including severe floods and droughts, and this had a marked effect on the spatial distribution, water levels and chemical composition of wetland ponds. Understanding the geochemical and hydrological processes under changing environmental conditions is needed in order to better understand the risk and mitigate the impacts of future soil and water salinization.

Here we explore salt dynamics in the prairies using field observations from St. Denis, Saskatchewan, taken over the last 40 yr. Measurements include meteorological data, soil salinity, groundwater levels and pond water volume, salinity, and chemical composition. The record includes periods of exceptional snow (1997, 2007) and periods of exception rainfall (2010, 2012), both of which resulted in unusually high pond water levels. However, severe salinization only occurred in response to extreme summer rainfall. We hypothesize that since rainfall and snowmelt activate different hydrological pathways, they have markedly different impacts on salinization. We propose that a wet condition associated with high snowmelt conditions does not pose a strong threat to salinization, which has important implications for agricultural planning. Whilst this hypothesis is consistent with our conceptual understanding of the system, it needs to be tested further at a range of field sites in the prairies.

1 Introduction

Surface water and shallow ground water salinization is a problem affecting agriculture, water resources and ecosystem health in many areas of the world, including Australia (Dehaan and Taylor, 2002; Rengasamy, 2006), the Aral Sea region (Micklin, 2007), playas and sabkhat environments (Tyler et al., 2006) and many other areas (Rengasamy, 2006). Salt dynamics are driven by hydrological processes, which cycle seasonally but also change over long time scales as a result of climate variability and change, and changes in land use and land management practices. The physical mechanisms that lead to salinization operate in different ways in different areas, largely as a function of local climatological, hydrological and geological conditions. This paper takes the salt rich glaciated plains of North America, known as the prairies, as a case study to explore how recent climate variability has dramatically affected the salinity of ponds. The prairies spread from Alberta, Saskatchewan and Manitoba in Canada, through Montana and the Dakotas in USA The landscape is characterized by hummocky terrain, where wetlands and ponds occupy the ubiquitous depressions (Winter, 1989; van der Kamp and Hayashi, 2009). The region supports a diverse community of wildlife species and major agricultural industry that are highly sensitive to the hydrological conditions (Wienhold et al., 1989). Hydrological processes in the prairies are complex and unusual, characterized by closed basins isolated from any regional drainage network, with drainage via a fill and spill sequence into terminal ponds (Shook and Pomeroy, 2011); Snowmelt runoff from surrounding uplands and precipitation on the wetlands are the dominant hydrological inputs into the wetland systems together with occasional runoff events due to high-intensity rainfall, and snow distribution over the landscape is an important control on the hydrology (Shaw et al., 2012; Spence and Woo, 2003); semi-arid conditions result in minimal shallow groundwater recharge from uplands, while ponds drain to shallow groundwater beneath the depressions (van der Kamp and Hayashi, 2009). Nachshon et al. (2013), based on previous work (e.g., Keller et al., 1991; Hayashi et al., 1998b; Berthold et al., 2004; Heagle et al., 2013)

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presented a conceptual model to describe the major salt dynamics occurring within the glacial till portion of the prairies under various land use and climatic conditions. In humid environments, streamflow is an integrated variable that aggregates the effect of the climatic boundary conditions and the land use/management mediated hydrological processes occurring within the watershed (Wheater et al., 1993). In the prairies, where there is often no regional streamflow, these factors are expressed in the pond storage, and therefore studying the ponds is a good way to understand the system sensitivity to a particular change. This work explores pond salinization using a 20 yr record of observations from a field site in the Canadian prairies, which in the past decade has been subject to extremely variable climatic conditions – including both extreme droughts and floods. In particular this study explores the impacts of extreme precipitation on salinity, considering separately summer rainfall and winter snowfall.

2 Methods

2.1 Field site

The field site is located at the St. Denis National Wildlife Area (106°05′0.20″ W, 52°12′31.32″ N), approximately 40 km east of Saskatoon, Saskatchewan, Canada (Fig. 1a). The St. Denis area has a hummocky topography and consists of a clayey glacial till, with typical hydrogeology for the glacial till portion of the prairies (Hayashi et al., 1998a). The main part of this study focuses on ponds 107, 108A, and 109 (Fig. 1). Pond 109 is a semi-permanent pond, i.e., the pond remains wet all year for most years. Ponds 107 and 108A are ephemeral ponds that dry out by the end of most summers. Most of the area is cultivated, with the exception of the area to the north of pond 109 which is a natural grassland, and the numerous wetlands (that is, the ponds and their riparian zones). Pond 109 has a substantial riparian zone with a "willow ring" with a width of ~ 10 m around the pond, covered by trembling aspen trees, balsam poplar and willows rising up to ~ 8 m. Ponds 107 and 108A willow rings are minor compared to pond

109, with a width of 1–2 m, and covered mainly by cat-tails (Typha) rising up to ∼ 1.5 m. A surface runoff flow path exists from pond 107 through pond 108A to pond 109, and pond 109 can be considered as the terminal pond of this local watershed, at least with respect to surface runoff.

2.2 Precipitation data

Precipitation data used in this study are from the nearest climate station with continuous records from 1993 to 2012, the Saskatchewan Research Council Saskatoon climate reference station (Beaulieu and Wittrock, 2013), about 35 km west of the St Denis site. The data are presented as cumulative amounts of summer rain (April–September) and winter precipitation (October–March), mostly snow. The winter data are not corrected for the wind under-catch effect, for example as described by Mekis and Vincent (2011). Comparison with the corrected precipitation data, available to 2007, indicates that the actual winter precipitation is under-reported by ∼ 40 to 50 %. However, for the purpose of this paper these readings are sufficient to indicate inter-annual variability; the focus is on the differences between the reported winter precipitation in every year (from the autumn of previous year, October, to the spring of the reported year, March) and the long-term average.

