



Supplement of

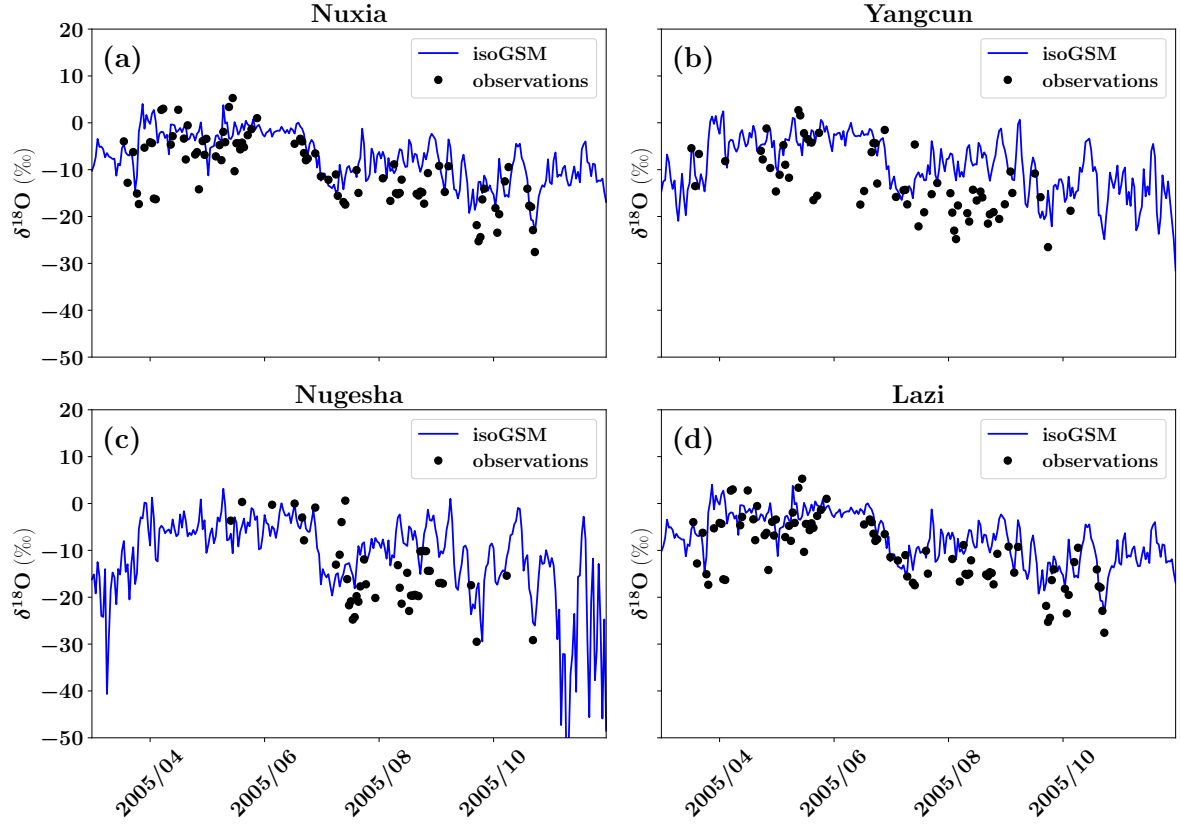
Reducing hydrological uncertainty in large mountainous basins: the role of isotope, snow cover, and glacier dynamics in capturing streamflow seasonality

Diego Avesani et al.

Correspondence to: Yi Nan (ny1209@qq.com)

