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Effects of mountain tea plantations on nutrient cycling at upstream watersheds

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Abstract. The expansion of agriculture to rugged mountains can exacerbate negative impacts of agricultural activities on ecosystem function. In this study, we monitored streamwater and rainfall chemistry of mountain watersheds at the Feitsui Reservoir Watershed in northern Taiwan to examine the effects of agriculture on watershed nutrient cycling. We found that the greater the proportion of tea plantation cover, the higher the concentrations of fertilizer-associated ions (NO₃⁻, K⁺) in streamwater of the four mountain watersheds examined; on the other hand, the concentrations of the ions that are rich in soils $(SO_4^{2-}, Ca^{2+}, Mg^{2+})$ did not increase with the proportion of tea plantation cover, suggesting that agriculture enriched fertilizer-associated nutrients in streamwater. Of the two watersheds for which rainfall chemistry was available, the one with higher proportion of tea plantation cover had higher concentrations of ions in rainfall and retained less nitrogen in proportion to input compared to the more pristine watershed, suggesting that agriculture can influence atmospheric deposition of nutrients and a system's ability to retain nutrients. As expected, we found that a forested watershed downstream of agricultural activities can dilute the concentrations of NO_3^- in streamwater by more than 70%, indicating that such a landscape configuration helps mitigate nutrient enrichment in aquatic systems even for watersheds with steep topography. We estimated that tea plantation at our study site contributed approximately $450 \text{ kg} \text{ ha}^{-1} \text{ yr}^{-1}$ of NO₃-N via streamwater, an order of magnitude greater than previously reported for agricultural lands around the globe, which can only be matched by areas under intense fertilizer use. Furthermore, we constructed watershed N fluxes to show that excessive leaching of N, and additional loss to the atmosphere via volatilization and denitrification can occur under intense fertilizer use. In summary, this study demonstrated the pervasive impacts of agricultural activities, especially excessive fertilization, on ecosystem nutrient cycling at mountain watersheds.

1 Introduction

Agriculture's expansion is taking place in some of the most rugged mountains in the world, including the Hindu Kush Himalaya (Brown and Shrestha, 2000; Tulachan, 2001), in India, China (Johda et al., 1992) and the Andes (Sarmiento and Frolich, 2002). It is well established that watershed nutrient cycling is tightly linked to land use and that conversion of natural forests to agricultural lands causes nutrient enrichment, especially of N and P, in streamwater (Omernik, 1976; Johnes, 1996; Tilman et al., 2001; Murty et al., 2002; Allan, 2004; Uriarte et al., 2011; Evans et al., 2014). The impacts are likely exacerbated by steep slopes and high precipitation as residence time is reduced and leaching potential increased under such conditions (Brouwer and Powell, 1998; Tokuchi et al., 1999). Thus, mountain agriculture in the tropics and subtropics characterized with high precipitation is likely to have a substantial negative impact on ecosystem function. Yet, empirical studies in tropical or subtropical mountain watersheds are very limited.

In addition to nutrient output in streamwater, cultivation and fertilization on agricultural lands could affect atmospheric deposition of nutrients (i.e., nutrient input via wet and dry deposition). Fine particles suspended from exposed lands and volatilized gases such as NH₃ from manure are scavenged by precipitation (van Breemen et al., 1982), which can then be deposited back to the watersheds. However, in contrast to the large number of reports on streamwater chemistry, few studies of watershed nutrient cycling have examined the effects of land use on precipitation chemistry.

Proper landscape configuration could potentially mitigate the negative effects of agriculture on watershed nutrient cycling. A study at the Hubbard Brook Experimental Forest demonstrated that watershed-level responses were most sensitive to areas of approximately 10–20 ha surrounding the drainage area, where much of the variation in element fluxes occurred (Johnson et al., 2000). Such understanding has led to the common practice of establishing riparian buffer zones as a way to remove pollutants and prevent nutrients from entering streamwater (reviewed by Muscutt et al., 1993). Through proper landscape configuration, negative impacts of agriculture on nutrient cycling in mountain watersheds may also be reduced without sacrificing socioeconomic benefits of agriculture. However, what constitutes a proper landscape configuration is likely to vary with climate and topography.

