



Supplement of

The ¹⁸O ecohydrology of a grassland ecosystem – predictions and observations

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Methods S1: Precipitation sampling

The sampling apparatus at Grünschwaige pasture paddock no. 8 consisted of a plastic funnel (94 mm in diameter) installed at 1 m above the soil surface and connected to a 1 L plastic collector bottle installed 1 m below ground by means of a silicone hose. A table tennis ball was placed inside the funnel to minimize evaporation losses of collected waters. The bottle was sampled and emptied regularly following rain events, i.e., at intervals of 3 to 61 days (average 14 d; n = 81).

Methods S2: MuSICA parameterisation

Parameter values for the 'standard' MuSICA runs were derived from data collected at the site (as explained in the main text and below) or taken from the literature (Table S1).

Soil

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- 10 Soil structural properties (proportion of quartz and organic matter) as well as hydraulic characteristics (water retention and hydraulic conductivity) were determined on soil core samples taken at a depth of 3 to 8 cm. Soil water retention and hydraulic conductivity properties were obtained by simultaneously measuring water tension and weight changes resulting from evaporative water loss on 250 mL soil core samples, according to the simplified evaporation method (Schindler, 1980; Peters et al., 2015) using a HYPROP apparatus (UMS, Munich, Germany). Drainage and hydraulic conductivity curves were
- 15 calculated from water tension and evaporative water loss data using the HYPROP software (Pertassek et al., 2015). Parameters of the van Genuchten-Mualem soil water retention model (van Genuchten, 1980; Mualem, 1976) and of the Brooks-Corey hydraulic conductivity model (Brooks and Corey, 1964), both used in MuSICA, were obtained by leastsquares fit to the drainage and conductivity curves (Fig. S5). Gravitational water flow was assumed at the bottom of the mineral topsoil, at 37 cm belowground. Estimated parameter values for the soil surface resistance to water vapour transport,
- 20 soil surface aerodynamic resistance and soil optical properties (albedo and emissivity) were taken from the literature (Table S1).

In the Moldrup et al. (2003) model for the water vapour effective diffusivity, the pore-size distribution parameter *b* was derived from the water retention curve parameters *m* and *n* as b = 1/m/n. In this work, we explore the consequences of using either the Penman or Moldrup soil diffusivity formulations on the prediction of the δ^{18} O signals of soil, xylem and leaf

25 waters (see sensitivity analysis in main text).

Soil respiration (the total of root and heterotrophic soil respiration) was predicted using a Q_{10} relationship with soil surface temperature, with basal soil respiration rate at 25°C (R_{25}) and the Q_{10} value obtained from open-top chamber respiration measurements performed at the site in September 2006, May 2007 and September 2007 (Gamnitzer et al., 2009; Ostler et al., unpublished).

Notes S1: Diel measurements and modelling of ¹⁸O enrichment of pasture vegetation

Leaf and soil water, and atmospheric moisture were sampled at intervals between 4 am on 4 August to 7 am on 5 August in 2005, in the centre of pasture paddock no. 8 at Grünschwaige. The procedures followed the same protocols as given in the Materials and Methods of the main text, except that soil water was collected at depths of 2, 12 and 22 cm. Leaf samples were

5 collected every hour with three replicates, soil samples every six hours with five replicates at 2 cm, three replicates at 12 cm and one replicate at 22 cm depth.

Fig. S7 shows the diurnal cycle of observed ¹⁸O enrichment of leaf water above soil water ($\Delta^{18}O_{leaf} = \delta^{18}O_{leaf} - \delta^{18}O_{soil 7}$), and of the $\Delta^{18}O$ predicted in the standard simulation (two-pool model with $\varphi = 0.39$) and in the Péclet simulation with L = 167 mm. Observed $\Delta^{18}O_{leaf}$ reached its minimum (1.9‰) at around 5 am (UTC) – pre-dawn – and then increased

10 progressively for about 5 h to approach a near-maximum value at around 10 am. The observed $\Delta^{18}O_{\text{leaf}}$ remained within 90% of maximum for about 5 h and then decreased continuously for about 12 h to reach another minimum (at ~0.1‰) at 2 to 5 am the next morning.

