

A framework for the quantitative assessment of climate change impacts on water-related activities at the basin scale

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Received: 17 December 2010 – Published in Hydrol. Earth Syst. Sci. Discuss.: 19 January 2011

Revised: 1 June 2011 – Accepted: 14 June 2011 – Published: 28 June 2011

Abstract. While quantitative assessment of the climate change impact on hydrology at the basin scale is quite addressed in the literature, extension of quantitative analysis to impact on the ecological, economic and social sphere is still limited, although well recognized as a key issue to support water resource planning and promote public participation. In this paper we propose a framework for assessing climate change impact on water-related activities at the basin scale. The specific features of our approach are that: (i) the impact quantification is based on a set of performance indicators defined together with the stakeholders, thus explicitly taking into account the water-users preferences; (ii) the management policies are obtained by optimal control techniques, linking stakeholder expectations and decision-making; (iii) the multi-objective nature of the management problem is fully preserved by simulating a set of Pareto-optimal management policies, which allows for evaluating not only variations in the indicator values but also tradeoffs among conflicting objectives. The framework is demonstrated by application to a real world case study, Lake Como basin (Italy). We show that the most conflicting water-related activities within the basin (i.e. hydropower production and agriculture) are likely to be negatively impacted by climate change. We discuss the robustness of the estimated impacts to the climate natural variability and the approximations in modeling the physical system and the socio-economic system, and perform an uncertainty analysis of several sources of uncertainty. We demonstrate that the contribution of natural climate uncertainty is rather remarkable and that, among different modelling uncertainty sources, the one from climate modeling is very significant.

1 Introduction

Climate change emerged as one of the major forces that will affect water availability in the future (Bates et al., 2008). In the last 20 years, a great research effort has been devoted to increasing our knowledge about atmospheric and ocean circulation and estimating future climatic scenarios. Unfortunately, the complexity and computational burden of circulation model do not allow for simulation at the local spatial scale where the impacts on water-related activities must be estimated. To fill the gap between global and local scale, many methods were developed to downscale General Circulation Model (GCM) and Regional Circulation Model (RCM) projections.

So far, most impact studies have focused on the hydrological response at the basin scale (e.g. Jasper et al., 2004; Bronstert et al., 2007; Groves et al., 2008; Abbaspour et al., 2009). Further evaluations on the ecosystem and human activities are qualitative and expert-based. Only recently new research efforts have been initiated to extend quantitative assessment from hydrological variables to the natural, economical and social sphere, e.g. hydropower production (Schaeffli et al., 2007; Christensen and Lettenmaier, 2007), floods, ecosystem and agriculture (Hingray et al., 2007). The purpose is to provide a transparent and reproducible evaluation of the potential impact of climate change and thus the essential knowledge base to support the planning of effective adaptation measures. It is of fundamental importance to increase public awareness, support water resource planners and promote stakeholders' participation in decision-making process (Wood et al., 1997). The need for increasing stakeholder participation in this type of analysis is well-recognized, for instance by the European Environmental Agency which claims that, "until now no reports on the impacts of climate change on the water resources of the European Alps have included specific stakeholder-oriented



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information on strategies to adapt to these impacts” (EEA, 2009, p.18, Sect. 1.2).

Quantitative assessment of climate change impacts on water-related activities, both in the biological and human sphere, is very complex, for several reasons.

First, if the analysis must account for the true expectations and needs of the water users, defining quantitative indicators requires a long and complex process of knowledge elicitation from experts and stakeholders’ representatives (Soncini-Sessa et al., 2007). This is not always straightforward, especially when not strictly economic issues are concerned.

Second, the system management must be modeled. Some authors (Schaeffli et al., 2007) try to reproduce the historical management by inferring it from historical time series; others (Ajami et al., 2008) propose and test different management strategies. The former approach is questionable because the system management is likely to change following changed meteo-hydrological conditions; the latter does not guarantee that the best adaptation policy has been considered, confounding the effect of climate change with that of using a sub-optimal policy.

Finally, uncertainty deeply affect the impacts quantification. The evolution of socio-economic drivers, e.g. population growth and economic and technological development, cannot be exactly predicted. For given driver scenario, the response of the climate and water system is estimated by simulation models that inevitably exhibit structural and parameter errors. All these uncertainties are propagated and possibly amplified in the modeling chain from the global climate to the impact assessment (Schaeffli et al., 2007). Uncertainty analysis must therefore be an integral part of any impact study.

Since taking into account all the uncertainty sources simultaneously requires a huge computational effort, impact studies usually analyse only the most relevant sources at the temporal and spatial scale of interest. For instance, Arnell (2004) assesses the hydrological implications of climate change using several consistent climate and socio-economic scenarios. Brekke et al. (2009) analyse projections from 17 different GCMs, while Lopez et al. (2009) use an ensemble of projections of the same GCM under different parametrizations or *perturbed physics ensembles*. Dèquè et al. (2007) compare the projection of many different RCMs on the European domain, while Bronstert et al. (2007) compare three different downscaling methods to estimate the long-term water availability, drought conditions and floods. Ajami et al. (2008) analyse uncertainty rising from different hydrological model structures and parametrizations.

Less attention is usually devoted to assessing the uncertainty due to the inner variability of climate or *multi-decadal variability* (Arnell, 2003), which limits the statistical significance of any impact quantification based on finite time series of climatic variables, either observed or obtained by model simulation. Even if relevant, this aspect is disregarded by

many authors, possibly bringing to misleading impact assessments.

The purpose of this paper is to present a framework for the quantitative assessment of the climate change impacts on water-related activities and the associated uncertainty analysis. The approach is demonstrated by application to Lake Como basin, Italy, a complex water system in the Southern Alpine region. Briefly, it is composed of an irrigation-fed agricultural district downstream of the lake, which is one of the largest irrigated area in Europe, and of a hydropower reservoir network located in the lake catchment, which provides nearly 25 % of the national hydropower production. Other interests in play are preventing floods on the lake shores and preserving ecosystems both in the lake and along the river.

The novelties of our approach are the following: (1) the quantification of the impacts is based on a set of performance indicators defined together with the stakeholders representatives, thus explicitly taking into account the water users preferences; (2) the multi-objective nature of the management problem is fully preserved by simulating a set of Pareto-optimal management policies under different climatic scenarios, which allows for evaluating not only variations in the indicator values but also tradeoffs among conflicting objectives; (3) uncertainty analysis results in deriving confidence bounds around the simulated Pareto frontiers.

The paper is organized as follows. In the next section, we will introduce the case study area, Lake Como basin. Then, we will discuss the general framework proposed for the quantitative assessment of climate change impacts and present application results. It will be shown that under the emission scenario A2, which determines the highest global temperature increase within the scenarios proposed by the Intergovernmental Panel on Climate Change (IPCC, 2000), climate change is likely to have a significant, negative impact both on agriculture and hydropower production, due to the temperature raise and the shift in precipitation pattern. We will highlight how the natural variability of the climate and the necessity to use finite time series (of observed data or climate simulations) affect the estimated impacts. The robustness of the result will be further discussed in the following section where we present an uncertainty analysis with respect to the modeling assumptions in the description of the physical and socio-economic system. Coherently with other studies (e.g. Hingray et al., 2007), we demonstrate that one major source of uncertainty lays in the climatic model. Nonetheless the indication that climate change will have negative impacts on water use clearly emerges. Finally, we show that there exists room for improvement by adapting the system management to the new hydrological conditions, however the gap between current water demand and future water supply remains large, meaning that structural interventions (e.g. change of the crop or measures for increasing irrigation efficiency) will be required.