2.3 Groundwater and surface water data

Ground water levels at a number of piezometers scattered throughout the site have been measured. For this work data from piezometers 94W7 (hereafter the "upland piezometer") and 802P1 (hereafter the "pond piezometer") were used (Fig. 1c). The upland piezometer, located north of pond 109, is screened from 1.8–5.3 m b.g.l. (below ground level) and the pond piezometer, located within Pond 109, is screened from 5.8–7.6 m b.g.l. Water table depths were measured continually from 1997 to 2012, on a daily basis. The depth at the deepest point of pond 109 was measured manually on a monthly basis during the summers from 1968 to the present (Conly et al., 2004).

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Vertical hydraulic gradients between the pond piezometer and the pond water were small. Here, a continuous pond water level time series was constructed from 2007 onwards, by using the daily logged piezometer data to temporally infill between the monthly manual surface water level measurements. The pond's water volume V (m^3) was computed based on the work of Hayashi and van der Kamp (2000):

$$V = 1420h^{2.24} \quad (1)$$

where h is the depth of water at the centre of the pond. These authors limited and validated Eq. (1) for maximal h of 1.2 m, since that was the deepest measurement of the water pond depth at that time. Since over recent years deeper pond water levels were recorded, the original bathymetric survey data of Hayashi and van der Kamp (2000), which is archived at Environment Canada, was used to compute V for $1.2 < h < 1.8$ m and to revalidate Eq. (1) for these depths. It was found (data not shown) that Eq. (1) is valid for these depths with errors smaller than 5 %.

2.4 Pond water salinity

Electrical Conductivity, EC, was measured at 14 ponds at St. Denis (Fig. 1b) several times a year during the summer months from 2009 to 2012. In addition, in pond 109 EC measurements were taken every few weeks in the summer months from 1993 to 2012.

Major ion analysis was conducted annually for pond 109 water from 1994 to 2009, and in 2012. Data from 2007–2009 and 2012 are used in this study.

Pond 109 EC and volume measurements permit an estimate of the mass of dissolved salts (M_{salt}) (kg) in the water, based on an approximate relationship between EC and dissolved salt mass (Rhoades, 1996):

$$M_{salt} = 0.00064 \cdot EC \cdot V \quad (2)$$

where M_{salt} is in kg, EC is in $\mu S\,cm^{-1}$ and V is in m^3.

2.5 Transect measurements

Spatially-detailed manual measurements of EC and water levels along a transect from pond 107 to pond 108A were obtained over a rainy 25 day period in July 2012. A series of 16 mini observation wells were installed to measure shallow water table changes beneath the upland between pond 107 and 108A (Fig. 1c). The transect was located along the shortest path between the ponds, in the lowest part of the landscape. The spill elevation of pond 107 is approximately 1 m higher than that of pond 108A and the lateral distance between the ponds is 35 m. Each mini observation well consists of a PVC tube, 75 cm long, inner diameter of 1.27 cm, perforated along its entire length. Each tube's lower end was sealed and they were inserted into the ground by direct push, to a depth of ∼ 60 cm. Water levels in the mini observation wells were measured manually with a ruler and EC measurements were taken in situ by a portable EC probe (Cole-Parmer, 1481-60, Canada). These measurements were taken on 5, 6, 11, 19 and 24 July. For this period, precipitation (rain) measurements were taken hourly at the climate station located within the St. Denis field site (operated by Environment Canada). In addition, on 24 July an EM38 probe (Geonics Ltd, Canada) was used to measure soil EC around pond 108A and pond 109. Measurements were taken at 1 m intervals along the pond's edge. Since the EM38 readings are sensitive both to water content and salinity, calibration had to be done to obtain an estimate of pore water EC. Since the soil around the pond is fully saturated, it was assumed that the changes in the EM38 readings were due to salinity differences. Manual EC measurements of the pore water at specific points around the pond were taken using the EC probe, and these data were used to calibrate the EM38 readings.

3 Results and discussion

3.1 Field scale changes in pond water levels

In the prairies, changes in pond storage are a response to a large number of often confounding processes, including rainfall amount, timing and intensity, snow spatial distribution and amount, timing of snowmelt, the spatiotemporal pattern of surface runoff vs. infiltration (strongly affected by soil freeze-thaw processes, as well as land use), and the spatiotemporal pattern of evapotranspiration (again strongly affected by land use). A particularly important factor is pre-freeze up soil moisture content: if the soils are very wet then when they freeze they have a very low infiltration capacity, and hence runoff over frozen soils during the subsequent melt period is more intense. This mechanism explains the widespread flooding in Saskatchewan and Manitoba in 2011, which was attributed to high rainfall in the summer of 2010 (see Fig. 2), leading to high antecedent soil moisture. Another factor is the timing of snowfall and accumulation. Earlier snowfall means that the relatively warm soils are more effectively insulated, and hence the extent of soil freezing is less, and a larger proportion of snowmelt is expected to go to infiltration vs. runoff. This discussion serves to demonstrate that the hydrological processes in the prairies are highly complex, and rainfall-runoff type of responses that apply elsewhere, do not tend to work in this environment. Figure 2 contrasts summer and winter precipitation (Fig. 2a) with an integrated measure of annual maximum pond level across the site (Fig. 2b). The pond level measurement was calculated by taking the mean of the normalized depths of 12 ponds (ponds 1, 2, 20, 25, 26, 35, 36, 37, 50, 60, 109, 120). Each pond depth time series (limited to the open-water season) was normalized by dividing the depth by the maximum depth observed in the period 1993–2012. The fact that the averaged level is never 100 % indicates that different ponds reached their maximum level in different years. For the purposes of this discussion, water levels of 70 % and above are assumed to represent unusually wet conditions, and these were measured in 2006, 2007, 2010, 2011 and 2012. It is clear from Fig. 2 that there is a considerable delay before precipitation extremes are translated into responses in the pond levels, and also that there is a different sensitivity to snow vs. rainfall. For all of the wet years, this is summarized, qualitatively, in Table 1. It is also useful to this discussion to consider the years 1994, when current and antecedent conditions were at or slightly below the average for the period of record, and 1997, where the snowpack was deepest, and the pond levels were high, though not up to the 70 % threshold.