These $\Delta^{18}O_{leaf}$ data were used to fine-tune the parameters controlling leaf water enrichment in MuSICA, mainly leaf water content, the Péclet effective length and stomatal conductance parameters such as nighttime and residual stomatal

15 conductance, within the known range for temperate grassland or cool-season grasses. Following these adjustments, modelled $\Delta^{18}O_{leaf}$ followed quite closely the temporal pattern of observed $\Delta^{18}O_{leaf}$ when a two-pool model was applied. In particular, the maximum of modelled $\Delta^{18}O_{leaf}$ was reached at approximately the same time as that observed. By contrast, when a Péclet model with a constant mixing length was applied in the simulation, predicted $\Delta^{18}O_{leaf}$ reached a maximum in the late afternoon and evening hours that was not present in the observed data (Fig. S7).

20 Notes S2: Testing the relevance of the Péclet effect in the pasture species *Lolium perenne* and *Dactylis glomerata* in controlled environments

Several recent studies (Roden et al., 2015; Song et al., 2015) have called into question the relevance of the Péclet effect to leaf water isotopes. Given this uncertainty, and the added complexity of including a Péclet effect in leaf water models, we tested the requirement for a Péclet effect in the pasture grasses *L. perenne* and *D. glomerata* – two of the co-dominant species in the grassland ecosystem study – with an aim to applying Occam's razor principle if appropriate (Figs. S12-13).

Lolium perenne

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Perennial ryegrass seeds (*L. perenne* L. cv. Bronsyn plus AR1 endophyte, 2 g per pot or 83 g m⁻²) were sown into 5-L pots containing 1700 g of seed-raising mix at field capacity and grown in a controlled-environment growth cabinet maintained at 20°C, 70% RH, 700 μ mol m⁻² s⁻¹ PAR during the 16-h light period, and 15°C, 70% RH during the 8-h dark period, for 17 d. The pots were then randomly allocated to either high (70%) or low (30%) relative humidity cabinets in which all other

settings were the same. All plants were clipped to 6 cm in height, and well-watered daily. Seven days after the humidity

treatments were applied, eight pots within each humidity treatment were allocated to either well-watered (field capacity) or droughted (midway between field capacity and oven-dried water content) treatments. Plants in these pots were again clipped to 6 cm in height. Water content was maintained in both treatments by daily gravimetric measurements, with water used replaced. Plants were grown for 21 days after the commencement of the water treatment and droughted pots took 2-3 days to

- 5 reach their target water content.
 - Leaf gas exchange measurements occurred between 8 and 16 days after the start of the water treatment, and leaf water sampling on day 20 of the treatment. Transpiration rate (*E*) was measured on a group of 10-20 leaves in each of 5 pots per treatment over a 24 hour period under growth conditions using a custom clear-top chamber fitted to a Li6400 (LiCor Inc., Lincoln, NE, USA) photosynthesis system (as described in Loucos et al., 2015, except that the incident light within the
- 10 growth cabinet was used rather than an external light source). Measurements were recorded every minute, averaged over 10 minutes, then a treatment average calculated to compare to leaf samples taken from randomly-assigned pots every two hours. Every 2 hours when the cabinet lights were on during a 29 hour period, three leaves (3 cm in length) were cut and immediately placed in small glass vials, then flushed with 2% CO₂ and sealed. The oxygen in leaf water was left to equilibrate with oxygen in CO₂ within the vial for 48 hours at 25°C, then the CO₂ was analysed for δ^{18} O on a tunable diode
- 15 laser absorption spectrometer (TDL, TGA100A, Campbell Scientific) as described by Song and Barbour (2016), with liquid water standards for correct isotope compositions of the leaf water relative to SMOW. The isotope composition of water vapour and irrigation water was measured on the TDL as described above. Water vapour and irrigation water was measured on the TDL as described above. Water vapour
- was collected by pumping air from each growth cabinet through a glass cold finger trap sitting in an ethanol-dry ice slurry. Air was pumped for 20 minutes for the low RH cabinet and 10-25 minutes for the high RH cabinet, and collections were made every 2 hours. The irrigation water had a δ^{18} O of -9.6‰, while the water vapour varied between -18.2 and -14.0‰ (the low RH cabinet had significantly less enriched water vapour than did the high RH; -16.0 ± 0.4‰ compared to -17.2 ± 0.3‰, P = 0.003). Irrigation water and vapour δ^{18} O were used to calculate $\Delta^{18}O_{e,ss}$ (using Eq. (2), main text) and measured leaf water enrichment, $\Delta^{18}O_{leaf}$.

