



*Supplement of*

## **Canopy structure, topography, and weather are equally important drivers of small-scale snow cover dynamics in sub-alpine forests**

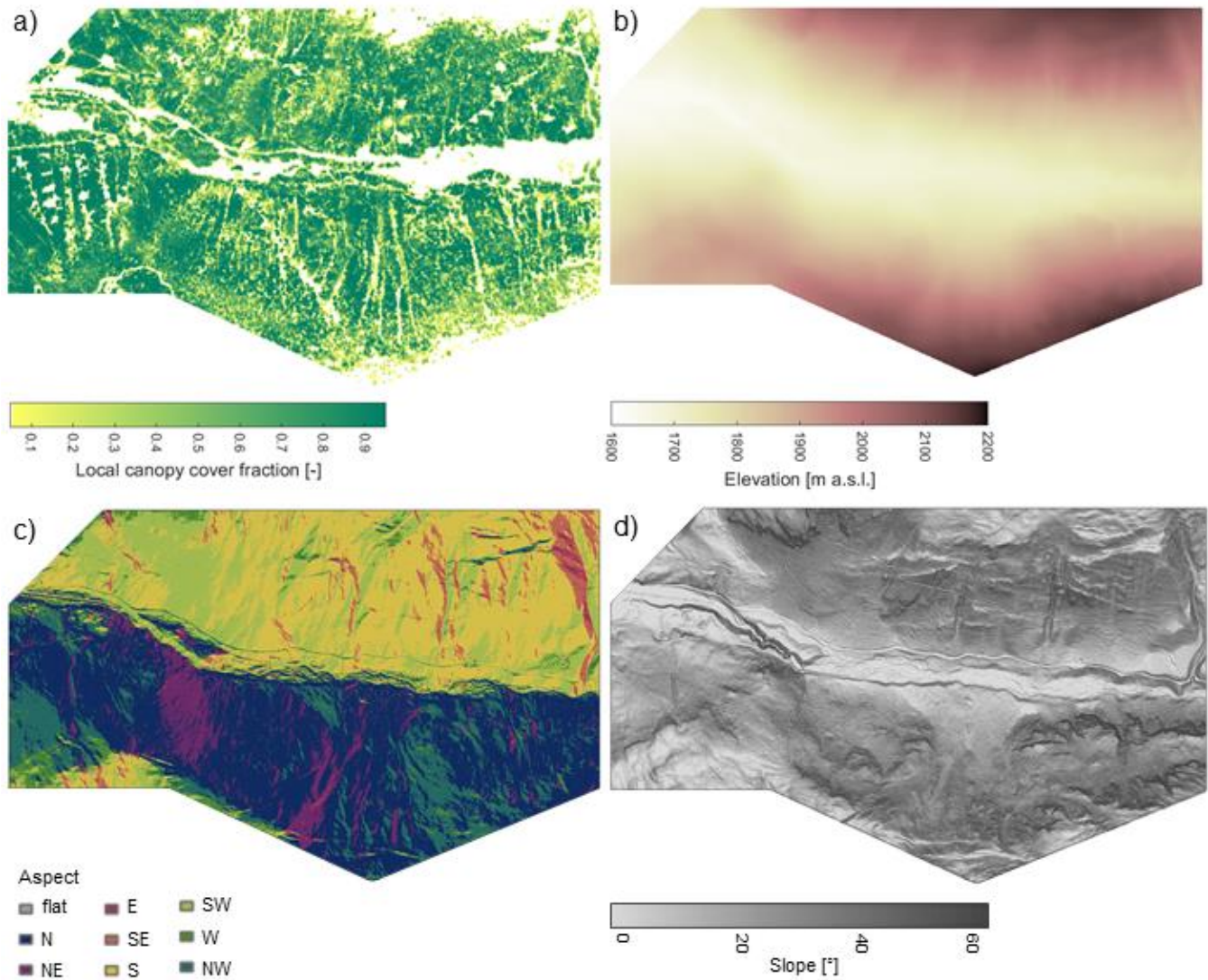
**Giulia Mazzotti et al.**

*Correspondence to:* Giulia Mazzotti ([giulia.mazzotti@slf.ch](mailto:giulia.mazzotti@slf.ch))

The copyright of individual parts of the supplement might differ from the article licence.

## S1: Physiographic maps over the full model domain

To complement the site description in the main article (Section 2.2 and Figure 1), Figure S1.1 reports additional maps of physiographic variables across the model domain, including local canopy cover fraction, elevation, aspect and slope.

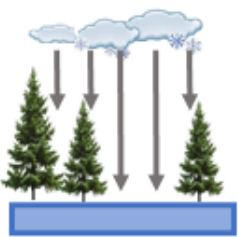

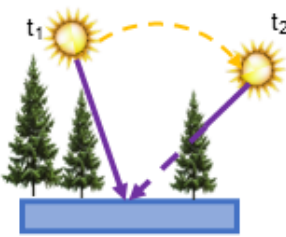
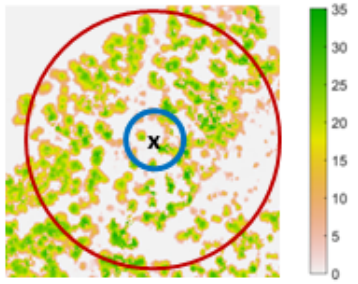

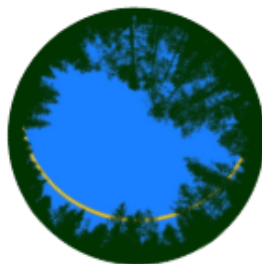


25 Figure S1.1: Local canopy cover fraction (a), elevation (b), aspect classes (c), and slope (d) across the full model domain.

## S2: Canopy structure representation in FSM2

30 The diversified representation of canopy structure of FSM2 with process-specific canopy metrics constitutes the model's principal asset. With this strategy, FSM2 considers that different canopy-mediated processes are regulated by different and potentially uncorrelated canopy characteristics. A point located in a canopy gap, as example, experiences the interception regime of a very sparse canopy, but may be affected by frequent shading, which implies conditions typical of a dense canopy. These independent and contrasting canopy controls on each of the various processes cannot be reproduced if all processes are parametrized with the same, bulk canopy descriptors. Bulk canopy metrics may thus be suited for intermediate-35 or coarse-resolution model applications targeting the estimation of spatially averaged fluxes and states, but hyper-resolution approaches that strive to reproduce meter-scale patterns require explicit representation of a wide range of detailed canopy features. The canopy metrics used within FSM2 comprise horizontal, vertical, local, and stand-scale canopy characteristics at each modelled location. Consequently, the canopy's structural diversity is captured even though canopy is represented with a so-called one-layer model. Figure S2.1 illustrates this concept, summarizing the different canopy perspectives, the40 corresponding metrics, approaches, and data sources used to derive them, and the processes employing them. The schematic complements the mathematical formulation of the process parametrizations provided in the respective model description papers, Essery (2015) and Mazzotti et al. (2020a,b).

Transmission of direct shortwave radiation through the canopy involves a particularly complex interaction with small-scale canopy elements. Because transmissivity is dictated by the presence of canopy elements in the path of the solar beam, it45 depends on the exact geometric arrangement of the point of interest, the canopy, and the sun, and is thus highly variable in space and time. By accepting transmissivity time series as model input, FSM2 forgoes simplification of the process representation that comes with any parametrization, and instead leverages the benefits of an (external) radiative transfer model that resolves the process explicitly. So far, FSM2.0.3 has been used in tandem with the radiative transfer model HPEval (Jonas et al., 2020), which comprises a traditional approach to compute transmissivity time series at a point based on50 a high-resolution hemispherical image. The radiative transfer model depicted in the schematic in Figure S2.1 illustrates the HPEval approach. In this study, transmissivity time series are computed following the HPEval methodology based on synthetic hemispherical images derived from LiDAR datasets instead of real ones. The workflow to create these synthetic images, detailed in Webster et al. (2020), integrates canopy segmentation algorithms to identify individual trees in a canopy height model and point cloud enhancing algorithms to densify the point cloud, aimed at mimicking opaque stems and55 branches. The resulting images feature more realistic and detailed tree shapes than images obtained with unprocessed LiDAR data (see Section S3).