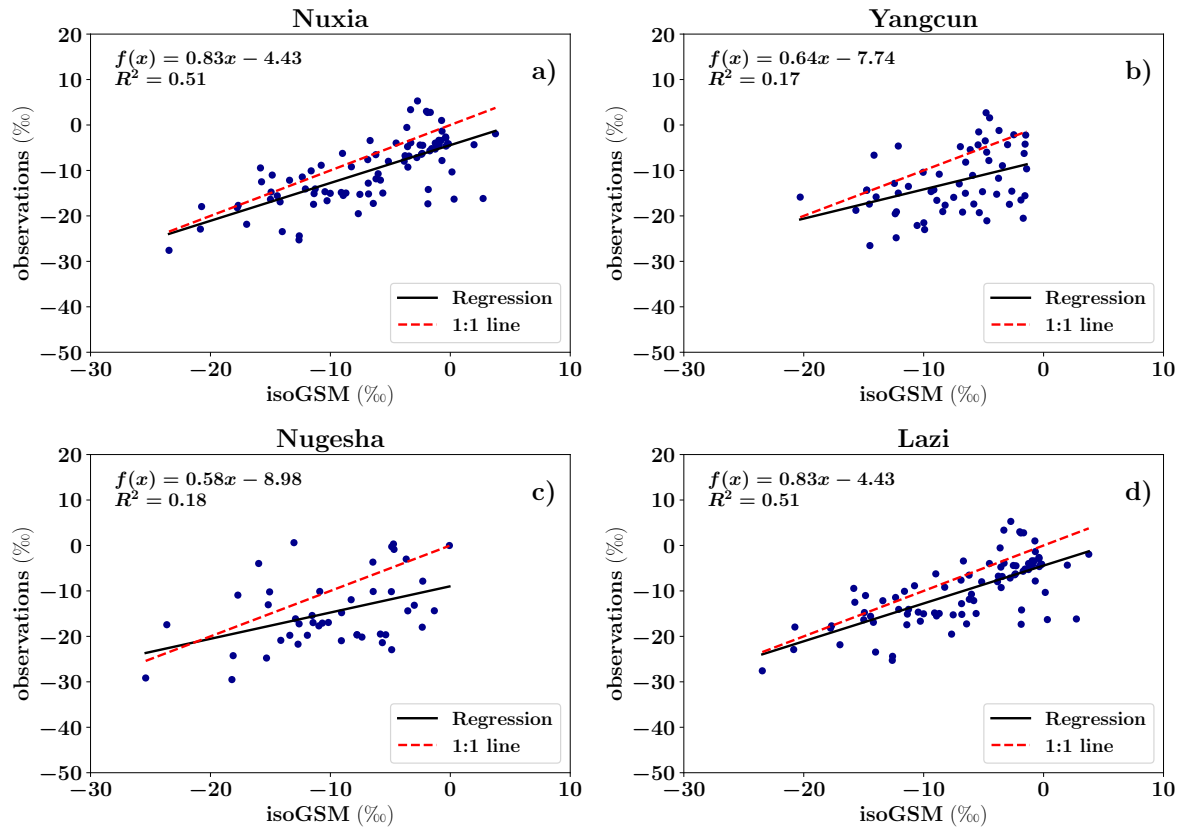
The copyright of individual parts of the supplement might differ from the article licence.

Figure S1: Precipitation $\delta^{18}\text{O}$ at four stations



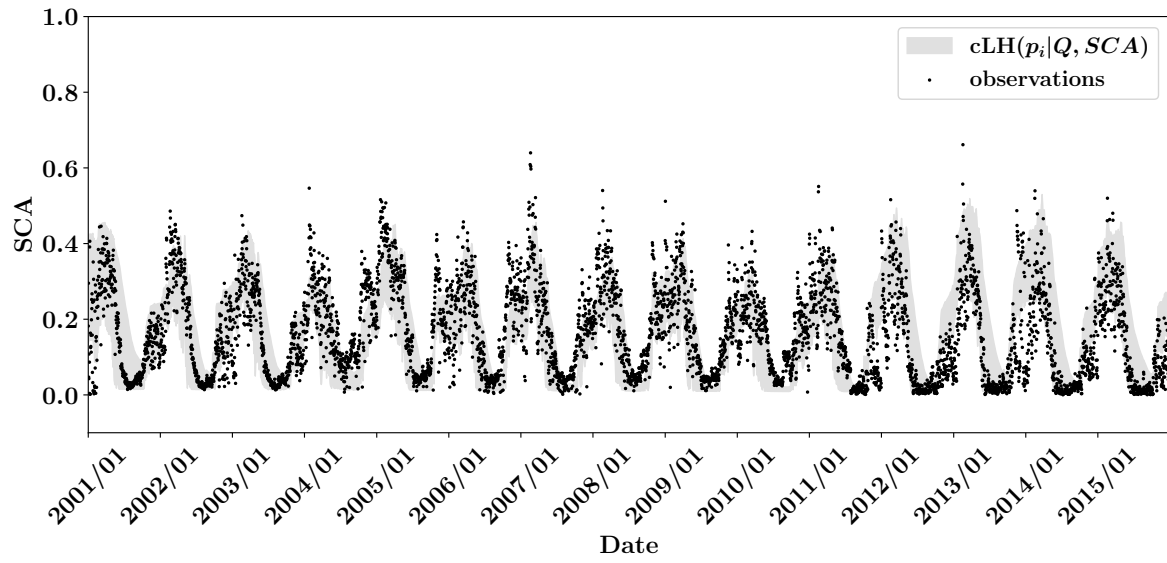
Temporal variations in the precipitation $\delta^{18}\text{O}$ (‰) derived from observation and isoGSM data at four stations.