The Péclet effect predicts a positive relationship between E and the proportional difference between $\Delta^{18}O_{leaf}$ and $\Delta^{18}O_{e}$, but it

25 can be seen from Figure S12 that variation in *E* explained very little variation in the proportional difference, suggesting that the Péclet effect was of limited relevance for *L. perenne*.

Dactylis glomerata

We also tested the relevance of the Péclet effect on a second, small stature grass species using the online gas exchange and equilibrated leaf water method developed by Song et al. (2015). *D. glomerata* L. plants were grown from seed in 7-L pots

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with potting mix amended with slow release fertiliser (Osmocote, Scotts Australia Pty Ltd., Sydney, NSW, Australia) in a controlled environment room set at day/night temperature of 28/20 °C, 75% air humidity in the day and night, 14 h day period and an approximate irradiance at the top of the canopy of 600 μ mol m⁻² s⁻¹. When the plants were 60 days old, 3-5 leaves were sealed in a 2 × 3 cm leaf chamber with a red-blue light source attached to a Li6400 photosynthesis system and

plumbed to a water vapour isotope analyser (L1102-i; Picarro Inc., Sunnyvale, CA, USA) for isotopologue measurement. Dry air entered the leaf chamber, so that all the water vapour measured by the analyser came from transpiration (*E*). The conditions inside the leaf chamber were manipulated to achieve a range in *E*, by altering flow rate through the chamber (between 250 and 700 μ mol s⁻¹) and CO₂ concentration (between 100 and 500 μ mol mol⁻¹), while temperature and irradiance

5 were held constant (30°C and 2000 μ mol m⁻² s⁻¹, respectively). Leaves remained in the chamber for 15-20 minutes, after which they were rapidly sampled into glass vials, flushed with 2% CO₂ and sealed prior to equilibration and subsequent isotope analysis as described above (following Song and Barbour, 2016).

There was no significant relationship between *E* and the proportional difference in *D. glomerata* using the online transpiration technique, consistent with the observation in *L. perenne* (Fig. S13).

10 References

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Table S1: Soil and plant parameters used in the standard MuSICA simulations.