3.2 Field scale changes in pond salinity

Salinity of ponds in the prairies is highly variable (Euliss et al., 2004). Stewart and Kantrud (1972) and Millar (1976) distinguished between fresh water ponds (EC < 500 $\mu S\,cm^{-1}$), moderately brackish ponds (EC between 500–5000 $\mu S\,cm^{-1}$) and brackish-saline ponds (EC > 5000 $\mu S\,cm^{-1}$). The reasons for these differences in salinity are understood to be a function of how the ponds interact with the groundwater and surface runoff. Fresh "recharge" ponds lose water to groundwater, brackish-saline "discharge" ponds gain water from groundwater and surface water (Nachshon et al., 2013), and the moderately brackish ponds are a more complex combination of input and output of surface and subsurface water of various degrees of salinity. The St. Denis field site is only 1.6 km × 2.4 km, yet contains ponds with salinities that cover this entire spectrum. To examine the impact of wet conditions on pond salinity, EC measurements from 14 ponds at St. Denis were taken from 2009–2012, shown in Fig. 3. A longer record from only one pond is explored in more detail below. There is an almost completely consistent pattern in the response, with fresh water ponds becoming salinized over the wet period from 2010 onwards, brackish-saline ponds becoming diluted, and moderately-brackish ponds having relatively stable EC values. The water flushed into the ponds may have a varying salinity over the landscape in the different ponds, but these data would suggest that the salinity of this water is of a similar order to the moderately-brackish ponds. Therefore note that while the brackish ponds are diluted, there may still be a considerable mass input of salt into the ponds. Heagle et al. (2013) showed this for pond 50 at St. Denis (a brackish-saline pond), where from 2009 to 2011

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the mass of SO_4 in the pond water increased by more than 50 % whilst the EC reduced by ∼ 20 %. Overall, the data in Fig. 3 suggest that all ponds in the landscape were enriched in salt mass under wet conditions, assuming that the pond volumes increased over the wet conditions period, as indicated by Fig. 2b.

Ponds 60 and 117 are fresh-water ponds that maintained a relatively constant EC over 2009–2012, in contrast to all of the other fresh water ponds where an increase in EC was observed. Ponds 60 and 117 are the highest and most isolated ponds in the site, and unlike all the other fresh water ponds, they do not have elevated ponds adjacent to them, from which they may receive salts, by surface or subsurface pathways.

3.3 Changes in water level in a terminal recharge pond

Pond 109 at St. Denis is a terminal recharge pond (i.e. it collects surface water but does not spill, and loses water and salt to groundwater), which is well-studied, with a good archived dataset. Figure 4 presents pond 109 water depths from 1969 to 2012. A depth of 1.27 m (70 % of the deepest recorded pond depth, shown in Fig. 4) was used to indicate unusually wet conditions. Consequently, and in good agreement with the precipitation data and the regional pond depths (Fig. 2), wet conditions in pond 109 persist in 1997, 2007 and 2010–2012. A distinct difference between the wet conditions associated with snowy winters (1997, 2007) and rainy summers (2010, 2012) is observed. In wet years associated with snowy winters, as in non-wet years, the pond maximal depths occurred in spring time, following the snowmelt. For wet years associated with rainy summers, the pond water levels continually rose up from spring to late summer. In 2010, pond levels rose from June to September. 2012 also had high water levels throughout the entire summer from March to October, with positive increases in the pond level from March to July. For 2011, as indicated in Table 1, the high water levels were due to the high antecedent water storage from 2010, with the (unexceptional) snowmelt event superimposed on this to produce what were then unprecedented water levels. This unprecedented pond level was matched and slightly exceeded in 2012, this time due to high summer rainfall.

3.4 Changes in salinity in a terminal recharge pond

EC measurements for pond 109 have been taken since 1993, excluding the drought years (1999–2004) when the pond dried out completely, and 1995 and 2006 due to logistical problems. Figure 5 presents pond 109 estimated pond water volume, based on the depth–volume relationship in Eq. (1) (A); measured pond water EC and estimated salt mass based on Eq. (2) (B); and changes in pond volume, EC, and salt mass over selected summer periods (C).