Figure S2: isoGSM and measured isotope data at four stations



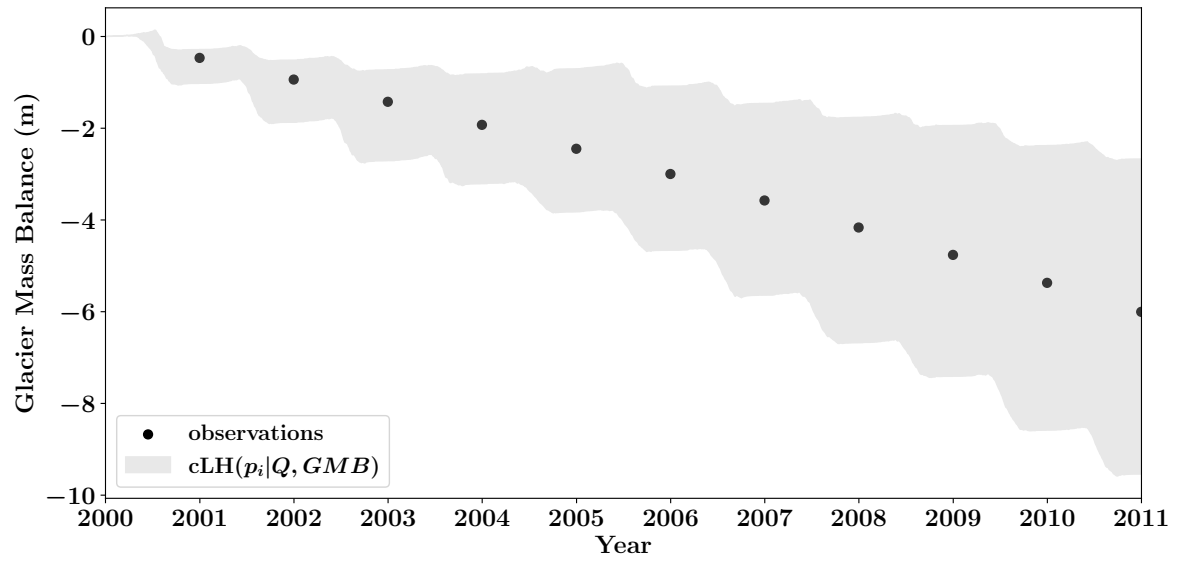
The scatter diagrams between the isoGSM and measured isotope data at four stations.