<p><i>perspective</i></p>	<p>vertically projected</p> 	<p>hemispherical</p> 	<p>canopy in path of solar beam</p> 
<p><i>data source / computation</i></p>	<p>canopy height model</p> 	<p>hemispherical image</p> 	<p>external radiative transfer model</p> 
<p><i>metrics</i></p>	<ul style="list-style-type: none"> <li>• Canopy cover fraction (CC)</li> <li>• Canopy height (H)</li> <li>• LAI = f(CC, H, species)</li> <li>• Computed at both <b>local</b> and <b>stand</b> scale (<b>5m</b> and <b>50m</b> radius around point)</li> </ul>	<ul style="list-style-type: none"> <li>• Sky view fraction (non-directional)</li> </ul>	<ul style="list-style-type: none"> <li>• Transmissivity for direct shortwave radiation for varying solar positions</li> </ul>
<p><i>processes</i></p>	<ul style="list-style-type: none"> <li>• Interception, preferential deposition, canopy snow depletion (local metrics)</li> <li>• Wind attenuation (stand scale metrics)</li> </ul>	<ul style="list-style-type: none"> <li>• Transmission of diffuse shortwave radiation</li> <li>• Longwave radiation enhancement</li> </ul>	<ul style="list-style-type: none"> <li>• Transmission of direct shortwave radiation</li> </ul>

60 **Figure S2.1: Overview of the concepts underlying the canopy structure representation in FSM2: different perspectives, relevant data sources and computation approaches, resulting metrics, and modelled processes that use the respective metrics.**

### S3: Summary of FSM2 validation at the process level

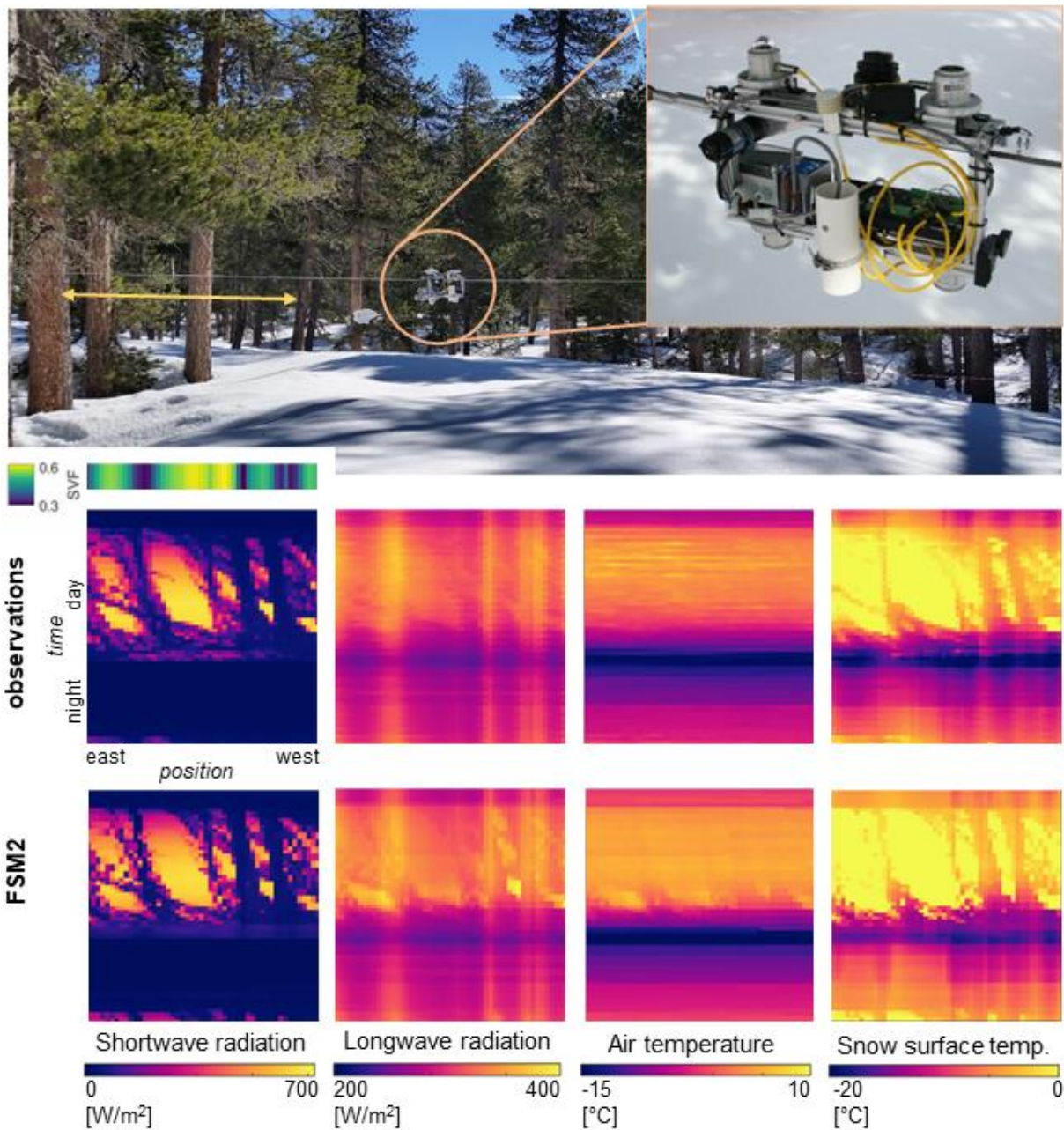
This study relies on FSM2's capability to accurately simulate the spatio-temporal patterns of individual fluxes and states that constitute forest snow processes, which was extensively tested in a previous study by Mazzotti et al. (2020a). Figure S3.1 reports an example of their validation efforts to illustrate the underlying methodology. Data were acquired with multiple meteorological sensors mounted on a motorized cable car platform (Figure S3.1). Over the course of a winter, the system was deployed along several forest transects of approx. 50 m length each in both Switzerland and Finland. Surveys covered variable canopy structures and forest species, weather conditions, as well as sub-alpine and boreal climates. Observations from the cable car setup were complemented by data from two additional handheld and stationary setups, yielding a dataset with an unprecedented variety of micro-meteorological and snow measurements. FSM2 simulations at high spatial and temporal resolution were directly compared to the acquired datasets. The example in Figure S3.1 features measurements from a full 24-hour cycle along a cable car transect and corresponding FSM2 simulations at 1.5m spacing and 10min temporal resolution. To our knowledge, FSM2 is the only forest snow model to have ever undergone such a rigorous validation.

FSM2's ability to replicate individual fluxes and states and their spatial and temporal variability was quantitatively assessed in terms of transect averages and standard deviations of all measured variables across all available datasets (Figure S3.2). Root means square errors (RMSE) of  $23 \text{ Wm}^{-2}$  and  $4 \text{ Wm}^{-2}$  resulted for average shortwave and longwave irradiances, respectively, while their standard deviations featured RMSEs of  $21 \text{ Wm}^{-2}$  and  $2 \text{ Wm}^{-2}$ . RMSEs of mean and standard deviation of canopy air space temperatures were  $1.6^\circ\text{C}$  and  $0.4^\circ\text{C}$ , for snow surface temperatures they amounted to  $0.6^\circ\text{C}$  and  $0.4^\circ\text{C}$ . Performance metrics all showed substantial improvements relative to the standard land surface model canopy implementation that was used as benchmark in Mazzotti et al. (2020a).

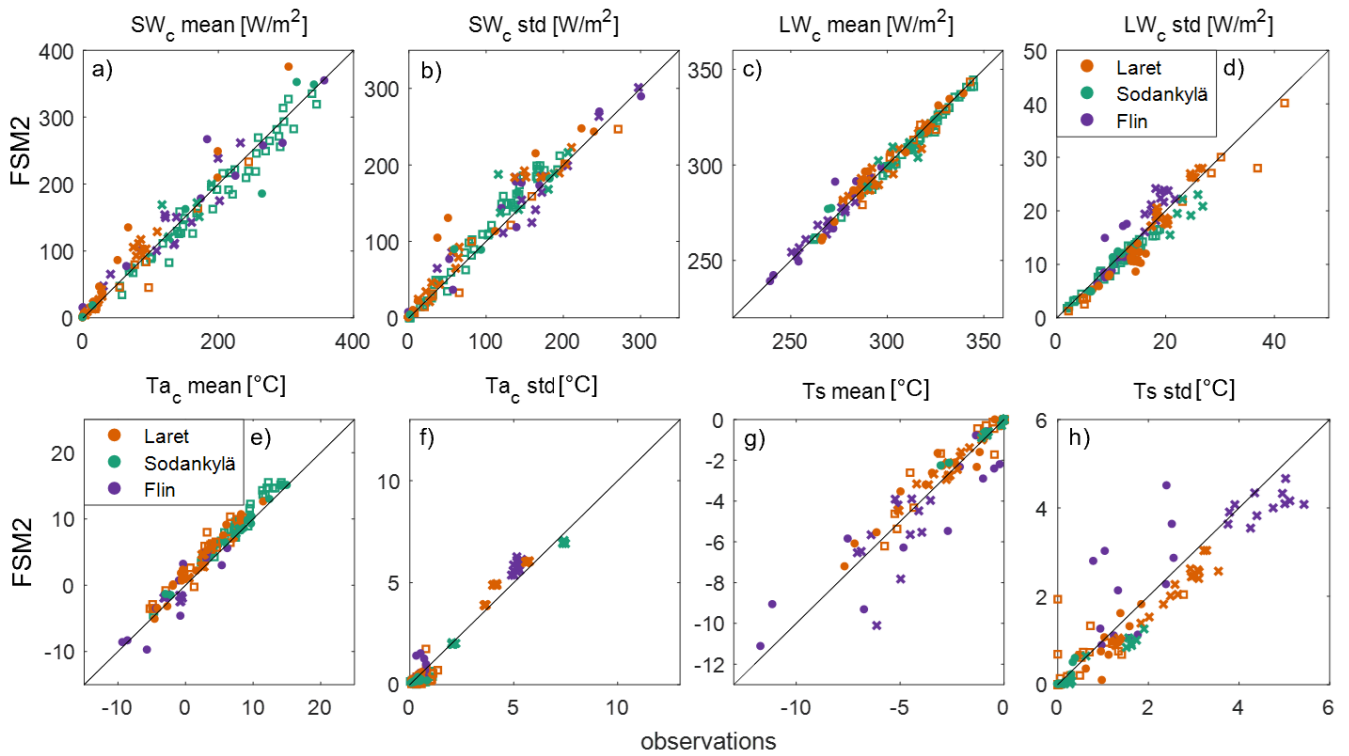
The accuracy of the shortwave radiative transfer calculations in FSM2 is dictated by the accuracy of the external radiative transfer model. Webster et al. (2020) showed that shortwave radiation estimates using synthetic images obtained with the LiDAR enhancing methodology were similar to estimates obtained with real images (Figure S3.3, a-e), and patterns of shadows on the snow surface (f) matched those captured with UAV-imagery (g) with high level of detail.