It can be seen that until 2010 the pond EC was of the order of 100–500 $\mu S\,cm^{-1}$, meaning the pond would be classified as fresh. The salt mass in the pond during this period was of the order of 500 kg. Average conditions are exemplified by 1994 in Fig. 5. Here, the pond water volume increases in spring time following the snowmelt runoff from the adjacent uplands which causes a strong dilution. Over the summer months, the water levels drop due to a combination of infiltration and evaporation, and the salinity increases, due to evaporation. The mass of salt in the ponds steadily decreases, due to infiltration, and salts accumulate in the soils, especially in the saline ring around the pond (Hayashi et al., 1998b; Heagle et al., 2007; Nachshon et al., 2013). This is the steady salt cycle that operates continuously in the closed drainage basins of the prairies. However, to complete the cycle, salts must re-enter the pond at some point. The re-entry mechanism is harder to observe, but is thought to be associated with the flushing of near surface salts from the soil by the snowmelt via surface or shallow subsurface pathways (Hayashi et al., 1998b). It is also possible that diffusion might play a role in returning salts to the ponds from the soils.

During the two wet years associated with exceptional snowmelt (1997 and 2007) the pond salinity was unexceptional, and pond salt mass was only marginally elevated. This is consistent with the snowmelt re-entry mechanism, but shows that this effect has a negligible and short-lived impact on the salt cycle.

2010 was the first wet year associated with extreme summer rainfall (which fell in that year), and here a dramatic increase in pond salinity (EC rose from 261 to 801 $\mu S\,cm^{-1}$)

and salt mass (M_{salt} rose from 260 to 1350 kg) was initiated, making the pond moderately brackish. In 2011, water levels were even higher, but this was largely due to the antecedent water levels from the previous year. After the melt event and through the summer, the water level dropped, whilst the salt mass increased, perhaps suggesting that the pond was functioning as a flow-through pond during this period (Nachshon et al., 2013), with a highly saline inflow at some point, and a more dilute outflow somewhere else. This will be explored further below. The EC during this period steadily rose from 527 to 846 $\mu S\,cm^{-1}$ due to evaporation. Finally in 2012, which had both high antecedent water levels and salt mass, and extreme high summer rainfall, the salt mass and salinity continued to rise rapidly to unprecedented levels (EC up to 1061 $\mu S\,cm^{-1}$, salt mass peaking at 3800 kg).

3.5 Changes in water chemistry in a terminal recharge pond

The major ion analysis of pond 109 water over the years of 2007–2009 and 2012 reveals a significant change in the cations composition of the pond water. In 2007–2009 the major cations in the pond water were Ca, Mg, and Na with molar fractions of approximately 60, 35, and 5 %, respectively (Fig. 6). In 2012, at the end of the measured wet period, the pond water cation composition was of 35, 50, and 15 % for Ca, Mg, and Na, respectively, indicating enrichment of the pond water with Mg and Na. This enrichment is likely due to dissolution and migration of mirabilite ($Na_2SO_4 \cdot 10H_2O$), bloedite ($Na_2Mg[SO_4]_2 \cdot 4H_2O$), and epsomite ($MgSO_4 \cdot 7H_2O$) which are the more soluble salts that persists in the prairie tills. Under normal conditions, these salts accumulate at the more distant parts of the uplands, away from the ponds, according to the Hardie–Eugster model (Miller et al., 1989; Miller and Brierley, 2011; Skarie et al., 1987; Timpson et al., 1986; Nachshon et al., 2013). The fact that under high summer rainfall these cations migrate into the pond indicates subsurface water flows from high and distant parts of the uplands that flush these salts from the subsurface into the pond.

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3.6 Groundwater-surface water interactions

The measurements obtained by the piezometers (Fig. 1) indicate the direction of subsurface exchanges between pond 109 and the upland to the north of the pond. Figure 7 contrasts differences in the magnitude and direction of this gradient with changes in the mass of salt (M_{salt}) in the pond water, for 2008–2012. It can be seen that most of the time the hydraulic head gradient is from the pond to the uplands, indicating flow out of the pond. However, on several occasions, most prominently in the summer of 2010, the hydraulic head gradients were reversed, indicating fluxes from the uplands back into the pond, with the potential of transporting dissolved salts from the uplands into the pond. For more than 85 % of the data points, these data behave consistently – that is an inflow to the pond is associated with an increasing salt mass and an outflow from the pond is associated with a reducing salt mass. However, for ~ 15 % of the data points, a positive buildup of salt mass coincided with apparent flow out of the pond. These anomalous measurements (red symbols in Fig. 7) were observed mainly in the summer of 2012. These points can be explained either as slow mixing within the pond, or, again, as evidence of non-uniform inflows and outflows to/from the pond.

3.7 Salinity of the riparian zone

Mini observation wells combined with EM38 readings provide estimates of EC around pond 108A (the end of the transect discussed below) and pond 109 (the terminal recharge pond discussed above) (Fig. 8). Around pond 109 average EC was 3342 µS cm^{-1} with a maximum EC of 5500 µS cm^{-1} in the eastern side of the pond. The soil was 3 to 5 times more saline than the pond water, which was 1020 µS cm^{-1} at the time of measurement. For pond 108A average EC at pore water around the pond was equal to 3200 µS cm^{-1} with maximum EC of 4000 µS cm^{-1} in the southern side of the pond. These values are ~ 1.5 times more saline than the pond water (equal to 2300 µS cm^{-1}). It follows that exchanges of water between the pond and soils/groundwater are far more efficient at advecting salts into the pond than out. For

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example, for pond 109, in order to remove the salts added to the pond for every unit of inflow (i.e. exfiltration of groundwater), 3–5 units of outflow (i.e. infiltration of pond water) would be required. It is also evident from Fig. 8 that the spatial distribution of salt in the saline ring is non-uniform with distinctive regions of high and low salt concentrations. We speculate that this may be related to the slope steepness and the spatial distribution of the elevated-adjacent ponds. However, this hypothesis and the associated processes should be explored in future studies.