Figure S3: Observed Snow Cover Area and posterior model ensemble



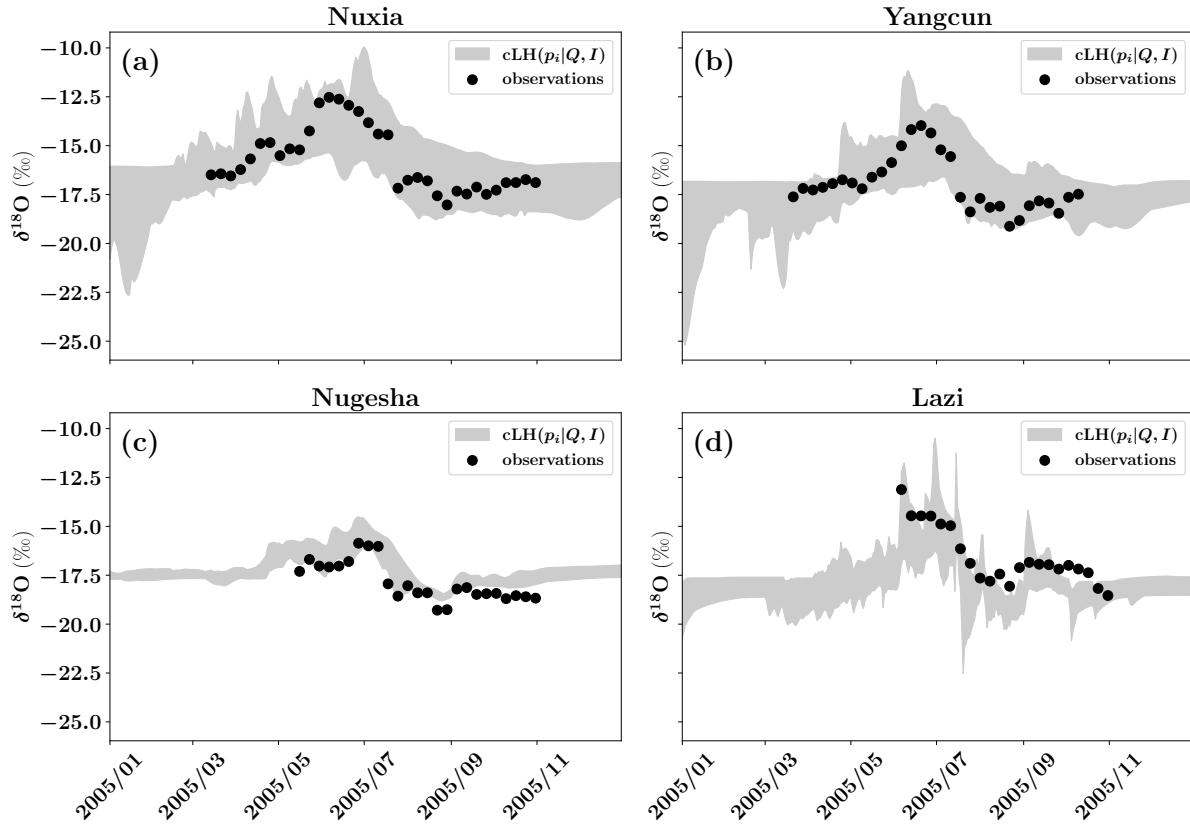
Observed snow cover area and posterior model ensemble $cLH(p_i|Q, SCA)$, based on joint conditioning with streamflow and snow cover area information.

Figure S4: Observed glacier mass balance and posterior model ensemble



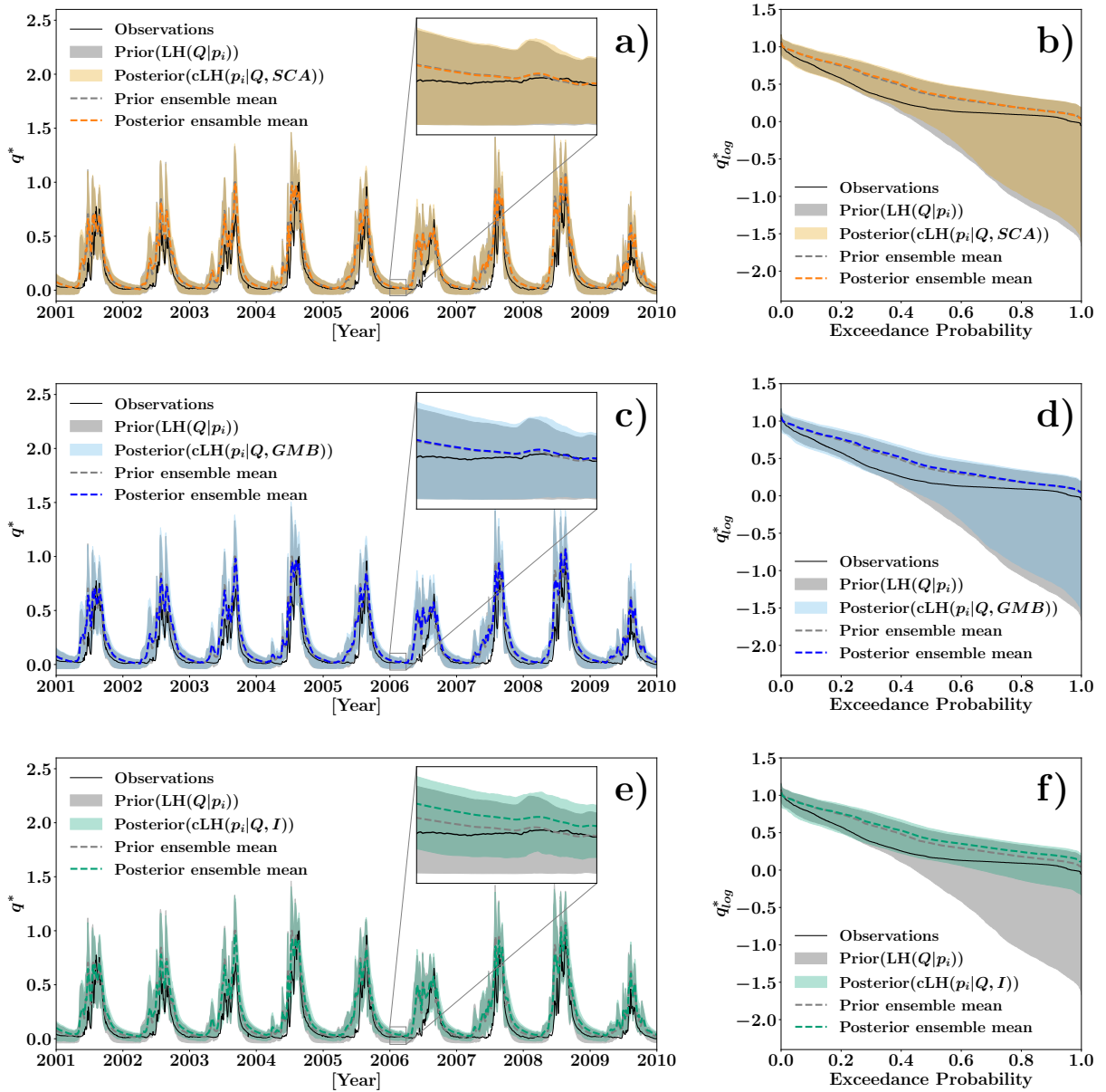
Observed glacier mass balance and posterior model ensemble $cLH(p_i|Q, GMB)$, based on joint conditioning with streamflow and mass balance information.