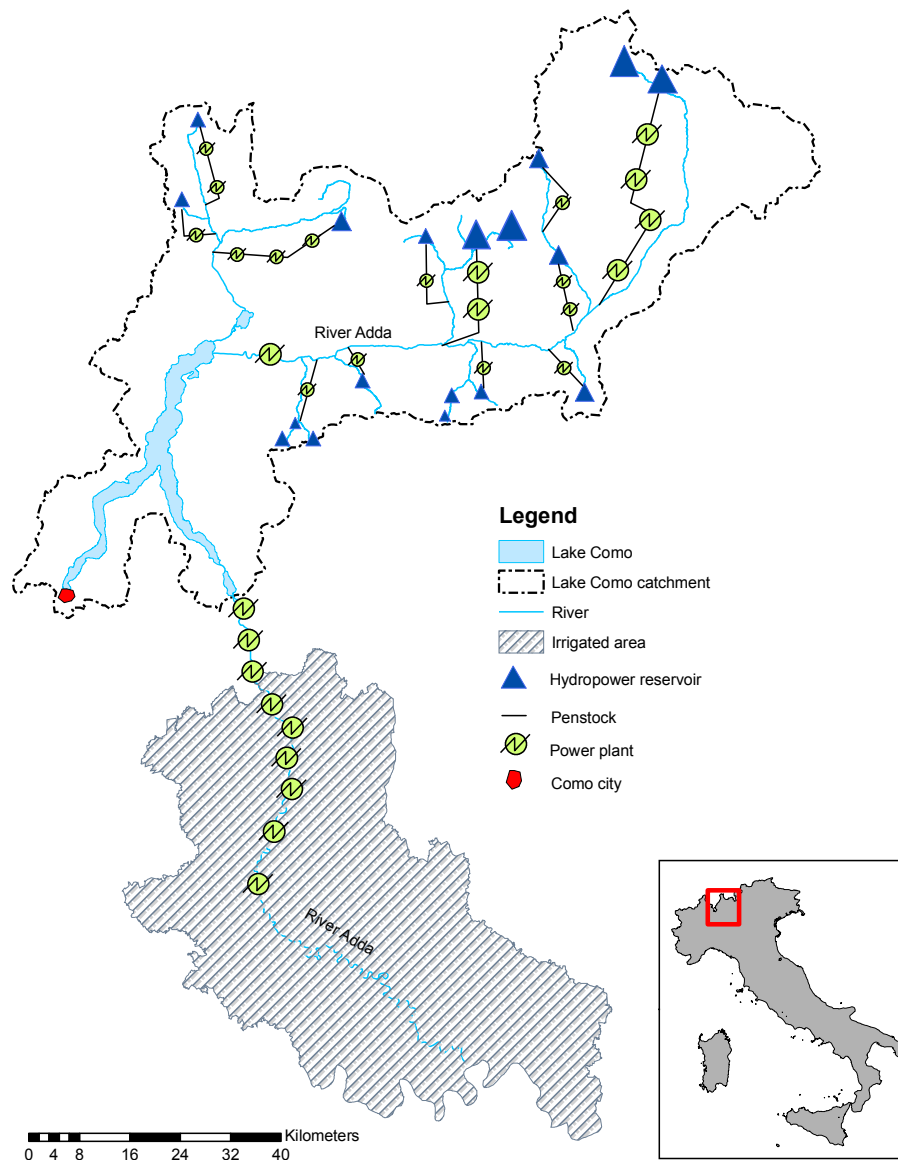


Fig. 1. Lake Como water system.

2 Lake Como case study

Lake Como is a regulated lake in Northern Italy (Fig. 1). Its operative storage is about 254 Mm^3 and it is fed by a catchment of about 4550 km^2 . The catchment is characterized by the typical Alpine hydrological regime with low discharge in winter and summer and high in late spring and autumn. The current regulation of the lake aims at attenuating flooding along the lake shores, especially in Como city, and to supply downstream users (5 irrigation districts and 9 run-of-river power plants) through a wide network of canals. The lake catchment area is covered by a dense network of smaller artificial lakes operated for hydropower production. However the overall storage of hydropower reservoirs is of 510 Mm^3 , more than twice the storage of the lake (OLL, 2005).

The management of lake Como has been intensively studied since from the first study by Guariso et al. (1986). Recently, the problem has been reframed in a wider perspective, including the management of both the lake and the hydropower reservoirs (Anghileri et al., 2011). However, despite its environmental, social and economic relevance and its potentially low resilience to climate change, few studies have addressed the problem of quantification of climate change impact on the water-related activities. Climate change may impact this complex water system in multiple ways. The average inflow to the reservoirs is expected to decrease, due to reduced snow melt in late spring and increased evaporation in summer. The subsequent reduced water availability may lead to water stress situations, possibly accrued

by increased water demand from the downstream irrigated areas due to the temperature raise.

In our analysis we will focus on two sectors: hydropower production in the upstream reservoirs and irrigation in the downstream areas. Presently, they are the main conflicting interest in the region: hydropower producers traditionally schedule their production in winter time and thus they store water in spring and summer when, on the contrary, the irrigation demand is at top. The conflict is worsened in case of water scarcity, as happened in the summer droughts of 2003 and 2005, two of the most severe droughts in Europe in the last decade, so that climate change is expected to aggravate the situation. Protection from flooding in Como city will not be considered since a defence system consisting of mobile gates is currently under construction and thus the lake management will soon become irrelevant to this purpose. Hydropower production in run-of-river plants downstream from the lake will also be neglected because its contribution is very limited.

From the modeling point of view, the Alpine hydropower reservoirs are described by one equivalent reservoir whose capacity is the sum of the total capacity of the actual reservoirs. The simplification is acceptable because the reservoir storages are strongly correlated with each other. The different water users downstream of the lake are lumped into one equivalent downstream user, whose benefit/cost is related to the total water amount that is diverted from the lake effluent (river Adda) to the main irrigation canals. The modeling time step is one day, that is the decision time step currently adopted by the lake manager. For more details on the model of the water system (catchments, reservoirs and river network) see Anghileri et al. (2011).

3 Assessment of the climate change impacts

Traditional approaches to climate change impact assessment at the basin scale rely on a modeling chain that usually includes the generation of future emission scenarios, the simulation of GCM to build global climate scenarios, the use of RCM and statistical downscaling to estimate climate scenarios at the basin scale, and the projection of climatic scenarios into discharge scenarios via simulation of hydrological models. The modeling chain often stops here, while further evaluation of hydrological scenarios is committed to experts.