3.8 Response to summer rainfall along a transect

Figure 9 shows the water levels and EC data observed along a transect between ponds 107 (shown on the left in Fig. 9b and c, where $x = 0$ m) and 108A (shown on the right in Fig. 9b and c, where $x = 35$ m) for a 20 day period in July 2012, which included a large rainfall event on 15 July. For the first two observation times (5 and 6 July, row 1 and 2 in Fig. 9), conditions were relatively dry. The water table pattern reflects the general conceptual model for the prairies (Fig. 6 in van der Kamp and Hayashi, 2009), whereby both ponds are close to equilibrium with the adjacent groundwater, which drops lower beneath the uplands. With respect to salinity; under these conditions the highest EC values were measured in the pore water close to pond 108A, which is the local saline ring of this pond. On 11 July (row 3 in Fig. 9), a minor rainfall event resulted in a moderate increase in the water table, which also allowed for more points to be sampled for EC. The only significant difference in EC was measured at $x = 31$ m (~ 5 m from pond 108A), where EC readings rose from $\sim 2000\,\mu S\,cm^{-1}$ to well over $6000\,\mu S\,cm^{-1}$. This may indicate leaching of salts from upper parts of the unsaturated profile by the infiltrating rain water, but the affect appears highly localized above the saline ring of Pond 108A. The next set of measurements on 19 July (row 4 in Fig. 9), followed a large rain event on 15 July and 16. The water table came to ground surface along much of the transect, and in these parts of the transect seepage of the water above ground surface was visibly observed. We can be certain that this was a saturation excess mode of runoff, and moreover no fill-and-spill of pond 107 into pond 108A was observed. Since

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the piezometers were screened along their entire length, it is not possible to determine whether this was local perching, or groundwater recharge bringing the water table up to the surface, but due to the high infiltration capacity of the fractured till, the latter is believed to be more plausible. During this period the EC was reduced across the transect, due to dilution, except on the right hand edge, closest to Pond 108A. This edge would have received much of the salts being flushed laterally from the transect, as the head gradients show, and hence here the EC rose up from ~ 2620 to 3310 $\mu S\,cm^{-1}$. For the final set of measurements on 24 July (row 5 in Fig. 9), taken after five days with no rain, the water table depth reduced along the entire transect, but not quite to depths as low as the antecedent levels (row 1 and 2 in Fig. 9). Hydraulic gradients from the earlier observations showed two ponds that were disconnected from one another, whereas the later observations show a more-or-less continuous hydraulic gradient from Pond 107 to Pond 108A, suggesting the ponds were then connected, albeit perhaps temporarily. As the pulse of rain water left the profile, the EC profile returned to levels very similar to the antecedent conditions. Adjacent to Pond 108A the salinity dropped again, probably in response to mixing with other fresher water in the pond.

4 Conclusions

Field measurements collected over the past 20 yr from St. Denis in central Saskatchewan shed new light on salt dynamics of prairie wetlands. It was shown that under wet conditions associated with rainy summers, large fluxes of salts from the subsurface are flushed into the ponds. The corresponding change in pond salinity (i.e. salt concentration) depends on the antecedent salinity, the volume of water flushed into the ponds and the salinity of the water flushed into the pond. However, the general tendency is quite uniform, with fresh-water ponds becoming more saline, moderately brackish ponds remaining largely unchanged, and brackish ponds becoming diluted. Most of the time, however, summer rainfall is small compared with potential evaporation, and thus the dominant (or only) source of infiltration is depression-focused

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recharge of snowmelt beneath the ponds. Significantly, it was seen that wet conditions associated with this kind of infiltration, i.e. unusually snowy winters, had a negligible impact on the salt dynamics and pond salinization. Moreover, the rain-associated wet conditions lasted for much longer durations (months) compare to the snow-associated wet conditions (weeks), therefore, increasing the potential impact of the unusual rain conditions on the subsurface salt transport processes.

The relatively high permeability of the top few meters of the prairies till results in a quick response in the upland subsurface storage that leads to relatively quick transport of salts from the upland subsurface to the ponds, on time scales of the order of few days only. Even quicker responses in water flows and pond levels may occur by surface fill and spill events, but usually the surface waters are less saline than the pore water, thus salt transport is more limited under these conditions.

It is hard to predict what would be the climatic conditions in the near and far future, but recent years' evidences, as well as climatic models indicate that the variability of the climate is expected to increase, with a high chance of extreme rain periods, as well as extreme droughts (Frelich and Reich, 2010; Semenov and Stratonovitch, 2010). If a series of extremely wet years, where salt is being accumulated in the ponds, is followed by a severe drought, the high evaporation rates of the drought period will result in further increase in the pond salinity and eventually, all of the salts within the pond water will be precipitated, following the complete drying of the pond. This scenario may result in salt concentrations in the upper soil horizons of the wetlands at concentrations that were never measured before. Since both the wildlife of the prairies, as well as agricultural activities, depend on the wetlands' physical and chemical conditions it is critical to improve our understanding of the geo-chemical processes and to be prepared to cope with salinization processes in the prairies.

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Table 1. Factors influencing the pond levels during wet conditions.