Figure S5: Observed $\delta^{18}O$ and the model posterior ensemble



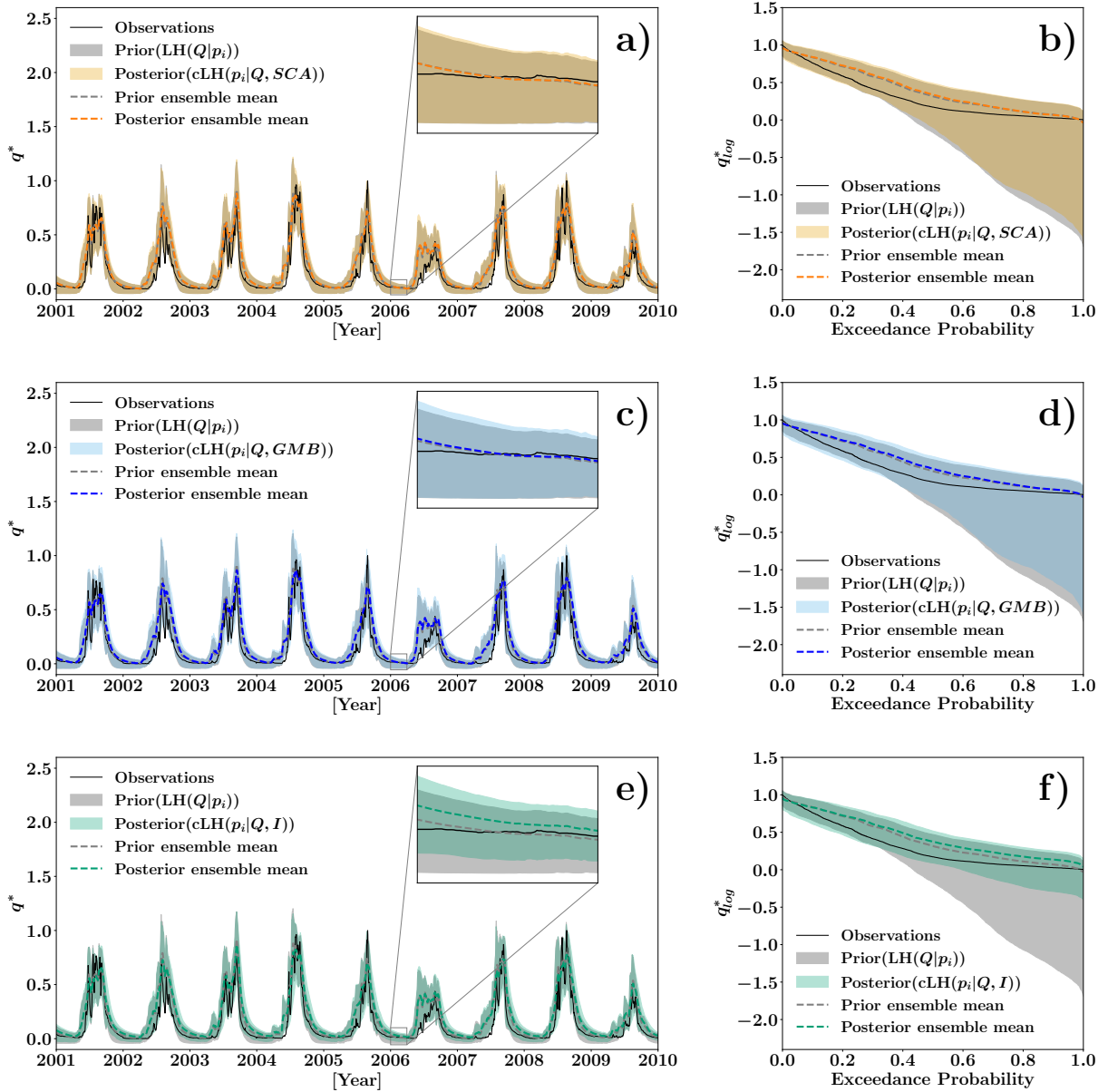
Comparison between observed $\delta^{18}O$ and the model posterior ensemble $cLH(p_i|Q, I)$, conditioned on streamflow and isotope information.