In this paper we extend quantitative assessment also to impacts on water-related activities like agriculture and hydropower generation. To this end, the modeling chain must be extended to include simulation of the water system management and evaluation of the impacts by means of performance indicators (Fig. 2). Both tasks are not trivial since they require a deep knowledge of the system functioning in all its aspects, from engineering to social and economic issues.

The definition of performance indicators is a challenging task, especially when not strictly economic issues are

concerned, e.g. impact of changed hydrological regime on the riparian ecosystems, or when the relation between water availability and economic outcome is complicated. For instance in the irrigation district downstream from Lake Como a reduction in the water supply from the canals can be partially compensated by pumping from groundwater, which saves the crop but is costly. Definition and validation of the indicators used in this study was performed by interacting with stakeholder representatives and deriving a set of criteria that reflects their judgments and expectations (Castelletti et al., 2007).

Simulating the system management is an issue because it requires modeling the behaviour of the managers of the reservoirs and distribution network. In this study, we formulate the decision-making problem faced by the human regulators as an optimal control problem, and use multi-objective optimization techniques to derive Pareto-optimal management policies (see right side of Fig. 2), thus obtaining an upper bound of system performances that may be achieved by a fully rational decision-maker (Soncini-Sessa et al., 2007). To link stakeholder expectations and decision-making process, we use the performance indicators defined by the stakeholder representatives as the objectives of the optimal control problem. Since the problem is a multi-objective one, the solution is not a unique optimal management policy but a set of Pareto-optimal policies, each providing a different trade-off between the conflicting objectives. Choosing one policy within this set is not a technical task but a political one, requiring subjective weighting of the objectives, and as such it must be left to stakeholders and decision-makers. Therefore our analysis will be conducted by considering the entire collection of Pareto-optimal policies.

In the next paragraph we will describe the modeling units developed for assessing the impact of climate change in the case study area of Lake Como basin. For the reader convenience, we define some of the terms that will be used in the following:

- *historical climate*: the time series of precipitation and temperature observed in the catchment (gauge records from 1967 to 1980)
- *historical inflow*: the time series of observed discharge from the catchment, flowing into the reservoirs (gauge records from 1967 to 1984)
- *historical inflow scenario*: the time series of simulated discharge obtained by feeding the catchment model with historical climate
- *backcast (forecast) climate scenario*: the time series of simulated precipitation and temperature provided by a circulation model over the backcast (forecast) period 1961–1990 (2071–2100)

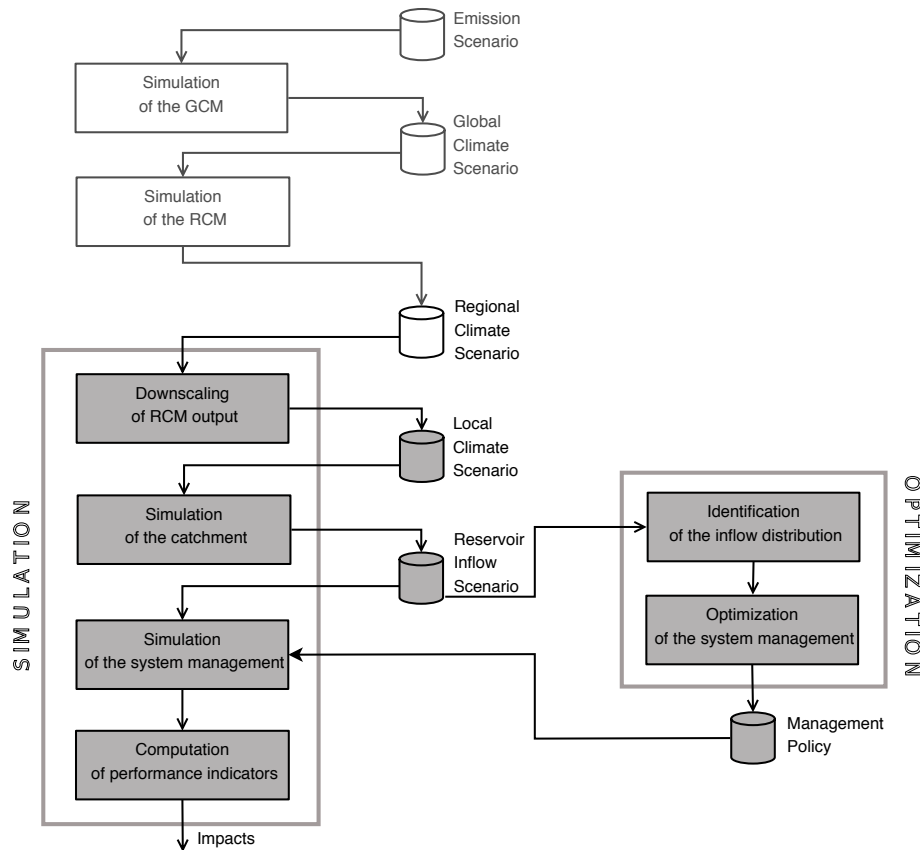


Fig. 2. The procedure for quantitative assessment of climate change impacts on water-related activities: simulation tools on the left side, optimization tools on the right side.

- *backcast (forecast) inflow scenario*: the time series of simulated discharge obtained by feeding the catchment model with the backcast (forecast) climate scenario.

3.1 Downscaling procedure

The climate of the Alps is strongly influenced by local phenomena (orographic forcing, rain-shadowing, etc.). In such cases, RCMs provide more realistic climatic forecast at the regional scale with respect to GCMs, since the mismatch of scale between the resolution of the climate models and the scale of interest for regional impacts is lower (Mearns et al., 2003; Fowler et al., 2007; Frei et al., 2006). The climatic time series considered in this study were derived as part of a larger multimodel ensemble in the framework of the European project PRUDENCE (see <http://prudence.dmi.dk/> and Christensen and Christensen, 2007). As backcast and forecast climate scenarios we considered the daily precipitation and mean temperature time series over the backcast period 1961–1990 and the forecast one 2071–2100 respectively. Each scenario was simulated using the emission scenario A2 (IPCC, 2000) and the GCM HadAM3H (Pope et al., 2000) as driving data.

Even if RCMs provide good estimate of the climate at the regional scale, some biases from the local climate of interest may still exist. In this study, RCMs' output were corrected via the statistical downscaling method known as Quantile Mapping. For a given variable, the cumulative density function (cdf) of the backcast is first matched with the cdf of the observations, thus generating a correction function depending on the quantile. The correction function is then used to unbias the variable from the forecast quantile by quantile. This method has been used in many hydrological impact studies, using a correction function at either annual or seasonal level (Dèquè, 2007; Boè et al., 2007).

One major limitation of statistical downscaling is that the goodness of the correction strongly depends on the quality of the available observations. To mitigate such effect, the backcast period was split into two sub-periods that were used for calibration and validation respectively. Both an annual and seasonal correction function were derived over the calibration period for both temperature and precipitation, and the one producing the smaller mismatch between down-scaled and observed data over the validation sub-period was adopted. This is an annual correction function for the precipitation time series, and a seasonal correction function for the temperature time series.