While plausibility checks as presented in Section 2.3 (main article) are important to ensure model applicability for a specific use case, we believe that quantitative validations at the level of individual fluxes and states such as in Mazzotti et al. (2020a) and summarized here are more conclusive than assessments based on snow extent and depth data alone.





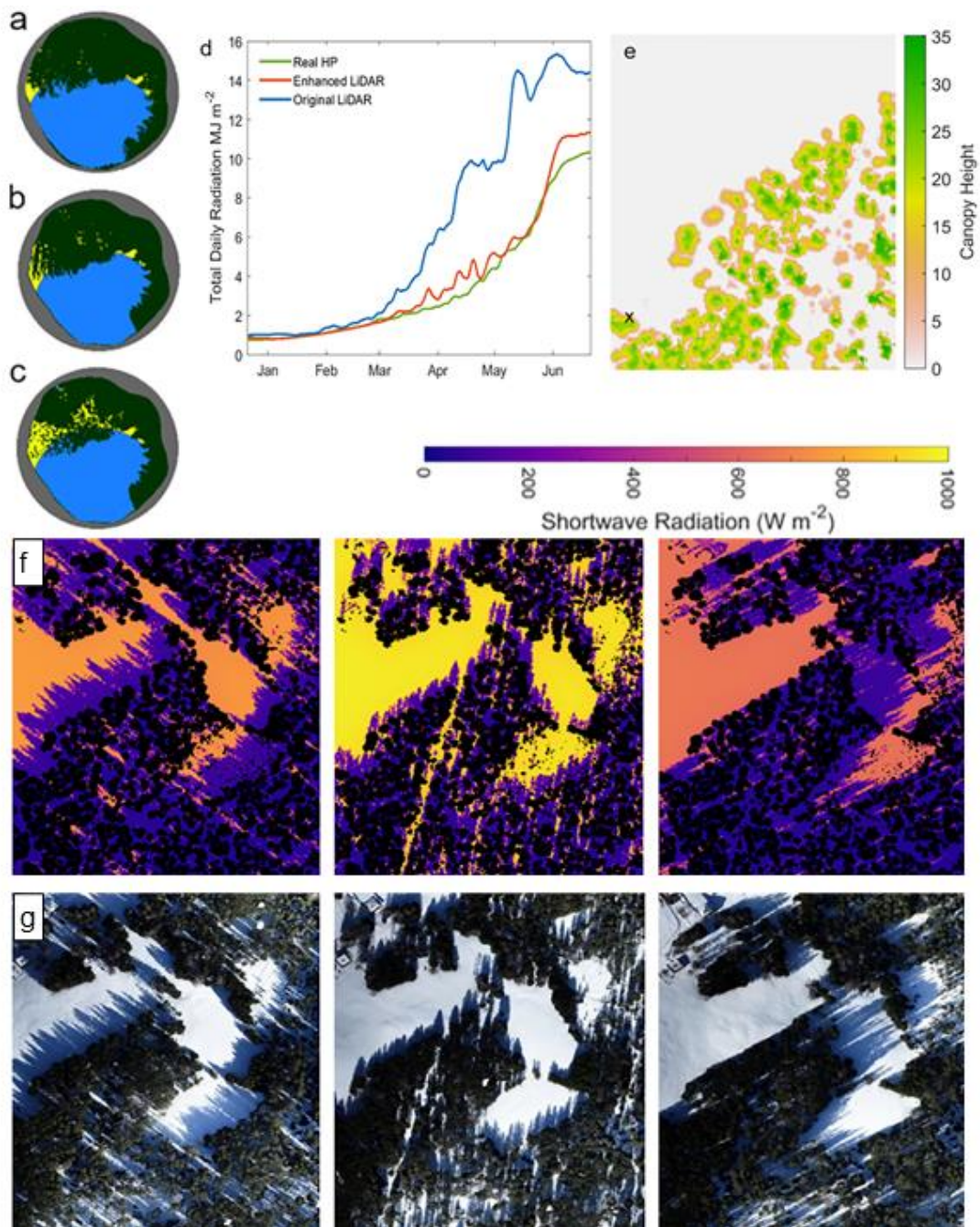
90 **Figure S3.1:** The multi-sensor platform in operation along a 50-m cable (yellow arrow) in a pine stand near Davos, Switzerland below variable canopy cover (sky-view fraction, SVF, obtained from hemispherical photographs acquired along the transect). Resulting model validation: observed spatiotemporal variability of measured variables during a 24-h cycle (19. to 20.03.2019) and corresponding FSM2 simulations. The direct comparison between model results and observations is achieved by extracting model output at each point based on the time stamp of the respective observation. Adapted from Mazzotti et al. (2020a), Fig. 1 and 3.



95

**Figure S3.2: Observed vs. modelled averages and standard deviations of subcanopy shortwave (a,b) and longwave (c,d) irradiances, canopy air space (e,f) and snow surface (g,h) temperatures. Data points correspond to temporal averages at point locations (crosses), spatial averages along the cable car transects (dots), and spatial averages over plots surveyed with a handheld setup (squares) at given points in time. Individual colors represent each study area, where Laret and Flin are located in the Eastern Swiss Alps and Sodankylä in sub-arctic Finland. Adapted from Mazzotti et al. (2020a), Figures 5 and 7.**

100



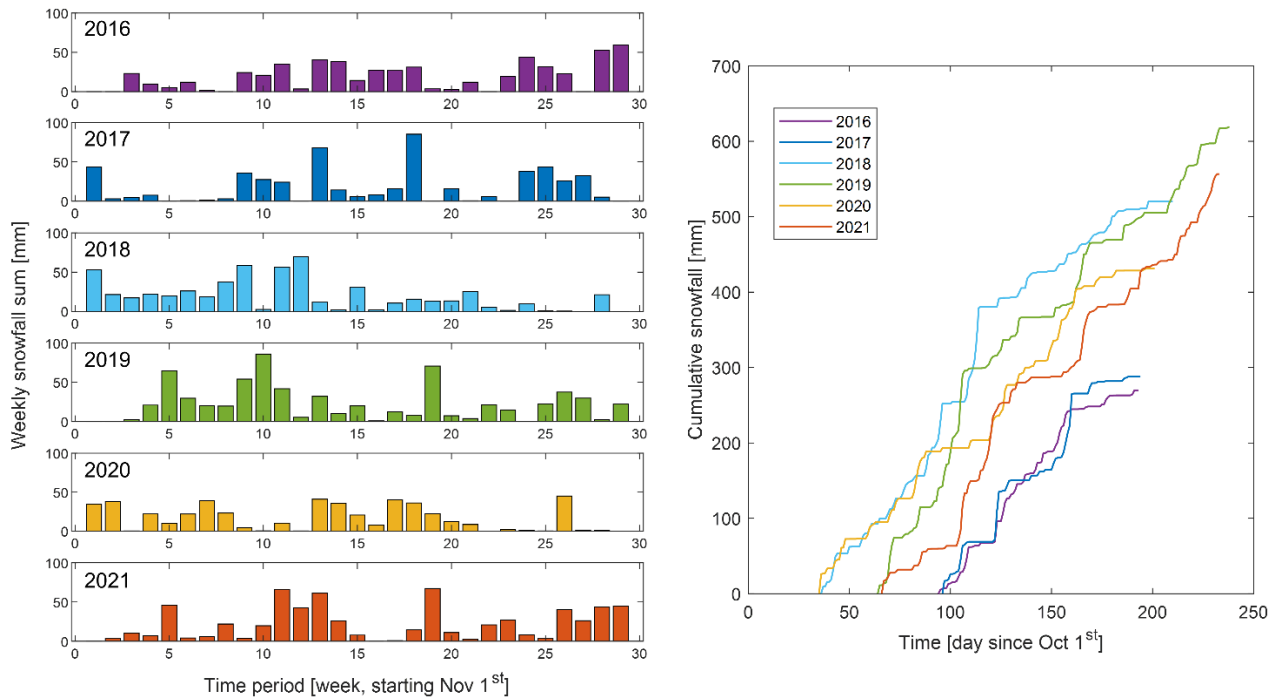
105 Figure S3.3: Changing solar position across a 6-month solar cycle in yellow within (a) a real hemispherical photograph, (b) synthetic hemispherical images from the enhanced LiDAR and (c) original LiDAR. Modelled total daily radiation across the same period (d) at location X in the canopy height model (e). Comparison of modelled radiation (f) and snow surface shadowing as photographed from a drone (g) at 09:30 (left), 12:30 (center) and 15:30 (right) on 28.03.2019. Adapted from Webster et al. (2020), Fig. 8 and 10.