Year	High pond level in previous year	High summer rainfall in previous year	High winter snowpack	High summer rainfall in current year
2006		X	X	X
2007	X	X	X	
2010				X
2011	X	X		
2012	X			X

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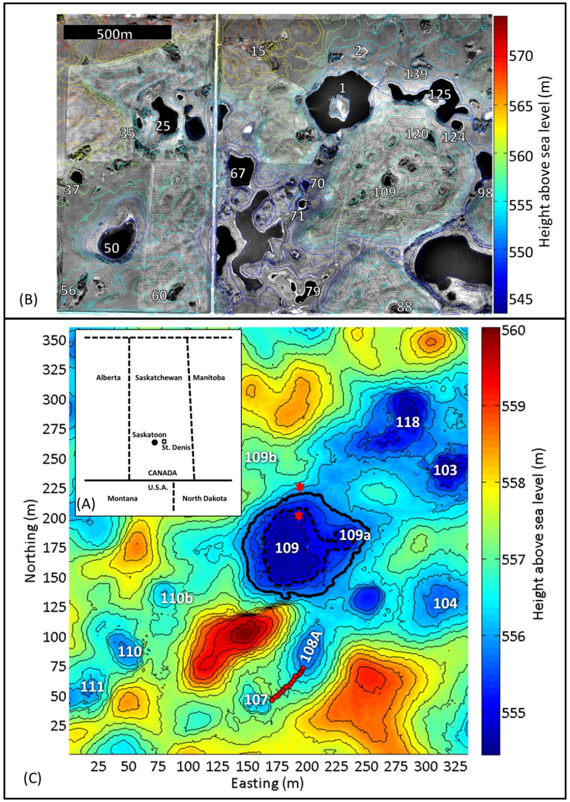

Fig. 1. (A) Regional location of St. Denis; **(B)** St. Denis site; **(C)** enlargement of pond 109 area. In **(C)** the dots between pond 107 and 108A indicate the location where salinity and ground water levels were measured along the transect. Stars represent piezometers; thick solid line around pond 109 indicates the water level of the pond in July 2012; and the dashed line at pond 109 indicates the average location of pond 109 water level. White numbers indicate selected ponds numbers.

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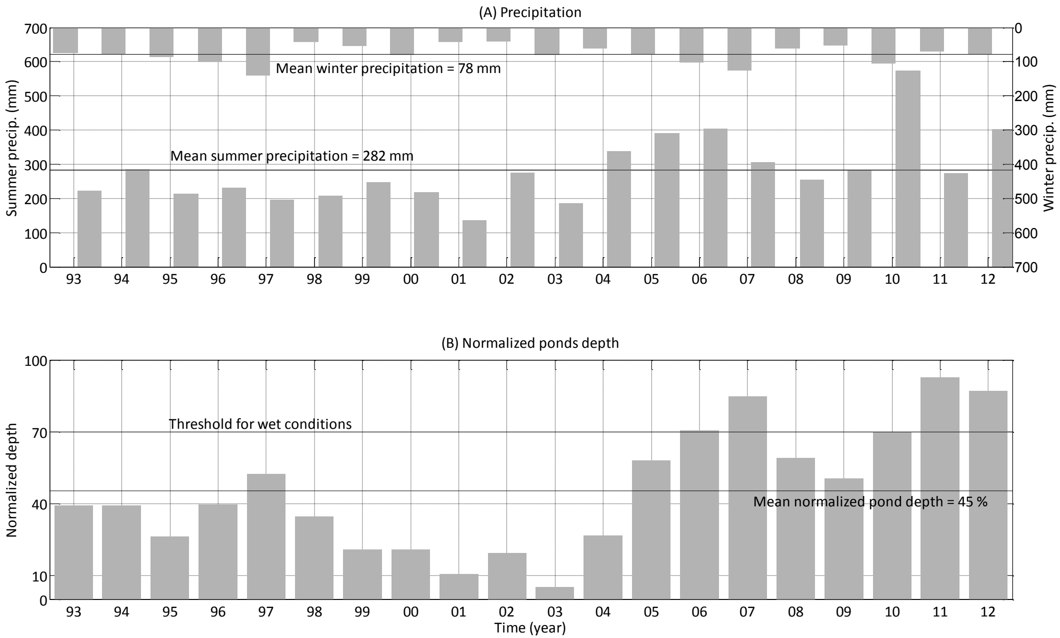

Fig. 2. (A) total summer rainfall plus winter snow per hydrologic year. **(B)** Averaged and normalized depths of 12 ponds at St. Denis.

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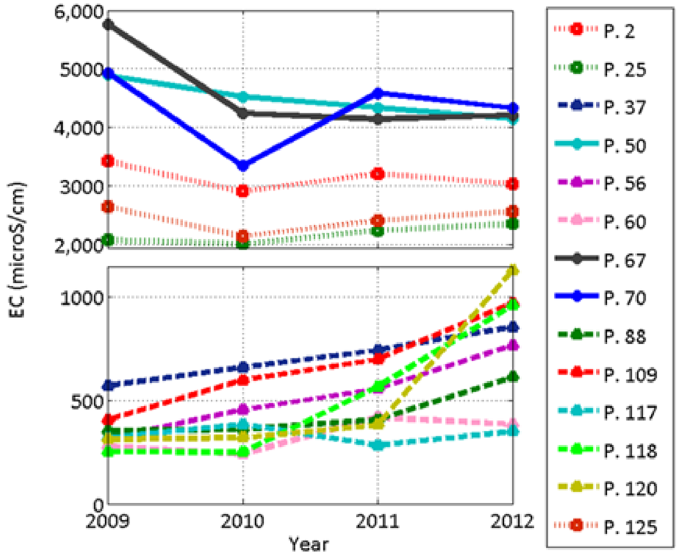

Fig. 3. Averaged EC of selected ponds in St. Denis from 2009 to 2012. Solid, dot, and dashed lines indicate on brackish-saline, moderately brackish, and fresh water wetlands, respectively.