Here we examined the effects of mountain agriculture, mainly tea plantations, on watershed nutrient cycling at the Feitsui Reservoir Watershed (FRW) in subtropical Taiwan. We first compared streamwater chemistry across four watersheds within the FRW, two with substantial agricultural land use and two primarily covered with natural forests. To assess the effects of agriculture on atmospheric deposition of nutrients and its role in watershed nutrient retention, we focused on the pair of watersheds with the highest and lowest tea plantation covers and compared their rainfall chemistry in relation to streamwater chemistry. The FRW is characterized by high rainfall (> 3000 mm; Taipei Feitsui Reservoir Administration), steep slopes (on average 42 %), and heavy use of fertilizers in tea plantations $(425-2373 \text{ kg N ha}^{-1} \text{ yr}^{-1} \text{ and}$ 99–551 kg P ha⁻¹ yr⁻¹; Water Resources Agency, 2010; see Sect. 2 for details). Many studies have demonstrated substantial nutrient efflux and sediment production from surrounding tea plantations to the reservoir over the past 2 decades (Chang and Wen, 1997; Lu et al., 1999; Kuo and Lee, 2004; Li and Yeh, 2004; Hsieh and Yang, 2006, 2007; Zehetner et al., 2008; Chiueh et al., 2011; Wu and Kuo, 2012). Yet, to our knowledge none examined both the effects of spatial configuration of agricultural lands on nutrient export and the effects of agriculture on atmospheric deposition. The FRW is rare among (sub)tropical mountain watersheds in that the effects of agriculture on its streamwater quality have been intensively studied. With the addition of this study, we believe that the FRW can serve as a classic case illustrating the effects of agriculture on nutrient cycling in watersheds with rugged topography and high precipitation, which can be very

informative to other less-studied (sub)tropical mountain watersheds.

We hypothesized that agriculture would increase nutrient output in streamwater (H_1) as well as atmospheric input of nutrients through rainfall (H_2) . We also hypothesized that through the disruption of natural vegetation, agriculture would increase nutrient leaching and decrease the retention ratio of essential nutrient elements (H_3) . Our specific predictions are that

- watersheds with higher proportion of tea plantation cover have higher concentrations and fluxes of fertilizer-associated ions in the streamwater than forested watersheds (H_1) ,
- watersheds with higher proportion of tea plantation cover have higher concentrations and fluxes of fertilizer-associated ions in the rainfall than forested watersheds (H_2) ,
- watersheds with higher proportion of tea plantation cover have a lower nitrogen retention ratio (in proportion to input) than forested watersheds (H_3) .

In addition, we explored (1) the role of landscape configuration in mitigating agricultural effects by quantifying the dilution effects of a forested watershed downstream from watersheds with substantial tea plantation cover, and (2) the N and P dynamics associated with tea plantations by quantifying the differences in their fluxes between a forested watershed (background values) and a nearby watershed with substantial tea plantation cover.

2 Materials and methods

2.1 Study site

The FRW is located along the Peishi Creek of northern Taiwan, with a drainage area of 303 km². The elevation of the FRW ranges from 45 to 1127 m, with a mean slope of 42 % (Fig. 1). The underlying geology of the FRW region is mainly argillite and slate with sandstone interbeds, and the soils are mostly Entisols and Inceptisols with high silt contents (Zehetner et al., 2008).

Annual precipitation is high and spatially varied, ranging from 3500 mm in the southwest portion of the FRW to 5100 mm in the northwest during 2001–2010 (J. C. Huang, unpublished data). The vegetation is primarily composed of secondary-growth, mixed broad-leaf forests dominated by Fagaceae and Lauraceae (Chen, 1993). Approximately 16% of the FRW is agricultural land with tea plantations covering an area of 1200 ha, or 25% of all agricultural lands (Chang and Wen, 1997; Chou et al., 2007). In 1986 the FRW was designated as a water resource protection area, followed by the construction of the Feitsui Reservoir in 1987. Today, the reservoir provides drinking water to the six million people



Figure 1. Location and land use distribution of the studied watersheds.

in the Taipei metropolitan area. The forests in the FRW have been protected (no cutting, thinning or converting to agricultural use) since 1986. Therefore, current agricultural activities are limited to private lands with a pre-existing agricultural use which still has an impact at the study site.

2.2 Sampling regime

Four watersheds of the FRW (A1, A2, F1, F2; Fig. 1) with varying proportions of tea plantation cover (22 % in A1, 17 % in A2, 2.9% in F1, 0.4% in F2; Table 1) were included in this study. Other crops make up only a small proportions of the watersheds (< 1 %), so they are not included in Table 1. Natural forests are the most dominant land cover for all four watersheds (68% in A1, 76% in A2, 93% in F1, 99% in F2; Table 1), making tea plantation the primary contributor to the differences in landscape across the four watersheds. Weekly samples of streamwater were collected from all four watersheds. In addition, weekly samples of rainwater were collected from the two watersheds with the lowest (F2) and highest proportions of agricultural lands (A1). A1, A2, and F2 are watersheds ($< 3 \text{ km}^2$) drained by first-order streams whereas F1 is a much larger watershed (86 km²) drained by a third order stream that drains through A1 and A2 (Fig. 1). We collected weekly rainfall and streamwater samples every Tuesday from September 2012 to August 2014. Rainfall samples were collected using a 20 cm diameter polyethylene (PE) bucket, from which a 600 mL subsample was taken and placed into a PE bottle for transportation back to the laboratory. Streamwater samples were collected by dipping a PE bucket into the stream and, similarly to rainfall sampling, a

Table 1. Basic information of the studied watersheds.