#### S4: Meteorological conditions during the simulation period

110 The following figures provide an overview of meteorological conditions during the six simulated water years to give context to the observed differences in snow cover dynamics outlined in the main article. We focus on the three variables that, as elaborated on in the main article, have the largest potential to cause differences between years: snowfall (affects total accumulation), shortwave radiation, and temperature (affect the timing and strength of melt, as well as whether it is mainly driven by short- or longwave radiation contributions). The left panel of Figure S4.1 reports weekly snowfall sums between the beginning of November and June, which represent the earliest and latest yearly median of the start of snow cover period and snow disappearance date. Some marked differences in how the precipitation is distributed throughout the snow season can be observed. The right panel of Figure S4.1 shows cumulative snowfall between the median start of snow cover and snow disappearance dates of each specific year. This visualization evidences the substantial differences in timing and length of snow cover period and total snowfall between the years, with WY2016 (lowest accumulation) amounting to only half of total snowfall of WY 2019 (highest accumulation).

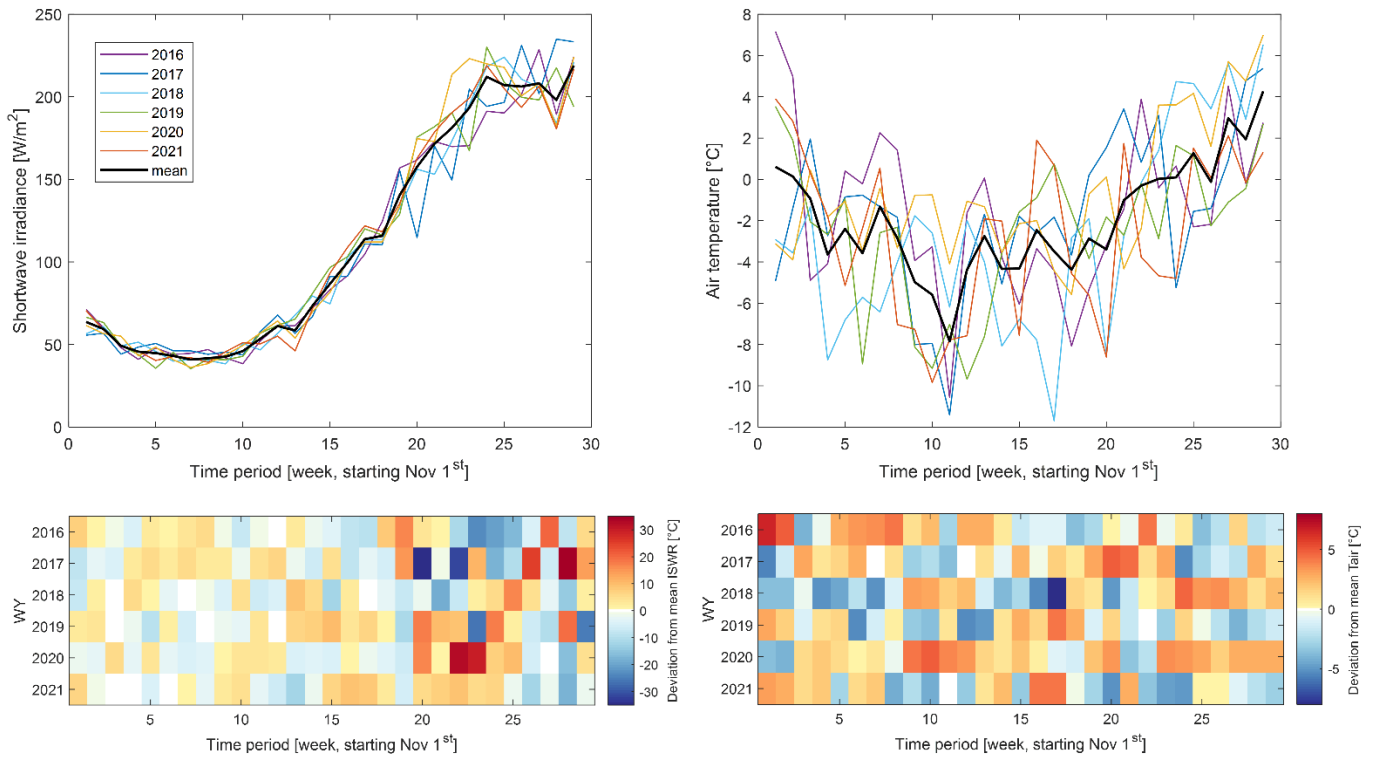
115



120

**Figure S4.1: Weekly snowfall during the six modelled winters between November and June (left) and cumulative snowfall over the median snow cover duration period of the respective year (right).**

125 Figure S4.2 contrasts shortwave radiation (left) and air temperatures (right) during the six simulated winters, with the upper row showing weekly values and the lower row the weekly deviation from the six-year average. While no year is clearly above or below average for any of the two variables for the entire season, some marked deviations from the averages can be detected in these plots, for instance the relatively warm period in the middle of the 2020 winter.



130 **Figure S4.2: Incoming shortwave radiation (left) and air temperature (top right) of the six simulated winters, and deviation from mean (lower row) for the same variables.**

## S5: Maps of snow cover and energy balance descriptors for WY2019 over the full model domain

While some figures in the main article include maps of a subdomain of the model domain to illustrate patterns of snow cover descriptors and energy fluxes, here we add the corresponding maps for the full model domain for interested readers. Figure S5.1 reports peak SWE, ablation rate, day of peak SWE and snow disappearance day for WY 2019, corresponding to Figure 5 of the main article.