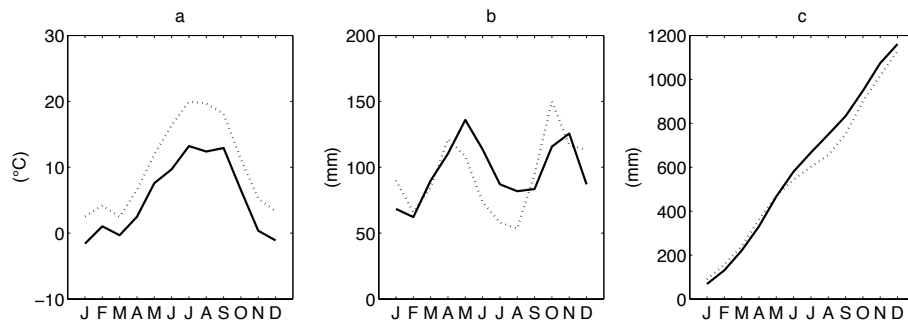


Fig. 3. Mean monthly temperature in the backcast (solid) and forecast (dotted) scenario (a); total monthly precipitation (b); and cumulate precipitation over the year (c) with downscaled RACMO RCM.

Figure 3 compares some statistics of the downscaled output of the RACMO RCM (Lenderink et al., 2003) over the backcast and forecast period. The forecast climate scenario shows an increase in monthly mean temperature (of about 4 °C) and a shift in the precipitation pattern (decrease in spring and summer and increase in autumn and winter) while the annual precipitation volume is only slightly lower than in the backcast scenario.

3.2 Catchment model

The catchment response to climatic input is simulated through a lumped, conceptual model, partially based on the HBV model (Bergstrom, 1976). The lumped modeling approach guarantees efficient parameterization even with limited historical time series. However spatial processes are neglected. In our case study spatial heterogeneity is not significant, but for elevation. Nonetheless comparison of our proposed model and an elevation-based model (Conorzio dell’Adda, 1986) shows that lumping does not induce significant loss of information.

Our model is composed of three modeling units. First, the precipitation input is splitted into snowfall and rainfall: average daily temperature in a reference station is used to determine the freezing level and snowfall is computed as a fraction of the total precipitation, through a proportionality coefficient that accounts for the catchment’s area located above the freezing level. Then, the snowpack dynamics is described by a mass balance equation, while a degree-day approach is used to determine the snowmelt. Finally, the HBV model is used to simulate the soil water balance and subsequent runoff, as a consequence of melt-water, rainfall, and evapotranspiration. The latter is computed through the Blaney-Criddle method (Brouwer and Heibloem, 1986).

Two different parametrizations were used for the two catchment, the one feeding the equivalent hydropower reservoir (catchment surface area of about 350 km²) and the other feeding Lake Como (4200 km²). They were derived using the Genetic Algorithm implemented in the Matlab Global Optimization Toolbox and time series of daily precipitation,

Table 1. Parameters of the optimally calibrated HBV model for the lake Como (LC) catchment and the hydropower reservoir (HR) catchment, and relevant performance indicators over the validation dataset (1977–1984).

model parameters		LC	HR
FC	maximum soil moisture content (mm)	251.9	238.5
LP	limit for potential evapotranspiration	1	0.9
ALFA	response box parameter	0.04	0.02
BETA	exponential parameter in soil routine	0.14	1.06
K	recession coefficient for upper tank (day ⁻¹)	0.29	0.11
K4	recession coefficient for lower tank (day ⁻¹)	0.04	0.04
PERC	maximum flow from upper to lower tank (mm day ⁻¹)	6.98	0
CFLUX	maximum value of capillarity flow (mm day ⁻¹)	0	0
MAXBAS	transfer function parameter (day)	1.01	1.21
performance indicators			
R ²	coefficient of determination (–)	0.654	0.799
MAE	mean absolute error (m ³ s ⁻¹)	57.5	3.22
RVE	relative volume error (–)	0.23	0.09

temperature and flow. Precipitation is the spatial average from several meteorological stations; temperature data come from two reference stations, one for each of the two catchments; flow data are derived by inversion of the reservoir mass balance equations. The objective function of the automatic calibration procedure is the coefficient of determination (one minus the ratio between error variance and measured flow variance). Table 1 shows the optimal parameter values of the HBV soil-moisture routine for the two catchments.

The calibrated model was evaluated by means of several performance indicators computed over the validation period 1977–1984 (last lines in Table 1) and graphical tools like scatter plot, duration curves and hydrographs of observed

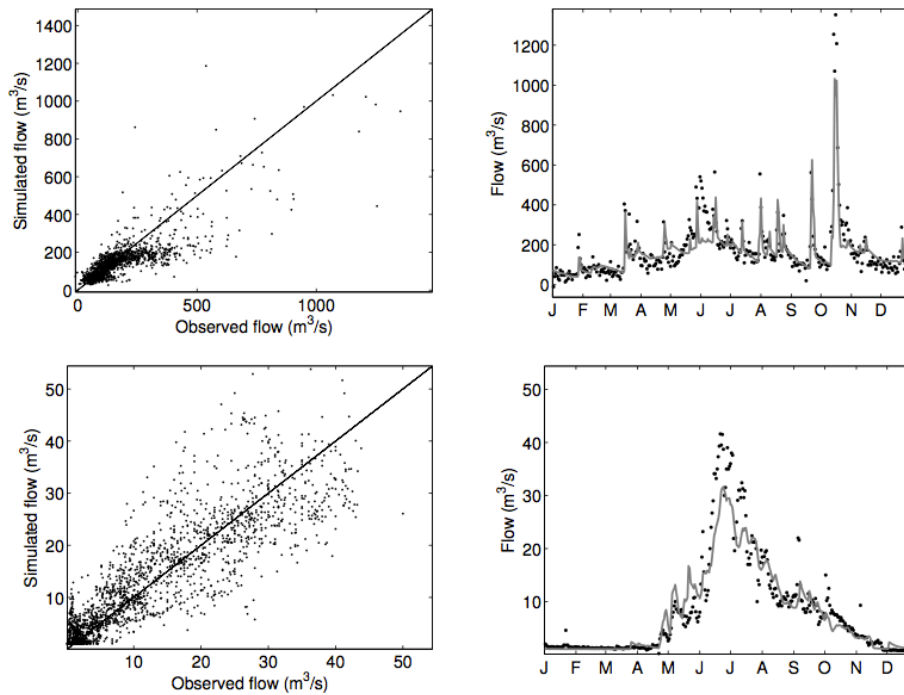


Fig. 4. Left: observed vs simulated flow from lake Como catchment (top) and hydropower reservoir catchment (bottom) in the validation period (1977–1980). Right: observed flow (dots) and simulated flow (grey line) in 1980.

and simulated flow (Fig. 4). They show that the model error is quite significant especially for high flow. Nonetheless the model accuracy is acceptable for the scope of this study, as we will show in Sect. 4.1.