	A1	A2	F1	F2
Area (km ²)	2.92	1.36	86.04	0.67
Slope (%)	39.3	34.8	38.7	48.1
	Land u	se (%)		
Natural forest	68.0	75.5	93.5	99.2
Agriculture	22.1	17.1	2.87	0.38
Road	3.61	2.96	0.77	0.00
Building	1.54	1.31	0.35	0.00
Water body	0.69	0.19	1.12	0.00
Others	4.11	2.96	1.44	0.38

600 mL subsample was taken and placed into a PE bottle for transportation back to the laboratory.

2.3 Water chemistry

All samples were transported back to the laboratory within 24 h. Conductivity and pH of the water samples were measured on the same day of collection. The samples were filtered through 0.45 µm filter paper. Major cations (Na⁺, K⁺, Ca²⁺, Mg²⁺, NH₄⁺) and anions (Cl⁻, SO₄²⁻, NO₃⁻) were analyzed by ion chromatography on filtered samples using Dionex ICS 1000 and DX 120 (Thermo Fisher Scientific Inc. Sunnyvale, CA, USA). PO₄³⁻ was measured using standard vitamin-C molybdenum-blue method with the detection limit of 0.01 µM (APHA, 2005). Prior to chemical analysis, samples were stored at 4 °C without preservatives.

Data on rainfall and streamflow quantity of the watersheds were estimated from the rain gauges and discharge gauges maintained by the Central Weather Bureau and Water Resource Agency of Taiwan, respectively. The distance between a watershed and its nearest rain gauges was 1.0-8.5 km, and that between a watershed and its nearest discharge gauges was 3.0-5.0 km. The weekly and monthly rainfall of a watershed was directly assigned to the values registered at the nearest rain gauge (i.e., COA530 for A1 and COA540 for F2; Fig. 1, S1a). The weekly and monthly streamflow of a watershed was estimated by the area ratio method in which the streamflow was assigned to the values registered at the nearest discharge gauge (i.e., 1140H099 for A1, A2, and F1, and 1140H097 for F2; Fig. S1b) and then adjusted by the area ratio of the studied watershed relative to the watershed where the discharge gauge was located. The validity of this method has been confirmed for several watersheds in Taiwan (Huang et al., 2012; Lee et al., 2014).

2.4 Element fluxes

Weekly element fluxes through rainfall and streamflow of A1 and F2 were derived by multiplying weekly concentrations by weekly rainfall/streamflow. Monthly fluxes were accumulated from weekly fluxes, and when a weekly sample spanned over 2 months, it was divided into the 2 months in proportion to the rainfall/streamflow quantity.

In order to provide a more comprehensive understanding on how mountain agriculture affects watershed nutrient cycling, we constructed and compared N and P fluxes for watersheds with the highest (A1) and lowest (F2) tea plantation cover. We made three assumptions in the calculation of watershed nutrient fluxes. First, we assumed the input from dry deposition is 28% of that from precipitation for both watersheds. This value was based on a study using the Na⁺ ratio method at the Fushan Experimental Forest (Lin et al., 2000), a natural hardwood forest 17 km south of the FRW. Second, the amount of fertilizer used is assumed to be close to $786 \text{ kg N} \text{ ha}^{-1} \text{ yr}^{-1}$ and $171 \text{ kg P} \text{ ha}^{-1} \text{ yr}^{-1}$, the values taken from a case study in which the management practices (e.g., applications of fertilizers and pesticides, time and yield of harvests) were carefully recorded by a farmer in the same region as the current study (Tsai and Tsai, 2008). Although only one farmer was involved in the case study, the values are consistent with those reported by FAO (2002) and very close to the mean values across 10 tea plantations in our study area $(743 \text{ kg N ha}^{-1} \text{ yr}^{-1} \text{ ranging from } 425 \text{ to}$ $2373 \text{ kg N ha}^{-1} \text{ yr}^{-1}$, and $179 \text{ kg P ha}^{-1} \text{ yr}^{-1}$ ranging from 99 to $550 \text{ kg P ha}^{-1} \text{ yr}^{-1}$; Water Resources Agency, 2010). Adjusting for the proportion of agricultural lands (22.1, 0.38 %), the amounts of fertilizers used in A1 were estimated to be $173.7 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ and $37.8 \text{ kg P ha}^{-1} \text{ yr}^{-1}$, and those in F2 to be $3 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ and $0.6 \text{ kg P ha}^{-1} \text{ yr}^{-1}$. There was very little change in biomass of tea plantation after 10 years because tea plants are regularly trimmed, with the litter left in the field, to maintain the same height optimal for harvest. Thus, our third assumption is that N and P is lost due to the uptake by tea trees being equivalent to the N and P in the harvested tea leaves. The amount of N removed through tea harvest (113 kg ha⁻¹ yr⁻¹) was taken from the same case study and the amount of P removed (7.35 kg ha⁻¹ yr⁻¹) was calculated using the median of P : N ratios (0.065) reported for tea trees in Taiwan (Tsai and Tsai, 2008). After adjusting for the proportion of tea plantation cover, A1 was estimated to have 25.0 kg N ha⁻¹ yr⁻¹ and 1.6 kg P ha⁻¹ yr⁻¹ removed through harvest, and F2 to have 0.43 kg N ha⁻¹ yr⁻¹ and 0.03 kg P ha⁻¹ yr⁻¹ removed through harvest. Using the following mass balance model, we constructed fluxes of N and P of the two watersheds:

$$Ratio_{ret} = 1 - \frac{OUT_{riv} + OUT_{harv}}{IN_{dep} + IN_{fer} + IN_{fix}}.$$
 (1)