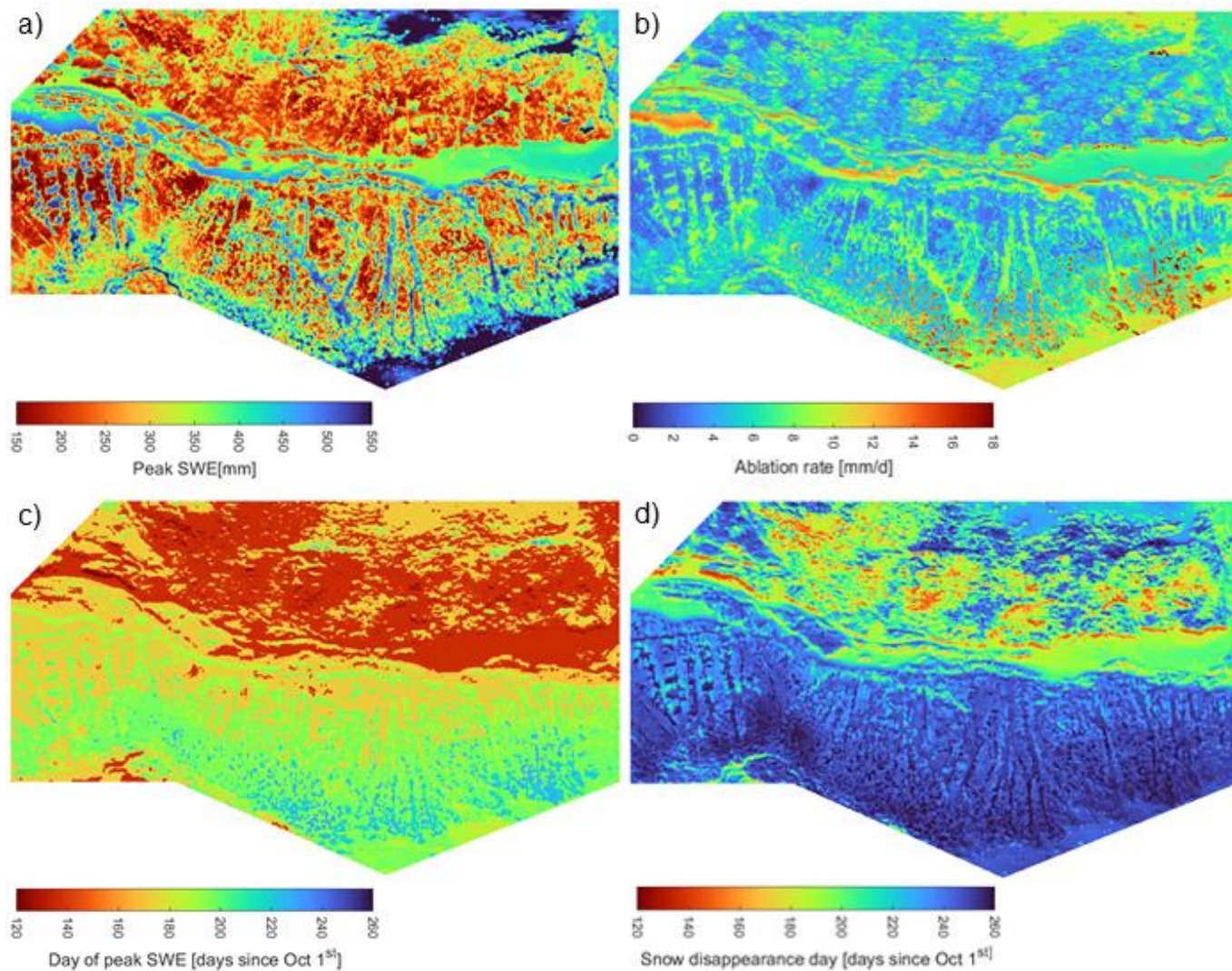
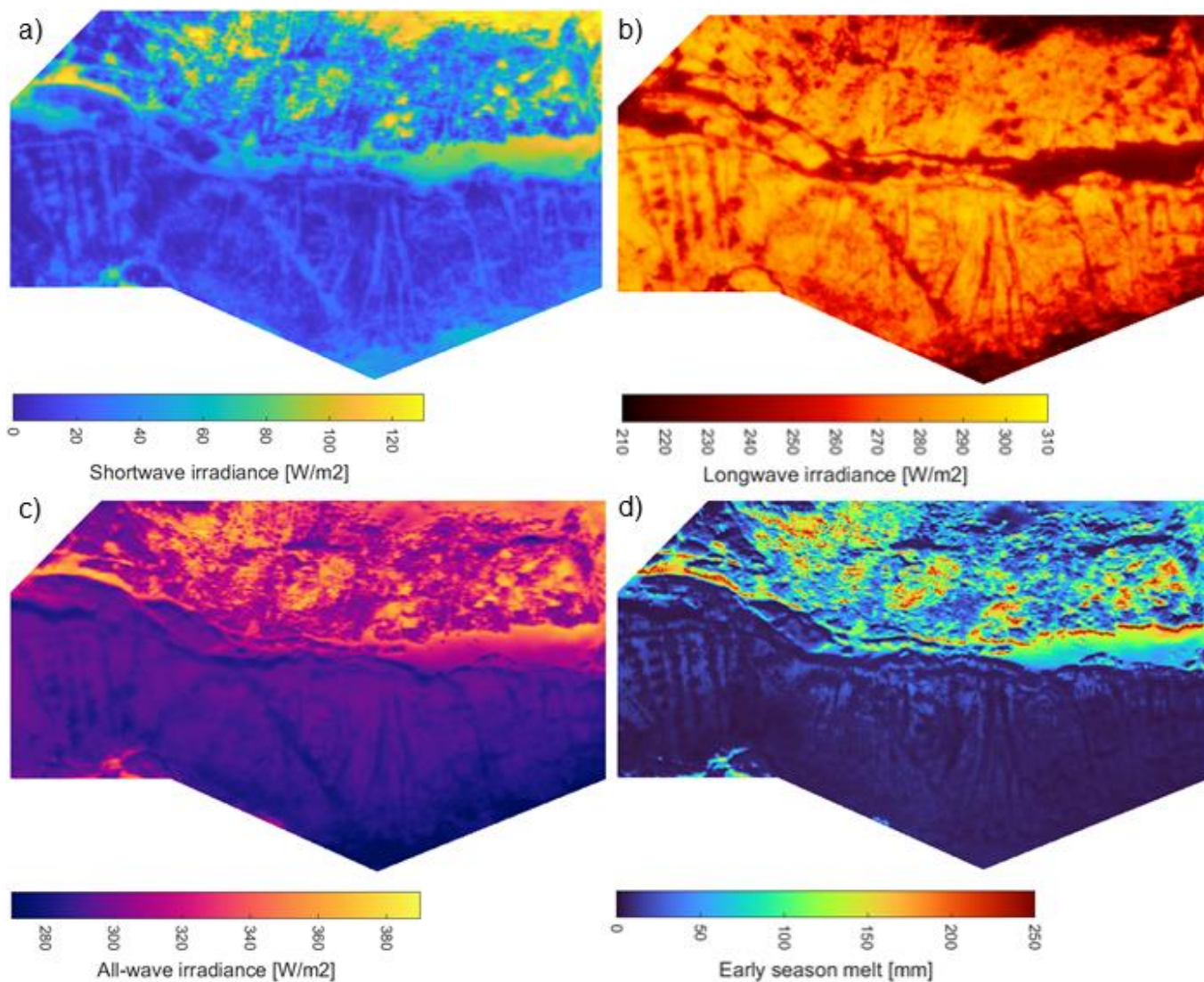


Figure S5.1: Peak SWE (a), ablation rate (b), day of peak SWE (c) and snow disappearance day (d) in WY 2019.

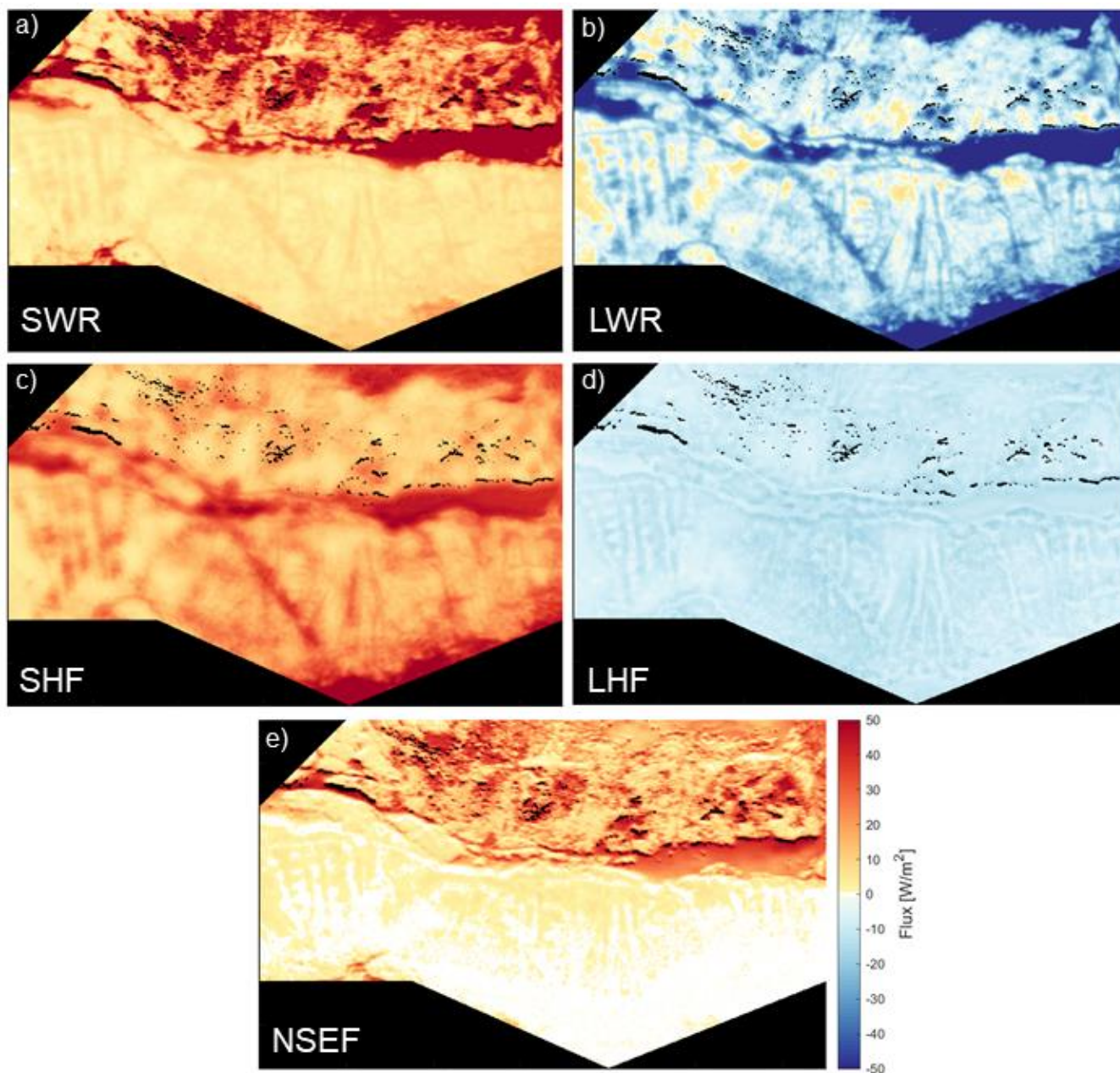


140 Figure S5.2 shows average incoming irradiances and early melt between mid-January and end of February 2019, analogous to Figure 7 of the main article. Note the addition of incoming longwave radiation which further underpins the statement that all-wave irradiance patterns are largely determined by patterns of incoming shortwave radiation.