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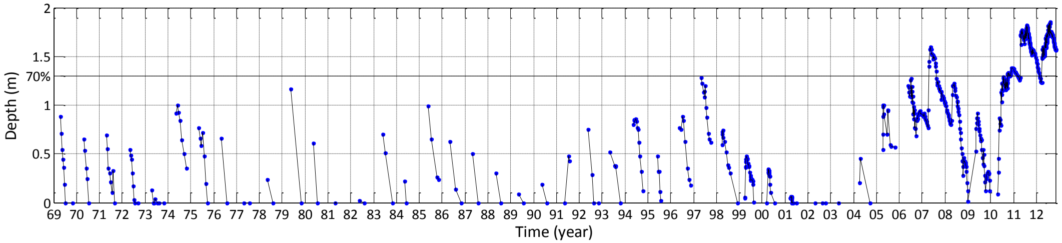

Fig. 4. Pond 109 water depths from 1969 to 2012.

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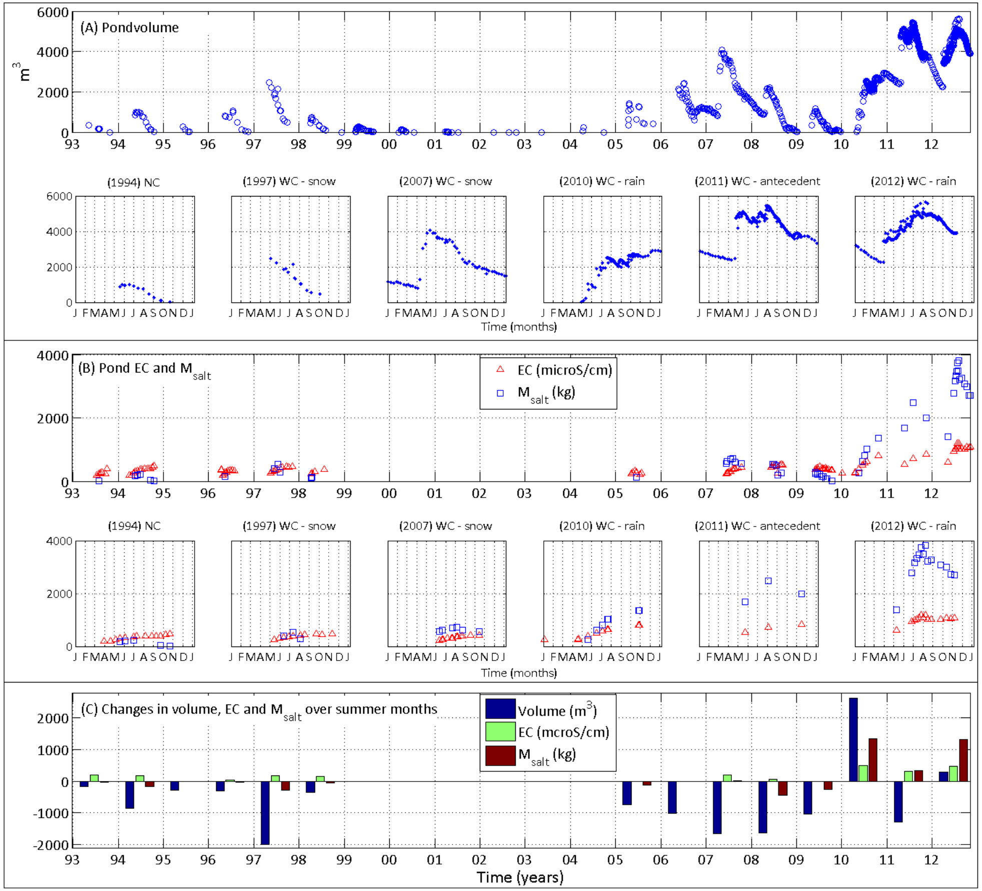

Fig. 5. (A) Pond 109 volume; **(B)** measured EC and calculated M_{salt}; and **(C)** summer-time changes in volume, EC and M_{salt} of the pond. Separated (lower) figures in **(A)** and **(B)** are zoom in for selected years with various climatic conditions.

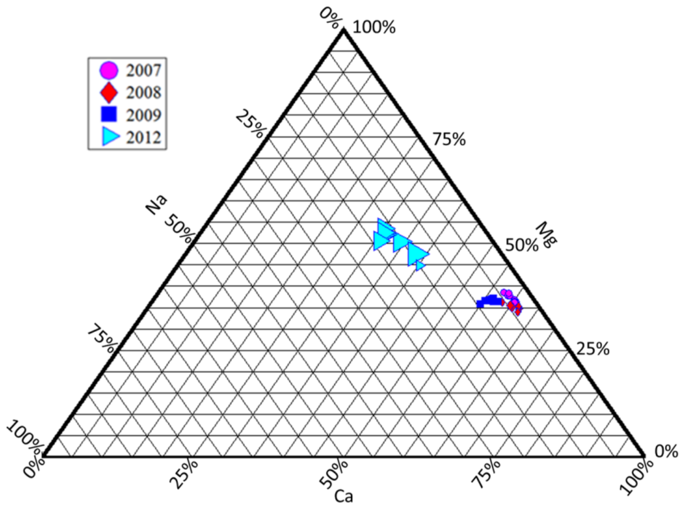

Fig. 6. Major cations molar fraction in pond 109 for the years of 2007–2009 and 2012. Marker size is proportional to the pond EC.