Note that in this study we did not include a model of the glacier dynamics. At present, the contribution of glacier melting is usually negligible but for extremely hot and dry summer periods, as for instance the 2003 drought. However, under future climate scenario of increased temperature, glacier melting may become relevant. Also, there exist multiple evidences of a constant glacier reduction since from the beginning of the 20th century (Smiraglia and Diolaiuti, 2006), which means that glacier melting may give a positive contribution to flow in the middle-term while disappearing in the long run. However, such an evolution cannot be reproduced in our study.

3.3 Reservoir and management model

The water system reservoir network is modeled by two reservoir in cascade: the equivalent hydropower reservoir and Lake Como. Each reservoir is described by the mass balance equation

$$s_{t+1}^i = s_t^i + q_{t+1}^i - R^i(s_t^i, u_t^i, q_{t+1}^i) \quad (1)$$

where s_t^i is the storage at time t of the i -th reservoir, q_{t+1}^i is the inflow from its catchment, u_t^i is the release decision

taken at time t and $R^i(s_t^i, u_t^i, q_{t+1}^i)$ is the release that actually occurred in the interval $[t, t+1)$, which may differ from u_t^i because of unintentional spill or other physical or legal constraints (Soncini-Sessa et al., 2007).

To simulate the system management, we need a model to compute the decisions u_t^i . At present, the regulation of Lake Como is committed to a water board composed of representatives of Como municipality and downstream farmers, while hydropower reservoirs are operated by different power companies. The latter aim at maximizing releases when the energy price is higher, while the lake operation aims mainly at supplying downstream users (especially irrigation), but also considering other concerns like flood control, recreational activities and ecosystem conservation. Managers act independently, with different purposes and considering different information: each manager makes his/her decision u_t^i in function of the storage s_t^i of the corresponding reservoir and considering his/her own purpose (either hydropower or irrigation supply). Recently, Anghileri et al. (2011) showed that the system performances could be largely increased by adopting an integrated approach where information and objectives are shared. This means that each decision u_t^i is made based on both storages s_t^1 and s_t^2 . Further, the function that maps storages into decisions is designed by solving a two-objective (irrigation and hydropower) optimal control problem. Anghileri et al. (2011) use Stochastic Dynamic Programming to solve the problem and demonstrate the room for improvement given by the integrated approach. In this

paper, we will use the set of management policies reported therein as the reference to evaluate how things may change under climate change. It consists of eight different policies, each corresponding to a different tradeoff of the two objectives, including the two extreme policies that consider either irrigation or hydropower only.

3.4 Performance indicators

The definition of indicators was developed together with the stakeholder representatives in a former research project (Castelletti et al., 2007). In that project, a representative person for each stakeholder group was identified. These are: the managers of the hydropower companies; the leaders of the irrigation consortia (representative for the farmers); officials from Como city and other towns along the lake shores (representative for the flooding, navigation, fishing and tourism issues); the manager of the Nature Park located along the lake effluent river. The indicators were identified by structured interviews and validated during a final meeting, where each stakeholder representative was asked to rank different situations and it was checked that the ranking was consistent with the one determined by the indicator.

In this study we focus only on the hydropower and irrigation indicators: a brief description of the two follows, while a detailed definition is given in Anghileri et al. (2011).

The hydropower indicator is the average daily revenue from hydropower production

$$J^{\text{hyd}} = \frac{1}{h} \sum_{t=0}^{h-1} \sum_{j=0}^{n_{t+1}} \theta_{t,j} G \quad (2)$$

where h is the length of the simulation horizon, used to evaluate the system performances; n_{t+1} is the number of working hours of the plant on day t (equal to the release r_{t+1}^1 from the equivalent hydropower reservoir divided by the power plant capacity); G is the energy produced per hour, when the plant is working at full capacity; and $\theta_{t,j}$ is the energy price in the j -th most profitable hour of day t . The energy price is a periodic parameter. Its weekly and annual pattern is estimated from time series of energy price over the period 2005–2006 provided by the national energy authority (see www.mercatoelettrico.org).

The irrigation indicator is the squared daily deficit in the water supply

$$J^{\text{irr}} = \frac{1}{h} \sum_{t=0}^{h-1} \left[\max \left(W_t - r_{t+1}^2, 0 \right) \right]^2 \quad (3)$$

where r_{t+1}^2 is the release from Lake Como in the time interval $[t, t + 1)$ and W_t is the water demand for irrigation on day t . The water demand is a periodic parameter. Its annual pattern is estimated combining the water requirement declared in the abstraction licenses and the historical time series of diverted flows. Power 2 in Eq. (3) is a means to implicitly select

management policies that reduce high-percentage deficit in a single time step while allowing for more frequent small shortages, which cause less damage to the crop.

3.5 Impact assessment

The performance indicators were used when designing the optimal management policies. In fact, the objective functions of the stochastic optimal control problem are the expected values of Eqs. (2) and (3) with respect to all the possible trajectories of the inflows (i.e. the inflow probability distribution) over an infinite horizon ($h \rightarrow \infty$). Note that, since the inflow probability distribution is estimated over historical time series, the result is optimal as long as the hydrological behaviour of the system remains stationary. For each Pareto-optimal policy reported in Anghileri et al. (2011), the expected values of the indicators can be assessed by Markov or Monte Carlo simulation. Alternatively, it is possible to use deterministic simulation and compute the indicator values over a single finite horizon. The latter approach is computationally less demanding and can provide a more informative output to stakeholders: for instance, using a historical horizon they can compare the simulated behaviour of the system with the historical one, which they directly experimented. The performance indicators under the historical inflow over the period 1967–1984 are shown in Fig. 5 (black dots). Note that even if produced by Pareto-optimal policies, they do not necessarily belong to the Pareto Frontier of the two-objective control problem, as they are obtained under historical inflow and not under the inflow probability distribution used in optimization. For this reason they will be called the Image of the Pareto Frontier (IPF). It can be noticed that the historical IPF (black dots in Fig. 5) can greatly improve the satisfaction of both the water users with respect to the historical management (cross in Fig. 5) and represent an effective tool to mitigate the conflict between upstream and downstream water users.

The historical IPF also constitutes a reference for comparison with system performances under climate change: Fig. 5 shows also the IPF under the forecast inflow scenario (2071–2100), as given by RACMO RCM and projected through our simulation procedure (magenta triangles). For all the policies, the system performances worsen with respect to both objectives, and particularly irrigation. In fact, the forecast climate scenario predicts a significant reduction of water availability just in late spring and summer, when the water demand for irrigation is higher, and the active storage of the lake is not sufficient to store the required volumes in anticipation of the dry period. The results is that failures in the water supply become more frequent. On the other hand, hydropower revenue is mainly sensitive to the total volume of available water, which is only slightly reduced in the forecast scenario, because the ratio of reservoir capacity to mean annual inflow is quite high (60 %, against 6 % for Lake Como).