Here, Ratioret indicates the ratio of input to the watershed that was retained within the watershed. The OUT_{riv} and OUT_{harv} are the riverine export and harvest, respectively. The IN_{dep}, IN_{fer}, and IN_{fix} indicate the atmospheric deposition, fertilizer application, and biological fixation. Note that the biologic fixation term was not used for P calculation. Since the tea plantation does not use leguminous crop as fertilizers and the biological fixation in tropical forest is known to be less than $10 \text{ kg N} \text{ ha}^{-1} \text{ yr}^{-1}$ (Sullivan et al., 2014), the IN_{fix} is assumed to be between 0 and $10 \text{ kg N ha}^{-1} \text{ yr}^{-1}$. We did not include the loss through denitrification and volatilization within tea field in the calculation of N retention ratio because we did not have good estimates. However, the effects of such uncertainties and omissions on estimating N retention ratio were discussed. We did not calculate the retention ratio for P because the majority of P in watersheds was in particulate forms (Smith et al., 1991) that were not analyzed in our study.

2.5 Statistical analysis

We used the general linear model with repeated measurements to compare monthly concentration and flux of ions in streamwater among the four watersheds (F1, F2, A1, A2), followed by Fisher's least significant difference (LSD) post hoc comparisons. NH_4^+ was excluded from streamwater analysis due to its low concentration. We used a onetail paired *t* test to examine if monthly ion concentration (volume weighted from weekly samples) and flux in rainfall were higher at the watershed with higher agricultural land cover (A1) than the more pristine watershed (F2). All statistical analysis was conducted using SPSS 22.0 (IBM Corporation, New York).

3 Results

3.1 Streamwater chemistry

The concentrations of all analyzed ions in streamwater differed significantly among the four watersheds (Table 2). A1, the watershed with the highest proportion covered by tea plantations, had significantly higher concentrations of all ions except H⁺ than the other three watersheds (Table 2, Fig. 2). In contrast, F2, the watershed with the lowest proportion covered by tea plantations, had the lowest concentrations of H⁺, Na⁺, K⁺, Cl⁻, and NO₃⁻. Furthermore, it is worth noting that F2, the watershed with the steepest slopes, had the second highest concentrations of ions rich in soils and soil solution, including Ca²⁺, Mg²⁺, and SO₄²⁻ (Table 2, Fig. 2).

Similar to ion concentration, the fluxes of all ions differed significantly among watersheds (Table 2). A1 had the largest fluxes of K⁺, Ca²⁺, Mg²⁺, NO₃⁻, and SO₄²⁻ and F2 had the smallest fluxes of H⁺, Na⁺, K⁺, Mg²⁺, Cl⁻, and NO₃⁻ (Table 2). PO₄³⁻ flux was significantly larger at A1 and A2, which were not so different from each other, than F1 and F2, which were also not so different from each other (Table 2). Although the fluxes of Na⁺ and Cl⁻ differed significantly among A1, A2, and F1, these differences were considerably smaller than the differences between the three watersheds and F2 (Table 2).

3.2 Rainfall chemistry

Five of the 10 measured ions had significant (p < 0.05) or marginally significant (p < 0.1) higher concentrations in A1 than in F2 (H⁺, Na⁺, Cl⁻, NO₃⁻, p < 0.05; NH₄⁺, p = 0.067; Table 3, Fig. 3). Furthermore, 7 of the 10 measured ions had significant or marginally significant higher fluxes in A1 than in F2 (H⁺, Ca²⁺, Cl⁻, p < 0.05; Na⁺, Mg²⁺, NH₄⁺, NO₃⁻, p < 0.1; Table 3).

3.3 N and P fluxes

Because the proportion of agricultural cover was very low at F2 (i.e., 0.38 %) and the resulting fertilizer input and harvest output were small and already accounted for (Table 4), we treated F2 as a background and attributed the differences between A1 and F2 to agricultural activities. We estimated stream N and P outputs from the tea plantation at A1 to be approximately 105.7 and 1.6 kg ha⁻¹ yr⁻¹, respectively (Table 4). Scaling up from 22 % of tea plantation cover to 100 %, the stream N and P outputs from A1 could reach as high as 450 and 7.3 kg ha⁻¹ yr⁻¹, respectively.

From our mass balance construction of element fluxes, N input exceeded output at both watersheds (Table 4, Fig. 4). At A1, 35% of the N input ($69 \text{ kg ha}^{-1} \text{ yr}^{-1}$) to the watershed was retained (Table 4, Fig. 4). At F2, 72% of the N input ($15 \text{ kg N ha}^{-1} \text{ yr}^{-1}$) was retained (Table 4, Fig. 4).