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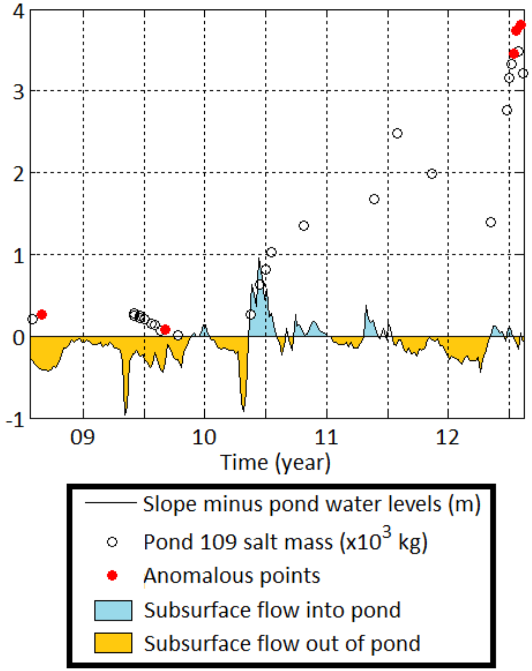

Fig. 7. Water level differences between upland and pond 109 and M_{salt}. Red symbols indicate on the anomalous measuring points, where a disparity was observed between the direction of the subsurface water flows and the changes in M_{salt}.

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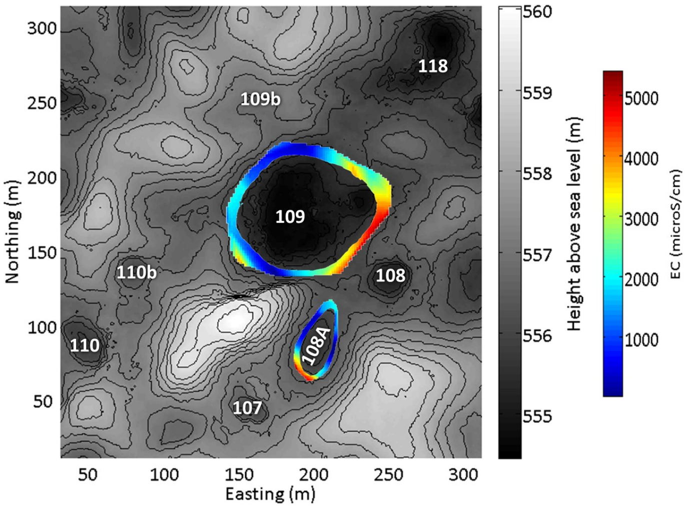

Fig. 8. Salinity around ponds 109 and 108A.

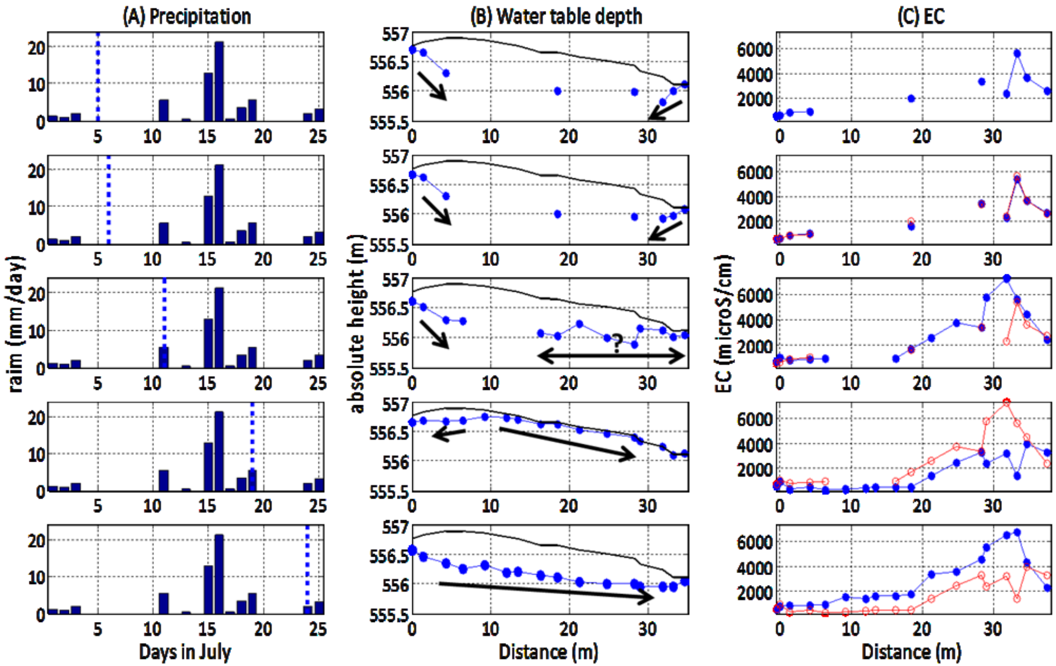

Fig. 9. Measured precipitation **(A)**; groundwater table depth **(B)**; and EC **(C)** along the 107–108A transect. The dashed lines in **(A)** indicate on the day at which measurements were taken for **(A)**, **(B)**, and **(C)** in the same row. In **(B)** the black line is ground surface and the blue curve is the location of the water table. Arrows indicate on flows directions. In **(C)** blue lines indicate on the measured EC along the transect. Red curves indicate on measured EC at previous measurement, to emphasize the change of EC with time.