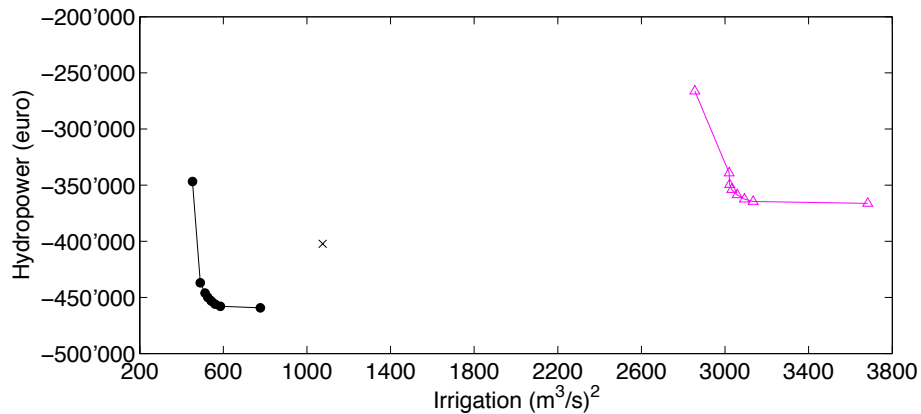


Fig. 5. Image of the Pareto Frontier (IPF) under historical inflow 1967–1980 (black dots) and forecast inflow scenario 2071–2100 by RACMO RCM (magenta triangles). The cross is the historical management. Hydropower revenue (on the vertical axis) is changed in sign.

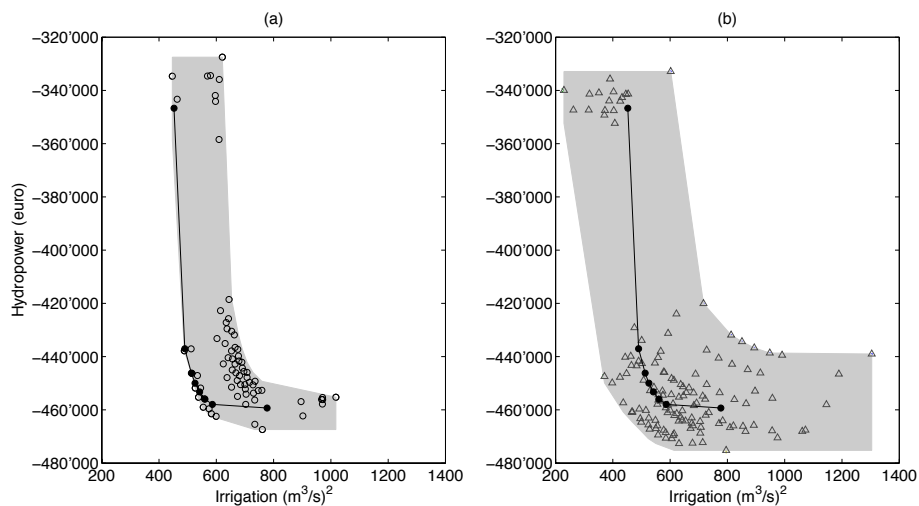


Fig. 6. Left panel: IPFs under historical inflow over a sliding window of 10 years between 1967 and 1984. Black dots are the IPF under historical inflow over the entire horizon 1967–1980. Right panel: IPFs under backcast inflow scenario over a sliding window of 14 years between 1961 and 1990. Black dots are the IPF under historical inflow scenario over the entire horizon 1967–1980.

3.6 Validation of the assessment procedure

The trustability of the results presented in the previous section depends on the robustness of the adopted simulation procedure, which is affected by two major sources of uncertainty. First, the comparison is based on indicator values computed over finite system trajectories. Would results be significantly different under a different choice of the simulation horizon? Second, as in any impact assessment analysis we shall consider what is the contribution of modeling errors. The latter problem will be discussed in the next section, here we will focus on the problem of using finite simulation horizon, which is an intrinsic issue of the assessment procedure independently of the modeling error issue.

To assess the uncertainty in the indicator values due to the choice of the simulation horizon, we computed seven

different IPFs with a sliding window of $h = 10$ years over the period 1967–1984 (grey dots in Fig. 6a). It can be seen that differences are generally small, exception made for two IPFs, which present a strongly lower irrigation cost: they correspond to simulation horizon that do not include the year 1973, characterized by one of the most severe droughts of the 20th century. The estimated indicator values are indeed sensitive to single extreme events occurring or not occurring in the selected horizon. The length of the horizon also affects the results. The historical IPF (black dots), although including the dry year 1973, shows lower irrigation costs because the same events are averaged over a longer simulation horizon (14 years instead of 10).

The same problem arises when using climate scenarios. Indeed, the problem is accrued because, due to the chaotic nature of the climate models, time series of simulated

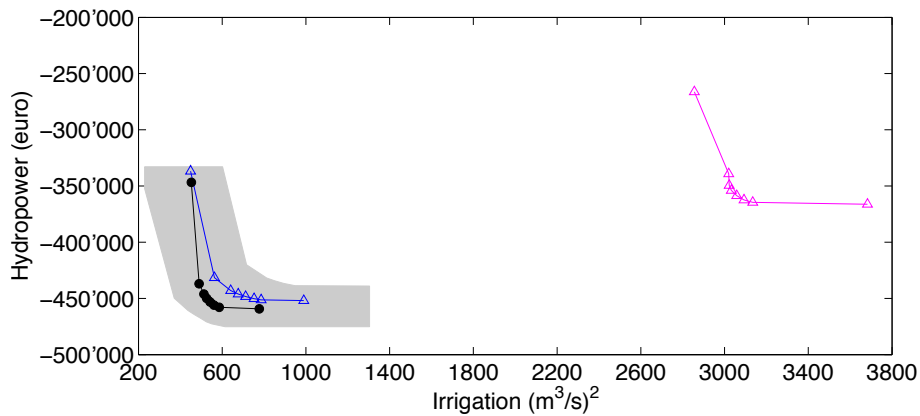


Fig. 7. IPF under historical inflow 1967–1980 (black dots), backcast inflow scenario 1961–1990 by RACMO RCM (blue triangles) and forecast inflow scenario 2071–2100 by RACMO RCM (magenta triangles). The grey region represents the natural variability of backcast climate scenario obtained as the envelope of the IPFs over a sliding window of 14 years reported in Fig. 6b.

precipitation and temperature, then projected into flows, can only be interpreted as equiprobable to observations (Royer, 2000). It follows that, even assuming that the RCM perfectly reproduced the climate dynamics (i.e. even neglecting the modeling error issue), we could not expect its output time series to perfectly overlap historical observations. Indeed, the observed climate over 1967–1980 is simply equiprobable to any 14-years long time series in the backcast period. To assess the uncertainty in the indicator values due to such statistical equiprobability, we computed several IPFs under the backcast scenario with a sliding window of $h = 14$ years. They are shown in Fig. 6b: as expected, none of these IPFs is superimposed to the historical one, but they are scattered around it. Finally, the IPF under the entire backcast scenario of 30 years from 1961 to 1990 can be computed: it is represented by the triangles in Fig. 7. This IPF may be used as a fair reference for comparison with the IPF under forecast scenario, in place of the historical IPF (1967–1984), since it is based on the same simulation model and horizon length h as the forecast IPF.