For P, the output through streamflow $(2.6 \text{ kg ha}^{-1} \text{ yr}^{-1})$ was smaller than the input through atmospheric deposition $(3.6 \text{ kg ha}^{-1} \text{ yr}^{-1})$ at F2. At A1, the output of P through streamflow and harvest $(5.8 \text{ kg ha}^{-1} \text{ yr}^{-1})$ was greater than the input through atmospheric deposition $(4.6 \text{ kg ha}^{-1} \text{ yr}^{-1})$, but when fertilization was taken into account, the total output of PO₄³⁻-P was trivial relative to the total P input $(42.4 \text{ kg ha}^{-1} \text{ yr}^{-1})$ (Table 4).

4 Discussion

4.1 Streamwater chemistry

The watershed with the highest proportion of tea plantation cover (A1) had the highest concentrations and fluxes of most ions in streamwater, suggesting the role of agriculture on increasing nutrient output. Furthermore, the fact that the output of fertilizer-associated ions (NO₃⁻ and K⁺) matched the proportion of tea plantation cover across the four watersheds (i.e., the rank of the proportion of tea plantation cover from high to low: A1, A2, F1, and F2; rank of ion concentration and flux from high to low: A1, A2, F1, and F2) strongly supports the effects of agriculture on streamwater chemistry (*H*₁).

However, streamwater chemistry is affected by complex processes beyond a single factor of land use. For example, P is also an important component of fertilizers but, unlike NO_3^- and K⁺, the concentration of PO_4^{3-} at F2 was not significantly different from that at A1 and A2, and all were significantly higher than F1. Erosion is known to enhance leaching loss of PO_4^{3-} (Gaynor and Findlay, 1995; Turtola and Jaakkola, 1995; Liu et al., 2006; Chang et al., 2008; Lee et al., 2013). The greater erosion and leaching associated with the steeper slopes of F2 may have matched the effect of fertilization and led F2 to have a PO_4^{3-} concentration as high as A1 and A2. To further illustrate this topographic effect, we compared streamwater chemistry between the two forested watersheds (F1 and F2), removing the potential confounding effect of land use. Indeed, the steeper F2 (48%) had a higher PO_4^{3-} concentration than the less steep F1 (39%) (Fig. 2, Table 2), despite that F2 has a higher proportion of natural forest cover. Soil erosion is arguably the greatest concern to most P mitigation programs because the concentration of P on the surface of soil particles is often orders of magnitude greater than that in a soil solution (Sharpley et al., 2002; Kleinman et al., 2011). Therefore, it is not surprising that topography may be a more important driver for riverine P than land use at our study site. The enhanced erosion/leaching associated with the steeper slope at F2 may also explain why F2 had the second highest concentration of SO_4^{2-} , Ca^{2+} , and Mg^{2+} , the ions that are abundant in soils.



Figure 2. Monthly ion concentration (volume-weighted from weekly samples) of streamwater of watersheds A1, A2, F1, and F2.



Figure 3. Monthly ion concentration (volume-weighted from weekly samples) of rainfall of watersheds A1 and F2.

Ion		Conce	entration (µeq L ⁻	¹)			Flux	$(\text{meq m}^{-2} \text{mo}^{-1})$)	
	A1	A2	FI	F2	diff	A1	A2	FI	F2	diff
H+	0.96 ± 0.006	1.22 ± 0.006	0.91 ± 0.007	0.76 ± 0.004	a, b, a, c	0.030 ± 0.001	0.038 ± 0.001	0.036 ± 0.001	0.016 ± 0.004	a, b, ab, c
Na^+	266 ± 4.88	254 ± 3.65	233 ± 4.45	231 ± 4.10	a, b, c, c	76.4 ± 1.74	73.0 ± 1.70	80.1 ± 1.68	46.7 ± 0.90	a, b, ab, c
\mathbf{K}^+	282 ± 0.87	213 ± 6.27	125 ± 0.49	108 ± 3.63	a, b, c, d	8.24 ± 0.20	6.14 ± 0.14	4.27 ± 0.50	2.19 ± 0.36	a, b, c, d
Ca^{2+}	306 ± 7.49	193 ± 5.41	170 ± 7.34	273 ± 8.04	a, b, c, d	87.0 ± 1.92	54.1 ± 1.17	55.8 ± 1.02	54.4 ± 1.00	a, b, b, b
Mg^{2+}	255 ± 5.10	188 ± 4.25	148 ± 4.72	206 ± 5.78	a, b, c, d	72.5 ± 1.62	52.8 ± 1.15	49.2 ± 0.94	41.0 ± 0.74	a, b, b, c
CI-	199 ± 4.00	182 ± 3.06	178 ± 4.76	145 ± 2.55	a, b, b, c	59.2 ± 1.51	53.2 ± 1.34	62.8 ± 1.49	29.8 ± 0.64	a, b, a, c
NO_3^-	209 ± 5.31	158 ± 2.80	28.3 ± 0.76	16.1 ± 0.95	a, b, c, d	62.9 ± 1.63	46.8 ± 1.19	10.2 ± 0.25	3.32 ± 0.078	a, b, c, d
SO_4^{2-}	212 ± 6.29	123 ± 3.96	116 ± 3.96	183 ± 6.45	a, b, c, d	59.2 ± 1.30	33.9 ± 0.74	39.1 ± 0.78	35.7 ± 0.66	a, b, b, b
PO_{2}^{2-}	1 50 ± 0 100		0.72 ± 0.114	129 ± 0.026	a, b, b, a	1.14 ± 0.0030	1.08 ± 0.0054	0.69 ± 0.028	0.69 ± 0.0030	a, a, b, b