145 **Figure S5.2: Average incoming short (a) and longwave (b) radiation, all-wave radiation (c) and snow melt (d) between mid-January and end of February 2019.**

Figure S5.3 includes average patterns of the individual surface energy balance components (net short- and longwave radiation, sensible and latent heat) as well as net surface energy flux computed as the sum of the four. Maps shown are for the second half of February 2019, analogous to the upper part of Figure 8 in the main article.



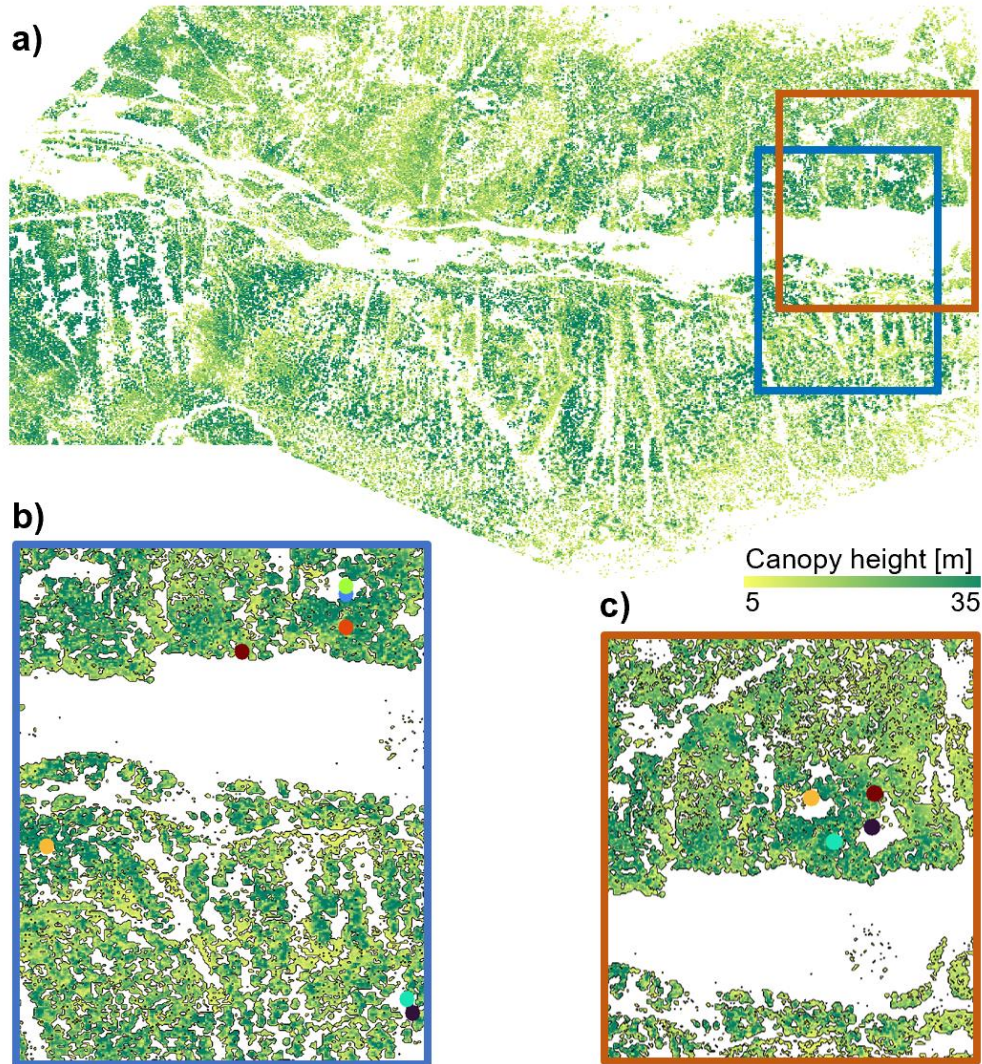
150

**Figure S5.3: Individual surface energy components (net short- and longwave radiation, SWR (a) and LWR (b), as well as sensible and latent heat fluxes, SHF (c) and LHF (d)) and net surface energy flux (NSEF, e) computed as sum of the four, for the second half of February 2019.**



## S6: Location of example points within the model domain

155 The location of example points used for the analysis of snow evolution and energy balance partitioning pathways presented in Section 3.3 and 3.5 are shown in Figure S6.1. In Section 3.3, points were chosen to cover the entire range of canopy structures and topographic locations (Figure S6.1b, corresponding to points shown in Figure 6 of the main article). In Section 3.5, we focused our analysis on points located at the south-exposed slope and sparse canopy or small gaps to explore differences between years (Figure S6.1c, corresponding to points shown in Figure 10 of the main article).

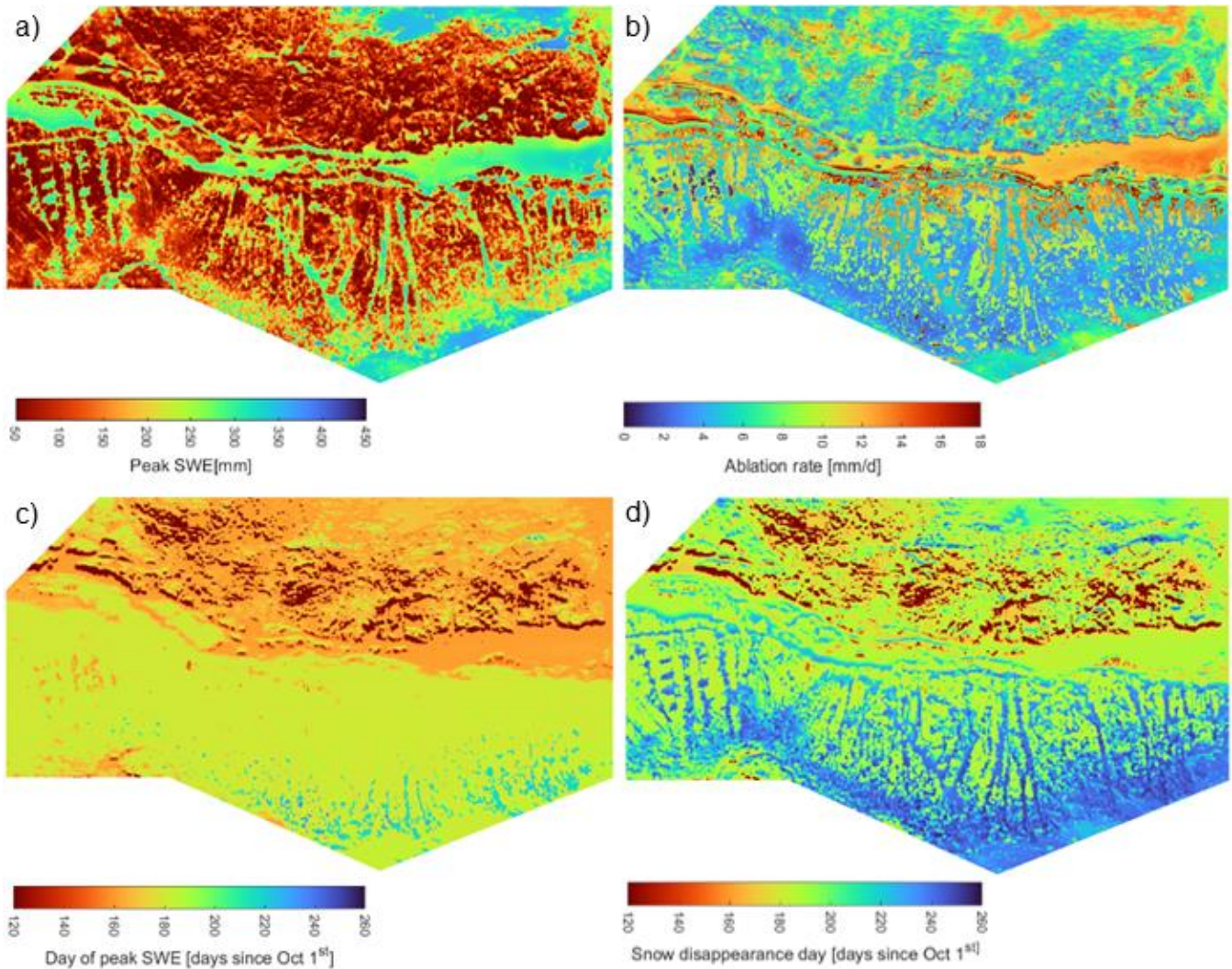


160

**Figure S6.1: Locations of example points. Overview (a), points presented in Section 3.3 / Figure 6 of the main article (b), and points used in Section 3.5 / Figure 10 of the main article. Colours marking individual points match the colour code of the figures in the main article.**

## S7: Maps of snow cover descriptors for the other simulated WYs

165 In Section S5, we presented full-domain maps of snow cover descriptors for WY 2019, which is used as main example throughout the article. To complement Figure S5.1, Figures S7.1 to S7.5 report snow cover descriptors for all other simulated WYs. These figures provide an additional visual impression of the between-year consistencies and differences of these snow metrics. Moreover, it should be noted that correlation coefficients presented in Section 3.2 of the main article rely on these data.