Figure S6: Streamflow simulation uncertainty range at Nugesha station



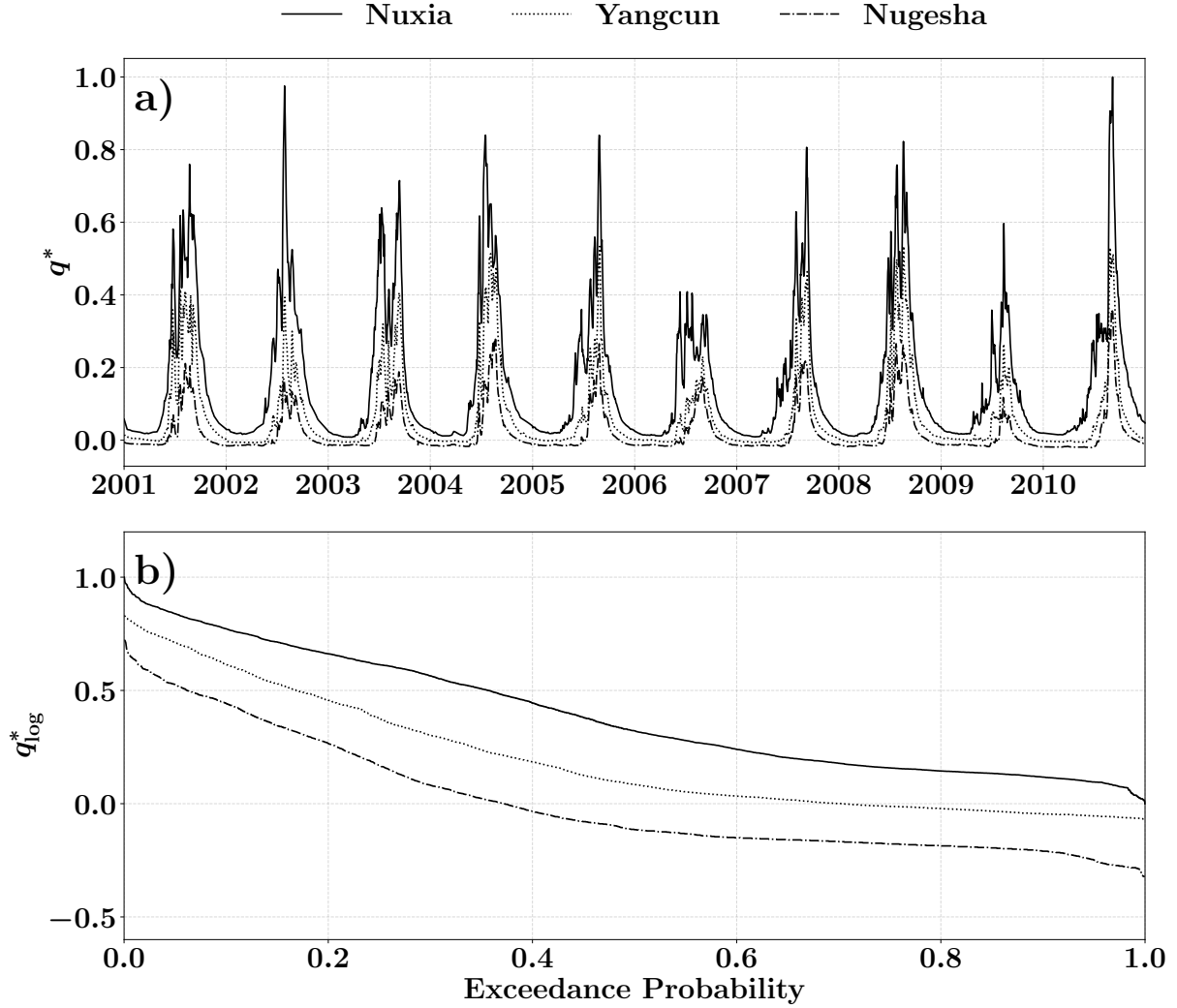
The 5–95% percentile prior, conditioned solely on streamflow, and the posterior predictive uncertainty ranges for streamflow are calculated under different conditions: snow cover area (SCA), glacier mass balance (GMB), and isotopes (I). Panels (a), (c), and (e) show daily streamflow time series for the period 2010–2015, with insets highlighting low-flow dynamics; panels (b), (d), and (f) show flow duration curves for the full period 2001–2015. Streamflow data are presented in dimensionless form due to dissemination restrictions imposed by the data provider. A linear normalization, $q^* = (Q - Q_{\min}) / (Q_{\max} - Q_{\min})$, is applied to the time series (panels a, c, e) to enable relative comparison of streamflow magnitude over time. In contrast, a logarithmic normalization, $q_{\log}^* = (\log Q - \log Q_{\min}) / (\log Q_{\max} - \log Q_{\min})$, is used for the flow duration curves (panels b, d, f), as is standard in FDC representation, to facilitate comparison across the full spectrum of discharges, including low-flow conditions. In both cases, $Q_{\min}$ and $Q_{\max}$ refer to the minimum and maximum observed discharges at the Nugesha station.