Although the use of a finite horizon and the statistical interpretation of the RCM output do not allow for a univocal quantification of the system performances over the past, this intrinsic variability is negligible with respect to the variation that is expected to be induced by climate change, as shown in Fig. 7. This is consistent with other research: for instance, Arnell (2003) demonstrates that changes in mean seasonal discharge in many basins in Britain are outside the range of natural climate variability by 2050s, but that climate change signal and natural variability could be difficult to distinguish when considering nearer horizons.

4 Uncertainty analysis

In the previous paragraph, we showed how the natural variability of the climate and the impossibility to use infinite time series (of observation or climate simulation) affect the robustness of the estimated impacts. Beside this intrinsic uncertainty in impact assessment, another source of uncertainty lays in the tools used to implement it, that is, the chain of simulation models shown in Fig. 2. We will distinguish two types of uncertainties: those introduced in modeling the physical system and those introduced in modeling the socio-economic system.

4.1 Uncertainty in modeling the physical system

The description of the physical system includes modeling the climate dynamics through the GCM, RCM and statistical downscaling; modeling the catchment response; and modeling the reservoirs.

Structural uncertainty is particularly high in the climate and hydrological modeling. For GCMs and RCMs, uncertainty rises from limited understanding of the processes occurring in the atmosphere, ocean, cryosphere, etc. ...; from the mismatch in scale between the grid resolution of the RCM and the catchment boundaries; and from error induced by using a coarse spatial resolution. Downscaling is not sufficient to restore all the characteristics of the climate time series observed at the basin scale: for instance, the Quantile Method used in this study cannot correct the temporal properties of the precipitation series (e.g. length of dry spells). For the catchment model, structural error is also significant because of the oversimplified description of the actual processes occurring in the basin and the lumping of all spacial processes into one average process. Structural error is much smaller in the reservoir models, which indeed are very accurate and

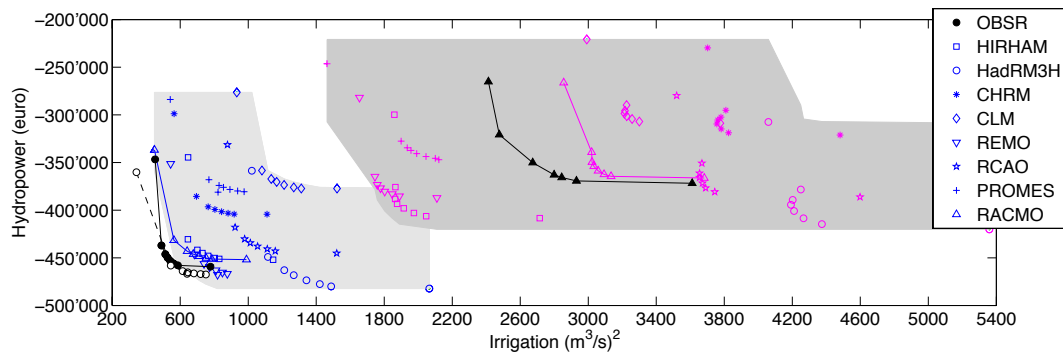


Fig. 8. IPF under historical inflow 1967–1980 (black dots), IPF under historical inflow scenario 1967–1980 (white dots), IPFs under backcast inflow scenarios (1961–1990) using eight different RCM models (blue symbols), IPFs under forecast inflow scenarios (2071–2100) with the same eight different RCMs (magenta symbols), IPF under forecast RACMO inflow scenarios (2071–2100) using optimal management policies for future climate (black triangles).

can be considered as exact at the spatial and temporal scale of interest.

Besides structural uncertainty, the simulation output is also affected by parameter uncertainty. In particular for downscaling and the catchment model, the problem is that parametrizations were selected by minimization of the simulation error over historical time series. This approach is highly questionable when the model is used for projecting future climate scenarios that are, by definition, violating the stationary assumption underlying calibration over historical time series. Unfortunately there is no solution to this paradox: the past is the only testing ground we have to assess the validity of our models.

Regardless of the distinction between structural and parameter uncertainty, the impact of the model error can be assessed, at least for the catchment model, by a simple experiment: to simulate the system under the historical inflow scenario, i.e. the discharge time series produced by the catchment model when fed by the historical climate. The corresponding IPF is shown in Fig. 8 (white dots). It can be seen that it does not perfectly overlap the historical IPF (black dots), as expected if there were no error in the catchment model. However the modeling error is rather limited compared to variability induced by the use of equiprobable climate scenarios and the full range of indicator values is reproduced.

As for the climate model, the impact of structural uncertainty may be assessed by simulating and comparing different circulation models. Generally, since the winter climate is mainly driven by global circulation while the summer climate is largely influenced by local phenomena, the choice of the GCM is the main source of uncertainty in winter time, while the RCM is more important in the summer (Jacob et al., 2007). Besides this distinction, some studies (Dèquè et al., 2007; Schaeffli et al., 2007) seem to indicate that the choice of the GCM is the most critical. However, as for the RCM scenarios generated in the PRUDENCE project

and used in this analysis, Hingray et al. (2007) show that variability among RCMs is comparable to the variability induced by the GCM choice. Following these considerations and for brevity's sake, in this paper we will focus only on the RCM variability. Starting from the climate scenario from seven different RCMs (beyond the RACMO model) provided by the PRUDENCE project, we applied the downscaling method to each of them and then projected climate input into inflow scenarios. Figure 8 shows the IPFs under these eight backcast (blue) and forecast (magenta) inflow scenarios. It can be seen that the spread of the IPFs is rather high even over the backcast scenario: the RACMO RCM, that we used so far as the reference model, produces an IPF quite close to the historical one, together with the REMO and HIRHAM, while other RCMs seem to be less accurate in reproducing the historical system performances. The spread of the IPFs strongly increases in the forecast scenario – although all future scenarios are derived from the same emission scenario, A2, and GCM boundary condition, HadAM3H.

To conclude, our study provides one more confirmation that circulation models, and specifically RCMs, are a major source of uncertainty in impact assessment studies, much more relevant than other sources like uncertainty from using finite simulation horizon or inner climate variability, as it can be seen by comparing the extent of the uncertainty regions (grey area) in Figs. 6.a, b and 8. Notwithstanding this high uncertainty, comparison of the uncertainty regions over backcast and forecast scenarios (Fig. 8) suggests that a significant worsening of the system performances can be expected: there is basically no overlapping between the backcast and forecast uncertainty region.

4.2 Uncertainty in modeling the socio-economic system

The description of the socio-economic system includes the definition of the emission scenario, the policies used to manage the reservoirs, and the definition of the performance

indicators. Uncertainty associated to these choices is rather different from uncertainty in modeling the physical system. In the latter case, uncertainty stems from our limited capacity of reproducing reality through models and, to some extent, it can be objectively quantified by comparison of model output with observations from the real system. Uncertainty in modeling the socio-economic system, instead, can rarely rely on observations or reference values. For instance, there is no exact choice of the emission scenario, and the only way to assess the impact of such choice is to repeat the entire simulation procedure under a different scenario.