170

Figure S7.1: Peak SWE (a), ablation rate (b), day of peak SWE (c) and snow disappearance day (d) in WY 2016.



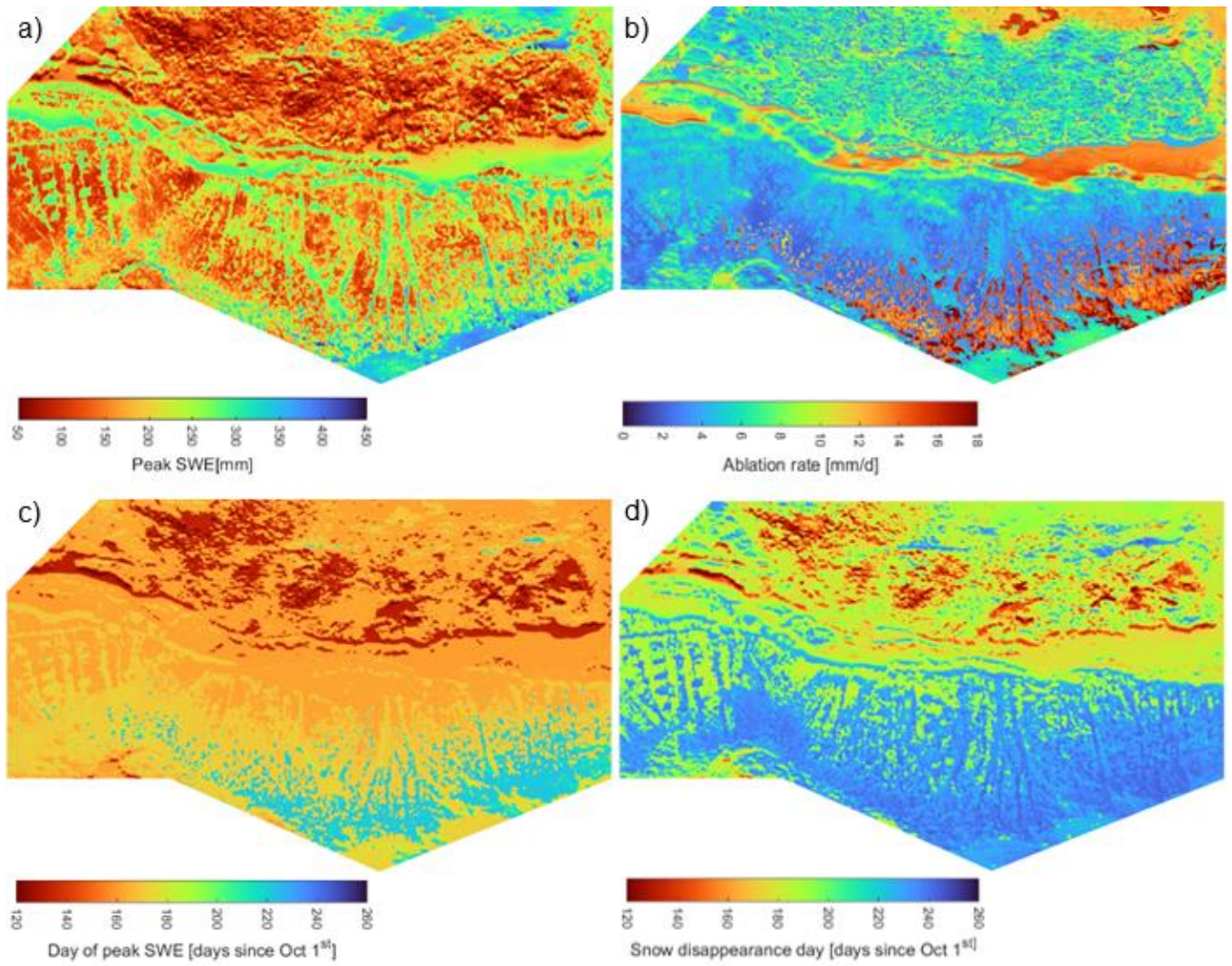
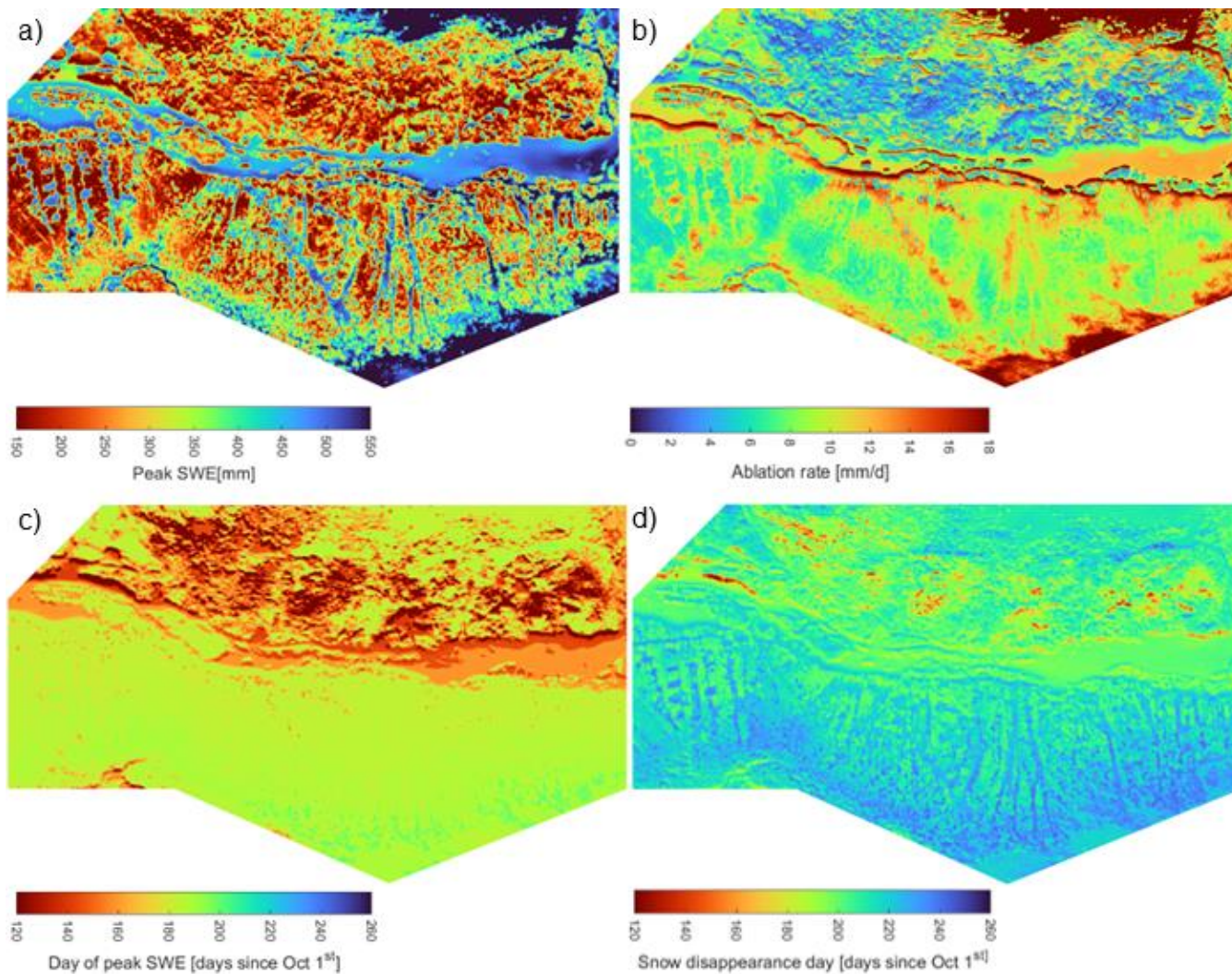


Figure S7.2: Peak SWE (a), ablation rate (b), day of peak SWE (c) and snow disappearance day (d) in WY 2017.