Parameter	Symbol	Value	Unit	Comment ^A
SOIL				
Structural characteristics				
Depth	$d_{\rm soil}$	0.37	m	measured
Quartz fraction	$f_{\rm quartz}$	0.16	% (w/w)	measured
Organic fraction	forganic	0.07	% (w/w)	measured
Remaining soil fraction	fremaining	0.77	% (w/w)	measured
Bulk density	$D_{\rm b}$	1.33	g cm ³	measured
Hydraulic characteristics			-	
Saturated water content	$ heta_{ m sat}$	0.49	$m^{3} m^{-3}$	calculated from water retention curve
Residual water content	$\theta_{\rm sat}$	0.01	$m^{3} m^{-3}$	calculated from water retention curve
Retention curve inflection point	α^{-1}	0.43	m	calculated from water retention curve
Retention curve shape factor	m _{ret}	0.13	-	calculated from water retention curve
Hydraulic conductivity at saturation	$K_{\rm sat}$	0.29	$m d^{-1}$	calculated from hydraulic conductivity
				measurement
Preferential flow				
Fraction of cracks	$f_{\rm crack}$	0.02	-	estimated
Depth of cracks	$d_{\rm crack}$	0.1	m	estimated
Resistance to water transport through soil s	urface pores			
Minimum resistance	r _{s,min}	800	s m ⁻¹	Kelliher et al. (1986)
Maximum resistance	r _{s,max}	16100	s m ⁻¹	Kelliher et al. (1986)
Threshold water content	$ heta_{ m tr}$	0.194	$m^{3} m^{-3}$	Schaap and Bouten (1997),
				Ogée and Brunet (2002)
Soil and root respiration				
Respiration at 25 °C	R_{25}	8.5	µmol m ⁻² s ⁻¹	Gamnitzer et al. (2009),
				Ostler et al. (unpublished)
Base for exponential soil respiration eqn.	Q_{10}	2.2	-	
Surface optical properties				
Surface albedo (of litter or mosses) for visible	$\alpha_{ m vis}$	0.15	-	Deardorff (1978)
light				
Surface albedo (of litter or mosses) for near-	$\alpha_{ m nir}$	0.60	-	
infrared light				
Surface emissivity	$\varepsilon_{ m soil}$	0.95	-	Deardorff (1978)
Soil surface aerodynamic resistance				
Aerodynamic coefficient	C_{u}	33	-	Ogée and Brunet (2002)
VEGETATION				
Canopy structure				
Canopy height	$h_{ m canopy}$	0.078	m	estimated from sward height measurements
Leaf area index	LAI	2.6		estimated from sward height measurements
Mean relative height of vertical leaf area	$\mu_{ m b}$	0.315	-	based on Wohlfahrt et al. (2003)

density profile				
Standard deviation of vertical leaf area	$\sigma_{ m b}$	0.21	-	based on Wohlfahrt et al. (2003)
density profile				
Leaf inclination index	LII	0	-	estimated from sward height measurements
Leaf photosynthesis				
Maximum rate of carboxylation at 25°C	$V_{\rm cmax}$	60	μ mol m ⁻² s ⁻¹	Rogers et al. (1998)
Potential rate of electron transport at 25 °C	$J_{ m max}$	100	μ mol m ⁻² s ⁻¹	calculated from $V_{\rm cmax}$ following
-				Medlyn et al. (2002)
Temperature optimum for $V_{\rm cmax}$	$T_{\rm opt,V}$	40	°C	Harley et al. (1992)
Temperature optimum for J_{max}	$T_{\rm opt.J}$	35	°C	Harley et al. (1992)
Curvature of <i>J</i> -PAR relationship	θ_{J}	0.85	-	-
Efficiency of light energy conversion	α_J	0.18	mol mol ⁻¹	Wullschleger (1993) and papers cited therein
(electrons per photon)				
Dark respiration rate at 25 °C	$R_{\rm d}$	0.86	μ mol m ⁻² s ⁻¹	Ostler et al. (unpublished)
Light inhibition factor for R_d	Ι	0.5	-	cf. Atkin et al. (1997)
Stomatal conductance				
Intercept	g_{\circ}	10	mmol $m^{-2} s^{-1}$	Collatz et al. (1991)
Slope	$m_{\rm gs}$	10	-	Miner et al. (2017), and references therein
Critical water potential	Ψ_{gs50}	-1.5	MPa	Braud et al. (1995)
Steepness parameter	v	4	-	Nikolov et al. (1995)
Minimum conductance for dawn and	g_{\min}	10	mmol $m^{-2} s^{-1}$	-
dusk conditions				
Maximum nocturnal conductance	$g_{ m night}$	30	mmol $m^{-2} s^{-1}$	fitted (see SI text)
VPD threshold for nocturnal	VPD _{thresh}	0.10	MPa	-
conductance				
Mesophyll conductance				
Maximum mesophyll conductance	$g_{ m m}$	0.35	mol $m^{-2} s^{-1}$	Warren (2008)
Leaf boundary-layer conductance				
Leaf size	d	8	mm	measured and estimated (see SI text)
Shoot size	$d_{\rm s}$	78	mm	calculated from sward height measurements
Shelter factor	P_{d}	1.3	-	Monteith and Unsworth, 1990
Root distribution				
Mean of the β -distribution	$\mu_{ m root}$	0.105		estimated
Standard deviation of the β -distribution	$\sigma_{ m root}$	0.06		estimated
Mean root length density		19	km m ⁻²	estimated (see Materials and Methods)
Root hydraulics				
Fine root radius	r	0.15	mm	Picon-Cochard et al. (2012)
Root hydraulic resistance	R _{root}	1	Ts m ⁻¹	estimated
Total internal storage capacity	$W_{\rm cap}$	0.01	kg m ⁻² MPa ⁻¹	estimated
Leaf optical properties				