The same holds for the choice of the performance indicators. Since they are aimed at reflecting the stakeholder preferences, stating whether they actually capture the stakeholder opinions is very difficult. For the hydropower producers, the choice of the revenue is rather straightforward, while for the farmers the definition of the indicator is more difficult. The proper choice would be the revenue from the crop production, however this indicator would need a model of the crop growth that is expensive to develop and often does not guarantee reliable results. The average squared deficit that we used in our analysis is a proxy indicator easy to compute and that received the approval of the farmers' representatives, and as such it is hardly questionable.

What can be argued is the value of the parameters inside the indicator formulation. So far, we implicitly assumed a business-as-usual scenario for the energy price and water demand. However, the pattern of energy price may change in the future following changed conditions in the energy market, while the water demand may be reduced thanks to improvement in the irrigation technique (e.g. from submersion to more efficient systems) or changes in the crop. Climate change itself will probably drive such changes. Therefore, the analysis so far must not be interpreted as a prediction of the future conditions, which would be unrealistic because the socio-economic system will certainly evolve and adapt to reduced water availability, but rather as the demonstration that the current socio-economic conditions cannot be maintained in the future.

The business-as-usual assumption involves also the system management. In fact, variations in the hydrological conditions, as well as potential variations in the energy price or water demand, will lead the reservoir managers to change their behaviour. In our analysis we simulated the management policies that proved Pareto-optimal over historical inflow statistics, energy price, etc. but they will not be Pareto-optimal any more if these conditions will change. Even if we set aside the issue of energy prices or water demand, still the results shown so far are overly pessimistic because based on sub-optimal policies, and there is room for improvement by re-optimizing the management policies under the new inflow scenarios. To explore this room for improvement, we ran the following experiment. We used the first years of the forecast inflow scenarios produced by the downscaled output of the RACMO RCM to re-estimate the probability

distribution of the reservoir inflows and, based on this new distribution, re-run Stochastic Dynamic Programming, thus obtaining eight Pareto-optimal policies for the new climate scenario. Then, we simulated the new policies under the entire forecasting horizon and derived the IPF represented by the black triangles in Fig. 8. It can be seen that the system performances improve with respect to the original, sub-optimal IPF (magenta triangles), especially for the irrigation objective. Nonetheless, the improvement is not sufficient to compensate for deteriorated hydro-climatic conditions, as it can be seen by comparison with the IPF under backcast inflow scenario (blue triangles). Notice that the comparison between these two IPFs is not affected by uncertainty in modeling the manager behaviour, since in both cases we assume the best possible behaviour, in Pareto-sense, that a rational decision-maker could follow for the corresponding inflow scenario and the selected performance indicators.

5 Conclusions

This paper presents a general framework for the quantitative assessment of the climate change impacts on the water-related activities at the basin scale. The proposed simulation procedure starts from the downscaling of regional circulation model output, and through the projection into the hydrological input and simulation of the system management, ends up with the computation of performance indicators. One major feature of our approach is the multi-objective perspective that is preserved throughout the entire simulation procedure. In fact, instead of simply reproducing the current system management, we first derive a set of Pareto-optimal policies and then simulate all of them over both the historical, backcast and forecast scenario. The advantage is that tradeoffs between different objectives can be explored under present and future climate conditions, and further that the comparison of past and future performances is not affected by subjective choice of the management policy.

The approach is demonstrated by application to the complex and intensively exploited system of Lake Como, Italy. It shows that climate change is expected to dramatically impact the water-related activities in the basin, both upstream of the lake, where hydropower reservoirs are located, and in the downstream irrigated areas. Re-optimization of the management policies following future hydrological conditions is not enough to compensate for such loss. The result was obtained using current energy price and water demand pattern, and thus it must not be interpreted as a prediction of the actual future conditions but rather as the demonstration of the unsustainability of the current ones. Evaluation of structural adaptation measures, for instance measures for reducing the irrigation water requirement, goes beyond the scope of this paper. However, the proposed approach may be easily used for this purpose: it will be sufficient to repeat the entire simulation procedure under a different system parameterization

characterizing the measure under study, for instance a different yearly pattern of the water demand in the definition of the irrigation performance indicator.

The results are highly affected by uncertainty. We analyzed both the uncertainty stemming from the inner variability of climate and the modeling uncertainty. The analysis proved that, although the contribution of the former is quite significant, it is negligible with respect to the latter. Also, among different sources of modeling uncertainty, the uncertainty in the climate modeling, specifically RCMs, seems to be the most significant. A reduction of this uncertainty may be expected when considering climate scenarios in the shorter term. As a result of these multiple uncertainties, the exact quantification of the impacts in terms of performance indicators is, at the state-of-the-art, not fully reliable. However the comparison of the uncertainty regions where current and future performances are expected to fall, clearly indicates that a significant loss will be induced by climate change, especially for the irrigation sector.

In the case study area, most of the stakeholders were already organized into associations with long-established procedures for selecting their delegates and resolving disputes. This was a clear advantage in developing our study. However, when stakeholders are fragmented and/or the analysis also includes issues where perceptions of individual stakeholders may be strongly different (like landscape changes or ecology), more sophisticated tools and spatial analysis should be considered to enhance true participation.

Several topics remain open for future research. For Lake Como system, the evaluation of climate change impacts should be extended to other important sectors like for instance ecosystem conservation. The system model should be improved, particularly the hydrological model of the catchment that, in its current version, does not include the glacier dynamics, and the uncertainty analysis may be further detailed considering the other sources of uncertainty (e.g. parameter uncertainty) mentioned in this paper but not fully analyzed yet.

While increasing complexity and accuracy of the simulation model will increase the trustability of the results, we question this will be sufficient to compensate for the large uncertainty that affects the assessment analysis, because of the inner variability of climate, our limited capacity in reproducing the complex circulation dynamics, and the errors induced by mismatches in scale. Therefore we think that the research effort to improve the model accuracy should be coupled with an equal effort towards developing effective methods to evaluate model uncertainty and its propagation through impacts assessment studies. This is especially true when dealing with multi-objective problems, where modelling and optimization is aimed at providing the knowledge base for political discussion and decision-making, not at replacing it. The role of uncertainty analysis in this process is very delicate. From the modeler standpoint, uncertainty analysis enhances the robustness of the assessment results,

while for political decision-makers it may be perceived as undermining their trustability. Communicating the information contained in the IPF graphs shown in this paper is difficult and time consuming: it requires the decision-makers to make an effort towards understanding at least the general principles of the underlying assessment methodology; and the willingness to assimilate a sophisticated message rather than simple answers. Effective communication of modelling results and their associated uncertainty should become integral part of the research in this area.

Acknowledgements. The authors are grateful to the anonymous reviewers for their comments that helped to improve the paper. The research presented in this paper was partially supported by *Fondazione Lombardia per l'Ambiente* (www.flanet.org).

Edited by: D. Solomatine

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