Table 2. Mean (± 1 SE – standard error) monthly ion concentration (volume-weighted from weekly samples) and flux of streamflow

Table 3. Mean $(\pm 1 \text{ SE})$ monthly ion concentration (volume-weighted from weekly samples) and flux of rainfall.

Ion	Concentration ($\mu eq L^{-1}$)		Flux (meq m ^{-2} mo ^{-1})		
	A1	F2	A1	F2	
H^+	39 ± 6.7	$31 \pm 5.4^{*}$	12 ± 3.9	$7.9 \pm 1.5^{*}$	
Na ⁺	107 ± 24	$84\pm18^*$	30 ± 8.5	$23 \pm 6.5(*)$	
K^+	8.0 ± 1.3	7.8 ± 1.3	2.2 ± 0.45	1.9 ± 0.32	
Ca ²⁺	21 ± 3.2	19 ± 4.4	5.7 ± 1.0	$4.2\pm0.61^*$	
Mg ²⁺	30 ± 5.8	26 ± 5.6	8.2 ± 2.1	$6.5 \pm 1.7(*)$	
NH_4^+	19 ± 2.9	$15 \pm 2.7(*)$	5.1 ± 1.3	$3.8 \pm 0.67(^*)$	
Cl-	140 ± 30	$100\pm22^*$	38 ± 11	$28\pm8.2^*$	
NO_3^-	24 ± 3.9	$18\pm3.0^*$	7.0 ± 2.0	$4.7 \pm 0.90(*)$	
SO_4^{2-}	58 ± 8.6	53 ± 7.7	15 ± 3.6	13 ± 2.4	
PO_4^{3-}	0.96 ± 0.03	0.63 ± 0.03	0.75 ± 0.30	0.51 ± 0.12	

A1 and F2 denote the two watersheds; an asterisk * indicates a significant difference between the two watershed (p < 0.05); an asterisk inside a parenthesis (*) indicates a marginally significant difference between the two watersheds (p < 0.1).



Figure 4. Schematic diagram of N fluxes of watersheds A1 and F2. A1 represents a watershed with 22% agricultural lands and 68% forests (**a**); F2 represents a watershed with 0.38% agricultural lands and 99% forests (**b**) (unit: kg Nha⁻¹ yr⁻¹).

4.2 Rainfall chemistry

We confirmed that agricultural activities can influence watershed nutrient cycling via atmospheric deposition in our study site (H_2). We found higher concentrations and fluxes of NO₃⁻ and NH₄⁺ in rainfall at A1, a watershed with 22 % of tea plantation cover, compared to F2, the watershed almost entirely covered by natural forests. Ammonium sulfate, urea and calcium ammonium nitrate [5Ca(NO₃)₂ · NH₄NO₃ · 10H₂O], which contain a high quantity of NO₃⁻ and NH₄⁺ are commonly used N fertilizers in Taiwan (Huang, 1994). Therefore, in tea plantations at FRW, substantial suspension and volatilization of ammonium sulfate, urea, and calcium ammonium nitrate likely contributed to the high concentrations and fluxes of NO₃⁻ and NH₄⁺ in rainfall at A1. On the other hand, the concentrations of PO₄³⁻ and K⁺ in rainfall were not higher at A1 compared to F2, which may be explained

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	Nitrogen $(kg ha^{-1} yr^{-1})$		Phosphorus $(kg ha^{-1} yr^{-1})$		
	A1	F2	A1	F2	
Input					
Wet deposition	20.4	14.3	3.6	2.8	
Dry deposition	5.7	4.0	1.0	0.8	
Fertilization	173.7	3.0	37.8	0.6	
Total	199.8	21.3	42.4	4.2	
Output					
Harvest	25.0	0.4	1.6	0.0	
Stream output*	105.7	5.6	4.2	2.6	
Total	130.7	6.0	5.8	2.6	

Table 4. Inputs and outputs of nitrogen and phosphorus of watersheds A1 and F2. See text for the assumptions made in the calculations of dry deposition, fertilization, and harvest.

* For stream output, only dissolved inorganic forms are considered.

by the low mobility of PO_4^{3-} and smaller quantity of P and K in fertilizers.