Reflectance for visible light	$ ho_{ m vis}$	0.105	-	Sellers (1985)
Reflectance for near-infrared light	$ ho_{ m nir}$	0.577	-	Sellers (1985)
Transmittance for visible light	$ au_{ m vis}$	0.07	-	Sellers (1985)
Transmittance for near-infrared light	$ au_{ m nir}$	0.248	-	Sellers (1985)
Leaf emissivity	$\varepsilon_{\text{leaf}}$	0.98	-	Nikolov et al. (1995); Braud et al. (1995);
				Jackson (1988)
Rain interception				
Water storage capacity	S	0.1	$mm m^{-2}$	
Exponent for power function		0.67		Deardorff (1978); Braud et al. (1995)
Wind attenuation				
Canopy drag coefficient	C_{d}	0.2	-	Massman and Weil (1999)
Leaf water isotope modelling				
Leaf water content	W	2	mol m ⁻²	fitted (see SI text)
Proportion of unenriched leaf water	φ	0.39	-	this work
Peclet effective length	L	0.162	m	this work

^A For details of parameter estimation or measurements, see Materials and Methods in main text and Supplemental Information



Time (h)

Figure S1: Comparison of latent heat flux obtained from eddy flux data (blue dots) and latent heat flux predicted by the MuSICA model in standard parameterisation (continuous black line). Panels show 10 d-long periods selected randomly from the first (left panels) and second half (right) of the vegetation periods of 2006 (top) to 2008 (bottom). The numbers above the diurnals indicate

5 the day of the year. Time is given in UTC. Both data sets were obtained at pasture paddock no. 8 of Grünschwaige Grassland Research Station.



Figure S2: δ¹⁸O of rain water (δ¹⁸O_{rain}) collected at the experimental site (black symbols), along with IsoGSM predictions (red symbols) and corrected IsoGSM predictions of δ¹⁸O_{rain} (grey symbols). The latter were obtained by subtracting the mean offset (-1.3‰; cf Fig. S3) between δ¹⁸O_{rain} observed at the site and IsoGSM predictions from the non-corrected IsoGSM data.



Figure S3: Relationship between the δ^{18} O of rainwater collected at the experimental site ($\delta^{18}O_{rain, observed}$) and the δ^{18} O of monthly IsoGSM predictions ($\delta^{18}O_{rain, IsoGSM}$). The solid line represents the 1:1 relation; the dashed line illustrates the mean difference between the two data sets (-1.3‰).





Figure S4: Relationship between the δ^{18} O of atmospheric water vapour as measured at the experimental site (δ^{18} O_{vapour, observed}) and predicted by IsoGSM (δ^{18} O_{vapour, IsoGSM}). The solid line represents the 1:1 relation; the dashed line gives the mean difference between the two data sets (-2‰).