175 **Figure S7.3: Peak SWE (a), ablation rate (b), day of peak SWE (c) and snow disappearance day (d) in WY 2018.**



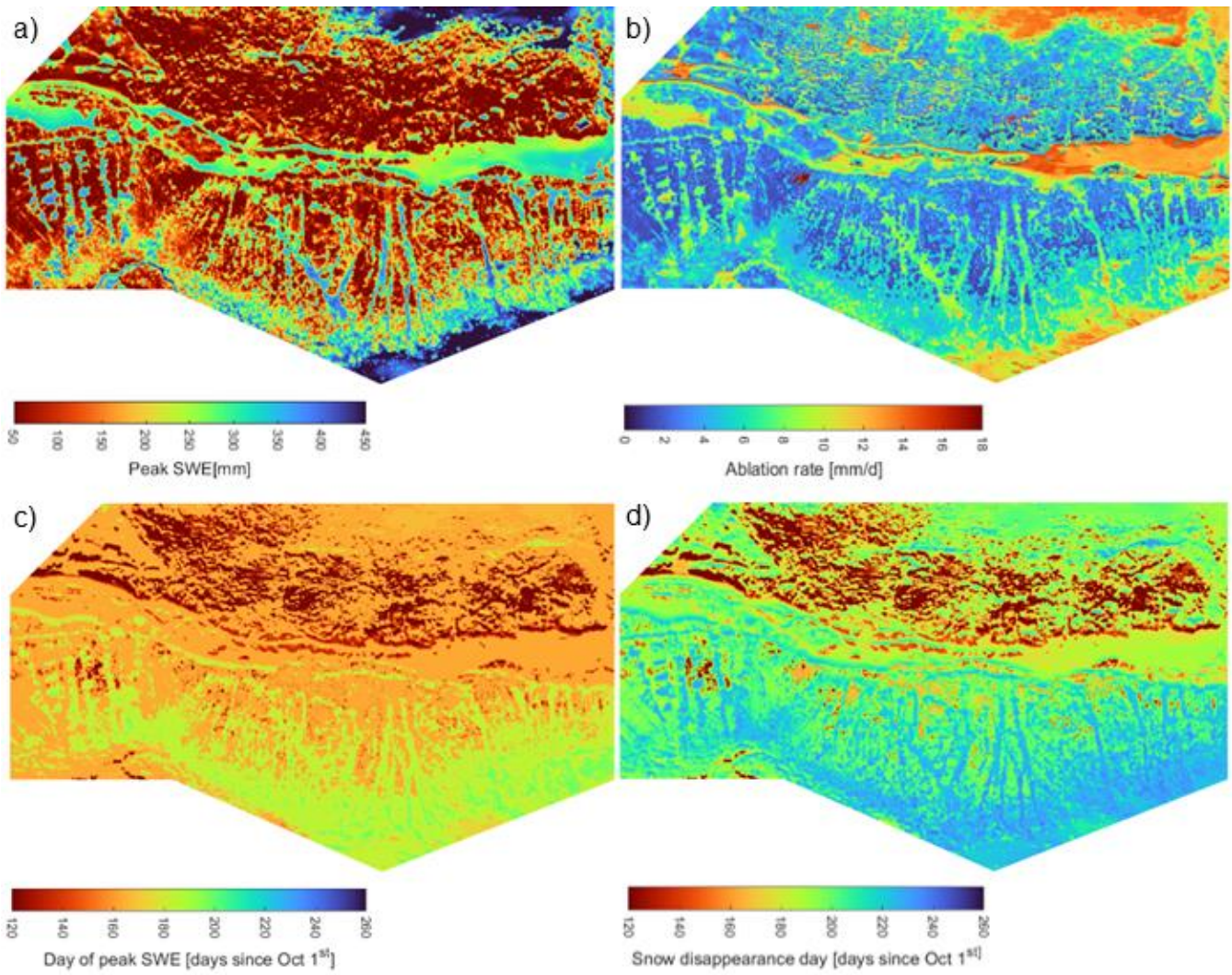


Figure S7.4: Peak SWE (a), ablation rate (b), day of peak SWE (c) and snow disappearance day (d) in WY 2020.



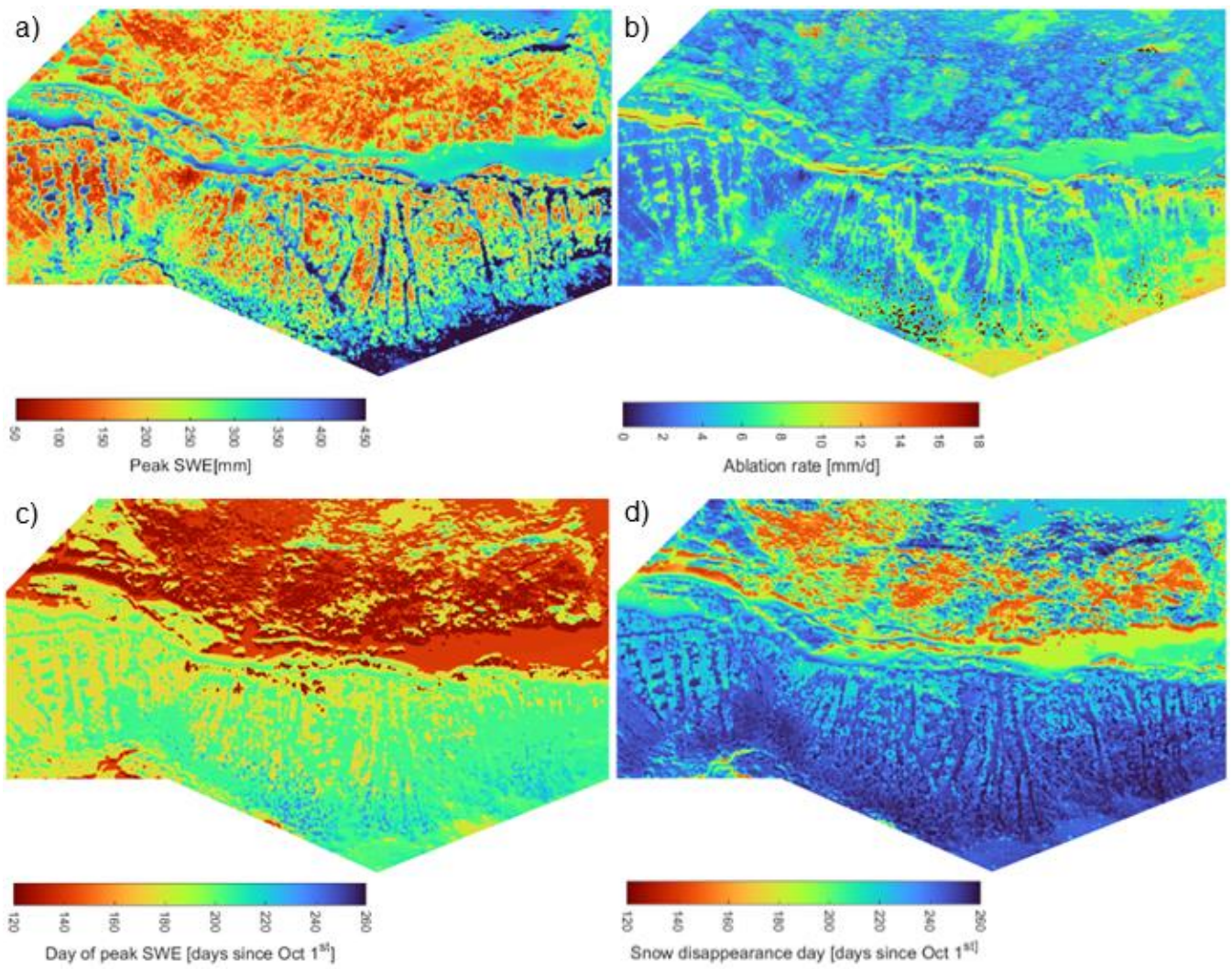


Figure S7.5: Peak SWE (a), ablation rate (b), day of peak SWE (c) and snow disappearance day (d) in WY 2021.

## References:

- Essery, R.: A factorial snowpack model (FSM 1.0), *Geoscientific Model Development*, 8, 3867–3876, <https://doi.org/10.5194/gmd-8-3867-2015>, 2015.
- Jonas, T., Webster, C., Mazzotti, G., and Malle, J.: HPEval: A canopy shortwave radiation transmission model using high-  
185 resolution hemispherical images, *Agricultural and Forest Meteorology*, 284, 107903, <https://doi.org/10.1016/j.agrformet.2020.107903>, 2020.
- Mazzotti, G., Essery, R., Webster, C., Malle, J., and Jonas, T.: Process-Level Evaluation of a Hyper-Resolution Forest Snow Model Using Distributed Multisensor Observations, *Water Resources Research*, 56, e2020WR027572, <https://doi.org/10.1029/2020WR027572>, 2020a.
- 190 Mazzotti, G., Essery, R., Moeser, C. D., and Jonas, T.: Resolving Small-Scale Forest Snow Patterns Using an Energy Balance Snow Model With a One-Layer Canopy, *Water Resources Research*, 56, e2019WR026129, <https://doi.org/10.1029/2019WR026129>, 2020b.
- Webster, C., Mazzotti, G., Essery, R., and Jonas, T.: Enhancing airborne LiDAR data for improved forest structure representation in shortwave transmission models, *Remote Sensing of Environment*, 249, 112017,   
195 <https://doi.org/10.1016/j.rse.2020.112017>, 2020.