Once in the atmosphere, aerosols/chemicals can be transported to other locations but most of them will be deposited in nearby ecosystems. In central Taiwan, the high NH_4^+ concentration in precipitation in a high elevation forest (2000 m) was attributed to mountain agriculture that occurred 10 km away (Ding et al., 2011). With the predicted expansion of agriculture to the mountains both in Taiwan and many other regions (Johda et al., 1992; Brown and Shrestha, 2000; Tulachan, 2001), even pristine ecosystems will not be free from the impacts (e.g., acidification and eutrophication associated with H⁺ and NO₃⁻) of agricultural activities.

Because Taiwan is a small island, sea salt aerosols are important components of rainfall (Lin et al., 2000). The distance to the coast, specifically, has been used to explain the variation of Na⁺ and Cl⁻ concentrations in precipitation among four sites in central Taiwan (Ding et al., 2011). The higher concentrations and fluxes of Na⁺ and Cl⁻, and to a lesser degree Mg²⁺, at A1 than at F2 likely reflected such oceanic influences. The watersheds receive winter rains, along with sea salt aerosols, from the north/northeast coasts (northeast monsoon). While A1 is located on the windward side, F2 is on the leeward side. Therefore, a substantial proportion of the sea salt aerosols may have been intercepted before they can reach F2. Although summer rains move from the opposite direction, the watersheds are relatively far from the west/southwest coasts (> 60 km), making summer rains less important to the input of sea salt aerosols to the watersheds.

In contrast to Na⁺ and Cl⁻, the differences in topographic position and distance to the ocean between A1 and F2 seemed to have a limited effect on SO_4^{2-} deposition. Many studies reported significant contributions of longrange-transported S and N from eastern China to Taiwan via the northeast monsoon (Lin et al., 2005; Junker et al., 2009). Because A1 is on the windward side of the northeast monsoon, it may experience a higher input of pollutants from long-range transport than F2, which is on the leeward side. The lack of significant differences in SO_4^{2-} between the two watersheds suggest that the two watersheds are too close to show differential influences of pollutants that are transported from sources several hundred kilometers away.

4.3 Landscape configuration and streamwater chemistry

The large differences in NO_3^- concentration and flux between F1 and A1, A2 highlight the role of landscape configuration on streamwater chemistry. Both A1 and A2 are subwatersheds of F1; however, the influence of tea plantation on A1 and A2 largely dissipated as water entered into forested F1. Specifically, the concentration of NO_3^- was 70% lower at F1 than at A1 and A2. Comparing to the difference in concentration and flux of NO_3^- between F1 and F2 (< 30%), that between F1 and A1, A2 is striking (> 300%; Fig. 2). Thus, by constraining agricultural activities away from the main stream and maintaining natural cover of its watershed, the impact of agriculture on nutrient enrichment could be reduced. Our result confirmed the importance of landscape configuration on streamwater chemistry (Dillon and Molot, 1997; Johnson et al., 1997; Palmer et al., 2004).

4.4 N and P output from agriculture

The per-hectare output of N from tea plantations reported here $(450 \text{ kg ha}^{-1} \text{ yr}^{-1})$ is extraordinary high compared to those reported for many agricultural watersheds around the globe. For example, a study from the Baltimore Ecosystem Study reported an annual output of NO3-N at $13-20 \text{ kg} \text{ ha}^{-1} \text{ yr}^{-1}$ for a 7.8 ha watershed that is completely covered by agricultural lands and has gentle slopes (Groffman et al., 2004). For the four watersheds that were 30-40% covered by row crops and received fertilization at $50-70 \text{ kg N} \text{ ha}^{-1} \text{ yr}^{-1}$ in the southeastern coastal plain of the US, nutrient output through streamflow was $< 6 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ (Lowrance et al., 1985). In the Great Barrier Reef, Australia, total output via streamflow was approximately $5 \text{ kg ha}^{-1} \text{ yr}^{-1}$ for NO₃-N from a watershed with 29 % of the land covered by pasture and 14 % by crop lands (Hunter and Walton, 2008).

High N output from agricultural lands is probably common in Taiwan and other regions under intensive fertilizer use. It has been reported that over-fertilization is common in Japan, Korea, and Taiwan, and despite an estimated 23– 63 % over-fertilization the use of fertilizers is still increasing in the region (Ahmed, 1996). In the Danshui River of northeastern Taiwan, the output of dissolved inorganic N ranged from 3 kg ha⁻¹ yr⁻¹ in relatively pristine headwaters covered mostly by natural forests to 100 kg ha⁻¹ yr⁻¹ in a populated estuary (Lee et al., 2014; Shih et al., 2015). In humid southeastern China, N output from a watershed with 17.5 % of agricultural lands, steep slopes (the watershed has a mean slope of 21 % and the site is located in the hilly upstream region), and very heavy application of N fertilizers $(300-1000 \text{ kg ha}^{-1} \text{ yr}^{-1})$ reached 73 kg ha⁻¹ yr⁻¹ (Chen et al., 2008), approximately the same magnitude as those reported here. Our study clearly demonstrated that high application of fertilizers in regions with high rainfall and steep slopes could lead to an extremely high output of N and, therefore, eutrophication risk for downstream watersheds. The misconception that heavy fertilization leads to high economic profit has resulted in the popular practice of heavy fertilization in tea plantations, commonly at a level similar to or higher than that in our study site $(740 \text{ kg N ha}^{-1} \text{ yr}^{-1})$. For example, conventional N fertilization in tea plantations is approximately $1100 \text{ kg ha}^{-1} \text{ yr}^{-1}$ in Japan, which is more than twice the suggested amount with the same tea yield (Oh et al., 2006).