Figure S5: Relationship between volumetric water content (m³ water m⁻³ soil) and pressure head, given as pF value (common logarithm of the pressure head in hPa), (left panel), and hydraulic conductivity (logarithmic scale) and pressure head (right panel), as derived from Hyprop measurements (open circles). The green curve in the left panel represents the Van Genuchten water retention curve fitted to the data, the green curve in the right panel shows the Brooks-Corey hydraulic conductivity curve fitted to







Figure S6: Beta distribution describing the assumed vertical leaf area density distribution at the experimental site (based on Wohlfahrt et al., 2003).



Figure S7: Diurnal time courses of ¹⁸O-enrichment of leaf water ($\Delta^{18}O_{leaf}$) observed (closed circles) on 4/5 August 2005 in pasture paddock no.8 at Grünschwaige and predicted using the two-pool model with a constant proportion of unenriched water ($\varphi = 0.39$; grey circles) and the Péclet model with a constant effective length (L = 0.162 m; open circles). Predicted and observed Δ^{18} O was calculated as the difference between δ^{18} O of leaf water and δ^{18} O of soil water at 7 cm depth. Observed δ^{18} O_{soil} at 7 cm depth was obtained from linear interpolation between the δ^{18} O_{soil} at 2 cm and 12 cm depth. Time is given in UTC.





10

Figure S8: Beta distribution of fine root length density versus soil depth. The black line, with highest root density at 7 cm belowground, represents that used in the standard MuSICA runs; blue and red lines give the low and high alternative root distributions used in the sensitivity analysis (see Fig. 6h in main text), with maxima of root length density at 2 and 30 cm depth, respectively. All distributions have the same total fine root length (19 km m⁻² soil surface).



Figure S9: Diurnal cycles of modelled δ^{18} O of leaf water (black dots) and measured δ^{18} O of the two replicates of leaf water for all sampling dates (light and dark green dots). Numbers in the panels give the day of the year and year. Time is given in UTC.



Figure S10 Correspondence between the δ^{18} O of stem water and soil water at 7 (upper panels) and 20 cm depth (lower) as observed (left) and predicted (right) in the first half (April to June; black squares) and in the second half of the vegetation period (July to October; red circles). The straight lines represent the 1:1 relationship.



Figure S11: Relationship between canopy transpiration rate and the proportional difference between observed leaf water enrichment ($\Delta^{18}O_{leaf}$) and $\Delta^{18}O$ at the evaporative site, as predicted by the Craig-Gordon model ($\Delta^{18}O_{e,ss}$).



Figure S12: Soil water content (SWC) and root water uptake (RWU) along the soil profile as predicted by MuSICA at midday (12:15) for the studied period (2006-2012). The year is indicated on the right hand side.



Figure S13: $\delta^{18}O_{soil}$ with soil depth as predicted by MuSICA (continuous lines) and mean uptake-weighted depth of root water uptake (dashed horizontal lines) on the different sampling dates. Closed circles: observations of $\delta^{18}O_{soil}$ at 7 and 20 cm depth. Sampling date is given by DOY and year, in the lower right corner of each panel.



Figure S14: The relationship between transpiration rate (*E*) and the proportional difference between measured leaf water and the Craig-Gordon predicted enrichment $(1 - \Delta^{18}O_{leaf}/\Delta^{18}O_{leaf})$ for *Lolium perenne*. The relationship in Fig. S13 is statistically significant, but very weak: $1 - \Delta^{18}O_{leaf}/\Delta^{18}O_{e} = 0.017 E + 0.035$; $r^2 = 0.11$; P = 0.045.



Figure S15: The relationship between transpiration rate (*E*) and the proportional difference between measured leaf water enrichment and that at the sites of evaporation $(1 - \Delta^{18}O_{eaf}/\Delta^{18}O_{e})$ within the leaf for *Dactylis glomerata*.