Figure S7: Streamflow simulation uncertainty range at Yangcun station



The 5–95% percentile prior, conditioned solely on streamflow, and the posterior predictive uncertainty ranges for streamflow are calculated under different conditions: snow cover area (SCA), glacier mass balance (GMB), and isotopes (I). Panels (a), (c), and (e) show daily streamflow time series for the period 2010–2015, with insets highlighting low-flow dynamics; panels (b), (d), and (f) show flow duration curves for the full period 2001–2015. Streamflow data are presented in dimensionless form due to dissemination restrictions imposed by the data provider. A linear normalization, $q^* = (Q - Q_{\min}) / (Q_{\max} - Q_{\min})$, is applied to the time series (panels a, c, e) to enable relative comparison of streamflow magnitude over time. In contrast, a logarithmic normalization, $q_{\log}^* = (\log Q - \log Q_{\min}) / (\log Q_{\max} - \log Q_{\min})$, is used for the flow duration curves (panels b, d, f), as is standard in FDC representation, to facilitate comparison across the full spectrum of discharges, including low-flow conditions. In both cases, $Q_{\min}$ and $Q_{\max}$ refer to the minimum and maximum observed discharges at the Yangcun station.

Figure S8: Normalized streamflow at the Nuxia, Yangcun, and Nugesha stations



(a) Time series of normalized streamflow between 2001 and 2010 at the Nuxia, Yangcun, and Nugesha stations. Discharge values are represented as

$$q^* = \frac{Q - Q_{\min}}{Q_{\max} - Q_{\min}}$$

where $Q_{\min}$ and $Q_{\max}$ refer to the minimum and maximum observed discharges at the calibration station (Nuxia). This normalization enables a relative comparison of streamflow magnitude across the three sites on a common scale.

(b) Flow duration curves (FDCs) for the same stations. Discharge values are shown as normalized logarithmic flow, calculated as

$$q_{\log}^* = \frac{\log(Q) - \log(Q_{\min})}{\log(Q_{\max}) - \log(Q_{\min})}$$

As is standard in FDC representation, the logarithmic scale allows for a more effective comparison across the full range of discharges, particularly improving the readability of low-flow variability.