In contrast to N, most of the P fertilizer was retained within the watershed or transported in particulate form so that dissolved P only accounts for a small proportion of the input. In most agricultural watersheds, the majority (>90%) of P leaves the watersheds in particulate form (Smith et al., 1991), and the loss in dissolved form (i.e., PO_4^{3-}) through runoff is relatively minor (Brady and Weil, 1999). Thus, while the dissolved form of P could respond to land use changes, a complete P budget at watershed scale still requires reliable estimates on the particulate P.

4.5 Watershed N fluxes

The 72 % N retention at F2 is likely an underestimate because the input from biological N fixation (BNF) was not included in the calculation. Based on a recent synthesis (Sullivan et al., 2014), BNF in tropical forests is not as high as previously reported and, on average, is slightly less than $10 \text{ kg ha}^{-1} \text{ yr}^{-1}$ for secondary forests. Thus, adding BNF to N input could increase the N retention ratio at F2 (assuming a BNF of $10 \text{ kg ha}^{-1} \text{ yr}^{-1}$, the N retention ratio at F2 would increase from 72 to 81%). The high N retention ratio of F2 suggests that the secondary natural forest is probably still growing. In contrast, because N fertilizers were applied at rates that are 1 order of magnitude greater than BNF at A1, and high N fertilization is known to negatively affect BNF (Sanginga et al., 1989; Fuentes-Ramírez et al., 1999), adding BNF to nutrient input has little effect on the N retention ratio at A1 (assuming a BNF of $10 \text{ kg ha}^{-1} \text{ yr}^{-1}$, the N retention ratio at A1 would increase from 35 to 37%).

In addition to BNF, the calculation of the N retention ratio did not take into account the loss through volatilization and denitrification. Because it rains frequently at the FRW, soil moisture is likely high throughout the year and, consequently, N loss through denitrification could be substantial. In addition, because fertilizers are applied in solid form, volatilization of NH_3 could also be high. Thus, if both denitrification and volatilization are taken into account, the N retention ratio at A1 is even lower. The return of N back to the atmosphere through denitrification and volatilization helps explain the higher atmospheric N deposition at A1 than at F2. The low retention ratio and the resulting high leaching loss of N at A1 impose a major threat to the streamwater quality that could lead to reservoir eutrophication.

Surprisingly, from our construction of the N fluxes, the loss of N through the annual harvest $(25 \text{ kg ha}^{-1} \text{ yr}^{-1})$ at A1 approximately equals the annual atmospheric deposition $(26 \text{ kg ha}^{-1} \text{ yr}^{-1})$, of which only a small portion should have come from fertilizers (atmospheric N deposition at F2 is only 8 kg lower than at A1, suggesting that less than 8 kg of atmospheric N deposition could potentially come from fertilizers). In other words, to maintain the current harvest, not much N fertilization is actually required, and most of the $173.7 \text{ kg N} \text{ ha}^{-1} \text{ yr}^{-1}$ from fertilization is simply lost through hydrological process (i.e., leaching) to the streams and the Feitsui Reservoir and/or returned to the atmosphere, both of which could have negative environmental impacts. Our construction of the element fluxes clearly showed that the N fertilizers are applied at rates that are neither ecologically nor economically sound, and such excessive fertilization may cause fundamental changes in watershed nutrient cycling (Fig. 4).

5 Conclusions

Agricultural and forested watersheds in tropical/subtropical mountains could have distinct patterns of nutrient cycling. Even a moderate proportion of tea plantation cover (17–22%) in mountain watersheds, when in combination with steep slopes and high precipitation, could lead to much higher ion concentrations in both streamwater (nutrient output) and rainwater (nutrient input) and much lower N retention ratios at watershed scale. Thus, mountain watersheds may be particularly vulnerable to agricultural expansion.

Topographic control is important in nutrient leaching from mountain watersheds, particularly for ions that are rich in soils, such as SO_4^{2-} , Ca^{2+} , and Mg^{2+} .

Proper spatial configuration of agricultural lands in mountain watersheds can mitigate the impact of agriculture on NO_3^- output by 70%, thus reducing the risk of eutrophication for streams and lakes.

The contribution of tea plantations to the N output in streamwater for one of the studied watersheds (i.e., A1) is estimated at approximately $450 \text{ kg N ha}^{-1} \text{ yr}^{-1}$. This level of fertilization exceeds previous reports around the globe and can only be matched in magnitude by one study in China where fertilizers were excessively applied.

The conservative construction of the N fluxes for the watersheds indicates over-fertilization at one of the studied watersheds (i.e., A1), which likely resulted in leaching of N and

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additional loss of N to the atmosphere via volatilization and denitrification.

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